

# EM: Natural Sources

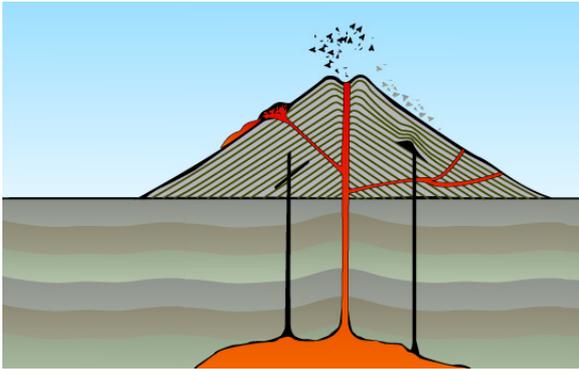


# Outline

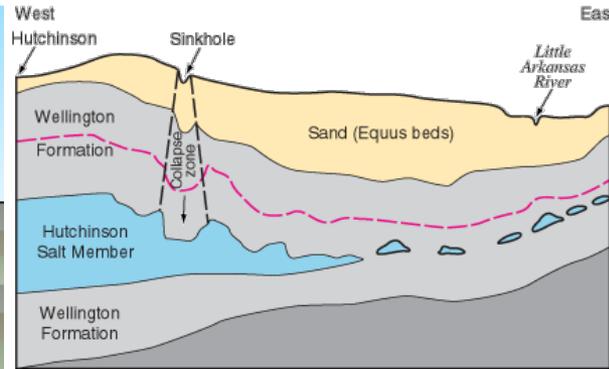
- Background on natural source EM methods
- Magnetotellurics
- MT case history
- Z-axis tipper electromagnetics
- ZTEM case history

# Motivation

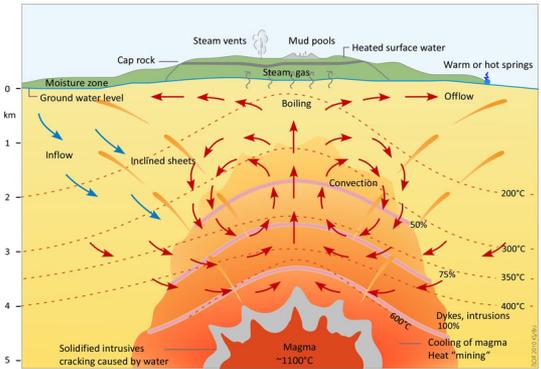
Volcanoes



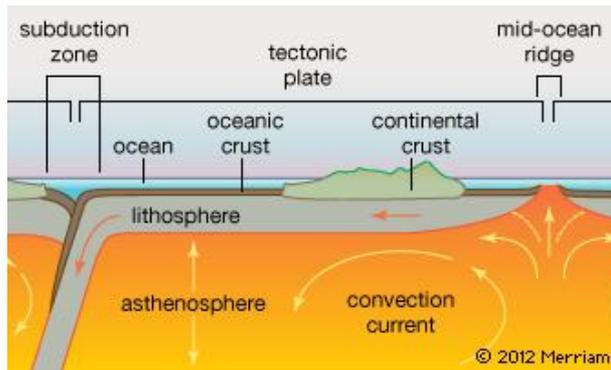
Base of salt



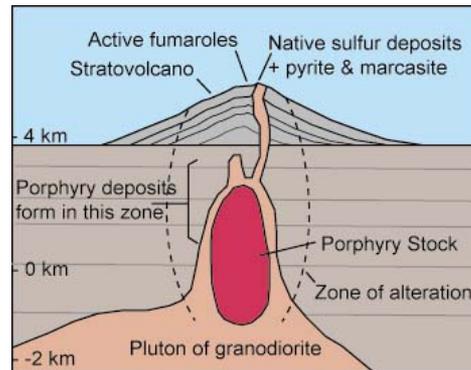
Geothermal



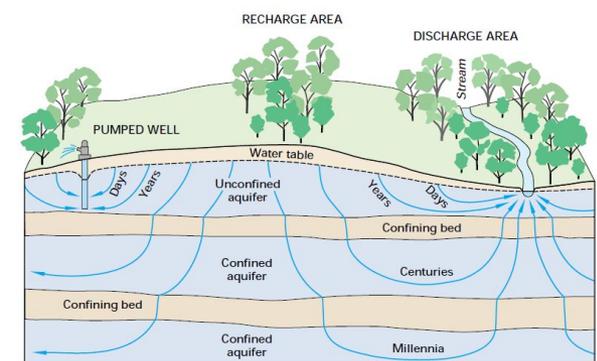
Tectonic settings of top few km



Mineral targets



Groundwater



Common challenge: getting enough energy into the ground

# What is required to see deeper?

- Penetration depth depends upon system power

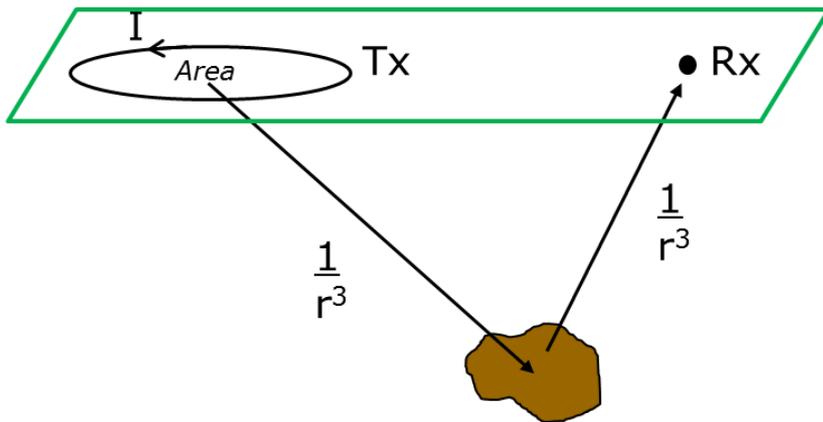
- Controlled source:

