

## LETTER TO THE EDITOR

## Reply to L. Kutzbach

By D. V. KHVOROSTYANOV<sup>1\*</sup>, G. KRINNER<sup>2</sup>, P. CIAIS<sup>3</sup>, M. HEIMANN<sup>4</sup> and S. A. ZIMOV<sup>5</sup>,  
<sup>1</sup>Laboratoire de Météorologie Dynamique, Palaiseau, France; <sup>2</sup>Laboratoire de Glaciologie et Géophysique de  
 l'Environnement, St Martin d'Hères, France; <sup>3</sup>Laboratoire des Sciences du Climat et l'Environnement, Saclay, France;  
<sup>4</sup>Max-Planck Institute of Biogeochemistry, Jena, Germany; <sup>5</sup>Northeast Science Station, Cherskii, Russia

(Manuscript received 24 February 2009; in final form 2 March 2009)

We thank Dr. Kutzbach for providing us with an opportunity to clarify a detail of our model recently presented in (Khvorostyanov et al., 2008, hereafter K08). The comment by Kutzbach (2009) is based on the assumption that the surface heat balance (SHB) in eq. (2) of K08 is a function of the prescribed atmospheric forcing only, thereby neglecting the diffusive ground heat flux  $\Phi_G = -k \partial T / \partial z|_{z=0+}$  (with  $z$  and  $\Phi_G$  being defined positive downwards). However this is not true.

Our model calculates the surface heat balance as:

$$B = S_{\downarrow}(1 - \alpha) + L_{\downarrow} - L_{\uparrow} - \Phi_L - \Phi_S + \Phi_G,$$

where  $S_{\downarrow}$  is the downwelling solar radiation,  $L_{\downarrow}$  and  $L_{\uparrow}$  are down- and upwelling longwave radiation fluxes,  $\Phi_S$  and  $\Phi_L$  are the turbulent sensible and latent heat fluxes, respectively (positive upwards), and  $\Phi_G$  is the ground heat flux (positive downwards).

The ground heat flux  $\Phi_G$  is calculated when solving eq. (1) of K08 for the soil temperature  $T(z, t)$ . In other words, the model is formulated as Kutzbach (2009) suggests it should be, taking into account the  $T|_{z=0}$  increase in case of the soil warming by microorganisms. We should probably have stated this explicitly; to us it went without saying that  $\Phi_G$  is to be taken into account, as is the case for climate models. The ground heat flux  $\Phi_G$  to the atmosphere does not increase by more than  $5 \text{ W m}^{-2}$  during the phase of intense soil warming in Fig. 3 of K08, compared to a simulation without biogenic soil heating. Given that soil warming occurs on very limited spatial scales, and the ground heat flux dissipates quickly in the atmosphere through advection, away from the heated soil patches, this microbial heat-

ing cannot significantly change the atmospheric temperature. At the same time neglecting  $\Phi_G$  in the SHB calculation *does* change the soil heat balance and the soil response to atmospheric warming.

However the effect of neglecting  $\Phi_G$  on the soil warming in our model is the opposite to that anticipated by Kutzbach (2009). In the additional simulation that does not take into account  $\Phi_G$  in the SHB calculation, the soil warming is smaller than in the original simulation. For the case corresponding to Fig. 3 of K08, permafrost thaws only above 2 m, with maximum soil temperatures at the surface layers, due to heat diffusion from the atmosphere in summer. With neglected  $\Phi_G$  the upper soil is a few degrees colder than with the ground heat flux taken into account, and the deep soil remains frozen.

The point is that the heat loss from the soil to the atmosphere is proportional to the vertical temperature gradient  $\frac{\partial T}{\partial z}$  in the upper soil. If the surface gets colder due to the SHB that neglects the ground heat flux (as assumed by Kutzbach, 2009), then  $\frac{\partial T}{\partial z}$  increases due to soil biogenic warming, and the soil loses more heat compared to the original case. Mathematically, taking into account the ground heat flux in the SHB calculation can be simplistically seen as adding another surface-layer resistance  $r_S$  between  $T|_{z=0}$  and the prescribed atmospheric forcing temperature  $T_a$ . The boundary condition at  $z = 0$  is then

$$r_S(T_a - T|_{z=0}) = -k \frac{\partial T}{\partial z}|_{z=0},$$

where  $k$  is the heat conductivity. Adding this additional resistance effectively insulates the soil more than if it were absent and thus confines the internal heating more to the soil and enhances its warming.

To summarize, the SHB in eq. (2) of K08 does take into account the ground heat flux  $\Phi_G$ , while neglecting the latter would

\*Corresponding author.

e-mail: dmitry.khvorostyanov@lmd.polytechnique.fr  
 Institut Pierre-Simon Laplace, Laboratoire de Météorologie  
 Dynamique, Ecole Polytechnique, 91128 Palaiseau, France.  
 DOI: 10.1111/j.1600-0889.2009.00420.x

require a greater atmospheric warming to trigger the intense deep-soil respiration and permafrost thaw.

### References

- Khvorostyanov, D. V., Krinner, G., Ciais, P., Heimann, M. and Zimov, S. A. 2008. Vulnerability of permafrost carbon to global warming. Part I: model description and role of heat generated by organic matter decomposition. *Tellus* **60B**, 250–264.
- Kutzbach, L. 2009. A comment on ‘Vulnerability of permafrost carbon to global warming. Part I: model description and role of heat generated by organic matter decomposition’ by D. V. Khvorostyanov et al. (2008). *Tellus* **61B**, doi:10.1111/j.1600-0889.2009.00419.x.