

# Distribution of sea salt components in snow cover along the traverse route from the coast to Dome Fuji station 1000 km inland at east Dronning Maud Land, Antarctica

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## ABSTRACT

We show and discuss the results of the regional distribution of  $\text{Cl}^-$  and  $\text{Na}^+$  concentrations in snow cover along the study route from the coast to Dome Fuji station, 1000 km inland, in east Dronning Maud Land, Antarctica. The concentrations of  $\text{Cl}^-$  and  $\text{Na}^+$  at the coast exponentially decreased with distance from the coast up to 200 km inland. Between 200–1000 km inland, the concentrations of  $\text{Na}^+$  were nearly constant. On the other hand, the concentrations of  $\text{Cl}^-$  increased with distance beyond 750 km inland. Furthermore, the contribution factors between  $\text{Cl}^-$  and  $\text{Na}^+$  decreased and the  $\text{Cl}^-/\text{Na}^+$  ratio and the  $\text{nssCl}^-/\text{Cl}^-$  ratio increased toward the interior of the continent. These results indicate that sea salt aerosols are transported from the coastal region toward the inland region of the Antarctic continent, and that a source of  $\text{Cl}^-$  other than sea salt exists in the inland plateau of Antarctica.

## 1. Introduction

Antarctic ice sheet records atmospheric deposition continuously in time and space. Therefore, vertical analyses of the ice sheet provide us with information about past climatic change (Delmas, 1992; Legrand and Mayewski, 1997) and horizontal analyses of the ice sheet provide us with knowledge regarding long-range transport of airborne materials (Kamiyama et al., 1989; Kreutz and Mayewski, 1999). However, studies on regional distribution of chemical tracers in the ice sheet with successive snow sampling are very scarce at present. The Japanese Antarctic Research Expeditions (JARE) have conducted many glaciological observations, including the Dome Fuji Project (Dome-F Coring Group, 1998) on the ice sheet in east Dronning Maud Land during recent decades. The

study route from the coast to Dome Fuji station, 1000 km inland, has been established through these observations (Ageta et al., 1989; Kamiyama et al., 1996). Here we show and discuss the results of regional distribution of  $\text{Cl}^-$  and  $\text{Na}^+$  concentrations in snow cover along the route, which were obtained by an inland expedition performed on the 40th JARE (1998–2000).

## 2. Samples and methods

A map of the study route is shown in Fig. 1. The traverse from the coast, S16 (69°02'S, 40°04'E, 591 m a.s.l.), to Dome Fuji station (77°19'S, 39°42'E, 3810 m a.s.l.) was carried out during the austral summer, from 27 December 1998 to 15 February 1999. In this expedition, a sample for ion analyses was taken every 10 km along the route by pushing a 100-ML pre-cleaned plastic bottle into the snow at a depth of 0–2 cm. The study route is approximately perpendicular to the prevailing

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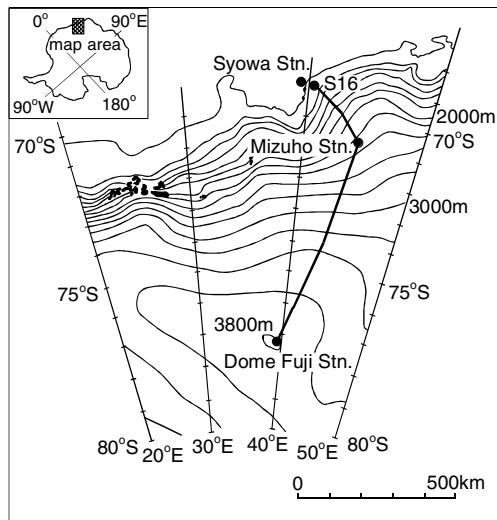


Fig. 1. Map showing the study route in East Dronning Maud Land, Antarctica.

wind direction. The sample collection was performed in the research area located to windward of the route. During the return traverse of the expedition, each sampling site was shifted 5 km further along the route than those of the outward traverse. The concentrations of  $\text{Cl}^-$  and  $\text{Na}^+$  were measured by portable ion chromatography (TOA Electronics, IA-100) at Syowa station ( $69^\circ 00' \text{S}$ ,  $39^\circ 35' \text{E}$ , 29 m a.s.l.) immediately after

the inland expedition. The coefficient of variation of this method is stated to be within 2% in the specification of the apparatus.