- Using a small loop
- Magnetic moment

$$m = IA$$

- Total geometric decay

$$\sim \frac{1}{r^6}$$

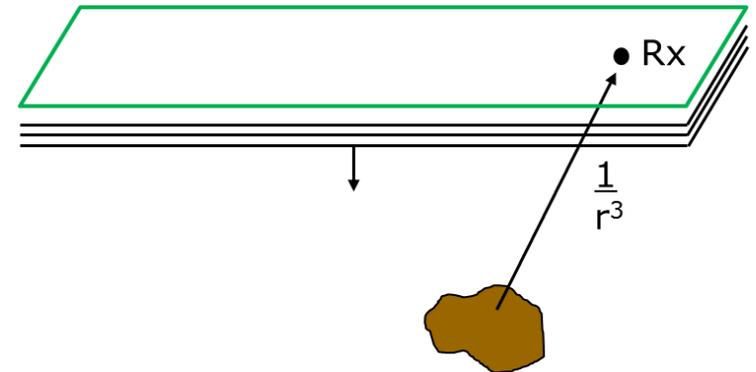


- Infinitely large loop source

- Sheet currents generate plane waves

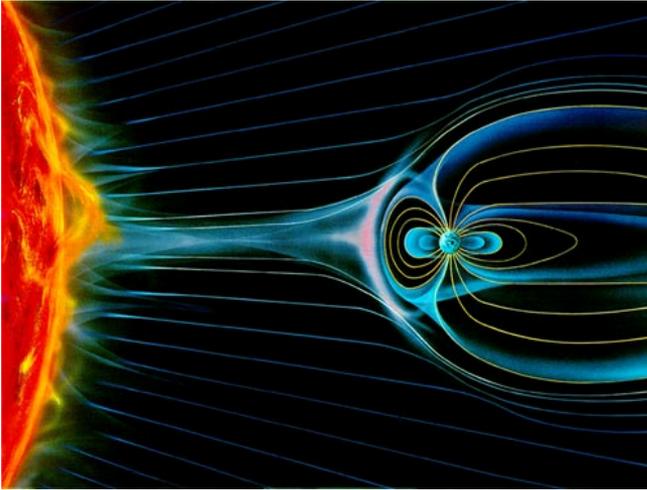
- Total geometric decay

$$\sim \frac{1}{r^3}$$

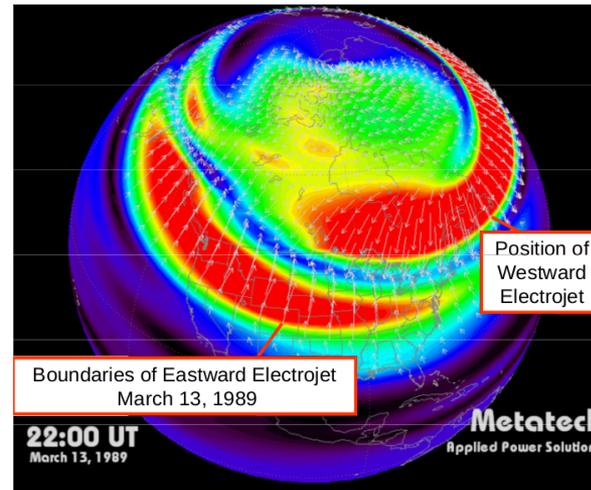


# Natural EM sources

Sun and magnetosphere, solar storms



Auroral electrojet; aurora



Lightning

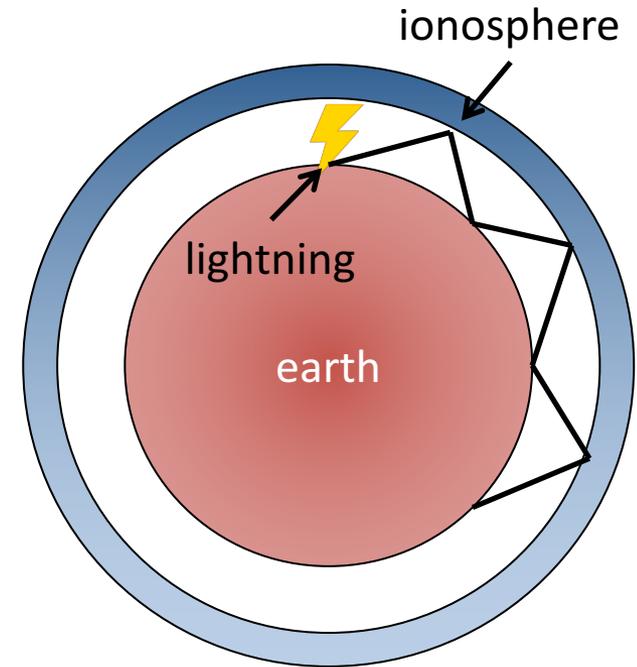


# Aurora movie

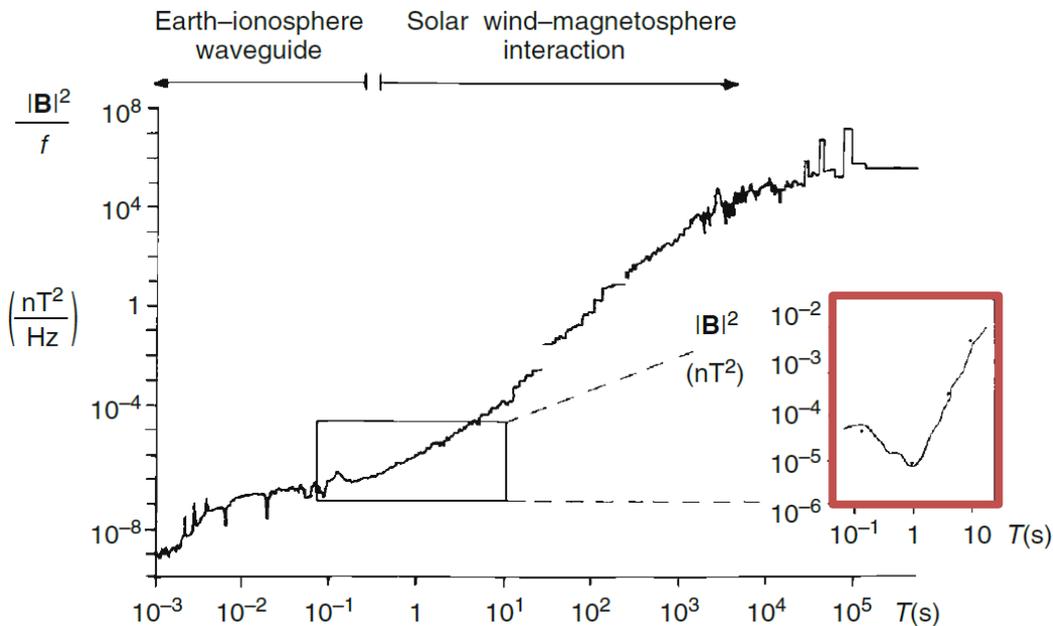


# Earth as a waveguide

- EM waves bounce between earth and highly conductive ionosphere
- Travel as plane waves



