

"All charged up"

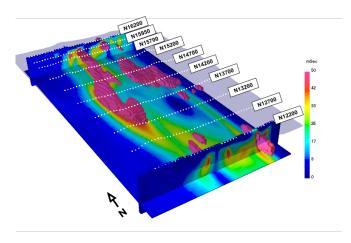
Advances and applications for IP surveys



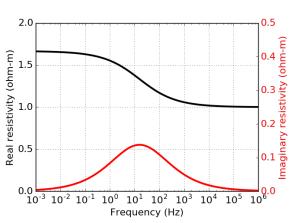
Douglas Oldenburg and Seogi Kang

Motivation

Minerals



Complex resistivity



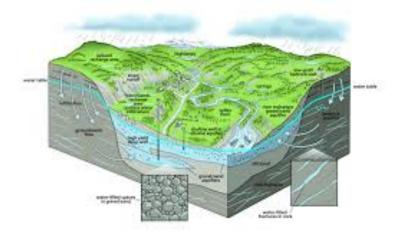
Permafrost



Geotechnical

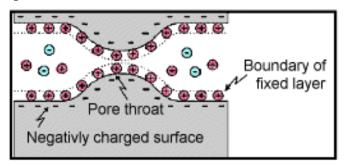


Groundwater

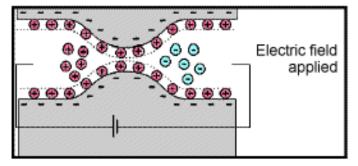


Chargeability

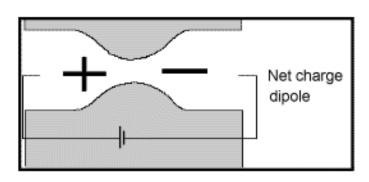
Initially - neutral



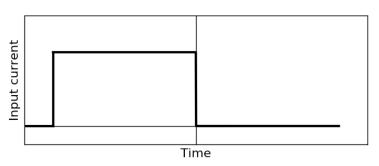
Apply electric field, build up charges



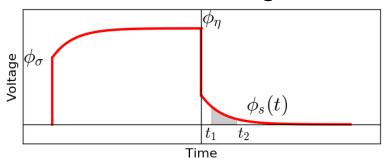
Charge polarization, Electric dipole



Input current

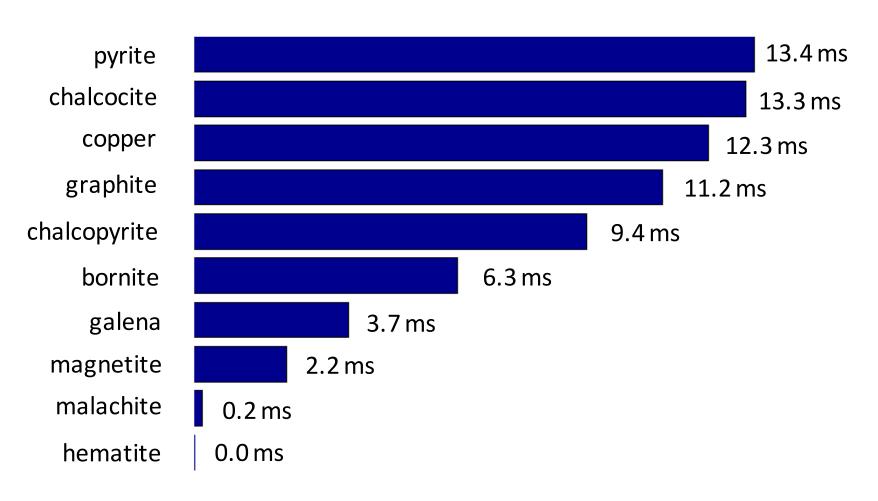


Measured voltage



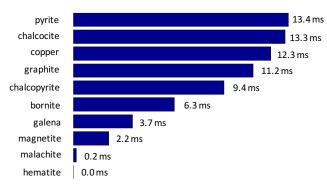
Chargeability

Minerals at 1% Concentration in Samples



Chargeability

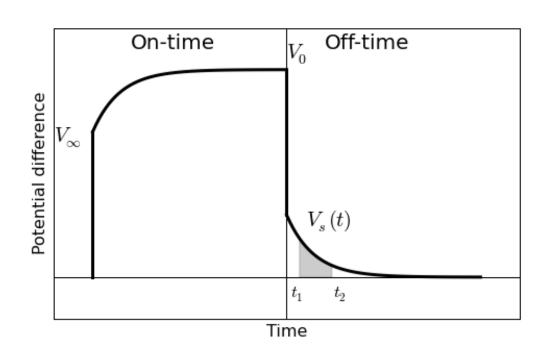


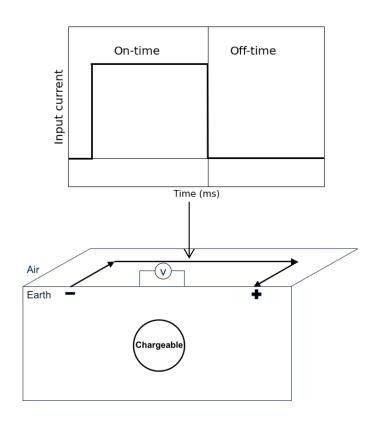


Material type	Chargeability (msec.)
20% sulfides	2000 - 3000
8-20% sulfides	1000 - 2000
2-8% sulfides	500 - 1000
volcanic tuffs	300 - 800
sandstone, siltstone	100 - 500
dense volcanic rocks	100 - 500
shale	50 - 100
granite, granodiorite	10 - 50
limestone, dolomite	10 - 20

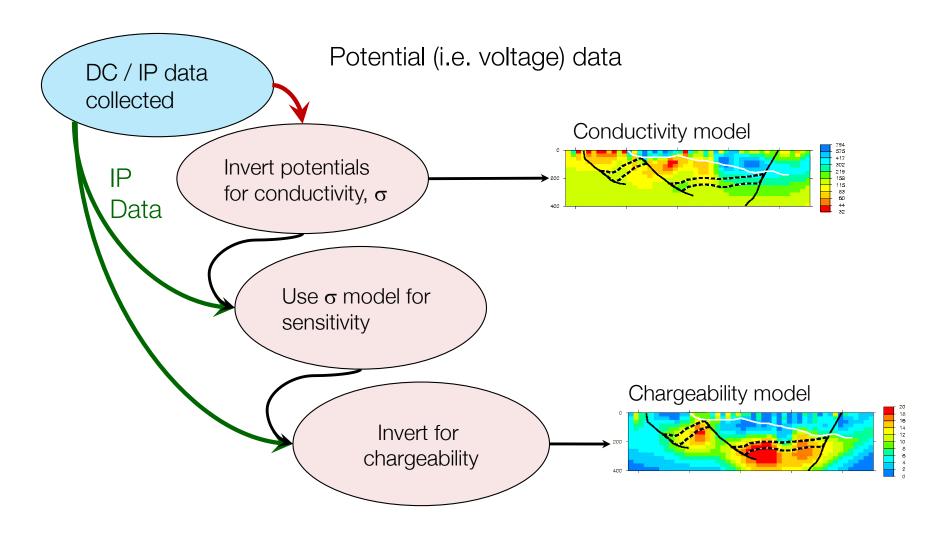
Material type	Chargeability (msec.)
ground water	0
alluvium	1 - 4
gravels	3 - 9
precambrian volcanics	8 - 20
precambrian gneisses	6 - 30
schists	5 - 20
sandstones	3 - 12

