# Impacts of additional HONO sources on $O_3$ and $PM_{2.5}$ chemical coupling and control strategies in the Beijing–Tianjin–Hebei region of China

By YING LI<sup>1†</sup>, JUNLING AN<sup>1\*</sup>, MIZUO KAJINO<sup>2,3</sup>, ISMAIL GULTEPE<sup>4</sup>, YONG CHEN<sup>1</sup>, TAO SONG<sup>1</sup> and JINYUAN XIN<sup>1</sup>, <sup>1</sup>State Key Laboratory of Atmospheric Boundary Layer Physics and Atmospheric Chemistry (LAPC), Institute of Atmospheric Physics, Chinese Academy of Sciences, Beijing, China; <sup>2</sup>Meteorological Research Institute, 1-1 Nagamine, Tsukuba, Japan; <sup>3</sup>RIKEN Advanced Institute for Computational Science, 7-1-26 Minatojima-minamimachi, Chuo-ku, Kobe, Hyogo, Japan; <sup>4</sup>Cloud Physics and Severe Weather Research Section, Environment Canada, Toronto, ON, Canada

(Manuscript received 27 January 2014; in final form 1 February 2015)

#### ABSTRACT

The objective of this work is to examine the impacts of additional HONO sources on the chemical interaction between ozone (O<sub>3</sub>) and particulate matter with a diameter  $\leq 2.5 \,\mu m$  (PM<sub>2.5</sub>). Three additional HONO sources, i.e. HONO emissions, the reaction of photo-excited nitrogen dioxide (NO<sub>2</sub>) with water vapour (H<sub>2</sub>O), and NO<sub>2</sub> heterogeneous reaction on aerosol surfaces, were inserted into the fully coupled Weather Research and Forecasting-Chemistry model to evaluate O<sub>3</sub> and PM<sub>2.5</sub> concentration enhancements in the Beijing-Tianjin-Hebei (BTH) region during August 2007. Results show that the additional HONO sources significantly increase  $O_3$  and  $PM_{2.5}$  concentrations during daytime. Up to 9 ppb enhancements of  $O_3$  and 32  $\mu$ g m<sup>-3</sup> increases in  $PM_{2.5}$ are found at seven urban sites over the BTH region. O3 increases are closely connected to PM2.5 increases over urban areas during daytime when the additional HONO sources are taken into account. PM<sub>2.5</sub> inorganic components of  $SO_4^{2-}$ ,  $NO_3^{-}$  and  $NH_4^{+}$  are increased by 5–18, 10–58 and 10–40%, respectively, over urban areas during daytime. The simultaneous increment of O<sub>3</sub> and PM<sub>2,5</sub> during daytime due to the additional HONO sources is related to the increasing oxidants (OH, H<sub>2</sub>O<sub>2</sub> and O<sub>3</sub>) that enhance the atmospheric oxidising capacity. The concentration variations of O<sub>3</sub> and PM<sub>2.5</sub> under a variety of NO<sub>x</sub>, volatile organic compound and ammonia (NH<sub>3</sub>) emission control scenarios show that the additional HONO sources increase the sensitivity of O<sub>3</sub> and PM<sub>2.5</sub> concentrations to the changes of NO<sub>x</sub> emissions. An increase of the PM<sub>2.5</sub> sensitivity to changes in NH<sub>3</sub> emissions is also found. This indicates that without considering the additional HONO sources, the effectiveness of emission control strategies in reducing O<sub>3</sub> and PM<sub>2.5</sub> concentrations would be significantly underestimated.

Keywords: HONO sources, O3 and PM2.5 chemical coupling, control strategy, WRF-Chem model

To access the supplementary material to this article, please see Supplementary files under 'Article Tools'.

## 1. Introduction

Nitrous acid (HONO) plays an important role in enhancing O<sub>3</sub> and particulate matter (PM) concentrations in the polluted atmospheric boundary layer due to its photolysis

into the hydroxyl radical (OH), which is a primary daytime oxidant (Alicke et al., 2002; Kleffmann et al., 2005). Jenkin et al. (2008) showed that increasing HONO emissions by 5% resulted in  $O_3$  enhancements of 11 ppb in the southern UK during a heat-wave period of August 2003. Besides direct emissions, HONO can also be formed through gasphase or heterogeneous reactions. The homogeneous gasphase reaction between excited nitrogen dioxide ( $NO_2^*$ ) and  $H_2O$  ( $NO_2^* + H_2O \rightarrow HONO + OH$ , hereafter referred to as  $NO_2^*$  chemistry) has gained interest since its reaction rate

\*Corresponding author. email: anjl@mail.iap.ac.cn

Responsible Editor: Kaarle Hämeri, University of Helsinki, Finland. <sup>†</sup>Now at: Multiphase Chemistry Department, Max Planck Institute for Chemistry, Mainz 55128, Germany.

Tellus B 2015. © 2015 Y. Li et al. This is an Open Access article distributed under the terms of the Creative Commons Attribution 4.0 International License (http://creativecommons.org/licenses/by/4.0/), allowing third parties to copy and redistribute the material in any medium or format and to remix, transform, and build upon the material for any purpose, even commercially, provided the original work is properly cited and states its license.

has been found at an order of magnitude faster than the previously estimated upper limit (Crowley and Carl, 1997), thus can substantially enhance OH production (Li et al., 2008). However, the atmospheric significance of this new mechanism is highly controversial. Carr et al. (2009) argued that the NO<sub>2</sub> chemistry would not be important under normal atmospheric conditions and they estimated the reaction rate similar to the result of Crowley and Carl (1997)  $(1.2 \times 10^{-14} \text{ cm}^3 \text{ molecule}^{-1} \text{ s}^{-1})$ . Amedro et al. (2011) claimed that no OH radicals were observed using an unfocused excitation laser beam at 565 nm, but OH formation was observed by focusing the laser beam. Recently, Wu and Chen (2012) concluded that the reaction of two photon absorbed NO<sub>2</sub> with H<sub>2</sub>O makes negligible contributions to the formation of OH radicals while single photon absorption at < 554 nm is a possible process.

The uncertainties in the NO<sub>2</sub> chemistry did not prevent the evaluation of its potential impact on tropospheric air quality using chemical transport models. Simulations conducted in the South Coast Air Basin of California during the summer of 1987 showed that the NO<sub>2</sub>\* chemistry increased  $O_3$  and  $PM_{2.5}$  (PM with a diameter  $\leq 2.5 \,\mu\text{m}$ ) concentrations by as much as 55 ppb and  $20 \,\mu g$  m<sup>-3</sup>, respectively (Wennberg and Dabdub, 2008). However, due to the lower ambient NO<sub>x</sub> and volatile organic compound (VOC) levels, the impact of the  $NO_2^*$  chemistry on  $O_3$ enhancement was found to be smaller in the U.S. after the year of 2000. The maximum increases of O<sub>3</sub> concentrations in urban areas in the U.S. less than 10 ppb (Sarwar et al., 2009; Ensberg et al., 2010) suggest that the effect of the NO<sub>2</sub> chemistry on pollution levels may be more significant in some industrialised regions with elevated levels of ambient NO<sub>x</sub> and VOCs, such as mega-cities in China. Thereafter studies reported that O<sub>3</sub> enhancements were more than 30 ppb in Beijing (Li et al., 2011) and Taiwan (Jorba et al., 2012), and 10 ppb over the urban areas of the Pearl River Delta region of China (Zhang et al., 2013).

Along with the direct emissions and the homogeneous formation of HONO, NO<sub>2</sub> heterogeneous reaction on wet surfaces (2NO<sub>2</sub>+H<sub>2</sub>O→HNO<sub>3</sub>+HONO) is commonly considered as one of the major HONO heterogeneous sources (Jenkin et al., 1988; Kleffmann et al., 1998; Finlayson-Pitts et al., 2003). The NO<sub>2</sub> heterogeneous reaction has been included previously in the Community Multiscale Air Quality (CMAQ) modelling system (Foley et al., 2010). The CMAQ simulations showed that the NO<sub>2</sub> heterogeneous reaction was found to be the most significant source of nocturnal HONO formation during the project of the Northeast Oxidant and Particle Study in the U.S. (Sarwar et al., 2008). The simulations conducted over the Beijing-Tianjin-Hebei region (referred as BTH region), as the most polluted areas in China, suggested that the largest enhancements of the monthly mean daytime O3 and  $PM_{2.5}$  during August 2007 were about 12 and 16%, respectively, and this was mostly due to the  $NO_2$  heterogeneous reaction on aerosol surfaces (An et al., 2013).

In addition to direct emissions, homogeneous gas-phase formation and heterogeneous reactions, a recent study has demonstrated that ammonia-oxidising bacteria in soil can directly release HONO (Oswald et al., 2013). The soil HONO emissions may play an important role in the biogeochemical and atmospheric nitrogen cycles. This finding helps understand how the biosphere affects the atmosphere and global climate. Before this new mechanism can be incorporated into chemical transport models, continuous and comprehensive field measurements are needed, including simultaneous monitoring of soil, ecosystem and atmospheric processes.