- Dead band: difficult to collect frequencies in notch (~1 Hz)



# Refraction of waves

- Snell's law

$$k_i \sin \theta_i = k_t \sin \theta_t$$

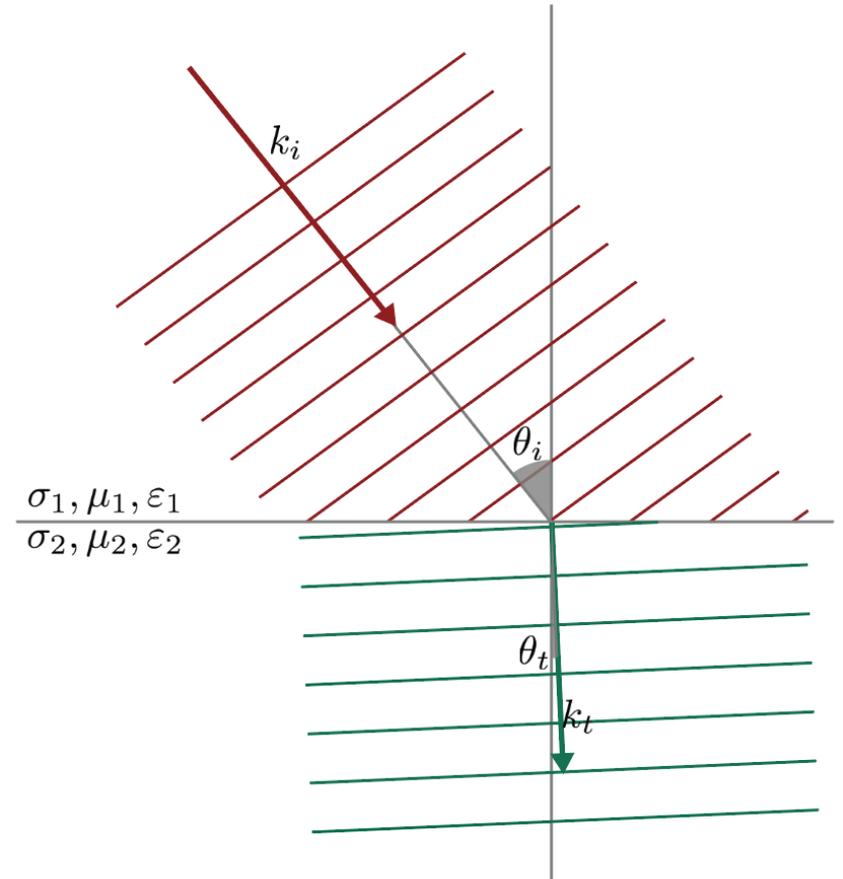
- $k$  is complex wave number

$$k^2 = \omega^2 \mu \epsilon - i \omega \mu \sigma$$

- Quasi-static:  $\frac{\omega \epsilon_0}{\sigma} \ll 1$

$$\sin \theta_t = \sqrt{\frac{2\omega \epsilon_0}{\sigma}} \sin \theta_i$$

- Angle of refraction is  $\theta_t = 0^\circ$  in almost every instance



Example for 10,000 Hz

$$\sigma = 10^{-3} \text{ S/m}$$

$$\theta_i = 89^\circ$$

$$\text{Then } \theta_t = 1.35^\circ$$

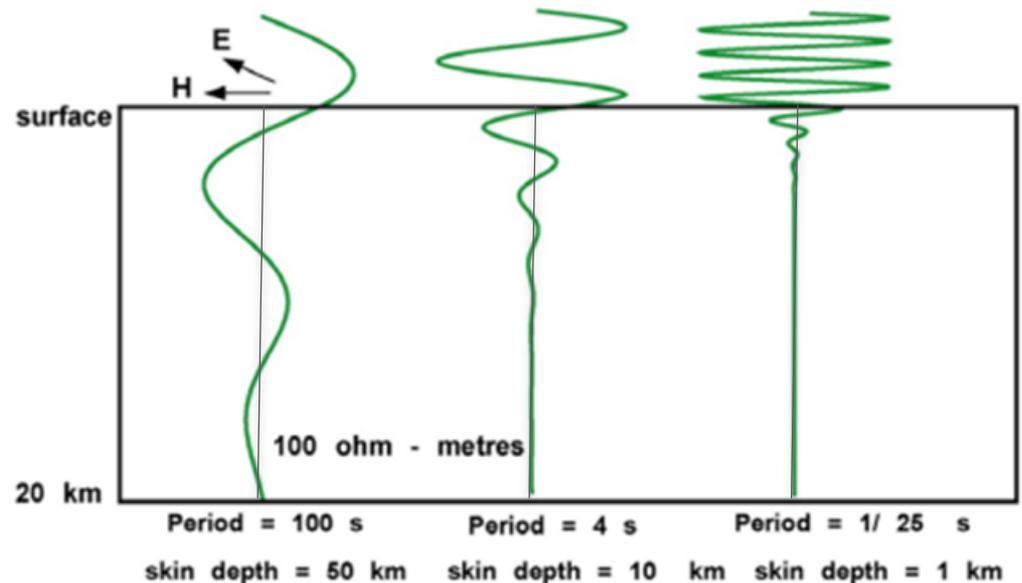
# Plane waves and skin depth

- Skin depth (meters)