Overvoltage

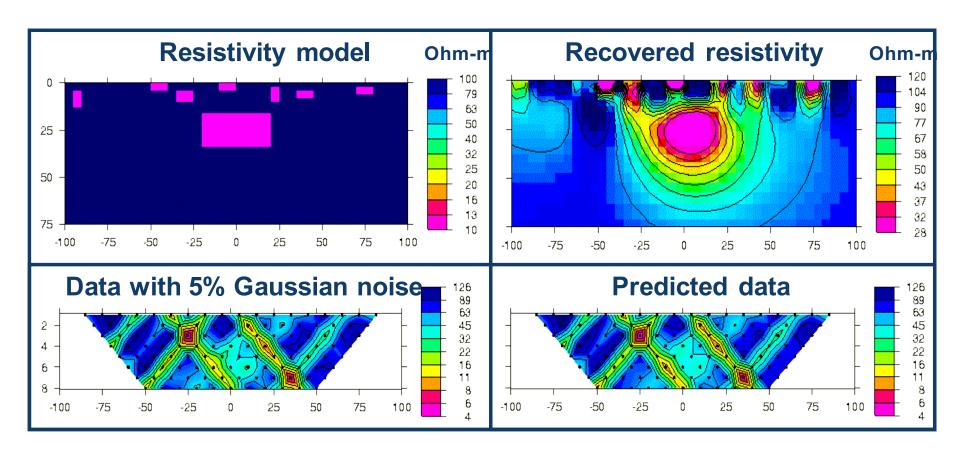




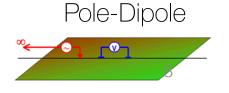
IP Inversion



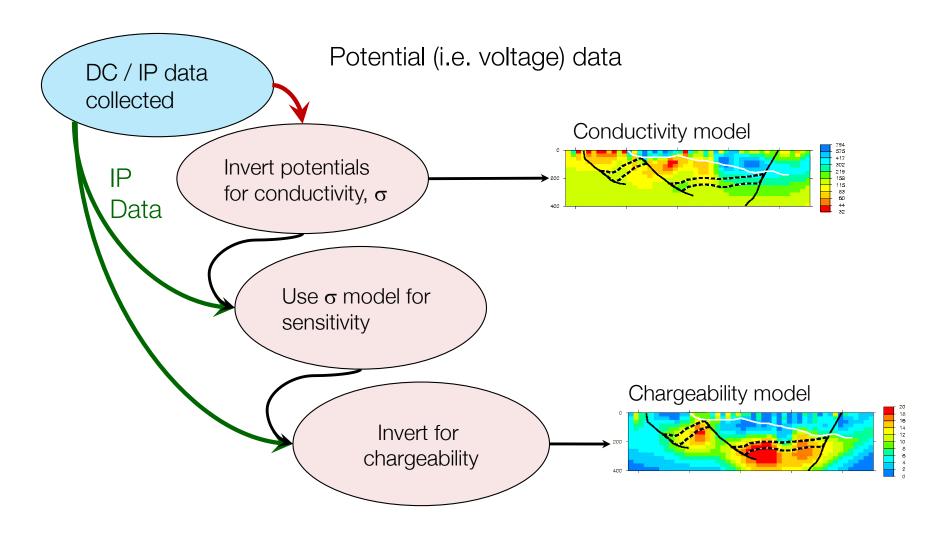
Example DC: prism with geologic noise



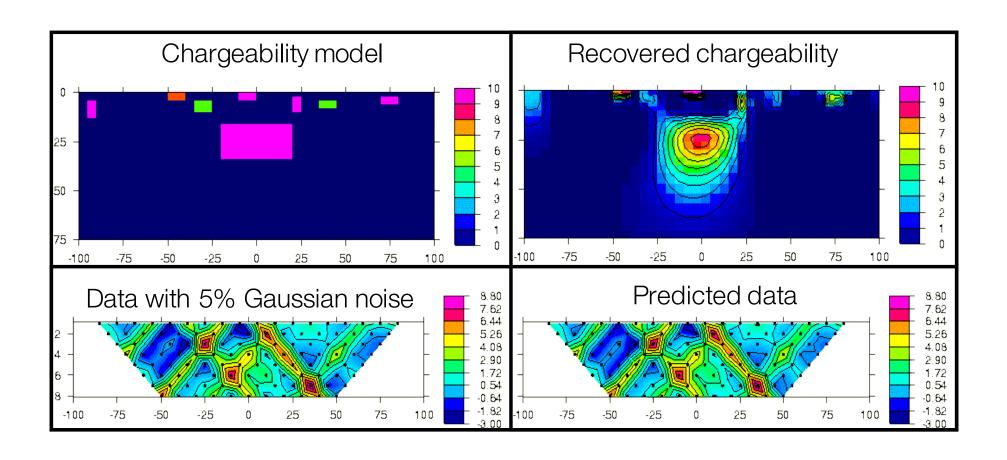
• Pole-dipole; n=1,8; a=10m; N=316; $(\alpha_s, \alpha_x, \alpha_z)$ =(.001, 1.0, 1.0)



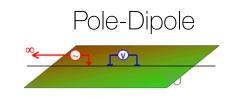
IP Inversion



Example IP: prism with geologic noise

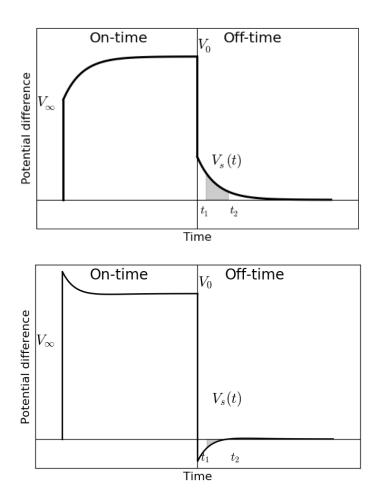


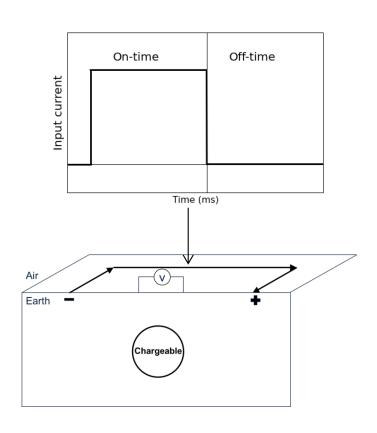
• Pole-dipole; n=1,8; a=10m; N=316; $(\alpha_s, \alpha_x, \alpha_z)$ =(.001, 1.0, 1.0)



Challenge: EM coupling, grounded sources

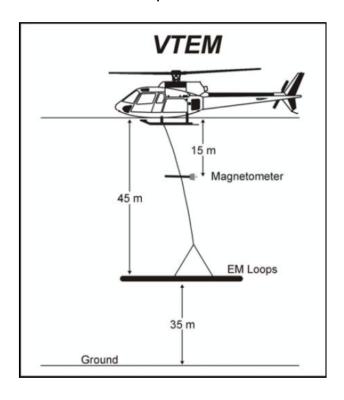
DC-IP: overvoltage



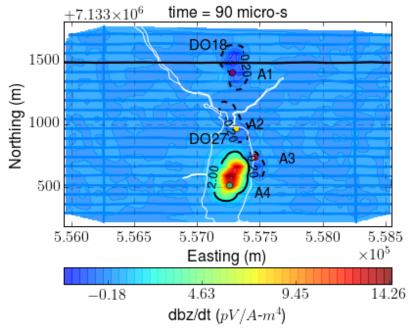


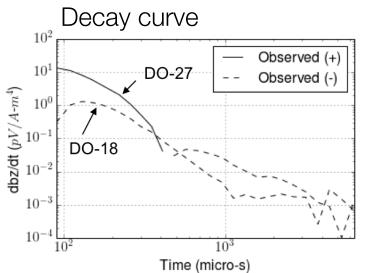
Challenge: inductive sources

VTEM set-up



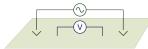
At 90 micro-s



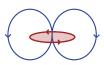


Outline

- Background
- Grounded IP



- How to remove EM contamination?
- How to extract valuable conductivity?
- Gradient array example
- Inductive source IP
 - Revisit physics
 - IP inversion workflow
 - Tli Kwi Cho kimberlites



Simulation of TEM data

Maxwell's equations:

Frequency domain

$$\vec{\nabla} \times \vec{E} = -\imath \omega \vec{B}$$

$$\vec{\nabla} \times \mu^{-1} \vec{B} - \vec{J} = \vec{J}_s$$

Ohm's law in **frequency** domain

$$\vec{J} = \sigma \vec{E}$$

Time domain

$$\vec{\nabla} \times \vec{e} = -\frac{\partial \vec{b}}{\partial t}$$

$$\vec{\nabla} \times \mu^{-1} \vec{b} - \vec{j} = \vec{j}_s$$

Ohm's law in time domain

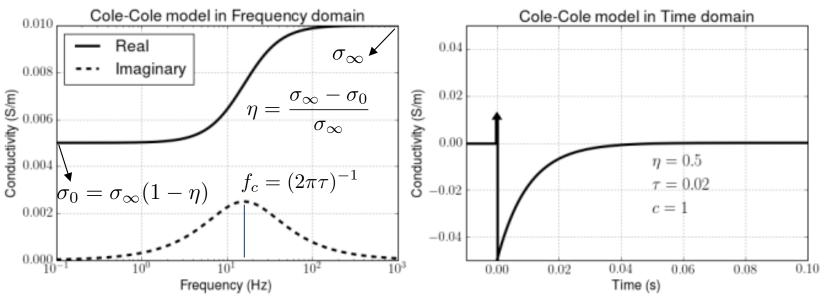
$$\vec{j} = \sigma \vec{e}$$

Complex conductivity

Cole-Cole model (Pelton et al., 1978)

Frequency domain

Time domain



Inverse Fourier transform

 σ_{∞} : Conductivity at infinite frequency

 σ_0 : Conductivity at zero frequency

 η : Chargeability

 τ : Time constant (s)

c: Frequency dependency

$$\sigma(\omega) = \sigma_{\infty} + \sigma_{\infty} \frac{\eta}{1 + (1 - \eta)(\imath \omega \tau)^{c}} \longrightarrow \mathcal{F}^{-1}[\sigma(\omega)] = \sigma(t)$$

Simulation of TEM data with IP

Maxwell's equations:

EMTDIP code (Marchant et al., 2015)

Frequency domain

$$\vec{\nabla} \times \vec{E} = -\imath \omega \vec{B}$$

$$\vec{\nabla} \times \mu^{-1} \vec{B} - \vec{J} = \vec{J}_s$$

Ohm's law in **frequency** domain

$$\vec{J} = \sigma(\omega)\vec{E}$$

Time domain

$$\vec{\nabla} \times \vec{e} = -\frac{\partial \vec{b}}{\partial t}$$

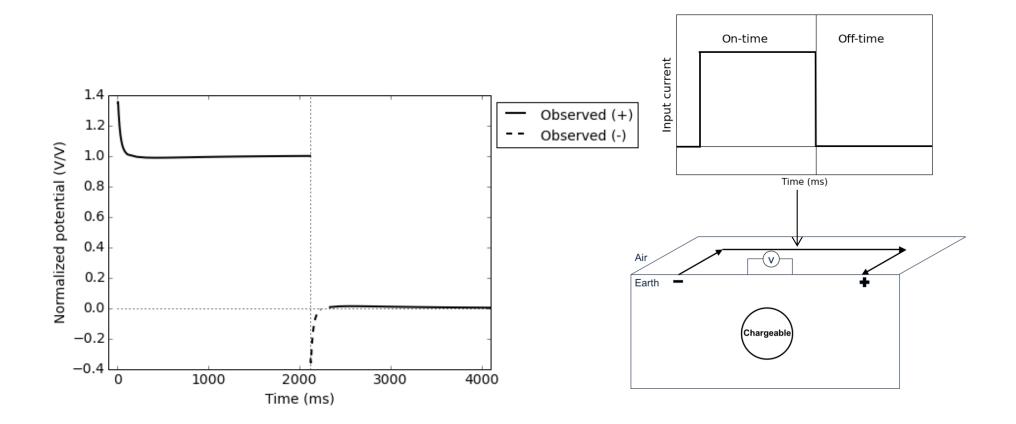
$$\vec{\nabla} \times \mu^{-1} \vec{b} - \vec{j} = \vec{j}_s$$