Previous studies mainly focused on the impacts of HONO sources on O<sub>3</sub> and PM<sub>2.5</sub> concentrations but rarely considered the relationship between O<sub>3</sub> and PM<sub>2</sub> 5 enhancements due to HONO sources. Meng et al. (1997) pointed out that there was a complex chemical coupling between O<sub>3</sub> and PM<sub>2.5</sub> components because they shared the same precursors as NO<sub>x</sub> and VOCs. In-situ observations also proved that severe PM pollution was often accompanied with high O<sub>3</sub> concentrations. Positive correlation between the daily PM<sub>2.5</sub> and maximum 8-h O<sub>3</sub> (O<sub>3 8h;max</sub>) concentrations was also observed over the BTH region during the Beijing 2008 Olympic Games (Xin et al., 2010). When additional HONO sources were considered in the O<sub>3</sub> and PM<sub>2.5</sub> interaction, another level of complexity was added into the chemical coupling system (Wennberg and Dabdub, 2008) and this impacted O<sub>3</sub> and PM<sub>2.5</sub> emission control strategies (Meng et al., 1997). Ensberg et al. (2010) studied the impacts of the NO<sub>2</sub> chemistry on O<sub>3</sub> and PM control strategies over the South Coast Air Basin of California during two emission episodes that occurred for 1987 and 2005. They showed that the NO<sub>2</sub> chemistry increased the sensitivity of O<sub>3</sub> and secondary PM concentrations with respect to changes in NO<sub>x</sub> emissions (Ensberg et al., 2010). Notable changes in O<sub>3</sub> responses to NO<sub>x</sub> and VOC emission controls were also found in the Pearl River Delta of China due to the NO<sub>2</sub> chemistry (Zhang et al., 2013). These studies suggested that the additional HONO sources had significant implications for polluted areas to achieve ozone and PM air quality standards.

In spite of considerable improvements in air quality indexes over the last decade, the BTH region continues to suffer from severe  $O_3$  and PM pollution as well as haze (Chan and Yao, 2008; Sun et al., 2013; Zhao et al., 2013). Significant  $O_3$  and PM<sub>2.5</sub> enhancements over the BTH region during August 2007 due to three additional HONO sources (including HONO emissions, the  $NO_2^*$  chemistry and  $NO_2$  heterogeneous reaction on aerosol surfaces) were also reported by Li et al. (2011) and An et al. (2013).

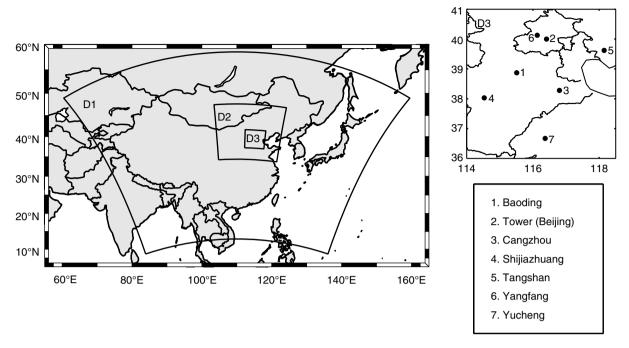


Fig. 1. Modelling domains and the monitoring sites over the BTH region.

These works suggested that more strict conditions on  $O_3$  and PM concentration standards were needed to be set up. Presently, the China's Grade-2 Standard list<sup>1</sup> includes the daily  $O_{3\_8h;max}$  at  $160\,\mu g$  m<sup>-3</sup> and 24-h average PM<sub>2.5</sub> at  $75\,\mu g$  m<sup>-3</sup>.

There are two main objectives in this study: (1) to analyse the relationship between the enhancements of daytime O<sub>3</sub> and PM<sub>2.5</sub> as a function of the additional HONO sources and (2) examine the reductions of daily O<sub>3\_8h;max</sub> and 24-h averaged PM<sub>2.5</sub> concentrations as functions of NO<sub>x</sub>, VOC and NH<sub>3</sub> emission reductions when the additional HONO sources are considered. Section 2 is introduced for the numerical experiments conducted in this study. Section 3 describes observation data used in the model evaluation. The effects of the additional HONO sources on O<sub>3</sub> and PM<sub>2.5</sub> chemical coupling and control strategies are discussed in Section 4 and the conclusions are presented in Section 5.

## 2. Model description

# 2.1. WRF-Chem model and new parameterisation of HONO

This study utilised a fully coupled Weather Research and Forecasting-Chemistry (WRF-Chem) model (Grell et al.,

2005; Fast et al., 2006) with the gas phase chemical Carbon-Bond Mechanism Z (CBM-Z; Zaveri and Peters, 1999) and aerosol module of the Model for Simulating Aerosol Interactions and Chemistry (MOSAIC; Zaveri et al., 2008). The physical schemes for simulations were identical to those chosen by Li et al. (2011).

Figure 1 shows the domains of the simulations. Domain 1, 2 and 3 primarily cover the Eastern Asia, Northern China and BTH region, respectively. Domain 3 contains 54 × 54 grids with a grid resolution of 9 km and centres at 40°N, 116°E. The 28 vertical layers extend from the surface to 50 hPa with the first level at 28 m above the ground. Meteorological initial and boundary conditions were obtained from the US National Centers for Environmental Prediction (NCEP; 6-h, 1° × 1°) final operation global analysis dataset. Chemical initial and boundary conditions were interpolated from the Model for OZone And Related chemical Tracers version 4 (MOZART-4; Emmons et al., 2010) outputs at every 6 h. Monthly anthropogenic emissions covering East Asia with a grid resolution of  $0.5^{\circ} \times 0.5^{\circ}$ were adopted from Zhang et al. (2009). Future emission inventory over this region requires a finer grid resolution to help air quality models better capture the localised photochemistry in the atmosphere. Biogenic emissions were calculated based on Guenther et al. (1993) and Simpson et al. (1995).

The following input for model reference simulation and test runs was considered. Besides HONO gas-phase production from OH and NO, three additional HONO sources

<sup>&</sup>lt;sup>1</sup>http://kjs.mep.gov.cn/hjbhbz/bzwb/dqhjbh/dqhjzlbz/<u>2012</u>03/W0 20120410330232398521.pdf

(i.e. HONO emissions, the NO<sub>2</sub> chemistry and NO<sub>2</sub> heterogeneous reaction on aerosol surfaces) were coupled into the WRF-Chem model. Li et al. (2011, 2014) and An et al. (2013) have described the parameterisation schemes in detail. Briefly, HONO emissions included direct emissions estimated as 0.8% of NO<sub>x</sub> emissions (Kurtenbach et al., 2001) and 2.3% of the NO<sub>x</sub> emitted in diesel exhaust gases converted to HONO via heterogeneous reaction with semivolatile organics (Gutzwiller et al., 2002). The ratio of HONO/NO<sub>x</sub> used in the model set up was about 1.18% in the urban centre of Beijing (Li et al., 2011; An et al., 2013). The rate constant for the reaction of NO<sub>2</sub> with H<sub>2</sub>O was estimated as  $9.1 \times 10^{-14}$  cm<sup>3</sup> molecule<sup>-1</sup> s<sup>-1</sup> which was similar to the mean values of Li et al. (2008) and Crowley and Carl (1997). The NO<sub>2</sub> heterogeneous reaction on aerosol surfaces followed Jacob (2000) recommendations. To compute the reaction rate, the total aerosol surface area  $(S_a)$  per unit volume of air is derived from aerosol mass concentrations and number density in eight size bins set by the MOSAIC aerosol module. The mean calculated  $S_a$  was 1367 μm<sup>2</sup> cm<sup>-3</sup> in Beijing during August 2007, within the observed range of 1000-2500  $\mu$ m<sup>2</sup> cm<sup>-3</sup> (Spataro et al., 2013).

#### 2.2. Experimental design

Five simulations given in Table 1 were conducted in this study. The first one was a reference case (Case R) that was performed using the standard CBM-Z mechanism and the MOSAIC module. Case Emis, Case NO<sub>2</sub>\* and Case Het were extensions of Case R by inclusion of HONO emissions, the NO<sub>2</sub>\* chemistry and the NO<sub>2</sub> heterogeneous reaction, respectively. The last simulation referred as an enhanced case (Case E) contained all the three additional HONO sources. The simulation period was August 2007, during which an extensive observed dataset was obtained, including highly time-resolved ambient gas phase species and aerosols (shown in Section 3).

Table 1. Design of WRF-Chem simulations to evaluate the effects of the additional HONO sources

Case ID	HONO sources included in the WRF-Chem
	simulations
Case R	Reference case with HONO gas-phase production
	from OH and NO
Case Emis	Case $R + HONO$ emissions
Case NO <sub>2</sub> *	Case $R + NO_2^*$ chemistry
Case Het	Case $R + NO_2$ heterogeneous reaction on aerosol
	surfaces
Case E	Enhanced case with all the three additional HONO
	sources

Additional sensitivity runs for the impacts of the three additional HONO sources on O3 and PM2.5 control strategies were conducted on 17 August 2007 when the BTH region experienced exceedingly high ozone and PM levels. The daily O<sub>3\_8h;max</sub> in Beijing violated China's Grade-2 Standard values. In order to save the computational time, four emission scaling factors, i.e. 1, 0.75, 0.5 and 0.25, are employed. The reference case emissions (Zhang et al., 2009) of NO<sub>x</sub>, VOCs and NH<sub>3</sub> were multiplied by the scaling factors individually, leading to 64 emission scenarios. The 128 (64 × 2) sensitivity simulations were conducted for Domain 3. The first 64 scenarios employed the standard CBM-Z and MOSAIC modules and the remaining 64 scenarios included the three additional HONO sources. All of the simulations started with the initial conditions at 00:00/16/8/2007 (UTC) using the standard CBM-Z and MOSAIC modules for the reference case emissions, then they were integrated from 00:00/16/8/2007 to 00:00/18/8/ 2007 (UTC) under different emission scenarios.