$$\delta = \sqrt{\frac{2}{\omega\mu\sigma}} = 503\sqrt{\frac{1}{\sigma f}}$$

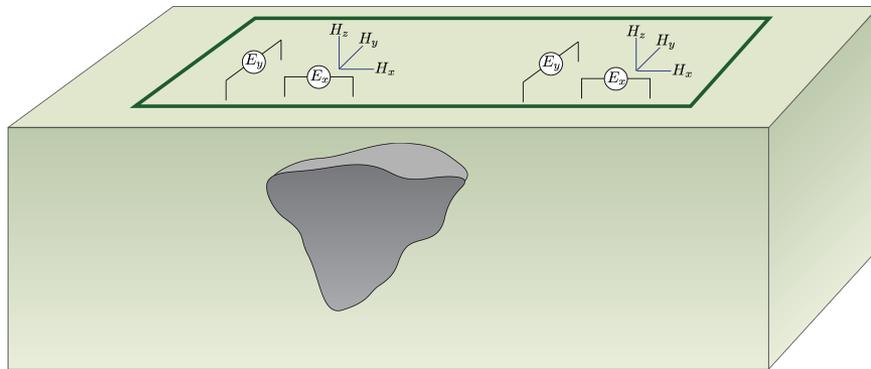
- Low frequency waves propagate further

- Depth of propagation
  - A few skin depths
  - Only a portion of a wavelength

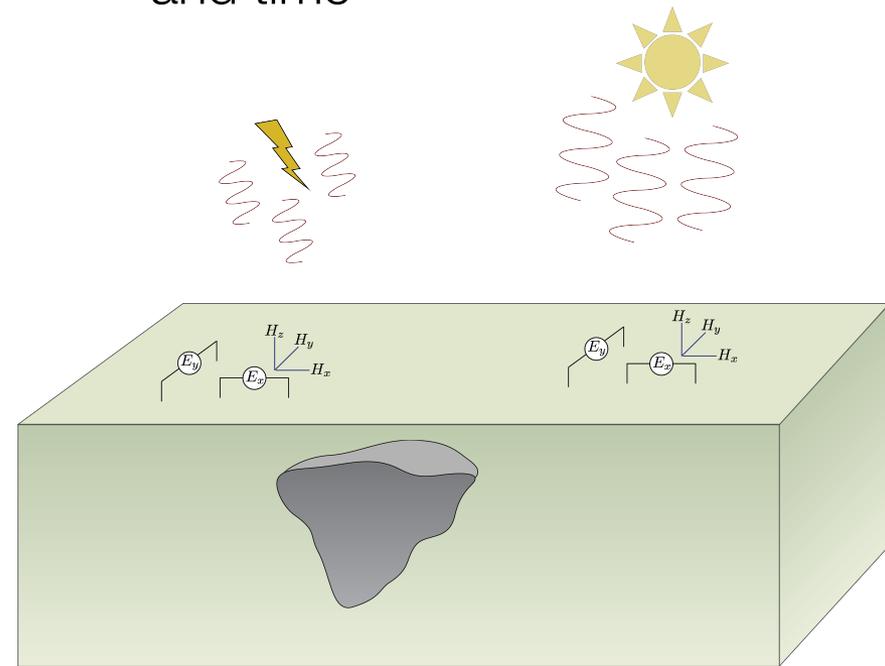


# Control source vs Natural source

- Controlled source
  - Well-defined location, geometry, and amplitude



- Natural sources
  - Sources are random in space and time



# MT Station

- Maxwell's equations:
  - Linear in  $J_s$
  - $E$  and  $H$  affected in the same way
- Effects of unknown source removed by taking ratio
- Transfer function

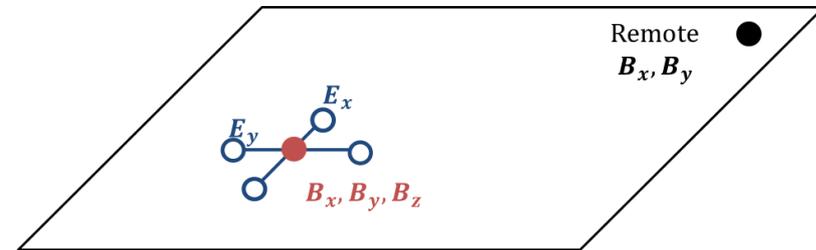
$$\nabla \times \mathbf{E} + i\omega\mu\mathbf{H} = 0$$

$$\nabla \times \mathbf{H} - \sigma\mathbf{E} = \mathbf{J}_s$$

$$\mathbf{E} = \mathbf{ZH}$$

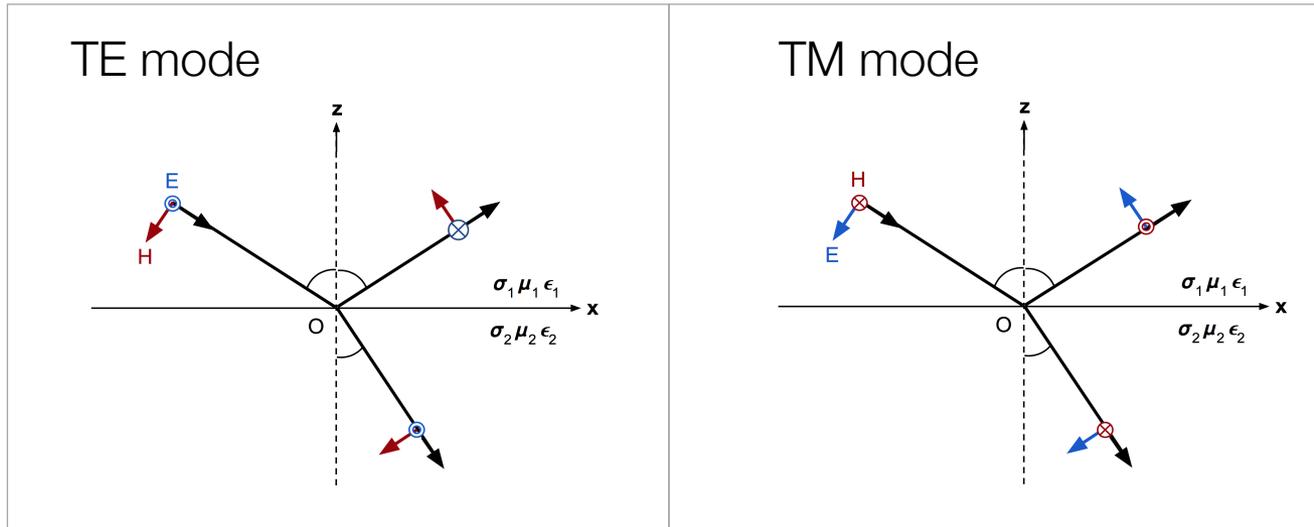
impedance (matrix)

$$\begin{pmatrix} E_x \\ E_y \end{pmatrix} = \begin{pmatrix} Z_{xx} & Z_{xy} \\ Z_{yx} & Z_{yy} \end{pmatrix} \begin{pmatrix} H_x \\ H_y \end{pmatrix}$$