Ohm's law in time domain

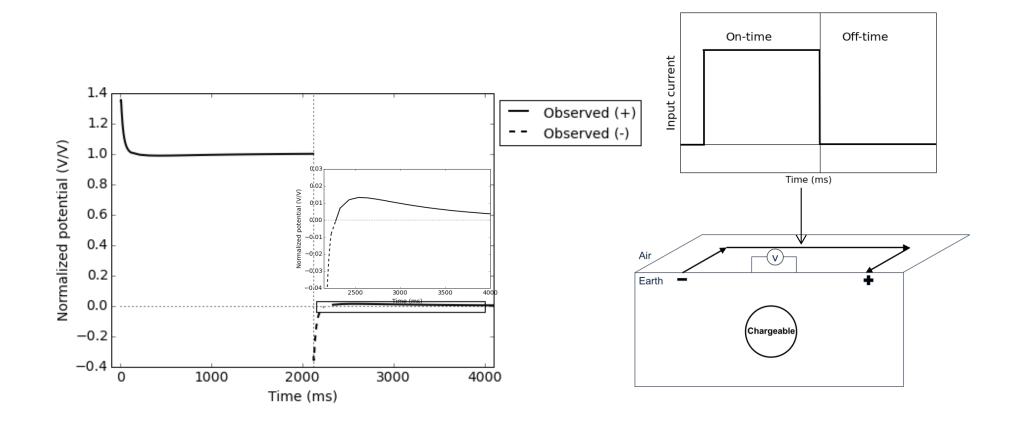
$$\vec{j} = \sigma(t) \otimes \vec{e}(t)$$

 $\otimes : convolution$

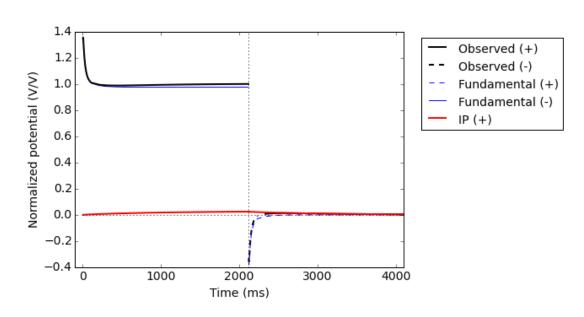
Observed response

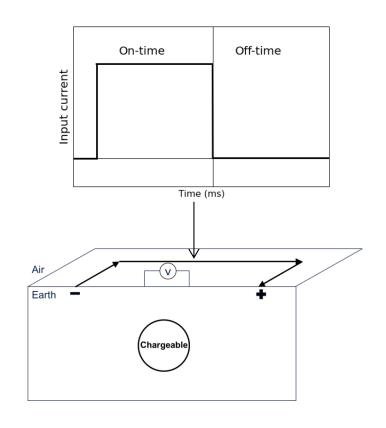


Observed response



Define IP datum





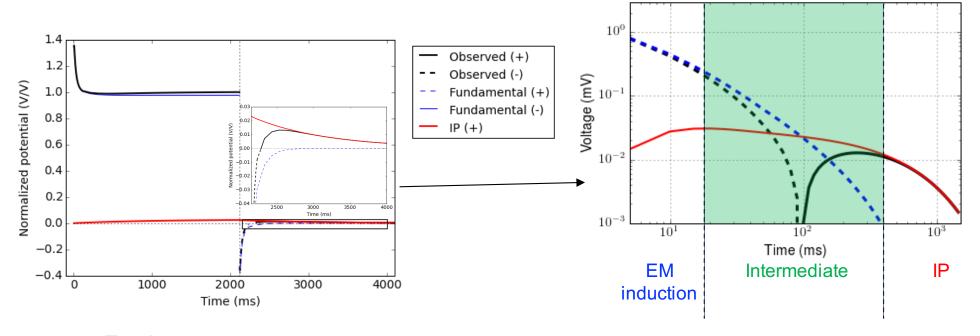
• IP datum:

IP = Observation - Fundamental
$$d^{IP}(t) = F[\sigma(t)] - F[\sigma_{\infty}]$$

 $F[\sigma(t)]$ Observation $F[\sigma_{\infty}]$ Fundamental

 $F[\cdot]$: Maxwell's operator

Define IP datum



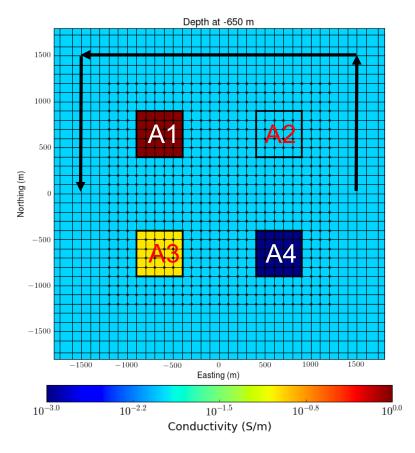
IP datum:

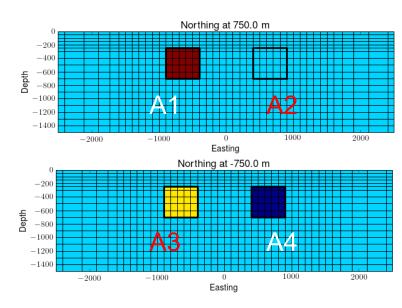
IP = Observation - Fundamental
$$d^{IP}(t) = F[\sigma(t)] - F[\sigma_{\infty}]$$

 $F[\sigma(t)]$ Observation $F[\sigma_{\infty}]$ Fundamental

 $F[\cdot]$: Maxwell's operator

Synthetic model





Conductivity at Infinite frequency:

halfspace =
$$0.01 \text{ S/m}$$

$$A1 = 1 \text{ S/m}$$

$$A2 = 0.01 \text{ S/m}$$

$$A3 = 0.1 \text{ S/m}$$

$$A4 = 0.001 \text{ S/m}$$

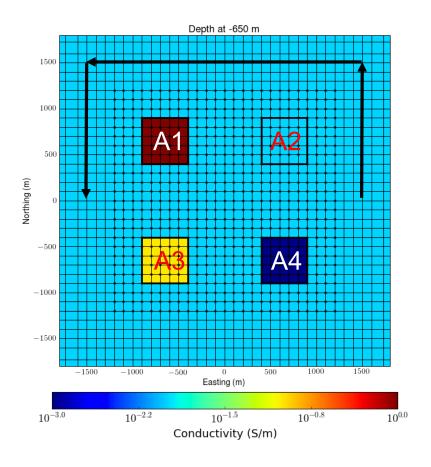
Chargeable objects: A2 and A3

$$\eta = 0.1$$

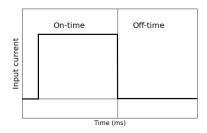
$$\tau = 0.5 \text{ s}$$

$$c = 1$$

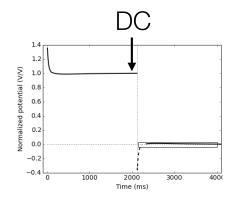
Forward modelling set up



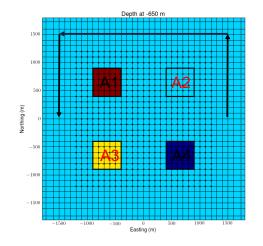
- Measure potential difference
 - 200 m bi-pole (625 mid points)
- Step-off waveform:

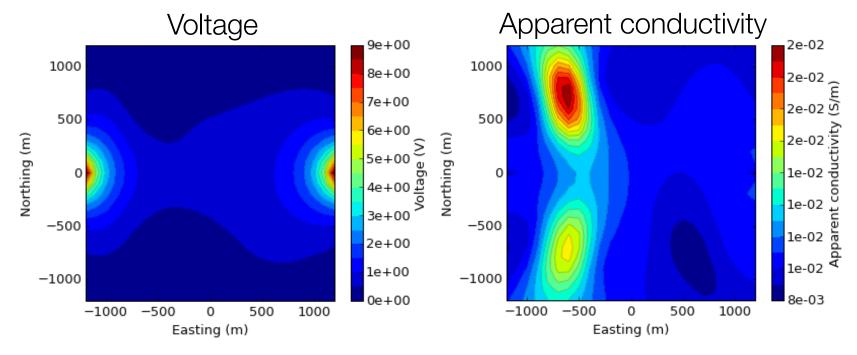


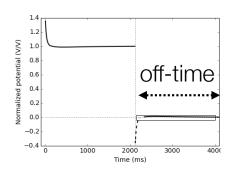
- Time range: 1 600 ms (off-time)
 - Chargeable objects: A2 and A3