### 3. Observations and data

The HONO concentrations were measured on a 24-h basis by the annular denuder system at the campus of Peking University (39.99°N, 116.28°E) from 2 to 31 August 2007 as part of the CAREBeijing-2007 Experiment (2007 Campaigns of Air Quality Research in Beijing; Spataro et al., 2013). Measurements of NO<sub>2</sub>, nitrate (NO<sub>3</sub><sup>-</sup>) and ammonium (NH<sub>4</sub>) were also used in this study (Ianniello et al., 2011). Hourly HONO concentrations were taken from Differential Optical Absorption Spectroscopy (DOAS) measurements in Beijing from 13 to 25 August 2007 (Zhu et al., 2009). The retroreflector array was placed on a platform at 15 m of the 325 m Meteorological Tower located at the Institute of Atmospheric Physics (IAP; 39.97°N, 116.37°E). The fixed light path was 470 m (one way) across an expressway (Li et al., 2011). Hourly measurements of meteorological parameters including air temperature (T), relative humidity (RH), wind speed (WS) and direction (WD) from 13 to 20 August were adopted from the Olympic Village (40.00°N, 116.37°E) site close to IAP because the observed meteorological data at the tower site were unavailable within the studied period. There was no precipitation happening during this period. Observations of the surface NO<sub>2</sub>, O<sub>3</sub> and PM<sub>2,5</sub> concentrations over the BTH region were obtained from the Beijing Atmospheric Environmental Monitoring Action which was supported by the Beijing Olympic Technological Project (Xin et al., 2010; Gao et al., 2011). Hourly measurements of NO<sub>2</sub>, O<sub>3</sub> and PM<sub>2.5</sub> at seven urban sites shown in Fig. 1 (Li et al., 2011) were used in the analysis.

#### 4. Results and discussion

#### 4.1. Model validation

Previously, the comprehensive model evaluations have been conducted for Case Emis, Case NO<sub>2</sub> and Case Het (Li et al., 2011; An et al., 2013; Tang et al., 2014). In this section, the simulation results for Case R and Case E are evaluated to examine the performance of the extended WRF-Chem model with the three additional HONO sources. Figure 2 shows the simulations of 24-h average concentrations of HONO, NO<sub>2</sub>, NO<sub>3</sub> and NH<sub>4</sub> compared to the measurements conducted at Peking University in Beijing during August 2007. The mean simulated HONO concentration is only 0.08 ppb in Case R. When the additional HONO sources are considered for Case E, the mean HONO concentration is improved up to 1.36 ppb. This value is very close to the observed mean of 1.43 ppb (Fig. 2a). The NO<sub>2</sub> concentrations in Case E are approximately 2-10 ppb lower than those in Case R (Fig. 2b), which is mostly due to the NO<sub>2</sub> heterogeneous reaction on aerosol surfaces. Although the simulated NO<sub>2</sub> agrees reasonably well with most of the observations (Fig. 2b), the simulated NO<sub>3</sub> overestimates the observations by 5-20% (Fig. 2c) for both Case R and Case E. This NO<sub>3</sub> overestimation was found in the whole Northeast Asian region (Kajino et al., 2012, 2013) and further investigation with additional observations is required. Both Case R and Case E well reproduce NH<sub>4</sub><sup>+</sup> variations and quantities (Fig. 2d). The correlation coefficient (R) is found as 0.59. The simulated mean NH<sub>4</sub><sup>+</sup> in Cases R and E are 8.55 and  $10.27 \,\mu g$  m<sup>-3</sup>, respectively.

The observed mean is  $12.27 \,\mu g \, m^{-3}$ . The additional HONO sources significantly enhanced  $NO_3^-$  and  $NH_4^+$  concentrations (Fig. 2c, d), and this resulted in an increase for  $PM_{2.5}$  concentrations (An et al., 2013).

The comparison between the simulated and observed hourly concentrations of HONO and OH at the Meteorological Tower in Beijing, NO2, O3 and PM2.5 at the seven urban sites over the BTH region are shown in the Supplementary file and our previous studies (Li et al., 2011; An et al., 2013; Tang et al., 2014). The simulations generally captured the diurnal variations of pollutants at different sites of the BTH region. The comparative statistics for Case E are given in Table 2. The variations of T, RH, WS and WD near the Meteorological Tower site are generally simulated well. During 13-20 August 2007 the high temperature and humidity values during low speed of southerly wind existed in Beijing. The meteorological conditions as well as the large emissions of NOx from the heavy traffic (An et al., 2009) nearby the sampling site resulted in the high HONO concentrations. The Case E average predicted HONO concentrations of 1.69 ppb agreed reasonably with the measured concentration of 1.27 ppb. The simulations of Case E also generally captured the hourly variations of NO<sub>2</sub>, O<sub>3</sub> and PM<sub>2.5</sub> at the seven urban sites over the BTH region, with the correlation coefficients being 0.52, 0.87 and 0.51, respectively. The FAC2 (fractions of simulated values within a factor of two of the observed values) for NO<sub>2</sub>, O<sub>3</sub> and PM<sub>2.5</sub> are 0.50, 0.62 and 0.64, respectively (Table 2).

For daytime (7:00–19:00, LT) and night – time (20:00–6:00, LT) conditions, HONO simulations for Case R and

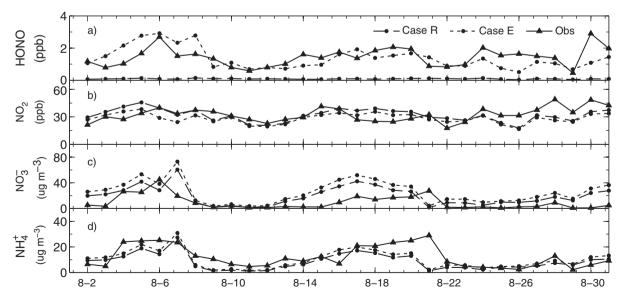


Fig. 2. Comparison of simulated 24-h average (a) HONO, (b)  $NO_2$ , (c)  $NO_3^-$  and (d)  $NH_4^+$  concentrations with observations at Peking University during August 2007 (Ianniello et al., 2011; Spataro et al., 2013). Case R is a reference case; Case E includes the HONO emissions,  $NO_2^*$  chemistry and  $NO_2$  heterogeneous reaction on aerosol surfaces.

Table 2	Statistics for c	comparison of t	he observed	(Obs.)	and simulated	(Sim.) h	ourly data	during Aug	rust 13-20.	2007

	T	RH	WS	WD	HONO	$NO_2$	$O_3$	$PM_{2.5}$
Units	°C	0/0	${\rm m~s}^{-1}$	deg	ppb	ppb	ppb	$\mu$ g m $^{-3}$
Number of data	184	184	184	184	117	1304	1194	1340
Median (Obs.)	28.00	68.00	1.40	135.0	0.89	17.02	37.23	94.78
Median (Sim.)	29.15	41.11	2.11	190.7	1.49	20.79	32.77	143.86
Mean (Obs.)	28.05	66.45	1.46	136.35	1.27	19.29	44.61	100.11
Mean (Sim.)	28.92	43.29	2.34	185.74	1.69	24.62	38.62	161.18
MB	0.87	-23.15	0.87	49.40	0.42	5.33	-5.99	61.07
$NMB^a$	3	-34	52	36	33	28	-13	61
RMSE	1.85	24.78	1.60	133.56	1.05	19.01	18.59	96.84
R	0.94	0.88	0.47	0.42	0.71	0.52	0.87	0.51
FAC2 <sup>b</sup>	1.0	0.98	0.55	0.58	0.60	0.50	0.62	0.64

Observations of HONO are from the Beijing Meteorological Tower and the meteorological parameters are from the Olympic Village near to the tower site (see Section 3).  $NO_2$ ,  $O_3$  and  $PM_{2.5}$  measurements are from the seven urban sites over the BTH region (Fig. 1). Model results are for Case E.

Case E are compared with the observations for the Meteorological Tower (Fig. 3a). The additional HONO sources significantly improved HONO simulations for both

daytime and night-time. The simulated mean concentration of HONO for daytime increased from 0.10 ppb in Case R to 0.51 ppb in Case E, and for night-time, it increased from

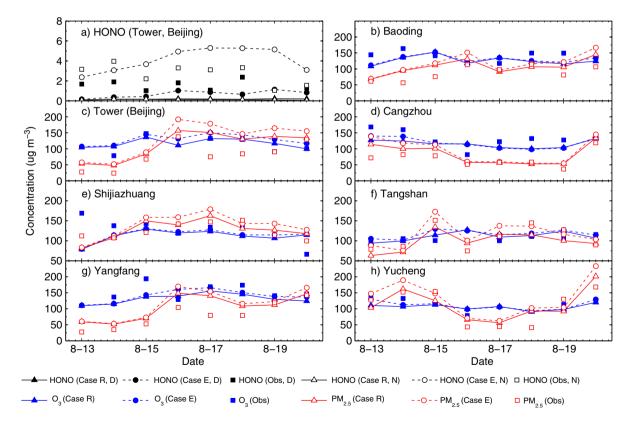


Fig. 3. Observed and simulated (a) daytime (D, 7:00–19:00) and night-time (N, 20:00–6:00) mean concentrations of HONO at the Beijing Meteorological Tower site and (b–h) daytime mean concentrations of  $O_3$  and  $PM_{2.5}$  at the seven urban sites over the BTH region during 13–20 August, 2007. The unit used is  $\mu g m^{-3}$  (at  $1.01325 \times 10^5 \, Pa$  and  $25^{\circ} C$ ,  $1 \, \mu g m^{-3}$  is  $\sim 0.52 \, ppb$  for HONO and  $\sim 0.51 \, ppb$  for  $O_3$ ).