### 3. Results and discussion

The concentrations of  $\text{Cl}^-$  and  $\text{Na}^+$  in snow cover along the study route are shown in Fig. 2. Variation of the altitude is also indicated in the figures. There were no remarkable differences between the results of outward and return traverses to and from Dome Fuji station. This implies that the temporal variation of  $\text{Cl}^-$  and  $\text{Na}^+$  concentrations in snow were not significant during the period of the inland expedition. The concentrations of  $\text{Cl}^-$  and  $\text{Na}^+$  at the coast were approximately 0.6 and 0.3  $\text{mg kg}^{-1}$ , respectively. These values decreased by one order of magnitude until about 200 km inland (2.0 km a.s.l.) from S16 point. This result indicates that sea salt aerosols are transported from the coastal region toward the inland region of the Antarctic continent (Kreutz and Mayewski, 1999). Delmas (1992) suggested that sea salt deposition in polar snow is very high in coastal areas but decreases rapidly inland as a function of elevation rather than of the distance from the sea, because the  $\text{Na}^+$  concentrations found at J9 on the Ross Ice Shelf (0.06 km a.s.l.) and at D80 in Adélie Land (2.5 km a.s.l.), both located about 430 km from the ice front, were 100  $\text{ng g}^{-1}$  (Herron and Langway, 1979) and 20  $\text{ng g}^{-1}$  (Legrand

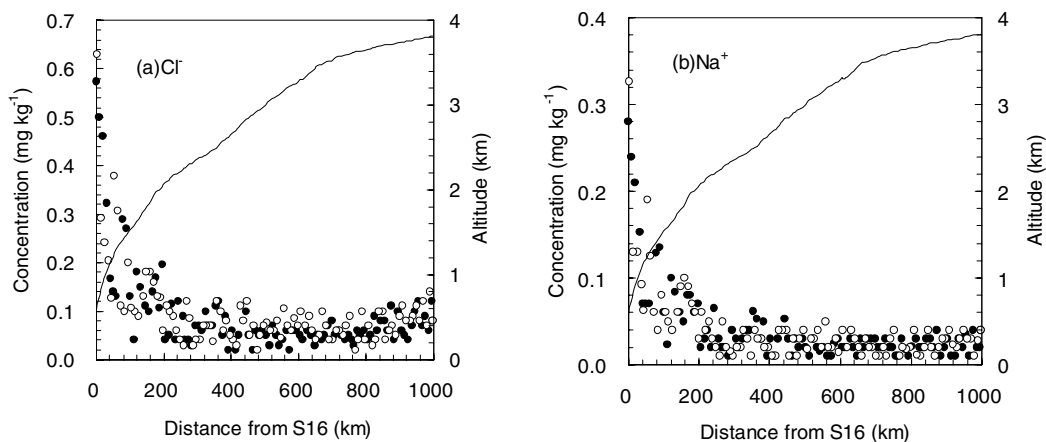


Fig. 2. Distribution of the concentrations of (a)  $\text{Cl}^-$  and (b)  $\text{Na}^+$  in snow cover along the study route. Solid circles and open circles express the data obtained on the outward and return traverses, respectively. Variation in altitude is indicated by the solid line.