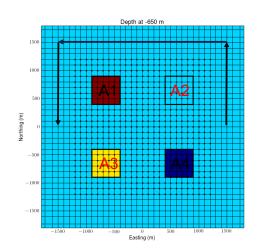


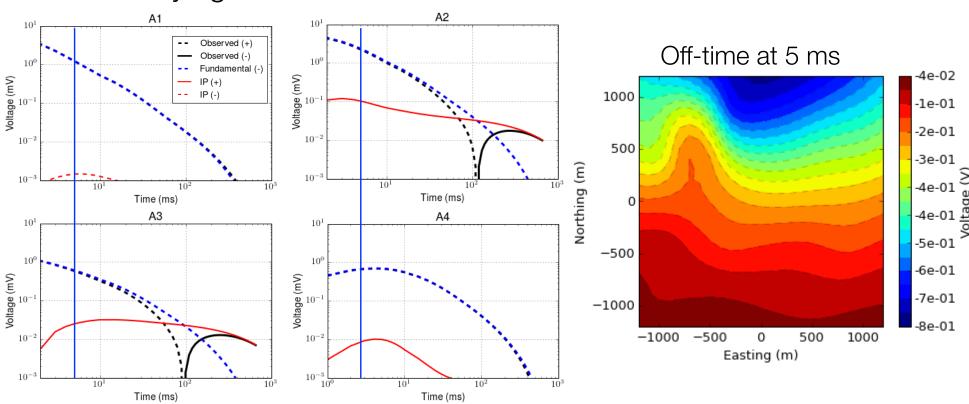
Observed DC data

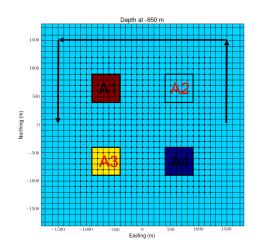


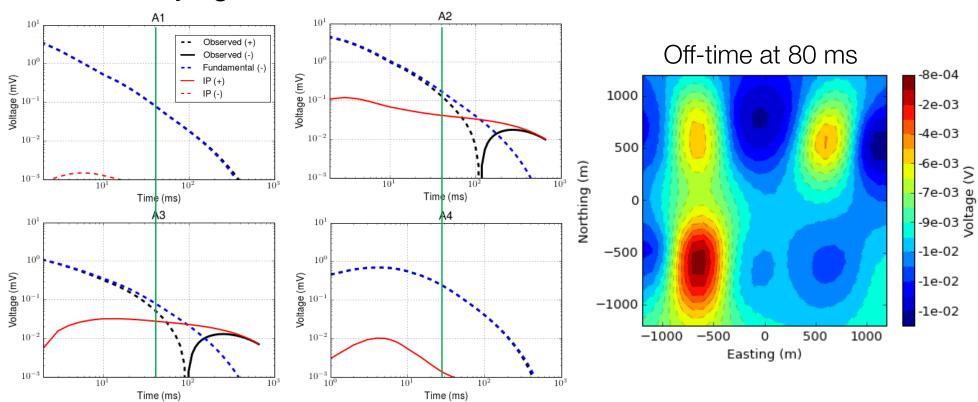


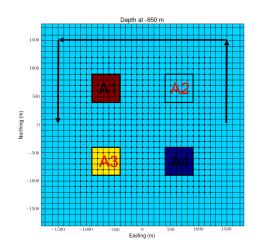


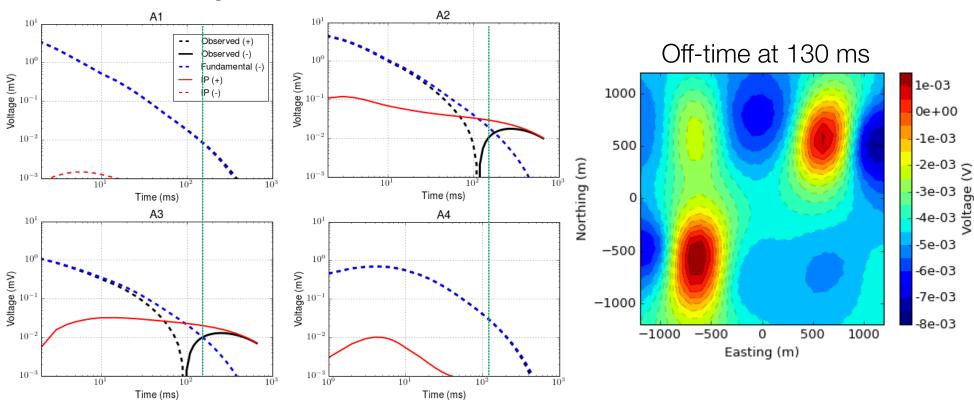


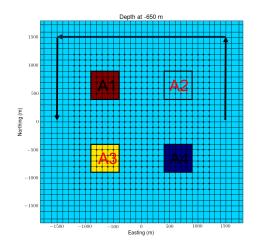


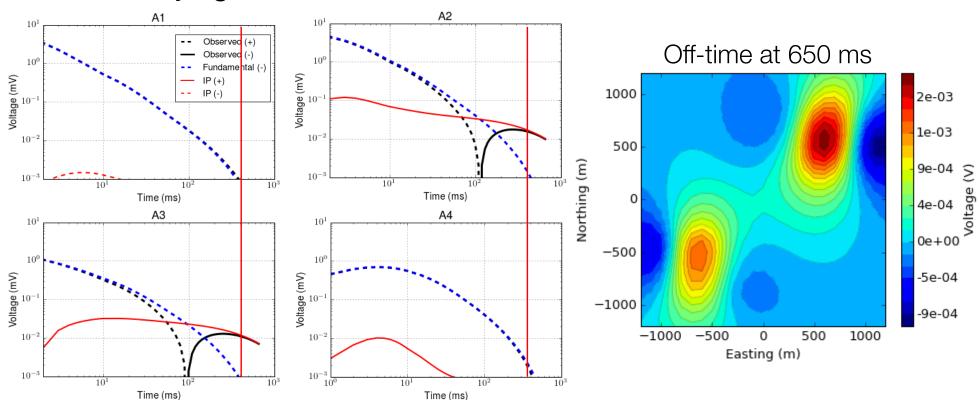




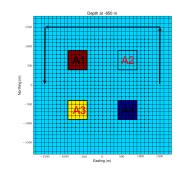




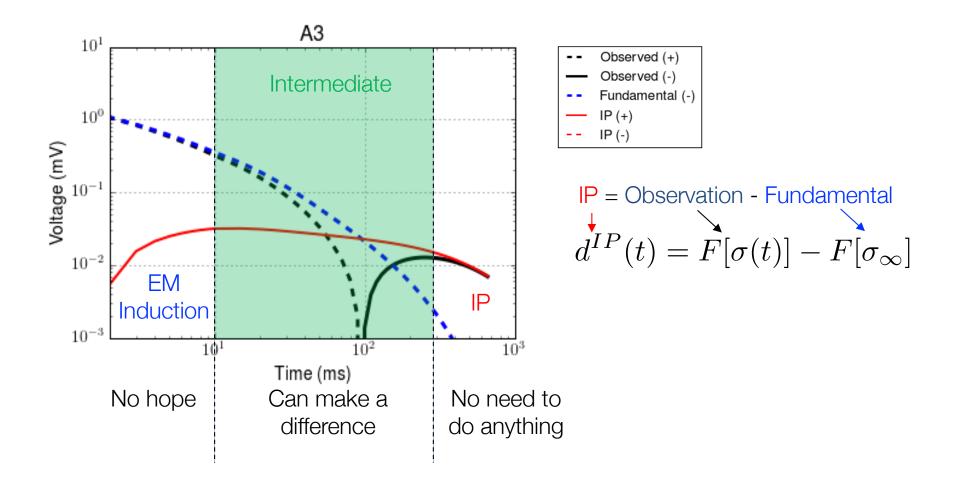




EM decoupling



Time decaying curves (off-time)



EM decoupling: true σ_{∞}

No hope

A3

Intermediate

IP

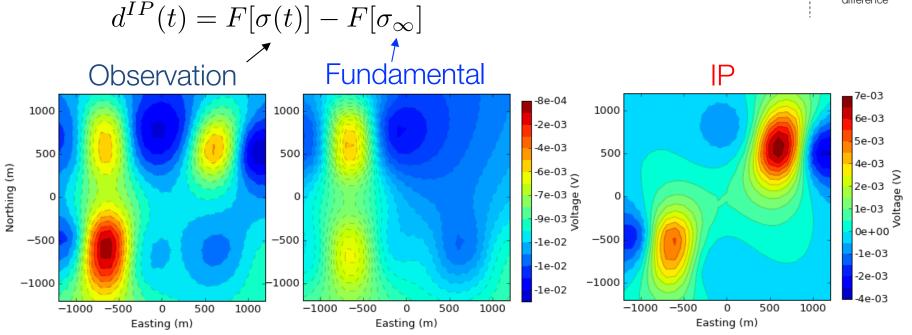
Induction

IO

Time (ms)

Can make a do anything

Off-time at 80 ms

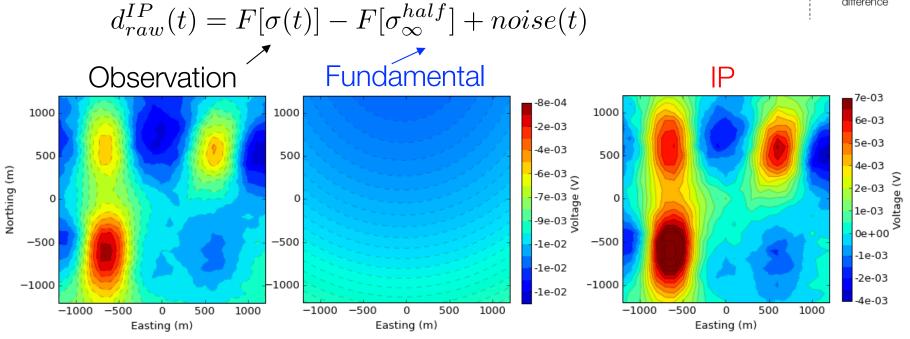


EM decoupling: σ_{∞}^{half}

Off-time at 80 ms

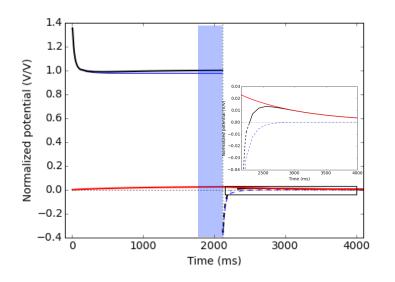
Intermediate

In

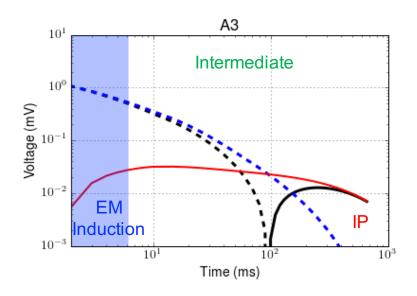


How do we estimate conductivity, σ_{∞} ?