# Impedance and resistivity

- Plane wave in homogenous media:
  - $E$  and  $H$  fields are perpendicular



Homogeneous half space

Impedance $Z_{xy} = \frac{E_x}{H_y}$	Resistivity $\rho = \frac{1}{\omega\mu}  Z_{xy} ^2$	Phase $\Phi = \tan^{-1} \left( \frac{\text{Im}(Z_{xy})}{\text{Re}(Z_{xy})} \right) = \frac{\pi}{4}$
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# MT soundings in 1D

- In general:

$$Z = \begin{pmatrix} Z_{xx} & Z_{xy} \\ Z_{yx} & Z_{yy} \end{pmatrix}$$

- Apparent resistivity:

$$\rho_a = \frac{1}{\omega\mu_0} |Z_{xy}|^2$$

- Phase:

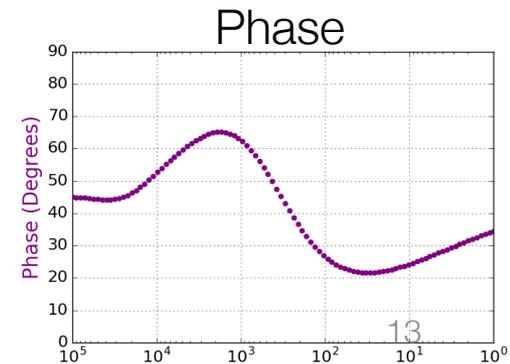
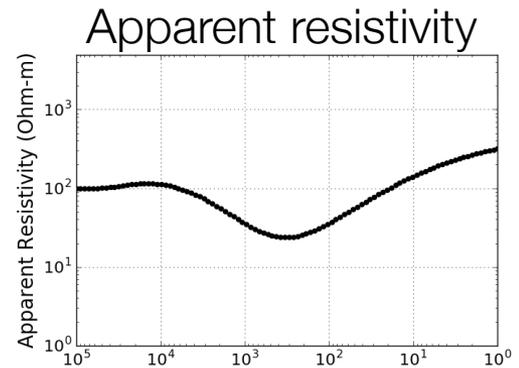
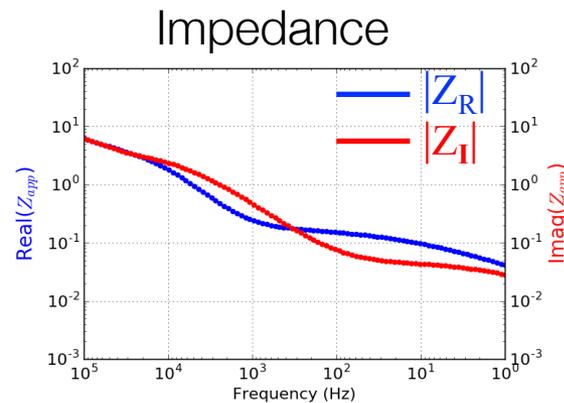
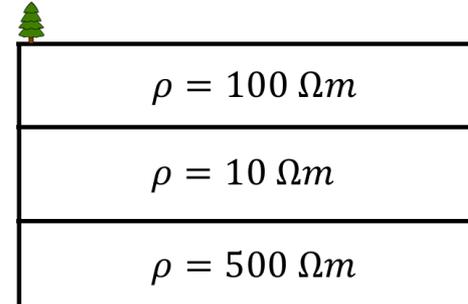
$$\Phi = \tan^{-1} \left( \frac{\text{Im}(Z_{xy})}{\text{Re}(Z_{xy})} \right)$$

- In 1D:

$$Z = \begin{pmatrix} 0 & Z_{xy} \\ Z_{yx} & 0 \end{pmatrix}$$

$$Z_{xy} = \frac{E_x}{H_y}$$

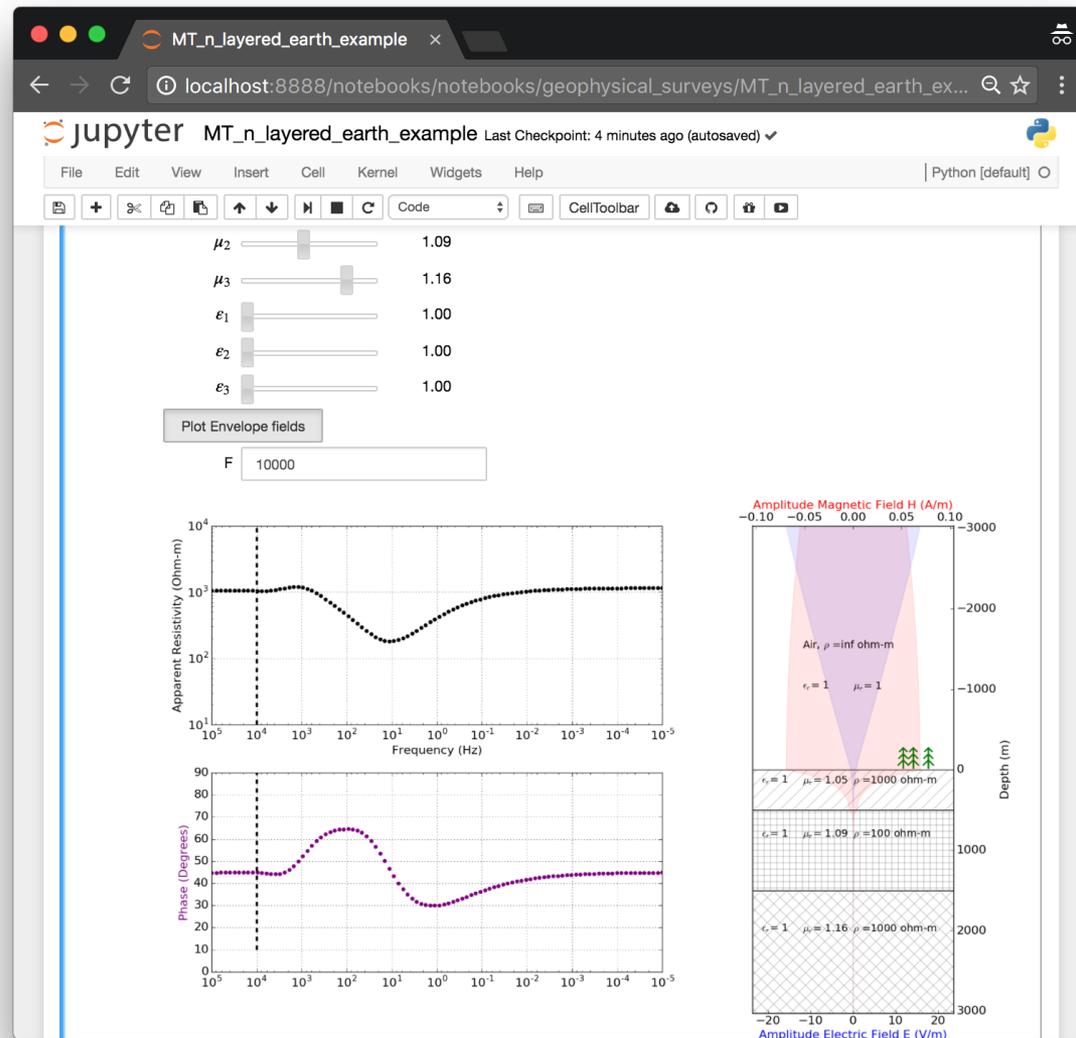
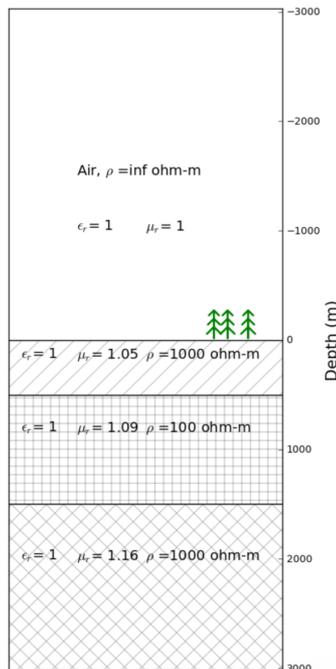
$$Z_{xy} = -Z_{yx}$$