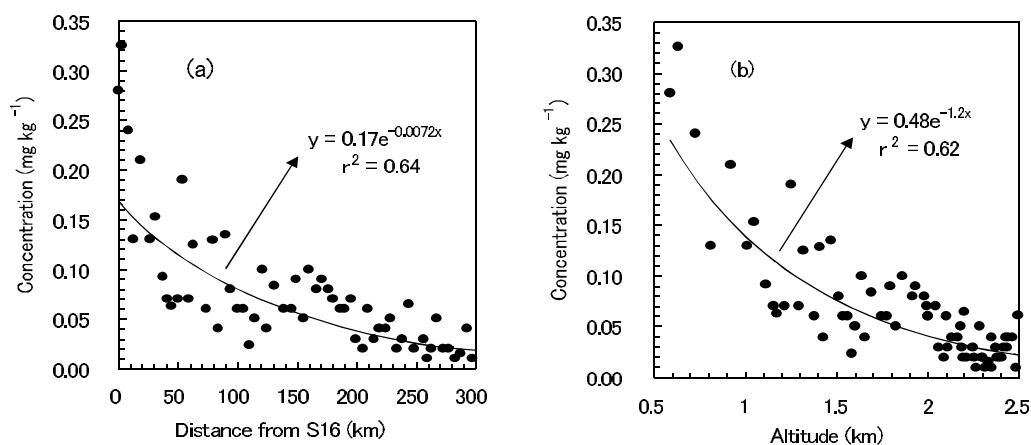


Fig. 3. Concentrations of  $\text{Na}^+$  in snow plotted against (a) the distance from S16 and (b) the altitude. The curves in the figures are obtained by a least-squares method.

and Delmas, 1985), respectively. On the other hand, Minikin et al. (1994) found an exponential decrease of the  $\text{Cl}^-$  concentration on the Filchner–Ronne ice shelf, which is essentially flat. These findings indicate that sea salt concentrations in snow at the coast of Antarctica are dependent both on altitude and distance from the coast. In Fig. 3 the concentrations of  $\text{Na}^+$  in snow are plotted against (a) the distance from S16 up to 300 km and (b) the altitude up to 2.5 km. Regression analyses using an exponential formula were performed on these data sets based on the assumption that sea salt aerosols in the atmosphere are removed to the ice sheet in proportion to their concentration in the atmosphere, i.e. a first-order removal process. The half-decrease distance of sea salt in the ice sheet expected from the exponential term of the equation in Fig. 3a is approximately 100 km. In the same way, from Fig. 3b we can recognize that the concentration of sea salt in the ice sheet decreases to half with increasing altitude every 0.6 km.

From 200 to 1000 km inland, the concentrations of  $\text{Na}^+$  were nearly constant (Fig. 2b). On the other hand, the concentrations of  $\text{Cl}^-$  increased with distance beyond 750 km inland (3.6 km a.s.l.) (Fig. 2a). These results suggest that a source of  $\text{Cl}^-$  other than sea salt exists in the inland plateau of Antarctica. The study route can be divided into three sections on the basis of the characteristics of snow surface features: Section I, a coastal region (0.6–2 km a.s.l.), characterized by a high frequency of small sastrugi and low frequency of dunes; Section II, a katabatic-wind region (2.0–3.6 km a.s.l.) characterized by the coexistence of

Table 1. Sample number, the correlation coefficients between  $\text{Cl}^-$  and  $\text{Na}^+$ , the  $\text{Cl}^-/\text{Na}^+$  and the  $\text{nssCl}^-/\text{Cl}^-$  ratios in three sections of the study route

Section	$n^a$	$r^{2b}$	$\text{Cl}^-/\text{Na}^{+c}$	$\text{nssCl}^-/\text{Cl}^-$
I	39	0.97	$2.0 \pm 0.3$	$0.09 \pm 0.13$
II	108	0.61	$2.6 \pm 1.0$	$0.22 \pm 0.24$
III	50	0.33	$3.4 \pm 1.5$	$0.38 \pm 0.24$

<sup>a</sup>Number of samples.

<sup>b</sup>Square of correlation coefficient between  $\text{Cl}^-$  and  $\text{Na}^+$ .