Late on-time data (DC)



Early off-time data (TEM)



3D inversion methodology

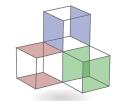
• Data misfit:

$$\phi_d = \|\mathbf{W_d}(\mathbf{Am} - d^{obs})\|_2^2$$

Model objective function:

$$\phi_m = \|\mathbf{W}_m(\mathbf{m} - \mathbf{m}_{ref})\|_2^2$$

Tikhonov inversion: minimize



$$\phi = \phi_d(\mathbf{m}) + \beta \phi_d(\mathbf{m})$$



Depth weight:

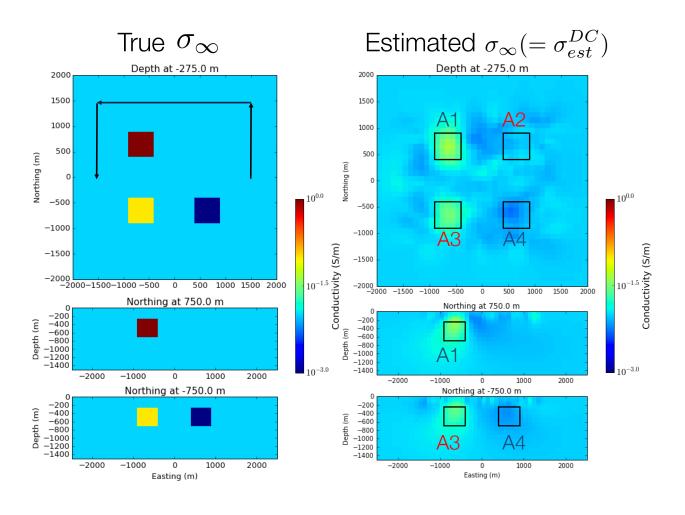
$$\frac{1}{(z-z_0)^3}$$

DC-IP inversion: SimPEG-DCIP

TEM inversion: UBC-H3DTD code

3D DC inversion

Recover 3D conductivity



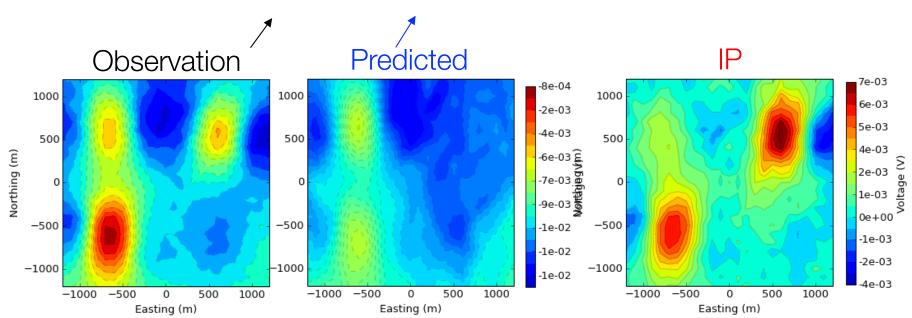
- Depth weighting
 - Compensate for high sensitivity near surface

$$\frac{1}{(z-z_0)^3}$$

EM decoupling: σ_{est}^{DC}

Off-time at 130 ms

$$d_{raw}^{IP}(t) = F[\sigma(t)] - F[\sigma_{est}^{DC}] + noise(t)$$



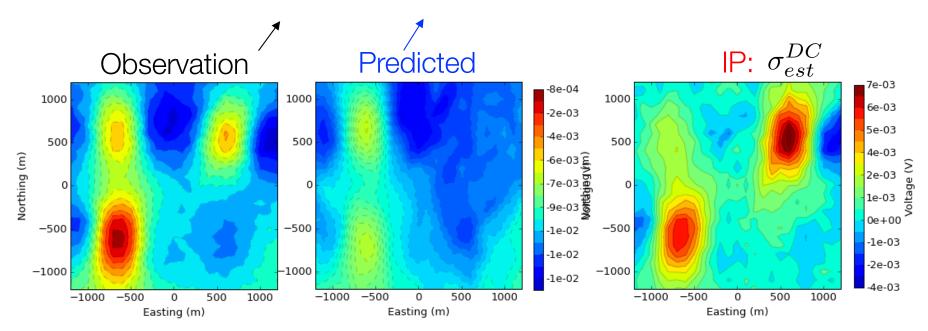
EM decoupling: σ_{est}^{DC}

IP: Half-space

1000
500
-1000
-1000
Easting (m)

Off-time at 130 ms

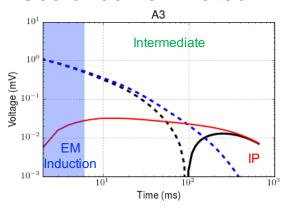
$$d_{raw}^{IP}(t) = F[\sigma(t)] - F[\sigma_{est}^{DC}] + noise(t)$$



3D TEM inversion

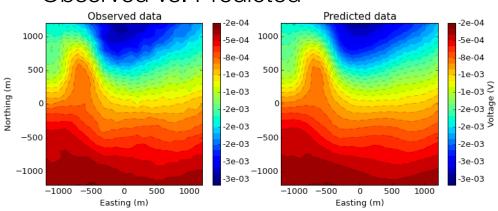
Recover 3D conductivity

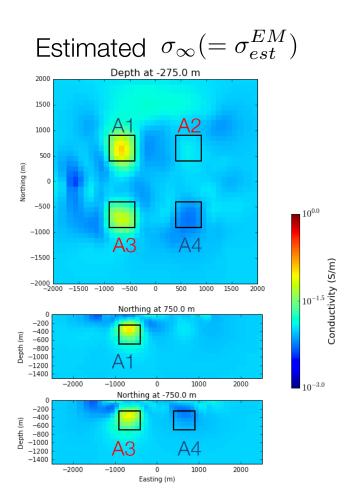
Use uncontaminated EM data



Time range: 1-6 ms (6 channels)

Observed vs. Predicted

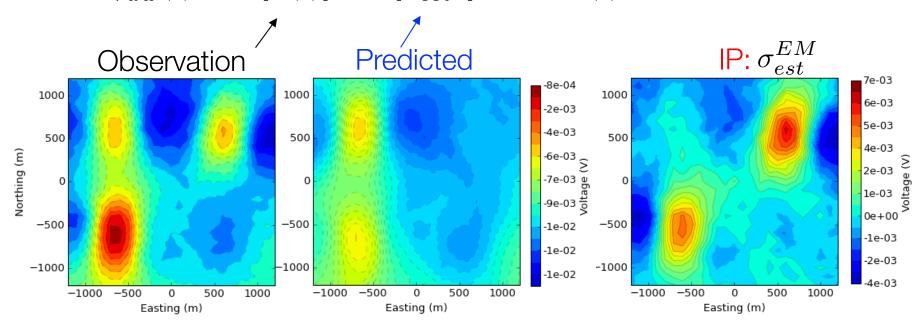




EM decoupling: σ_{est}^{EM}

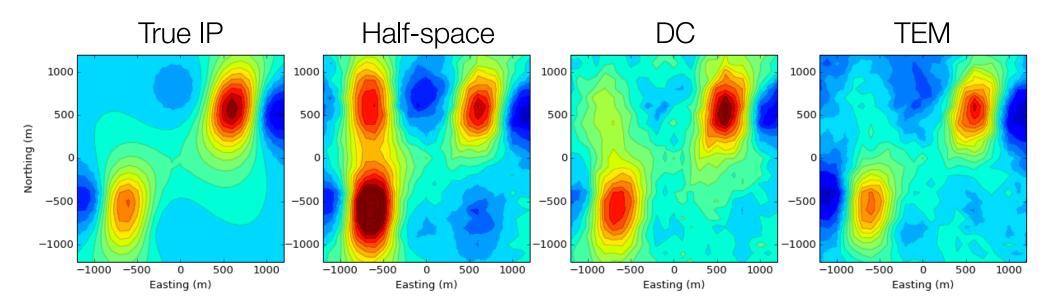
Off-time at 80 ms

$$d_{raw}^{IP}(t) = F[\sigma(t)] - F[\sigma_{est}^{EM}] + noise(t)$$



Comparisons of IP

IP data at 80 ms



IP inversion workflow

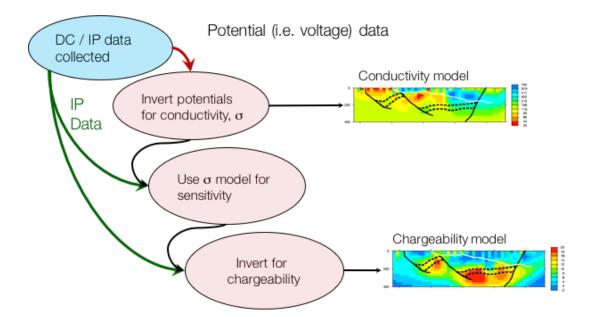
Workflow

Invert TEM data, to recover σ_{∞}

Compute IP datum Remove EM responses

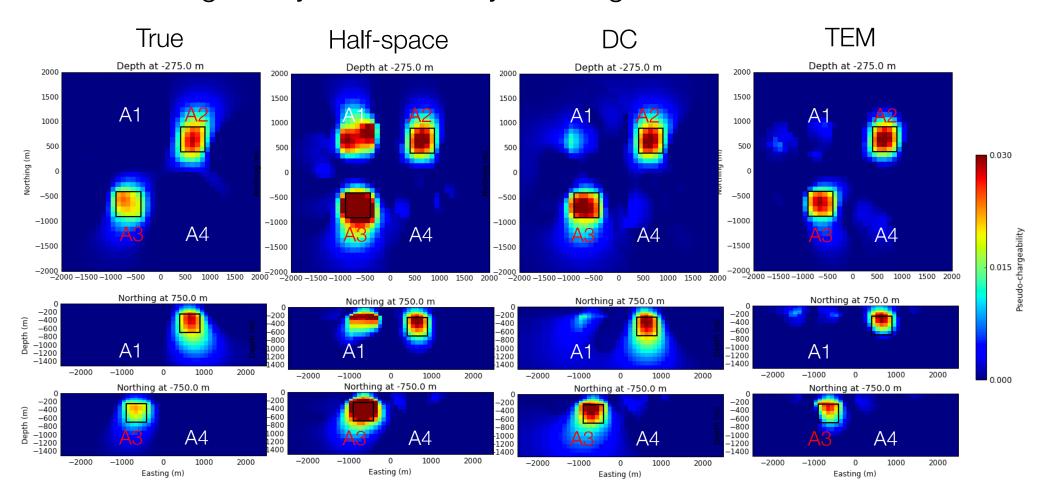
Linearized equations

Invert d^{IP} data, recover pseudo-chargeability



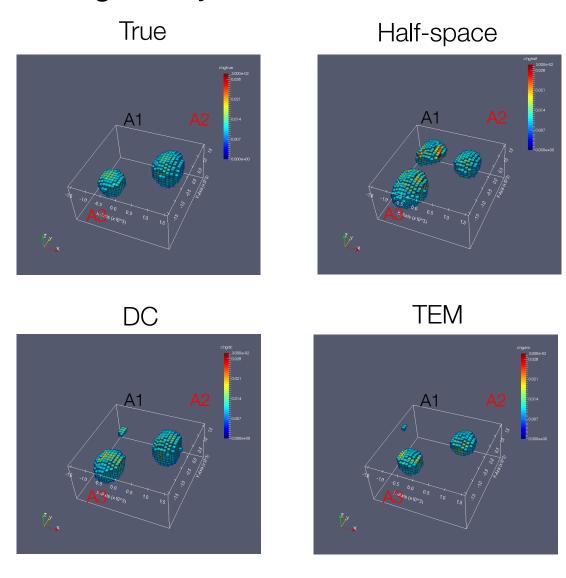
3D IP inversion

Chargeability: recovered by inverting:



3D cut-off volume

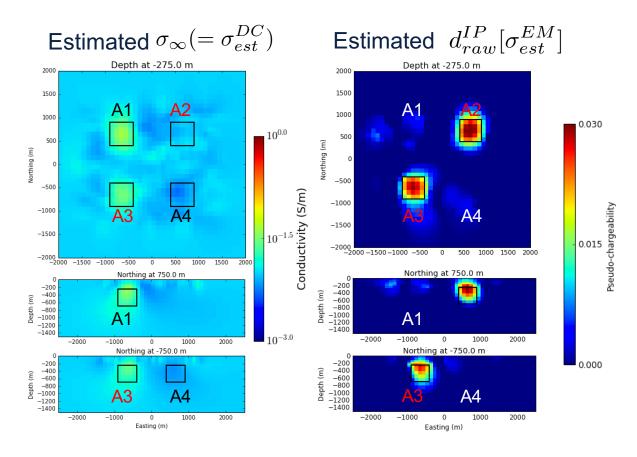
Pseudo-chargeability > 0.015



Take home

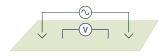
Traditionally, early time TEM data has been discarded

By using these discarded TEM signals we can better estimate both 3D conductivity and chargeability

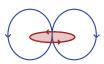


Outline

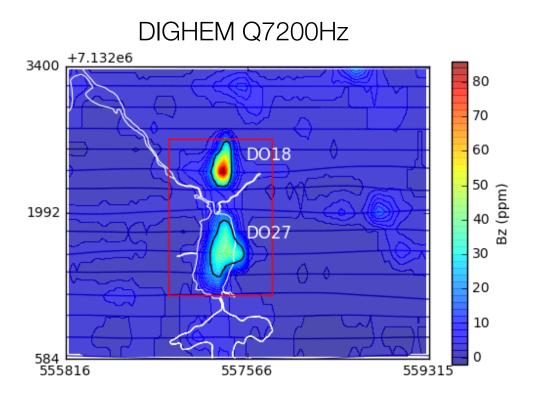
- Backgrounds
- TEM-IP inversion workflow
- Galvanic source IP
 - Synthetic example: gradient array



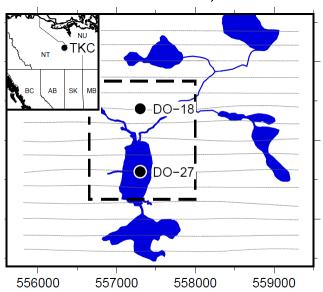
- Inductive source IP
 - Field example: Tli Kwi Cho kimberlites



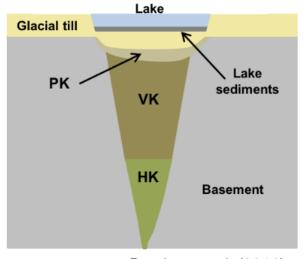
Discovery of Tli Kwi Cho (TKC)



Location of TKC, NWT

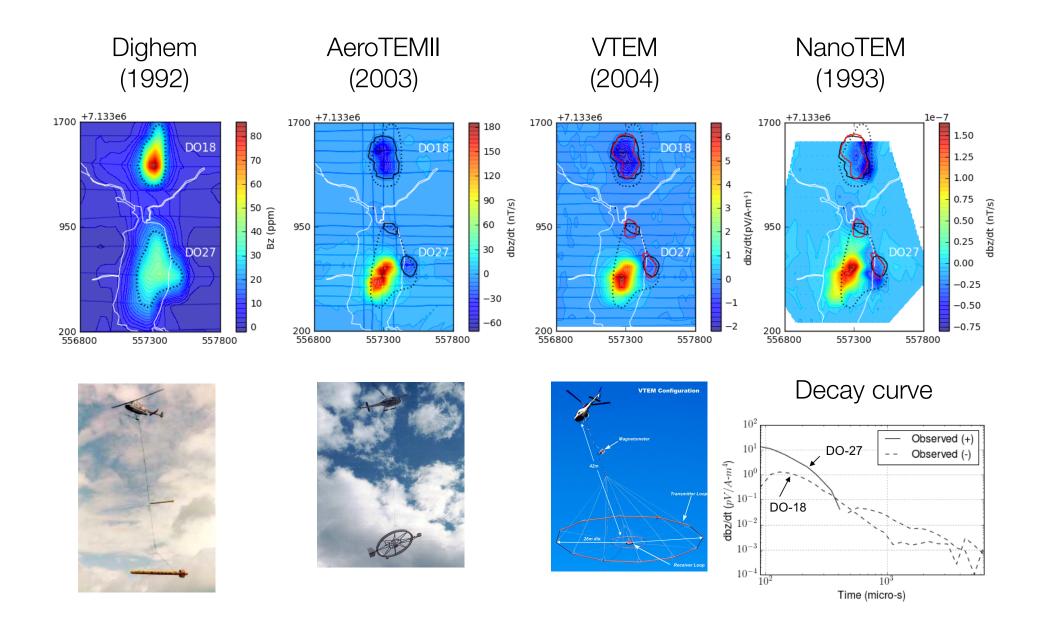


Kimberlite pipe structure

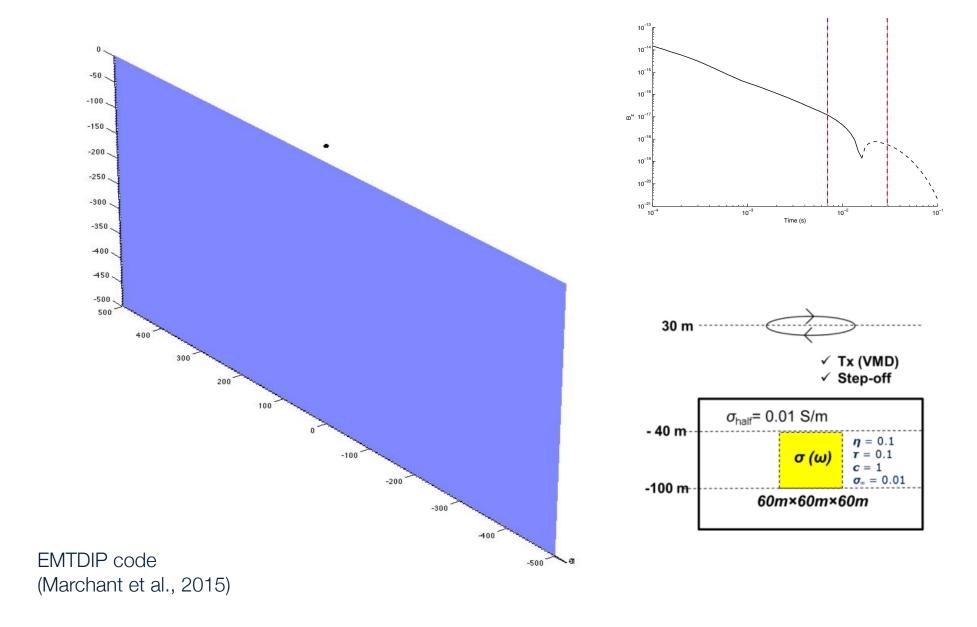


Devriese et al. (2016)