# 1D MT app

$$\nabla \times \mathbf{E} + i\omega\mu\mathbf{H} = 0$$

$$\nabla \times \mathbf{H} - \sigma\mathbf{E} = \mathbf{J}_s$$



# MT soundings in 2D

- In general:

$$Z = \begin{pmatrix} Z_{xx} & Z_{xy} \\ Z_{yx} & Z_{yy} \end{pmatrix}$$

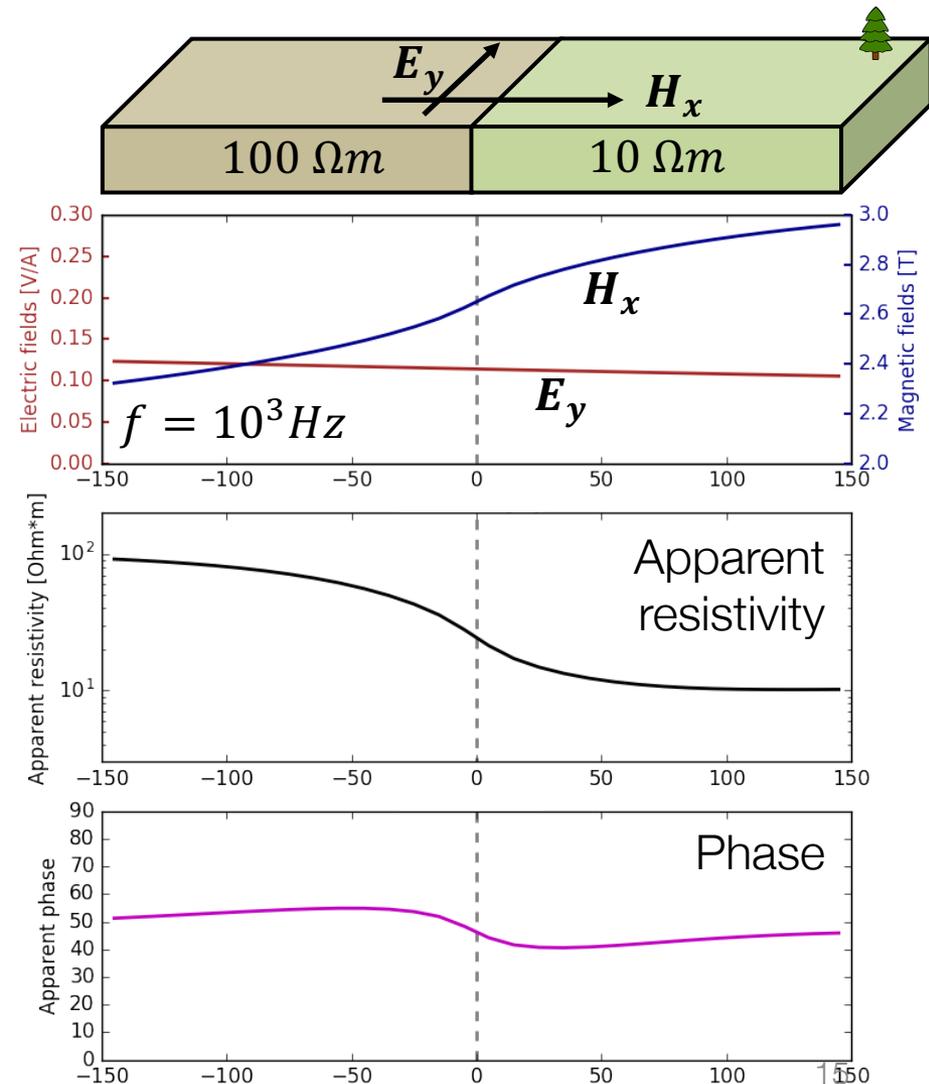
- In 2D:

$$Z = \begin{pmatrix} 0 & Z_{xy} \\ Z_{yx} & 0 \end{pmatrix}$$

$$Z_{xy} \neq Z_{yx}$$

- TE mode
  - E-field parallel to structure

$$Z_{yx} = \frac{E_y}{H_x}$$



# MT soundings in 2D

- In general:

$$Z = \begin{pmatrix} Z_{xx} & Z_{xy} \\ Z_{yx} & Z_{yy} \end{pmatrix}$$

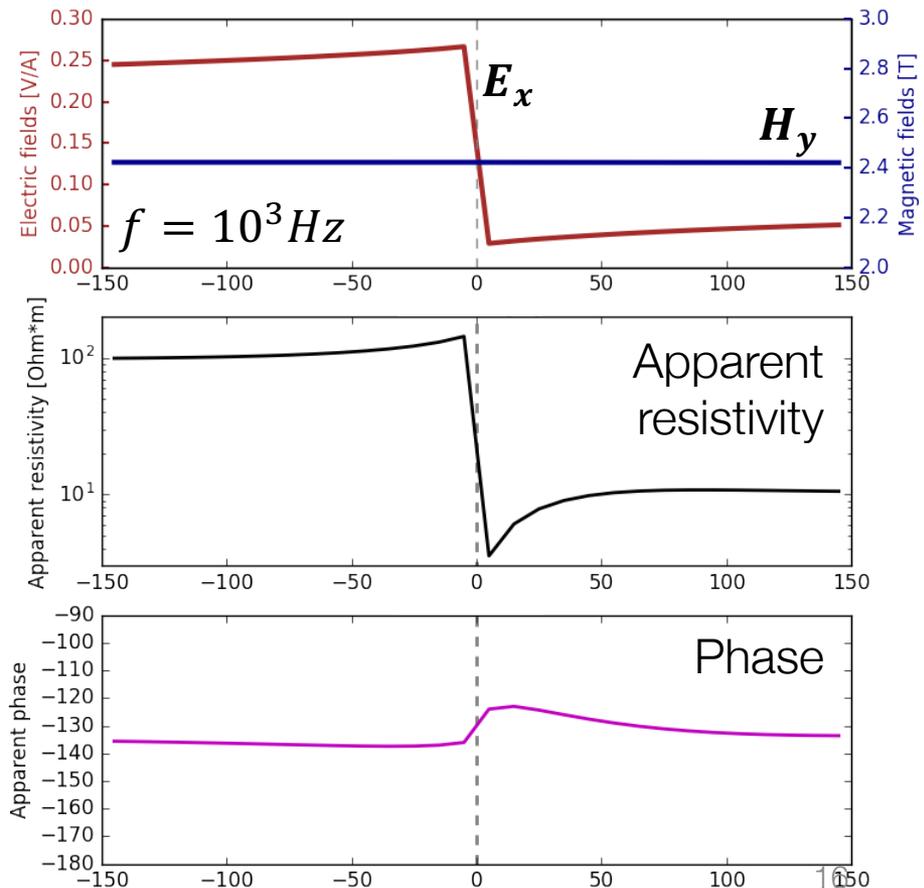
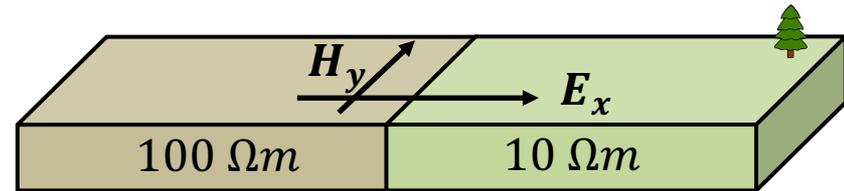
- In 2D:

$$Z = \begin{pmatrix} 0 & Z_{xy} \\ Z_{yx} & 0 \end{pmatrix}$$

$$Z_{xy} \neq Z_{yx}$$

- TM mode
  - H-field parallel to structure
  - $E_x$  discontinuous

$$Z_{xy} = \frac{E_x}{H_y}$$



# MT soundings in 3D

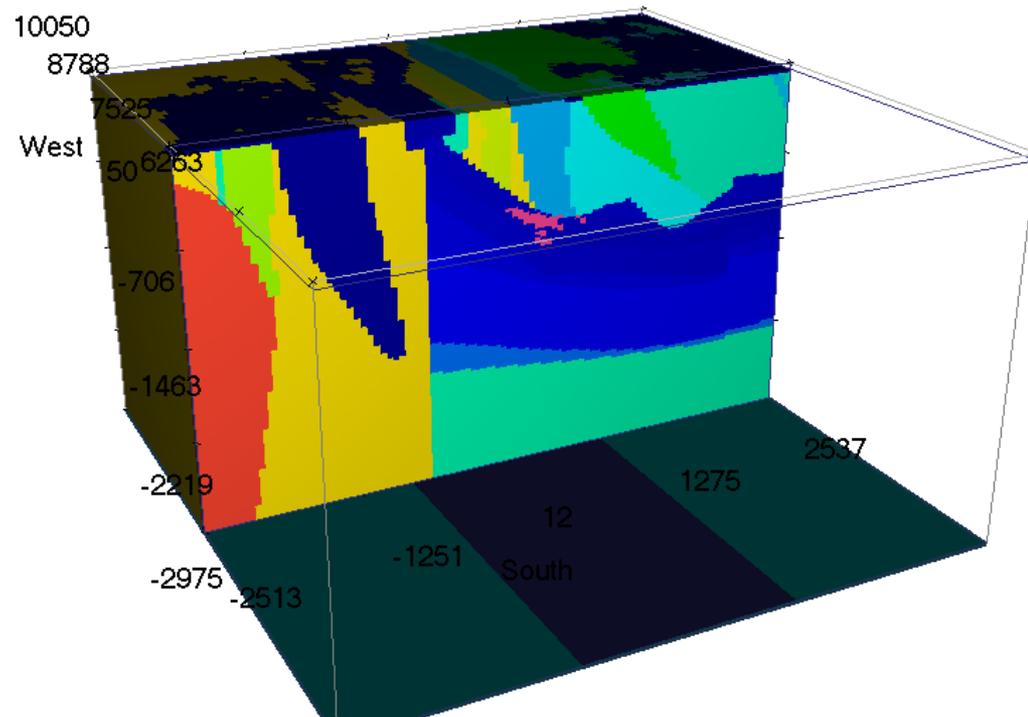
- In general:

$$Z = \begin{pmatrix} Z_{xx} & Z_{xy} \\ Z_{yx} & Z_{yy} \end{pmatrix}$$

- In 3D:

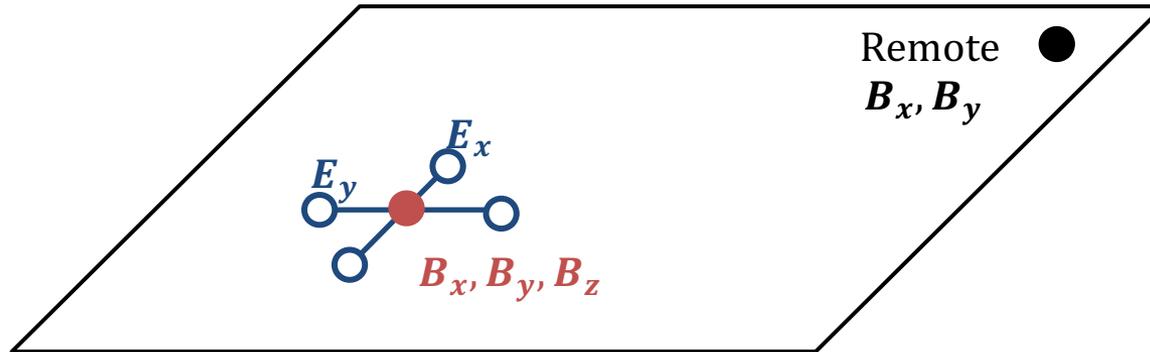
$$Z = \begin{pmatrix} Z_{xx} & Z_{xy} \\ Z_{yx} & Z_{yy} \end{pmatrix}$$