<sup>&</sup>lt;sup>a</sup>The unit of NMB is in %.

<sup>&</sup>lt;sup>b</sup>Fraction of simulations within a factor of two of observations.

0.06 ppb to 2.05 ppb. Figure 3b-h shows the daytime simulations of  $O_3$  and  $PM_{2.5}$  at the seven urban sites over the BTH region. Both Cases R and E performed well for predicting  $O_3$  and  $PM_{2.5}$  daytime concentrations. The observed and simulated means, the mean bias (MB), the normal mean bias (NMB), the root mean square error (RMSE) and FAC2 of  $O_3$  for Case E are 64.97 ppb, 63.05 ppb, -1.92 ppb, -2.95%, 12.48 ppb, and 0.98, respectively; corresponding  $PM_{2.5}$  values for Case E are 94.01  $\mu$ g m<sup>-3</sup>, 119.24  $\mu$ g m<sup>-3</sup>, 25.23  $\mu$ g m<sup>-3</sup>, 26.83%, 36.05  $\mu$ g m<sup>-3</sup> and 0.91, respectively.

# 4.2. Daytime enhancements of $O_3$ and $PM_{2.5}$ due to the additional HONO sources

4.2.1. The chemical coupling between  $O_3$  and  $PM_{2.5}$ . The chemical coupling between O3 and PM2.5 was emphasised in previous studies (Meng et al., 1997; Nguyen and Dabdub, 2002; Wennberg and Dabdub, 2008). This is of profound importance in understanding processes that control the concentrations of both O<sub>3</sub> and PM<sub>2.5</sub>. As shown in Fig. 4, OH is the essential oxidant during daytime. The OH oxidises VOC to produce O<sub>3</sub> in presence of NO<sub>x</sub> and oxidises SO<sub>2</sub> and NO<sub>x</sub> to produce secondary inorganic particulate in presence of NH<sub>3</sub>. OH and O<sub>3</sub> attack on organic molecules can generate semivolatile secondary organic aerosol (SOA). HONO sources add another level of complexity to this mechanism. When the HONO emissions, NO<sub>2</sub> chemistry and NO2 heterogeneous reaction on aerosol surfaces are considered, the OH formation is increased via the HONO photolysis and reaction of  $NO_2^*$  with  $H_2O$ . The increases in OH concentrations simultaneously enhance O<sub>3</sub> and PM<sub>2.5</sub>

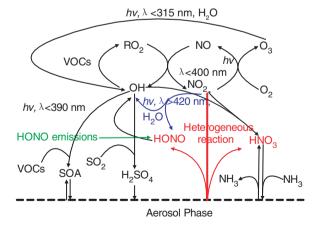


Fig. 4. Impacts of the additional HONO sources on the chemical coupling between ozone and particulate matter. HONO sources increase OH concentrations and subsequently promote O<sub>3</sub> and PM production. HNO<sub>3</sub> produced from the NO<sub>2</sub> heterogeneous reaction on aerosol surfaces enhances nitrate concentrations in presence of NH<sub>3</sub>.

production. In addition, increasing  $O_3$  concentration further increases the daytime formation of OH. The OH in turn reacts with  $NO_2$  to form gas nitric acid (HNO<sub>3</sub>). The  $NO_2$  heterogeneous reaction on aerosol surfaces also contributes to  $HNO_3$  enhancement, which eventually increases nitrate concentrations.

Although O<sub>3</sub> and PM<sub>2.5</sub> are subjected to a complex series of common emissions and photochemical production pathways, the coupling between O<sub>3</sub> and PM<sub>2,5</sub> is in a non-linear manner. O<sub>3</sub> production is related mainly to the VOCs and NO<sub>x</sub> gas-phase chemistry while PM<sub>2.5</sub> production is contributed mostly to the gas-to-particle conversion processes, involving chemical reactions and mass transport in the gas and particle phases (Shiraiwa et al., 2014). The impacts of the additional HONO sources on O<sub>3</sub> and PM enhancements are found to be more significant over urban and coastal areas with high NO<sub>x</sub> emissions. On the other hand, the impacts are small in rural and background areas where NO<sub>x</sub> emissions are relatively low (Wennberg and Dabdub, 2008; Sarwar et al., 2009; An et al., 2011; Li et al., 2011; Elshorbany et al., 2012, 2014; Jorba et al., 2012; Zhang et al., 2013). In the regions with high VOCs to NO<sub>x</sub> ratios (e.g. rural areas), O<sub>3</sub> production is sensitive to NO<sub>x</sub> concentrations. The NO<sub>2</sub> reductions due to the NO<sub>2</sub> heterogeneous reaction can decrease O<sub>3</sub> production in rural areas. Therefore, the additional HONO sources are more likely to enhance O3 and PM<sub>2.5</sub> simultaneously in urban areas with a low ratio of VOCs to NO<sub>x</sub>. In those areas, decreasing NO<sub>2</sub> alone likely leads to increasing O<sub>3</sub> concentrations (Wennberg and Dabdub, 2008; Ensberg et al., 2010).

In order to investigate the impacts of the additional HONO sources on O<sub>3</sub> and PM<sub>2.5</sub> chemical interactions, the enhancements of daytime mean concentrations of O3 and PM<sub>2.5</sub> at the seven urban sites during 13-20 August 2007 shown in Fig. 3 are demonstrated in scatter diagrams (Fig. 5). Both O<sub>3</sub> and PM<sub>2.5</sub> daytime enhancements (Fig. 5a) due to the  $NO_2^*$  chemistry are mostly less than  $5 \mu g \text{ m}^{-3}$  (for  $O_3$ , 1 µg m<sup>-3</sup> = 0.51 ppb at 1.01325 × 10<sup>5</sup> Pa and 25°C). Among the three additional HONO sources, the NO2 heterogeneous reaction is the largest contributor to O<sub>3</sub> and PM<sub>2.5</sub> enhancements. Up to  $14 \mu g \text{ m}^{-3}$  (7 ppb) and  $28 \mu g \text{ m}^{-3}$ increases are found for O<sub>3</sub> and PM<sub>2.5</sub>, respectively (Fig. 5b). The contributions of HONO emissions to O<sub>3</sub> and PM<sub>2.5</sub> enhancements are relatively small ( $<2 \mu g \text{ m}^{-3}$ ; not shown). The largest enhancements of daytime concentrations of O<sub>3</sub> and PM<sub>2.5</sub> due to all the three additional HONO sources are estimated to be  $17.27 \,\mu g$  m<sup>-3</sup> (8.79 ppb) for  $O_3$  and 31.54 µg m<sup>-3</sup> for PM<sub>2.5</sub> (Fig. 5c). Figure 5 also clearly shows that the daytime O3 increases are closely related to the daytime PM<sub>2.5</sub> increases. The correlation coefficients between O<sub>3</sub> and PM<sub>2.5</sub> enhancements are 0.71 and 0.75, respectively, when the NO<sub>2</sub> chemistry and NO<sub>2</sub>

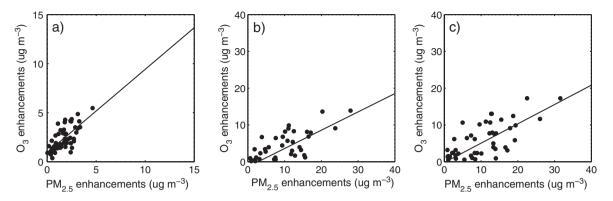


Fig. 5. Enhancements of daytime mean concentrations of  $O_3$  and  $PM_{2.5}$  due to (a) the  $NO_2^*$  chemistry (b) the  $NO_2$  heterogeneous reaction on aerosols and (c) all the three additional HONO sources at the seven urban sites demonstrated in Fig. 3 during 13–20 August, 2007.

heterogeneous reaction are individually included in the WRF-Chem model.

The correlation coefficient between the daytime enhancement of O<sub>3</sub> and PM<sub>2.5</sub> over the BTH region for the whole month was estimated. As shown in Fig. 6a, the correlation coefficients between O<sub>3</sub> and PM<sub>2.5</sub> daytime enhancements due to the additional HONO sources are greater than 0.6 in many urban areas and greater than 0.8 over Beijing and

Tianjin areas. The high correlation is found mainly over the urban areas where the values of the linear-fitting slope for  $O_3$  and  $PM_{2.5}$  enhancements are above zero (Fig. 6b), indicating that the additional HONO sources increase  $O_3$  and  $PM_{2.5}$  concentrations simultaneously. The increases in both  $O_3$  and  $PM_{2.5}$  are because of the enhanced OH concentrations as demonstrated in Fig. 4. The significant increases in  $O_3$  concentrations occur in urban areas also

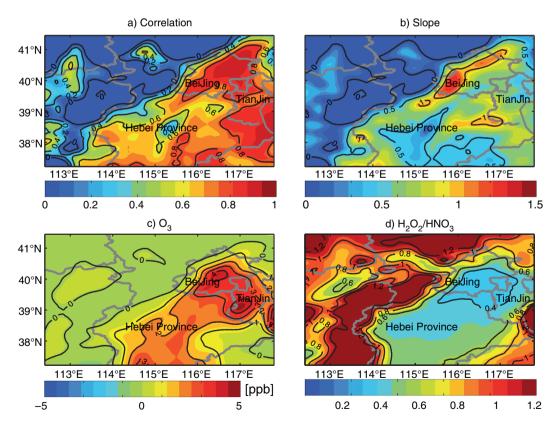


Fig. 6. Values of (a) the correlation coefficient and (b) the linear-fitting slope for  $O_3$  and  $PM_{2.5}$  daytime enhancements due to the three additional HONO sources during August 2007; monthly mean daytime (c)  $O_3$  enhancements due to the three additional HONO sources and (d)  $H_2O_2/HNO_3$  ratio over the BTH region.

because  $O_3$  production is mainly sensitive to VOC concentrations in those areas, where the  $NO_2$  decreases due to the  $NO_2$  heterogeneous reaction increase  $O_3$ . As shown in Fig. 6c, 2–5 ppb increases in the monthly mean daytime  $O_3$  concentrations are found in Beijing, Tianjin and south of Hebei Province.