Time domain EM data



Reversed currents



IP inversion workflow

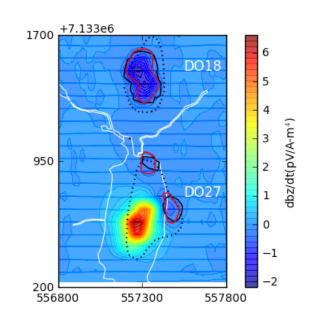
Workflow

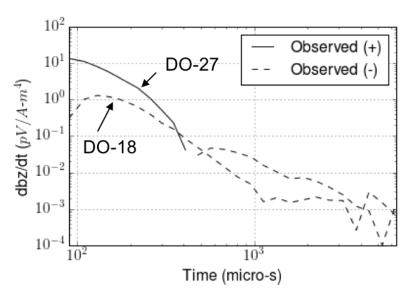
Invert TEM data, to recover σ_{∞}

Compute IP datum
Remove EM responses

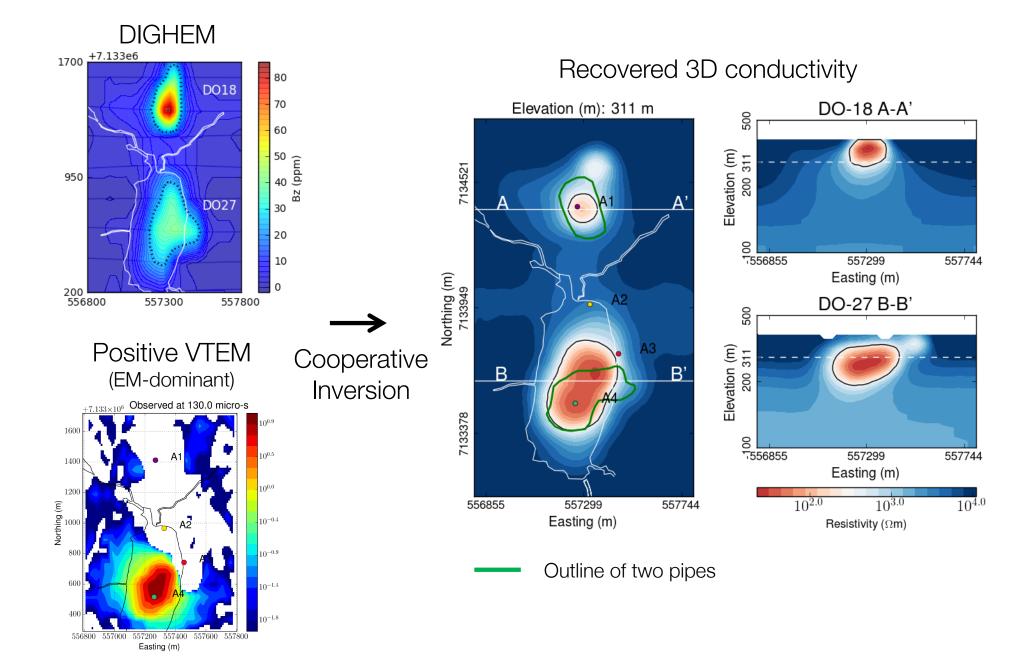
Linearized equations

Invert d^{IP} data, recover pseudo-chargeability





Step 1: Conductivity inversion



IP = Observation - Fundamental

$$d^{IP} = F[\sigma(t)] - F[\sigma_{\infty}]$$
 $F[\cdot]$: Maxwell's operator

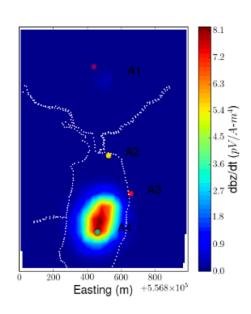
130 micro-s

Observed

1600
1400
1400
1400
1400
A1
6.0
4.8
3.6
2.4
1.2
0.0
-1.2
-2.4

Easting (m) $+5.568 \times 10^5$

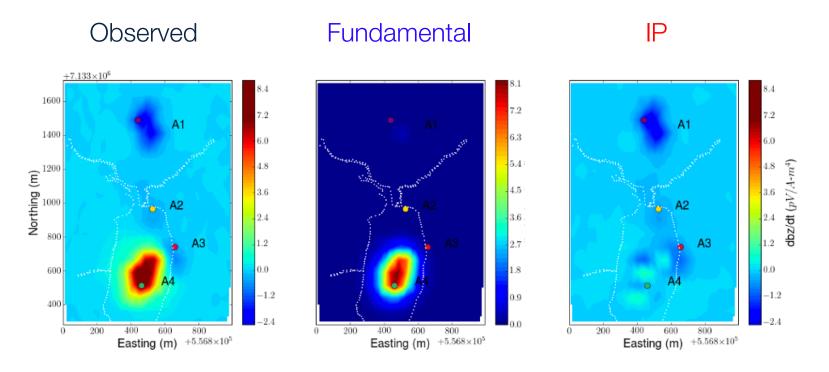
Fundamental



IP = Observation - Fundamental

$$d^{IP} = F[\sigma(t)] - F[\sigma_{\infty}]$$
 $F[\cdot]$: Maxwell's operator

130 micro-s



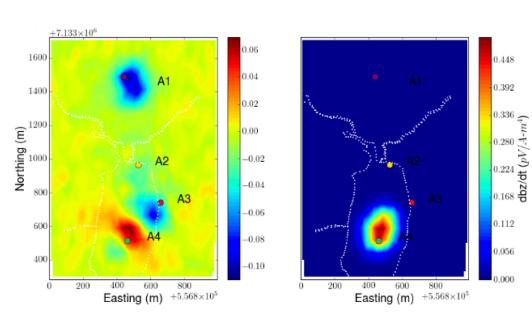
IP = Observation - Fundamental

$$d^{IP} = F[\sigma(t)] - F[\sigma_{\infty}]$$
 $F[\cdot]$: Maxwell's operator

410 micro-s

Observed

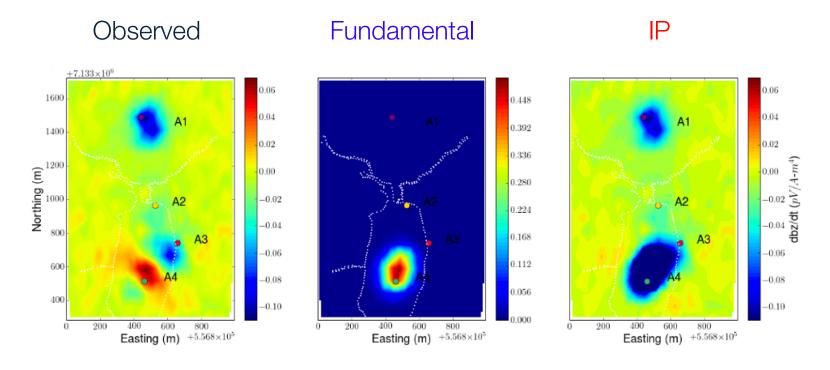
Fundamental



IP = Observation - Fundamental

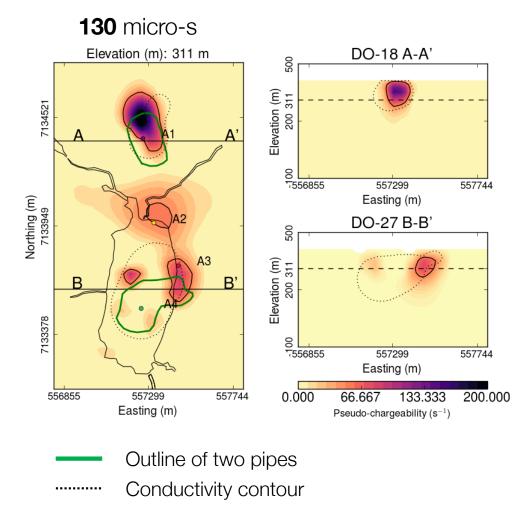
$$d^{IP} = F[\sigma(t)] - F[\sigma_{\infty}]$$
 $F[\cdot]$: Maxwell's operator

410 micro-s



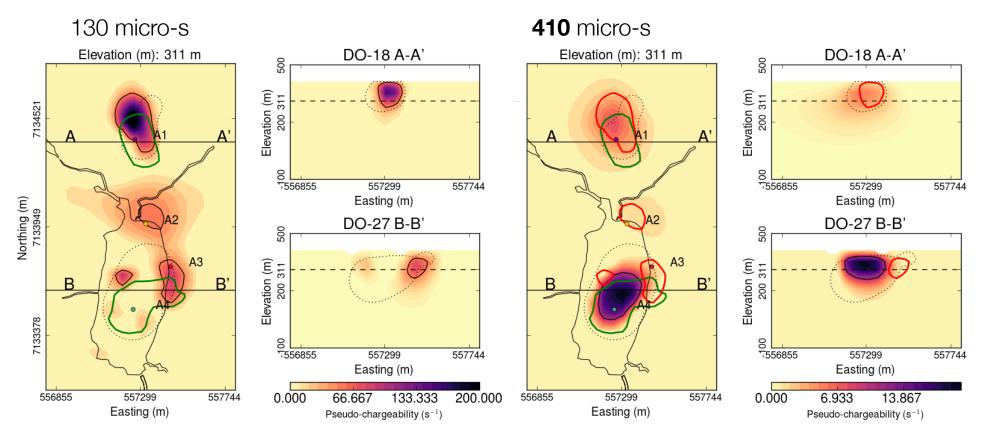
Step 3: 3D IP inversion

Recovered 3D pseudo-chargeability



Step 3: 3D IP inversion

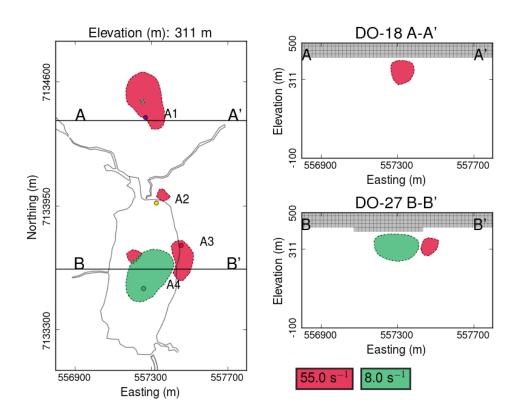
Recovered 3D pseudo-chargeability



Outline of two pipes
Conductivity contour

Step 4: Estimate η and τ

Anomaly contours



- A1-A3 has small time constant
- A4 has greater time constant

Cole-Cole model

$$\sigma(\omega) = \sigma_{\infty} + \sigma_{\infty} \frac{\eta}{1 + (1 - \eta)(\imath \omega \tau)^{c}}$$

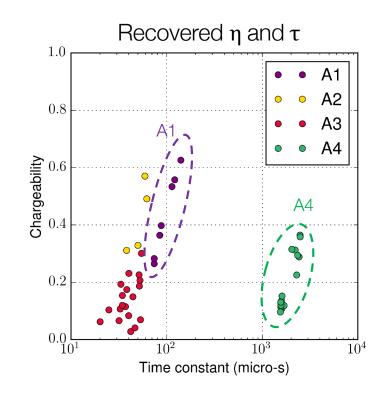
 σ_{∞} : Conductivity at infinite frequency

 σ_0 : Conductivity at zero frequency

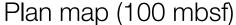
 η : Chargeability

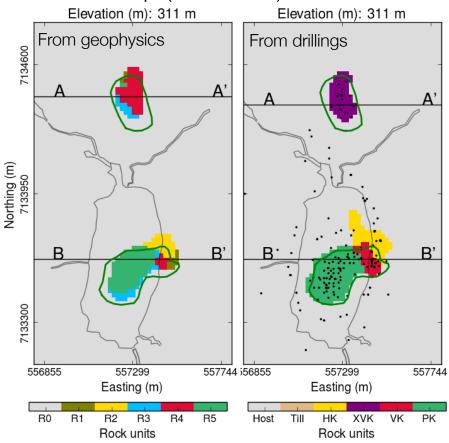
 τ : Time constant (s)

c: Frequency dependency



Impact on kimberlite exploration





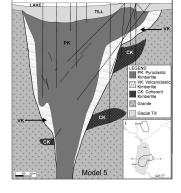
200 Elevation (m) Elevation (m) 557299 557744 556855 557299 557744 556855 Easting (m) Easting (m) R0 R3 HK XVK VK R1 R2 R4 Host Till Rock units Rock units