<sup>c</sup>Errors are  $\pm 1$  standard deviation.

small and large sastrugi, dunes and a glazed surface; and Section III, an inland plateau region (3.6–3.8 km a.s.l.) characterized by low frequencies of small sastrugi and dunes (Furukawa et al., 1996). Number of samples, the correlation coefficients ( $r^2$ ) between  $\text{Cl}^-$  and  $\text{Na}^+$ , the average  $\text{Cl}^-/\text{Na}^+$  ratios and the average  $\text{nssCl}^-/\text{Cl}^-$  ratios in each section are summarized in Table 1. The concentrations of  $\text{nssCl}^-$  (non-sea-salt  $\text{Cl}^-$ ) were obtained as follows:

$$[\text{nssCl}^-] = [\text{Cl}^-] - 1.8[\text{Na}^+] \quad (1)$$

where 1.8 is the  $\text{Cl}^-/\text{Na}^+$  ratio in bulk sea water (Broecker and Peng, 1982), considering that  $\text{Na}^+$  is the best reference element for sea salt. The correlation coefficients between  $\text{Cl}^-$  and  $\text{Na}^+$  were 0.97 in the coastal region, 0.61 in the katabatic-wind region and 0.33 in the inland plateau, respectively. The decrease in the correlation coefficients indicates that the

correlation between the concentration of  $\text{Cl}^-$  and  $\text{Na}^+$  is weakened toward the inland region. In the coastal region, the average value of the  $\text{Cl}^-/\text{Na}^+$  ratio,  $2.0 \pm 0.3$ , was close to the ratio in bulk sea water, 1.8, and the proportion of  $\text{nssCl}^-$  to total  $\text{Cl}^-$  was less than 10%. On the other hand, the  $\text{Cl}^-/\text{Na}^+$  ratio in the inland plateau region,  $3.4 \pm 1.5$ , was approximately twice the value of the ratio in sea water, and roughly 40% of  $\text{Cl}^-$  in this section was the non-sea-salt fraction. Not only the average values of  $\text{Cl}^-/\text{Na}^+$  and  $\text{nssCl}^-/\text{Cl}^-$ , but also their standard deviations, increased toward the interior of the continent. These results indicate that the intensity and instability of unknown source of  $\text{Cl}^-$  become strong toward the inland plateau. The increase in the concentration of  $\text{Cl}^-$  in the snow inland may be due to an inflow of  $\text{nssCl}^-$  into the Antarctic boundary layer from the stratosphere and the upper troposphere, along with atmospheric convergent flow in the interior of the Antarctic continent (James, 1989; Parish and Bromwich, 1991). Although we cannot identify the origin of  $\text{nssCl}^-$  in this study, the anthropogenic  $\text{Cl}^-$  which is injected into the upper atmosphere in the mid-latitude region and/or naturally occurring  $\text{Cl}^-$  which is produced by the chemical reaction in the Antarctic atmosphere, e.g. HCl production from sea salt aerosol as suggested by Legrand and Delmas (1988), are possibilities. Wagon et al. (1999) clarified that volatile chemical species in the upper snow layers of the inland region suffer very serious post-depositional change. If the influence of post-depositional migration and re-deposition process increases with altitude or distance from the coast, the enrichment of  $\text{Cl}^-$  in the surface snow by this phenomenon will be important to the interpretation of our results. However, Kamiyama and Watanabe (1994) reported that not only the concentrations of  $\text{Cl}^-$  and  $\text{NO}_3^-$  but also the concentration of  $\text{SO}_4^{2-}$ , non-volatile species, in the surface snow have also increased on the inland plateau of east Dronning Maud Land. Further observations of the regional distribution of the upper atmospheric tracers and the extent of post-depositional migration of the volatile species

are needed to identify the source of  $\text{nssCl}^-$  in the interior of Antarctica.

#### 4. Conclusions

We have shown the regional distribution of  $\text{Cl}^-$  and  $\text{Na}^+$  concentrations in snow cover along the study route from S16 to Dome Fuji station in east Dronning Maud Land, Antarctica, and obtained the following results.

(1) The concentrations of  $\text{Cl}^-$  and  $\text{Na}^+$  at the coast rapidly decreased up to the point of 200 km inland (2.0 km a.s.l.). This result may be due to the fact that sea salt aerosols are transported from the coast toward the interior of the continent and are removed onto the Antarctic ice sheet.