The ratio of H<sub>2</sub>O<sub>2</sub>/HNO<sub>3</sub> is often used as an indicator to identify VOC-sensitive or NO<sub>x</sub>-sensitive regimes (Sillman, 1995; Lam et al., 2005). The transition value of 0.4-0.8 is adopted in this study according to the field measurements conducted in Beijing (He et al., 2010): a value < 0.4 indicates a VOC-sensitive regime, a value > 0.8 indicates a NO<sub>x</sub>sensitive regime, and otherwise it indicates an intermediate regime (Fig. 6d). Comparing Fig. 6a and d, it is found that the low correlation (<0.2) between  $O_3$  and  $PM_{2.5}$  daytime enhancements occurs mainly over the rural areas, such as Taihang Mountains to the west of Hebei Province, and Yan Mountains to the north of Beijing. Those rural areas are within the NO<sub>x</sub>-sensitive regimes where NO<sub>2</sub> decreases due to the NO<sub>2</sub> heterogeneous reaction (see Fig. 2b and Li et al., 2011) will reduce O<sub>3</sub> (Fig. 6c). Daytime enhancements of O<sub>3</sub> and PM<sub>2.5</sub> for all the simulated grids within the BTH region during August 2007 are presented in Supplementary Fig. 5. The correlation coefficients between O<sub>3</sub> and PM<sub>2.5</sub> daytime enhancements are estimated as 0.83, 0.91 and 0.90, for the NO<sub>2</sub> chemistry, NO<sub>2</sub> heterogeneous reaction and the three additional HONO sources included in the simulations, respectively.

4.2.2. Impacts on concentrations of  $PM_{2.5}$  and its major components. The additional HONO sources enhance O<sub>3</sub> production mainly through increasing OH concentrations which will enhance PM<sub>2.5</sub> production as well. However, the coupling between O<sub>3</sub> and PM<sub>2.5</sub> production is in a non-linear way. As stated in Fig. 4, PM<sub>2.5</sub> production also involves complex gas precursors that go through multiphase chemical evolution. Figure 7 shows that the impacts of the additional HONO sources on individual PM<sub>2.5</sub> components during daytime. NO<sub>3</sub> has the largest enhancements  $(10-58\%, 4-9 \,\mu g \, m^{-3})$ , followed by  $NH_4^+$  (10-40%, $1-3 \,\mu \text{g m}^{-3}$ ) and  $SO_4^{2-}$  (5–18%, 0.5–1.5  $\mu \text{g m}^{-3}$ ). The large increment is found mainly in areas with high NO<sub>x</sub> levels, e.g. Beijing, Tianjin and south of Hebei Province. The HONO-related enhancement of NO<sub>3</sub> is due to the enhanced OH level (10-40%, Fig. 7e) and the NO<sub>2</sub> heterogeneous reaction that produces more HNO<sub>3</sub> (Fig. 7f). The reason for the enhancement of SO<sub>4</sub><sup>2-</sup> is that the increasing OH can oxidise SO<sub>2</sub> to produce gas-phase H<sub>2</sub>SO<sub>4</sub> and enhance H<sub>2</sub>O<sub>2</sub> (Fig. 7g) and O<sub>3</sub> (Fig. 6c), which can oxidise dissolved  $SO_2$  to produce  $SO_4^{2-}$  through aqueous phase reactions. The increases in HNO3 and H2SO4 can be neutralised by NH<sub>3</sub> to form NH<sub>4</sub>. The NO<sub>2</sub> heterogeneous reaction and H<sub>2</sub>O<sub>2</sub> enhancements are major reasons for night-time enhancements of PM<sub>2.5</sub> and its major components (Supplementary Fig. 6) and the contributions of the individual HONO sources to PM<sub>2.5</sub> and its major inorganic components can be found in An et al. (2013).

Monthly-mean concentrations of PM<sub>2.5</sub> and its major components characterised by the MOSAIC aerosol module in the WRF-Chem are shown in Supplementary Fig. 7. It should be noted that the formation of SOA is not taken into account in this study. Organic aerosol, especially SOA could contribute significantly to PM25. In-situ observations showed that OA on average accounted for approximate 40% of submicron aerosols in Beijing during summertime, and SOA dominated OA, contributing 64% (Sun et al., 2012 and references therein). However, the complexity involved in SOA formation makes SOA difficult to be simulated well. Current air quality models have generally underpredicted observed SOA concentrations, with model-measurement discrepancies of a factor of two or more. More recent modelling efforts incorporating a number of additional factors such as new SOA precursors and particle phase chemistry, have begun to close the gap between predicted and measured SOA concentrations. But even for cases when simulations show reasonable agreement with measurements, the agreement may not be for the right reasons (Hallquist et al., 2009). The impacts of the additional HONO sources on SOA concentrations will be considered in following studies since the enhancement of the oxidants augments the production of semi-VOCs, which are apt to partition more to particle phase and further escalates the SOA concentrations (Li et al., 2010).

# 4.3. Sensitivity of $O_3$ and $PM_{2.5}$ to the additional HONO sources

Impacts of the additional HONO sources on the  $O_3$  and  $PM_{2.5}$  chemical coupling require rethinking of their control strategies. Meng et al. (1997) suggested that attempts to reduce PM levels required control of the  $NO_x$  and VOC emissions that are also precursors to  $O_3$  formation. The  $O_3$  control strategies rely primarily on  $NO_x$  and VOC emission variations. The urban areas of Beijing are often in a VOC-sensitive regime for ozone chemistry (see Fig. 6d and Wang et al., 2006). This section mainly presents the simulation results for Beijing Meteorological Tower, where the traffic emissions are large (An et al., 2009).

The resulting peak 8-h average ozone isopleths are shown in Fig. 8. Figure 8a shows that with decreased VOC emissions,  $O_{3\_8h;max}$  concentrations are reduced steadily, whereas a decrease in  $NO_x$  emissions over 50% is necessary to obtain the  $O_3$  reduction. Figure 8b shows the differences in  $O_{3\_8h;max}$  concentrations between Cases E and R.

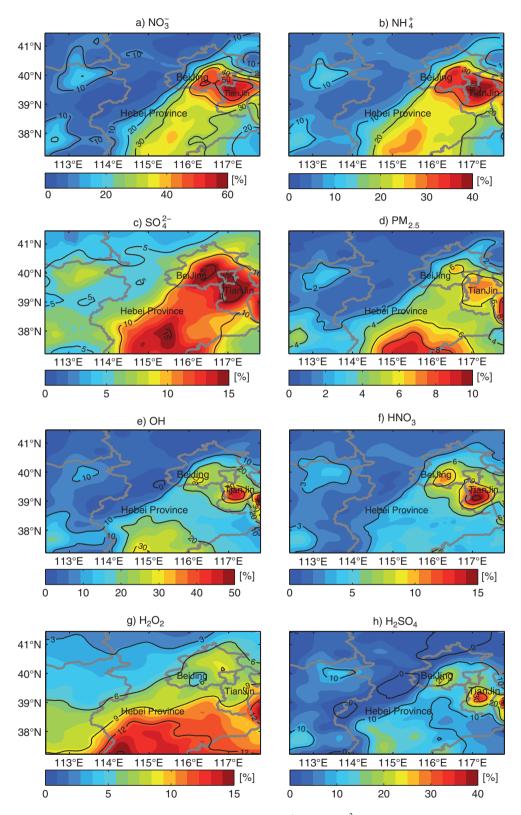


Fig. 7. Percentage increases of monthly-mean daytime (a)  $NO_3^-$ , (b)  $NH_4^+$  and (c)  $SO_4^{2-}$  in  $PM_{2.5}$  and (d)  $PM_{2.5}$ , (e) OH, (f) HNO<sub>3</sub>, (g)  $H_2O_2$  and (h)  $H_2SO_4$  concentrations over the BTH region during August 2007 due to the three additional HONO sources.

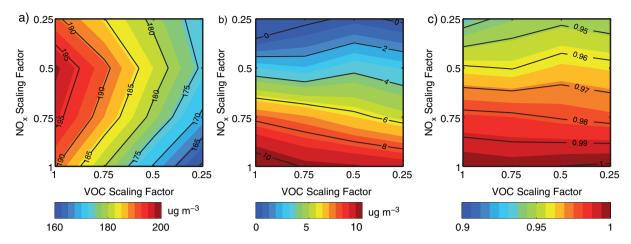


Fig. 8. Impacts of the additional HONO sources on  $O_{3\_8h;max}$  concentrations resulting from the control of  $NO_x$ , VOC and  $NH_3$  emissions. (a)  $O_{3\_8h;max}$  concentrations for Case R in Beijing on 17 August 2007; (b)  $O_{3\_8h;max}$  concentrations for Case E minus those for Case R; (c) ratio of Relative Reduction Factors (RRF) for Case E divided by RRF for Case R.