From **drillings**

DO-27 B-B'

From **geophysics**

DO-27 B-B'



Harder et al. (2009)

HK, PK, and VK are delineated in 3D

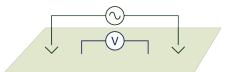
Workflow

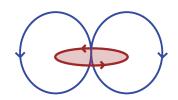
Invert TEM data, to recover σ_{∞}

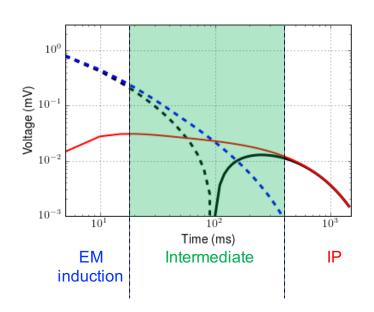
Compute IP datum
Remove EM responses

Linearized equations

Invert d^{IP} data, recover pseudo-chargeability







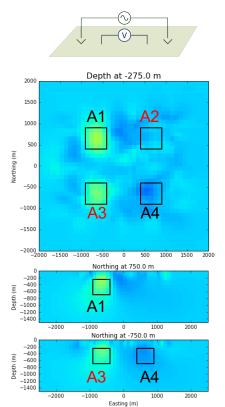
Workflow

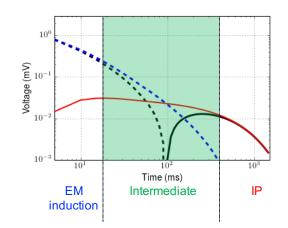
Invert TEM data, to recover σ_{∞}

Compute IP datum
Remove EM responses

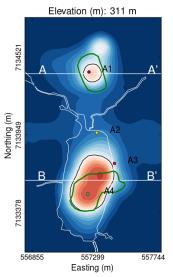
Linearized equations

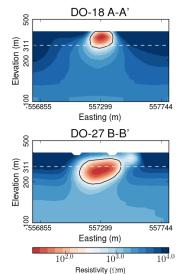
Invert d^{IP} data, recover pseudo-chargeability

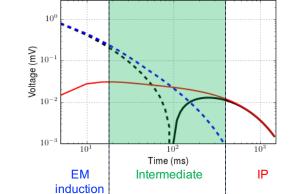












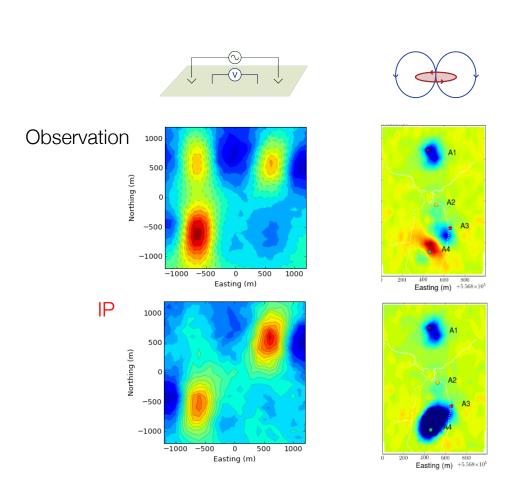
Workflow

Invert TEM data, to recover σ_{∞}

Compute IP datum
Remove EM responses

Linearized equations

Invert d^{IP} data, recover pseudo-chargeability



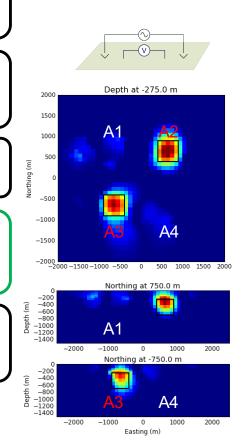
Workflow

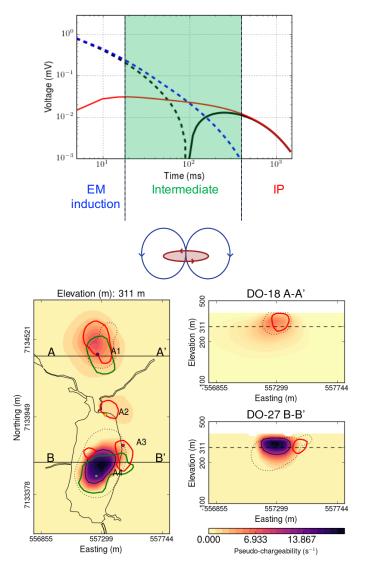
Invert TEM data, to recover σ_{∞}

Compute IP datum
Remove EM responses

Linearized equations

Invert d^{IP} data, recover pseudo-chargeability







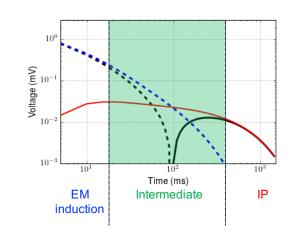
Invert TEM data, to recover σ_{∞}

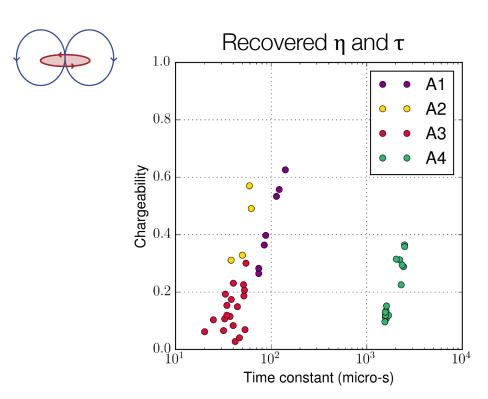
Compute IP datum
Remove EM responses

Linearized equations

Invert d^{IP} data, recover pseudo-chargeability

Estimate intrinsic IP parameters

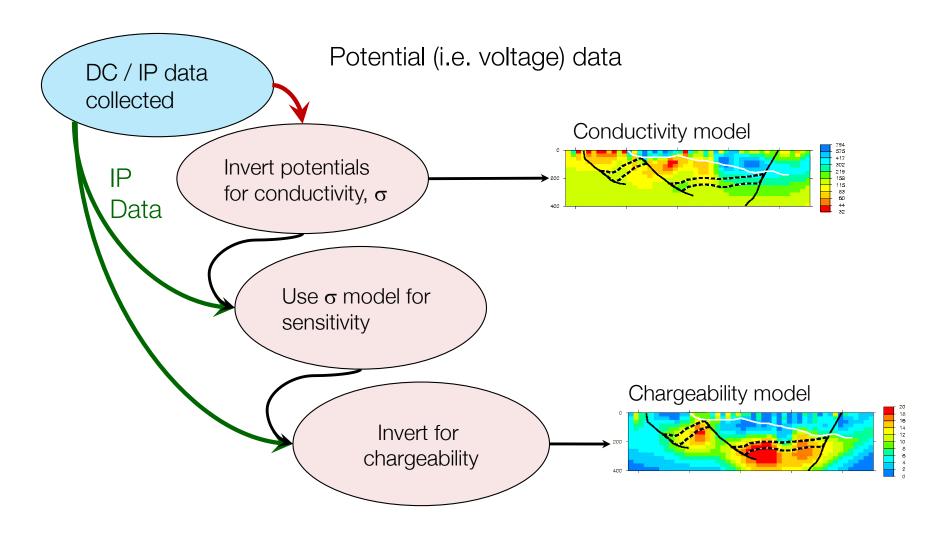




We may distinguish different rock types

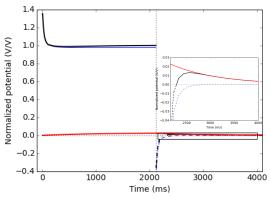
Thank you

IP Inversion



TEM-IP inversion workflow

Kang and Oldenburg (2016)



10⁰
10⁻¹
10⁻³

EM Intermediate IP

Invert TEM data, to recover σ_{∞}

Compute IP datum Remove EM responses

Linearized equations

Invert d^{IP} data, recover pseudo-chargeability

Estimate intrinsic IP parameters

IP = Observation - Fundamental $d^{IP}(t) = F[\sigma(t)] - F[\sigma_{\infty}]$

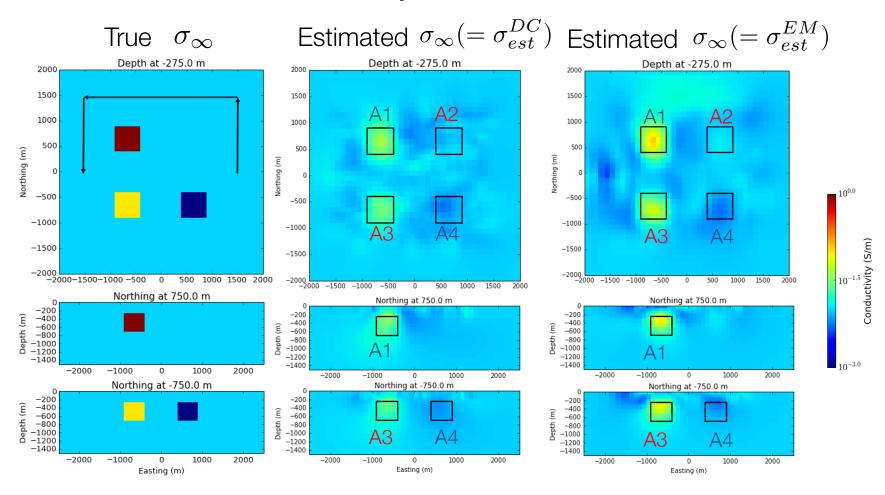
 $F[\cdot]$: Maxwell's operator

$$d^{IP}(t) = G\tilde{\eta}(t)$$

 $G(\sigma_{\infty})$: Sensitivity function $\tilde{\eta}$: Pseudo-chargeability

Comparison of 3D conductivities

Recovered 3D conductivity



Comparisons of Fundamental

Fundamental data at 80 ms