(2) Between 200 and 1000 km inland, the concentrations of  $\text{Na}^+$  were nearly constant. On the other hand, the concentrations of  $\text{Cl}^-$  increased with distance beyond 750 km inland. Furthermore, the correlation between  $\text{Cl}^-$  and  $\text{Na}^+$  decreased and the  $\text{Cl}^-/\text{Na}^+$  and the  $\text{nssCl}^-/\text{Cl}^-$  ratios increased toward the interior of the continent. These results suggest that a source of  $\text{Cl}^-$  other than sea salt exists in the inland plateau of Antarctica.

(3) The concentrations of  $\text{Na}^+$  in snow cover in the ice sheet within the coastal region decreased to half with every 100 km distance from the coast and 0.6 km altitude.

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#### REFERENCES

- Ageta, Y., Kamiyama, K., Okuhira, F. and Fujii, Y. 1989. Geomorphological and glaciological aspects around the highest dome in Queen Maud Land, East Antarctica. *Proc. NIPR Symp. Polar Meteorol. Glaciol.* **2**, 88–96.
- Broecker, W. S. and Peng, T.-H. 1982. *Tracers in the sea*. Eldigio Press, Palisades, New York, 26–27.
- Delmas, R. J. 1992. Environmental information from ice core. *Rev. Geophys.* **30**, 1–21.
- Dome-F Coring Group. 1998. Deep ice-core drilling at Dome Fuji and glaciological studies in east Dronning Maud Land, Antarctica. *Ann. Glaciol.* **27**, 333–337.

- Furukawa, T., Kamiyama, K. and Maeno, H. 1996. Snow surface features along the traverse route from the coast to Dome Fuji Station, Queen Maud Land, Antarctica. *Proc. NIPR Symp. Polar Meteorol. Glaciol.* **10**, 13–24.
- Herron, M. M. and Langway, C. C. Jr., 1979. Dating of Ross Ice Shelf cores by chemical analysis. *J. Glaciol.* **24**, 345–357.
- James, I. N. 1989. The Antarctic drainage flow: implications for hemispheric flow on the Southern Hemisphere. *Antarc. Sci.* **1**, 279–290.
- Kamiyama, K., Ageta, Y. and Fujii, Y. 1989. Atmospheric and depositional environments traced from unique chemical compositions of the snow over an inland high plateau, Antarctica. *J. Geophys. Res.* **94**, 18515–18519.
- Kamiyama, K., Motoyama, H., Fujii, Y. and Watanabe, O. 1996. Distribution of hydrogen peroxide in surface snow over Antarctic ice sheet. *Atmos. Environ.* **30**, 967–972.
- Kamiyama, K. and Watanabe, O. 1994. Substances deposited on an inland plateau, Antarctica. *Nankyoku Shiryo (Antarctic Record)* **38**, 232–242 (in Japanese with English abstract).
- Kreutz, K. J. and Mayewski, P. A. 1999. Spatial variability of Antarctic surface snow glaciochemistry: implications for palaeoatmospheric circulation reconstructions. *Antarc. Sci.* **35**, 105–118.
- Legrand, M. and Delmas, R. J. 1985. Spatial and temporal variations of snow chemistry in Terre Adélie (Antarctica). *Ann. Glaciol.* **7**, 20–25.
- Legrand, M. and Delmas, R. J. 1988. Formation of HCl in the Antarctic atmosphere. *J. Geophys. Res.* **93**, 7153–7168.
- Legrand, M. and Mayewski, P. A. 1997. Glaciochemistry of polar ice cores: a review. *Rev. Geophys.* **35**, 219–243.
- Minikin, A., Wagenbach, D., Graf, W. and Kipfstuhl, J. 1994. Spatial and seasonal variations of the snow chemistry at the central Filchner–Ronne Ice Shelf, Antarctica. *Ann. Glaciol.* **20**, 283–290.
- Parish, T. H. and Bromwich, D. H. 1991. Continental-scale simulation of the Antarctic katabatic wind regime. *J. Climate* **4**, 135–146.
- Wagnon, P., Delmas, R. J. and Legrand, M. 1999. Loss of volatile acid species from upper firn layers at Vostok, Antarctica. *J. Geophys. Res.* **104**, 3423–3431.