The additional HONO sources increase  $O_{3\_8h;max}$  concentrations but the enhancements decline as  $NO_x$  emissions decrease, indicating that the impacts of the additional HONO sources on  $O_3$  are significant in areas with high  $NO_x$  emissions (Sarwar et al., 2009; Ensberg et al., 2010).

In regulatory applications, the relative reduction factors (RRF) can be used to reflect control strategy effectiveness. RRF for a given pair of VOCs and NO<sub>x</sub> reduction factors, RRF ( $f_{VOC}$ ,  $f_{NOx}$ ), are calculated as the ratio of O<sub>3 8h;max</sub> resulting from that particular pair  $(f_{VOC}, f_{NOx})$  divided by  $O_{3 \text{ 8h;max}}$  from  $f_{\text{VOC}} = 1.0$  and  $f_{\text{NOx}} = 1.0$  (Hogrefe et al., 2008). Figure 8c presents the ratio of the RRF obtained in Case E divided by that obtained in Case R. The RRF ratios are below one under nearly all control scenarios, illustrating that the enhanced HONO sources increase the effectiveness of the control strategies in reducing O<sub>3</sub> concentrations, especially through NO<sub>x</sub> emission decreases. For example, when the additional HONO sources are considered, a reduction of 75% in emitted NO<sub>x</sub> leads to an O<sub>3</sub> reduction of 6.5% more than the O<sub>3</sub> reduction obtained in a case without the additional HONO sources. The RRF in Case E is up to 7% lower than that in Case R, similar to the results obtained in Los Angeles where the NO<sub>2</sub> chemistry produced RRF 5% lower than that in the reference case (Ensberg et al., 2010).

The additional HONO sources also affect PM<sub>2.5</sub> control strategies. Figure 9 shows isopleths of the 24-h average concentrations of PM<sub>2.5</sub> for Cases R and E. Although PM<sub>2.5</sub> decreases with any reductions of NO<sub>x</sub>, VOCs or NH<sub>3</sub> emissions in both cases, the additional HONO sources induce more reductions of PM<sub>2.5</sub> in Case E than those in Case R. For example, when the NO<sub>x</sub> reduction factor declines from 1 to 0.25, PM<sub>2.5</sub> is decreased from 143.1  $\mu$ g m<sup>-3</sup>

to 123.8  $\mu$ g m<sup>-3</sup>, with a reduction of 13.5% in Case R (Fig. 9a), whereas PM<sub>2.5</sub> is decreased from 163.1  $\mu$ g m<sup>-3</sup> to 128.6  $\mu$ g m<sup>-3</sup>, with a reduction of 21.1% in Case E (Fig. 9d).

The present results are found to be in agreement with previous studies. Ensberg et al. (2010) showed that the NO<sub>2</sub>\* chemistry increases the sensitivity of particle formation to changes in NO<sub>x</sub> emissions. In addition, the results in this study show that the three additional HONO sources also increase the sensitivity of PM<sub>2.5</sub> formation to changes in NH<sub>3</sub> emissions. Comparing Fig. 9b and e, the number of isopleths in Case E is more than that in Case R as NH<sub>3</sub> emissions are reduced, indicating that a decrease in NH<sub>3</sub> emissions leads to more PM<sub>2.5</sub> reduction in Case E than that in Case R. Figure 9c and f show that reducing NO<sub>x</sub> and NH<sub>3</sub> emissions simultaneously results in significant decreases in PM<sub>2.5</sub> concentrations. When NO<sub>x</sub> and NH<sub>3</sub> reduction factors decrease from 1 to 0.25, the PM<sub>2.5</sub> reduction in Case E is found to be 8% larger than that in Case R.

To better understand the effects of the additional HONO sources on  $PM_{2.5}$  and  $O_3$  control strategies, Fig. 10 illustrates how many times the concentration reductions of  $PM_{2.5}$  and  $O_3$  fall in a certain range due to the emission reductions of  $NO_x$ , VOCs and  $NH_3$ . Among the 64 sensitivity simulations (see Section 2.2), the number of large reductions of  $PM_{2.5}$  or  $O_3$  in Case E is more than that in Case R. For example, the number of  $PM_{2.5}$  reductions in the range of -25 to -20% is 17 in Case E, whereas that in Case R is 4. The number of  $O_3$  reductions in the range of -15 to -10% is 20 in Case E, much larger than that of 8 in Case R. Figure 11 further shows that within the same range of  $PM_{2.5}$  or  $O_3$  reductions in Cases R and E, the frequency of large decreases in pollutant emissions is much

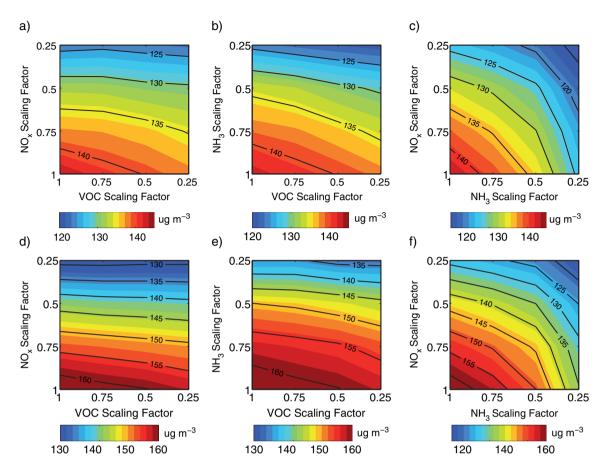


Fig. 9. Twenty four-hour average PM<sub>2.5</sub> concentrations for Case R (a-c) and for Case E (d-f) in Beijing on 17 August 2007.

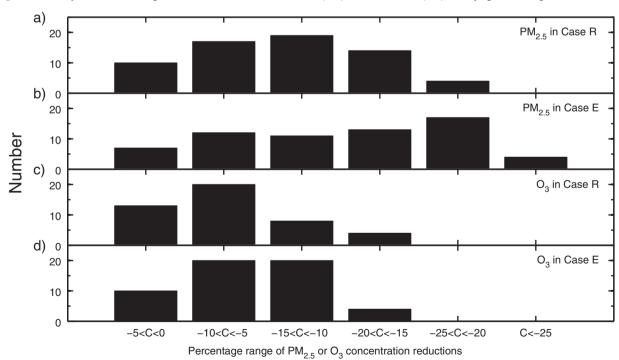


Fig. 10. The number of concentration reductions (C, with unit of %) of (a) PM2.5 in Case R, (b) PM2.5 in Case E, (c) O3 in Case R, and (d) O3 in Case E in a certain reduction range among the sensitivity simulations using the 64 emission scenarios (see Section 2.2).

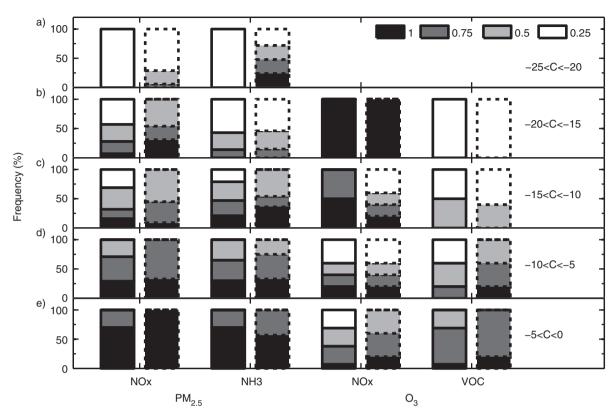


Fig. 11. Comparison of  $NO_x$ , VOC and  $NH_3$  emission reductions between Case R (bars with solid baselines) and Case E (bars with dotted baselines) to gain the same range of  $PM_{2.5}$  or  $O_3$  concentration reductions (C, in unit of %) of (a) -25% < C < -20%; (b) -20% < C < -15%; (c) -15% < C < -10%; (d) -10% < C < -5%; and (e) -5% < C < 0. The frequency for each scaling factor is computed by  $T_i/T$ , where T is the total times for  $PM_{2.5}$  or  $O_3$  concentration reductions falling in a certain reduction range;  $T_i$  is the occurrence times within T for a specific scaling factor (1, 0.75, 0.5 or 0.25 shown in Section 2.2).

less in Case E than that in Case R. For example, the frequency of  $NO_x$ , VOC and  $NH_3$  reduction factors of 0.25 is less in Case E than that in Case R in a certain range of  $PM_{2.5}$  or  $O_3$  concentration reductions. The frequency of 50% decreases in  $NO_x$  and  $NH_3$  emissions in Case E for the  $PM_{2.5}$  reductions between -10 and -5% is also much smaller than that in Case R (Fig. 11d). These results further indicate that the additional HONO sources enhance the effectiveness of control strategies for reducing  $PM_{2.5}$  and  $O_3$  concentrations.

# 5. Conclusions

Three additional HONO sources, i.e. the HONO emissions,  $NO_2^*$  chemistry and  $NO_2$  heterogeneous reaction on aerosol surfaces, were added into a fully coupled meteorology–chemistry model called WRF-Chem to examine how the additional HONO sources influence  $O_3$  and  $PM_{2.5}$  chemical coupling and their control strategies. The following main conclusions are obtained from the current study:

(1) The three additional HONO sources significantly increased daytime O<sub>3</sub> and PM<sub>2.5</sub> concentrations over

- the urban areas of the BTH region. The enhancements of up to 9 ppb for  $O_3$  and  $32 \,\mu g$  m<sup>-3</sup> for  $PM_{2.5}$  are found at the seven urban sites during August 2007. The  $NO_2$  heterogeneous reaction on aerosol surfaces is the largest contributor to the enhancements of  $O_3$  and  $PM_{2.5}$  during daytime.
- (2) The O<sub>3</sub> increases are found to be closely associated with PM<sub>2.5</sub> increases over VOC-sensitive regimes during daytime when the additional HONO sources are considered. The correlation coefficient between O<sub>3</sub> and PM<sub>2.5</sub> daytime enhancements is estimated to be greater than 0.6 in many urban areas of the BTH region and > 0.8 over Beijing and Tianjin areas.
- (3) The additional HONO sources enhance PM<sub>2.5</sub> components of SO<sub>4</sub><sup>2-</sup>, NO<sub>3</sub><sup>-</sup> and NH<sub>4</sub><sup>+</sup>, which are increased by 5–18, 10–58 and 10–40%, respectively, over urban areas during daytime. The simultaneous increment of O<sub>3</sub> and PM<sub>2.5</sub> during daytime due to the additional HONO sources is related to the increasing oxidants (OH, H<sub>2</sub>O<sub>2</sub> and O<sub>3</sub>) that enhance the atmospheric oxidising capacity. The 20–40%

enhancements of the monthly mean daytime OH concentrations are found over Beijing, Tianjin and south of Hebei Province.

- (4) The changes in O<sub>3</sub> and PM<sub>2.5</sub> concentrations under a variety of NO<sub>x</sub>, VOC and NH<sub>3</sub> emission reduction scenarios revealed that the additional HONO sources enhanced the effectiveness of control strategies for reducing O<sub>3</sub> and PM<sub>2.5</sub> concentrations.
- (5) A decrease in NO<sub>x</sub> emissions produced more decreases in O<sub>3</sub> concentrations when the additional HONO sources were included. The additional HONO sources resulted in the O<sub>3</sub> RRF 7% lower than that in the reference case and this is consistent with the recent study conducted for the South Coast Air Basin of California (Ensberg et al., 2010).
- (6) The additional HONO sources significantly increased  $PM_{2.5}$  sensitivity to the changes in  $NO_x$  and  $NH_3$  emissions.

The results presented in this study suggest that the complex chemical interactions between O<sub>3</sub> and PM<sub>2.5</sub> need better explanations when the additional HONO sources are considered. Without considering the additional HONO sources, the effectiveness of emission control strategies in reducing O<sub>3</sub> and PM<sub>2.5</sub> concentrations would be underestimated significantly. The effects of HONO sources on SOA concentrations will be evaluated in the following studies.

## 6. Acknowledgements

The research was partly supported by the National Natural Science Foundation of China (Grant No. 41405121 and 41175105), the Beijing Municipal Natural Science Foundation (Grant No. 8144054), the Key Project of Chinese Academy of Sciences (XDB05030301), and the Carbon and Nitrogen Cycle project of Institute of Atmospheric Physics, Chinese Academy of Sciences. Special thanks are given to Prof. Pinhua Xie from Anhui Institute of Optics and Fine Mechanics, Chinese Academy of Sciences for providing hourly HONO observations in Beijing and Dr. Bo Hu and Dongsheng Ji from the Chinese Ecosystem Research Network (CERN). NO<sub>2</sub>, O<sub>3</sub> and PM<sub>2.5</sub> observed data at seven sites over the BTH region were obtained from the CERN.

#### References

- Alicke, B., Platt, U. and Stutz, J. 2002. Impact of nitrous acid photolysis on the total hydroxyl radical budget during the Limitation of Oxidant Production/Pianura Padana Produzione di Ozono study in Milan. J. Geophys. Res. 107(D22), 8196. DOI: 10.1029/2000JD000075.
- Amedro, D., Parker, A. E., Schoemaecker, C. and Fittschen, C. 2011. Direct observation of OH radicals after 565 nm multi-

- photon excitation of  $NO_2$  in the presence of  $H_2O$ . Chem. Phys. Lett. **513**, 12–16.
- An, J., Li, Y., Chen, Y., Li, J., Qu, Y. and co-author. 2013. Enhancements of major aerosol components due to three HONO sources in the North China Plain (NCP) and implications for visibility and haze. Adv. Atmos. Sci. 30(1), 57–66.
- An, J., Li, Y., Wang, F. and Xie, P. 2011. Impacts of photoexcited NO<sub>2</sub> chemistry and heterogeneous reactions on concentrations of O<sub>3</sub> and NO<sub>y</sub> in Beijing, Tianjin and Hebei Province of China. In: *Air Quality-Models and Applications* (ed. N. Mazzeo). InTech, Rijeka, pp. 197–210.
- An, J., Zhang, W. and Qu, Y. 2009. Impacts of a strong cold front on concentrations of HONO, HCHO, O<sub>3</sub>, and NO<sub>2</sub> in the heavy traffic urban area of Beijing. *Atmos. Environ.* **43**, 3454–3459.
- Carr, S., Heard, D. and Blitz, M. 2009. Comment on "Atmospheric hydroxyl radical production from electronically excited NO<sub>2</sub> and H<sub>2</sub>O". *Science* **324**(5925), 336.
- Chan, C. K. and Yao, X. 2008. Air pollution in mega cities in China. Atmos. Environ. 42(1), 1–42.
- Crowley, J. N. and Carl, S. A. 1997. OH formation in the photo-excitation of NO<sub>2</sub> beyond the dissociation threshold in the presence of water vapor. *J. Phys. Chem. A.* 101(23), 4178–4184.
- Elshorbany, Y. F., Crutzen, P. J., Steil, B., Pozzer, A., Tost, H. and co-author. 2014. Global and regional impacts of HONO on the chemical composition of clouds and aerosols. *Atmos. Chem. Phys.* 14, 1167–1184.
- Elshorbany, Y. F., Steil, B., Brühl, C. and Lelieveld, J. 2012. Impact of HONO on global atmospheric chemistry calculated with an empirical parameterization in the EMAC model. *Atmos. Chem. Phys.* **12**, 9977–10000.
- Emmons, L., Walters, S., Hess, P., Lamarque, J., Pfister, G. and co-authors. 2010. Description and evaluation of the Model for Ozone and Related chemical Tracers, version 4 (MOZART-4). *Geosci. Model Dev.* 3, 43–67.
- Ensberg, J. J., Carreras-Sospedra, M. and Dabdub, D. 2010. Impacts of electronically photo-excited NO<sub>2</sub> on air pollution in the South Coast Air Basin of California. *Atmos. Chem. Phys.* 10, 1171–1181.
- Fast, J. D., Gustafson, W. I. Jr., Easter, R. C., Zaveri, R. A., Barnard, J. C. and co-authors. 2006. Evolution of ozone, particulates, and aerosol direct radiative forcing in the vicinity of Houston using a fully coupled meteorology-chemistry-aerosol model. *J. Geophys. Res.* 111, D21305. DOI: 10.1029/2005JD006721.
- Finlayson-Pitts, B., Wingen, L., Sumner, A., Syomin, D. and Ramazan, K. 2003. The heterogeneous hydrolysis of NO<sub>2</sub> in laboratory systems and in outdoor and indoor atmospheres: an integrated mechanism. *Phys. Chem. Chem. Phys.* **5**(2), 223–242.
- Foley, K., Roselle, S., Appel, K., Bhave, P., Pleim, J. and coauthors. 2010. Incremental testing of the Community Multiscale Air Quality (CMAQ) modeling system version 4.7. *Geosci. Model Dev.* 3(1), 205–226.
- Gao, Y., Liu, X., Zhao, C. and Zhang, M. 2011. Emission controls versus meteorological conditions in determining aerosol concentrations in Beijing during the 2008 Olympic Games. *Atmos. Chem. Phys.* 11, 12437–12451.

- Grell, G. A., Peckham, S. E., Schmitz, R., McKeen, S. A., Frost, G. and co-authors. 2005. Fully coupled "online" chemistry within the WRF model. *Atmos. Environ.* 39, 6957–6975.
- Guenther, A. B., Zimmerman, P. R., Harley, P. C., Monson, R. K. and Fall, R. 1993. Isoprene and monoterpene emission rate variability: model evaluations and sensitivity analyses. *J. Geophys. Res.* 98(D7), 12609–12617.
- Gutzwiller, L., Arens, F., Baltensperger, U., Gaggeler, H. W. and Ammann, M. 2002. Significance of semivolatile diesel exhaust organics for secondary HONO formation. *Environ. Sci. Technol.* 36(4), 677–682.
- Hallquist, M., Wenger, J., Baltensperger, U., Rudich, Y., Simpson, D. and co-authors. 2009. The formation, properties and impact of secondary organic aerosol: current and emerging issues. *Atmos. Chem. Phys.* 9, 5155–5236.
- He, S. Z., Chen, Z. M., Zhang, X., Zhao, Y., Huang, D. M. and co-authors. 2010. Measurement of atmospheric hydrogen peroxide and organic peroxides in Beijing before and during the 2008 Olympic Games: chemical and physical factors influencing their concentrations. *J. Geophys. Res.* 115, D17307. DOI: 10.1029/2009JD13544.
- Hogrefe, C., Civerolo, K. L., Hao, W., Ku, J. Y., Zalewsky, E. E. and co-author. 2008. Rethinking the assessment of photochemical modeling systems in air quality planning applications. J. Air Waste Manage. Assoc. 58(8), 1086–1099.
- Ianniello, A., Spataro, F., Esposito, G., Allegrini, I., Hu, M. and co-author. 2011. Chemical characteristics of inorganic ammonium salts in PM<sub>2.5</sub> in the atmosphere of Beijing (China). Atmos. Chem. Phys. 11, 10803–10822.
- Jacob, D. J. 2000. Heterogeneous chemistry and tropospheric ozone. Atmos. Environ. 34, 2131–2159.
- Jenkin, M., Utembe, S. and Derwent, R. 2008. Modelling the impact of elevated primary NO<sub>2</sub> and HONO emissions on regional scale oxidant formation in the UK. *Atmos. Environ.* **42**(2), 323–336.
- Jenkin, M. E., Cox, R. A. and Williams, D. J. 1988. Laboratory studies of the kinetics of formation of nitrous acid from the thermal reaction of nitrogen dioxide and water vapour. *Atmos. Environ.* (1967) 22(3), 487–498.
- Jorba, O., Dabdub, D., Blaszczak-Boxe, C., Pérez, C., Janjic, Z. and co-authors. 2012. Potential significance of photoexcited NO<sub>2</sub> on global air quality with the NMMB/BSC chemical transport model. *J. Geophys. Res.* 117, D13301. DOI: 10.1029/2012JD017730.
- Kajino, M., Inomata, Y., Sato, K., Ueda, H., Han, Z. and coauthors. 2012. Development of an aerosol chemical transport model RAQM2 and predications of Northeast Asian aerosol mass, size, chemistry, and missing type. *Atmos. Chem. Phys.* 12, 11833–11856.
- Kajino, M., Sato, K., Inomata, Y. and Ueda, H. 2013. Source-receptor relationships of nitrate in Northeast Asia and influence of sea salt on the long-range transport of nitrate. *Atmos. Environ.* 79, 67–78.
- Kleffmann, J., Becker, K. and Wiesen, P. 1998. Heterogeneous NO<sub>2</sub> conversion processes on acid surfaces: possible atmospheric implications. *Atmos. Environ.* 32(16), 2721–2729.

- Kleffmann, J., Gavriloaiei, T., Hofzumahaus, A., Holland, F., Koppmann, R. and co-authors. 2005. Daytime formation of nitrous acid: a major source of OH radicals in a forest. *Geophys. Res. Lett.* 32, L05818. DOI: 10.1029/2005GL022524.
- Kurtenbach, R., Becker, K. H., Gomes, J. A. G., Kleffmann, J., Lörzer, J. C. and co-authors. 2001. Investigations of emissions and heterogeneous formation of HONO in a road traffic tunnel. *Atmos. Environ.* 35, 3385–3394.
- Lam, K. S., Wang, T. J., Wu, C. L. and Li, Y. S. 2005. Study on an ozone episode in hot season in Hong Kong and transboundary air pollution over Pearl River Delta region of China. *Atmos. Environ.* 39, 1967–1977.
- Li, G., Lei, W., Zavala, M., Volkamer, R., Dusanter, S. and coauthors. 2010. Impacts of HONO sources on the photochemistry in Mexico City during the MCMA-2006/MILAGO Campaign. Atmos. Chem. Phys. 10, 6551–6567.
- Li, S., Matthews, J. and Sinha, A. 2008. Atmospheric hydroxyl radical production from electronically excited NO<sub>2</sub> and H<sub>2</sub>O. Science 319, 1657–1660.
- Li, Y., An, J. and Gultepe, I. 2014. Effects of additional HONO sources on visibility over the North China Plain. Adv. Atmos. Sci. 31(5), 1221–1232.
- Li, Y., An, J., Min, M., Zhang, W., Wang, F. and co-author. 2011. Impacts of HONO sources on the air quality in Beijing, Tianjin and Hebei Province of China. Atmos. Environ. 45(27), 4735– 4744
- Meng, Z., Dabdub, D. and Seinfeld, J. H. 1997. Chemical coupling between atmospheric ozone and particulate matter. *Science* 277, 116–119.
- Nguyen, K. and Dabdub, D. 2002. NO<sub>x</sub> and VOC control and its effects on the formation of aerosols. *Aerosol Sci. Tech.* **36**(5), 560–572
- Oswald, R., Behrendt, T., Ermel, M., Wu, D., Su, H. and co-author. 2013. HONO emissions from soil bacteria as a major source of atmospheric reactive nitrogen. *Science* 341, 1233–1235.
- Sarwar, G., Pinder, R. W., Appel, K. W., Mathur, R. and Carlton, A. G. 2009. Examination of the impact of photoexcited NO<sub>2</sub> chemistry on regional air quality. *Atmos. Environ.* 43, 6383– 6387.
- Sarwar, G., Roselle, S. J., Mathur, R., Appel, W., Dennis, R. L. and co-author. 2008. A comparison of CMAQ HONO predictions with observations from the Northeast Oxidant and Particle Study. *Atmos. Environ.* 42(23), 5760–5770.
- Shiraiwa, M., Berkemeier, T., Schilling-Fahnestock, K.A., Seinfeld, J.H. and Pöschl, U. 2014. Molecular corridors and kinetic regimes in the multiphase chemical evolution of secondary organic aerosol. Atmos. Chem. Phys. 14, 1–19.
- Sillman, S. 1995. The use of NO<sub>y</sub>, H<sub>2</sub>O<sub>2</sub>, and HNO<sub>3</sub> as indicators for ozone-NO<sub>x</sub>-hydrocarbon sensitivity in urban locations. J. Geophys. Res. 100(D7), 14175–14188.
- Simpson, D., Guenther, A., Hewitt, C. N. and Steinbrecher, R. 1995. Biogenic emissions in Europe: 1. Estimates and uncertainties. J. Geophys. Res. 100(D11), 22875–22890.
- Spataro, F., Ianniello, A., Esposito, G., Allegrini, I., Zhu, T. and co-author. 2013. Occurrence of atmospheric nitrous acid in the urban area of Beijing (China). Sci. Total Environ. 447, 210–224.

Sun, Y., Wang, Z., Dong, H., Yang, T., Li, J. and co-authors. 2012. Characterization of summer organic and inorganic aerosols in Beijing, China with an aerosol chemical speciation monitor. Atmos. Environ. 51, 250–259.

- Sun, Y., Wang, Z., Fu, P., Yang, T., Jiang, Q. and co-authors. 2013. Aerosol composition, sources and processes during wintertime in Beijing, China. Atmos. Chem. Phys. 13(9), 4577–4592.
- Tang, Y., An, J., Li, Y. and Wang, F. 2014. Uncertainty in the uptake coefficient for HONO formation on soot and its impacts on concentrations of major chemical components in the Beijing-Tianjin-Hebei region. *Atmos. Environ.* 84, 163–171.
- Wang, T., Ding, A., Gao, J. and Wu, W. S. 2006. Strong ozone production in urban plumes from Beijing, China. *Geophys. Res.* Lett. 33, L21806. DOI: 10.1029/2006GL027689.
- Wennberg, P. O. and Dabdub, D. 2008. Rethinking ozone production. Science 319, 1624–1625.
- Wu, N. and Chen, X. 2012. Two photon dissociation dynamics of NO<sub>2</sub> and NO<sub>2</sub>+H<sub>2</sub>O. *J. Phys. Chem. C* **116**(25), 6894–6900.
- Xin, J., Wang, Y., Tang, G., Wang, L., Sun, Y. and co-author. 2010. Variability and reduction of atmospheric pollutants in Beijing and its surrounding area during the Beijing 2008 Olympic Games. *Chinese Sci. Bull.* 55(18), 1937–1944.

- Zaveri, R. A., Easter, R. C., Fast, J. D. and Peters, L. K. 2008. Model for simulating aerosol interactions and chemistry (MOSAIC). J. Geophys. Res. 113, D13204. DOI: 10.1029/ 2007JD008782.
- Zaveri, R. A. and Peters, L. K. 1999. A new lumped structure photochemical mechanism for large-scale applications. *J. Geophys. Res.* 104, D23, 30387–30415.
- Zhang, Q., Streets, D. G., Carmichael, G. R., He, K. B., Huo, H. and co-authors. 2009. Asian emissions in 2006 for the NASA INTEX-B mission. *Atmos. Chem. Phys.* 9, 5131–5153.
- Zhang, R., Sarwar, G., Fung, J. C. H. and Lau, A. K. H. 2013.
  Role of photoexcited nitrogen dioxide chemistry on ozone formation and emission control strategy over the Pearl River Delta, China. *Atmos. Res.* 132–133, 332–344.
- Zhao, X., Zhao, P., Xu, J., Meng, W., Pu, W. and co-authors. 2013. Analysis of a winter regional haze event and its formation mechanism in the North China Plain. *Atmos. Chem. Phys.* 13(11), 5685–5696.
- Zhu, Y., Hiu, W., Xie, P., Dou, K., Liu, S. and co-authors. 2009. Observational study of atmospheric HONO in summer of Beijing. *Environ. Sc.* 30(6), 1567–1573. (in Chinese)