- No symmetry or special conditions



# Measuring MT data

- Basic acquisition

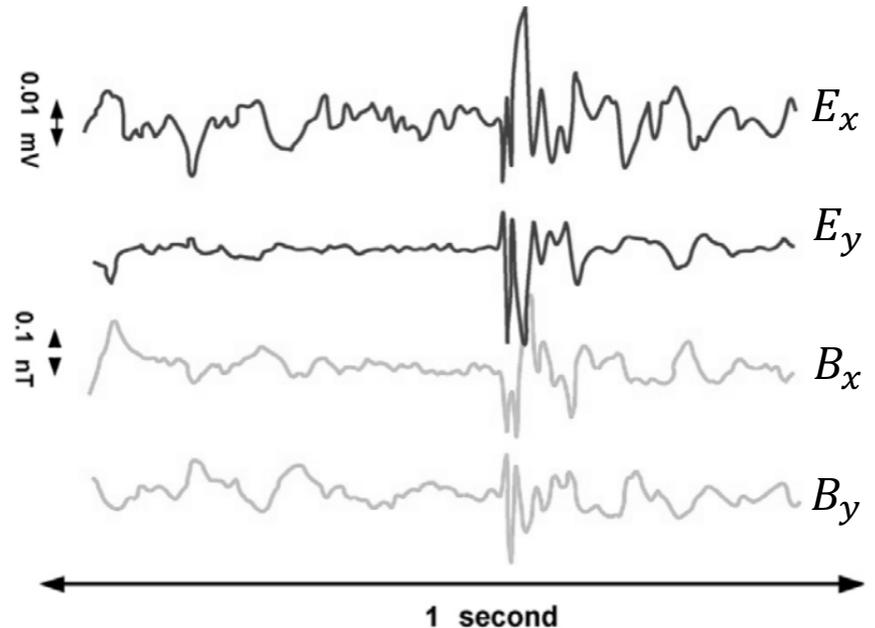


- At each station, measure:

$$E_x, E_y, B_x, B_y, B_z$$

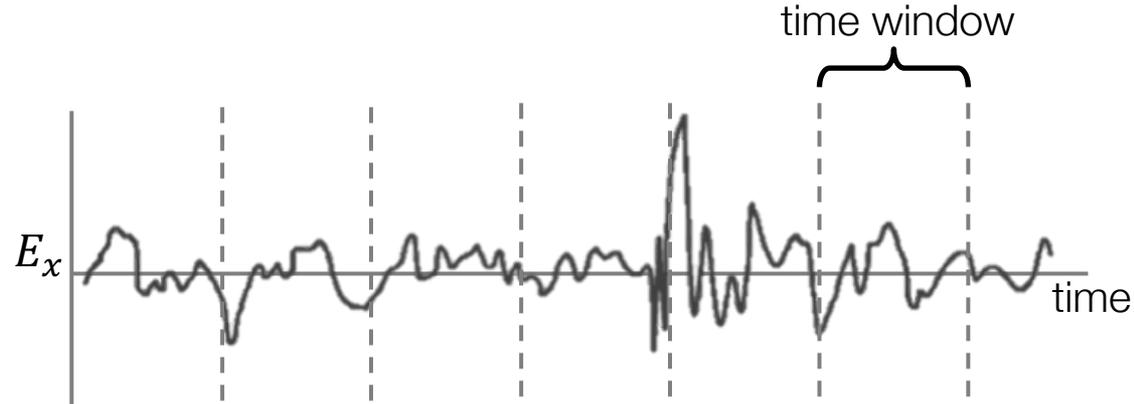
- At remote reference, measure:

$$B_x, B_y$$



# Processing MT data

- Divide time series into time windows



- Apply Fourier transform
  - For each station:

$$\begin{aligned} e_x(t) &\rightarrow E_x(\omega) \\ h_y(t) &\rightarrow H_y(\omega) \end{aligned}$$

- For the remote reference:

$$h_y^R(t) \rightarrow H_y^R(\omega)$$

- Form the impedance tensor:

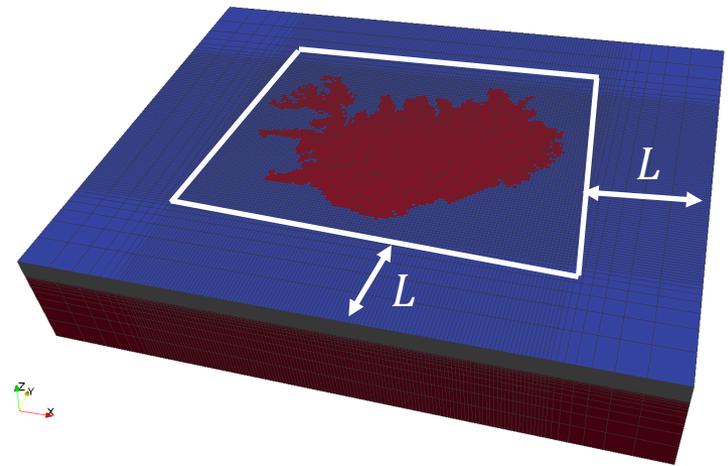
$$Z_{xy}(\omega) = \frac{\langle E_x(\omega) H_y^{R*}(\omega) \rangle}{\langle H_y(\omega) H_y^{R*}(\omega) \rangle}$$

(\*) complex conjugate

<> average over multiple samples

# Inverting MT data

- Boundary conditions important for modelling
- Mesh size:
  - MT: extended grid
  - $L$ : a few skin depths from data area
- Challenge: Unknown boundary conditions
  - Possible channeled currents
  - Data can be affected by distant structures
- Otherwise, inversion of MT is essentially same as CSEM data



# Outline

- Background on natural source EM methods
- Magnetotellurics
- Questions?
- MT case histories
- Z-axis tipper electromagnetics
- ZTEM case histories

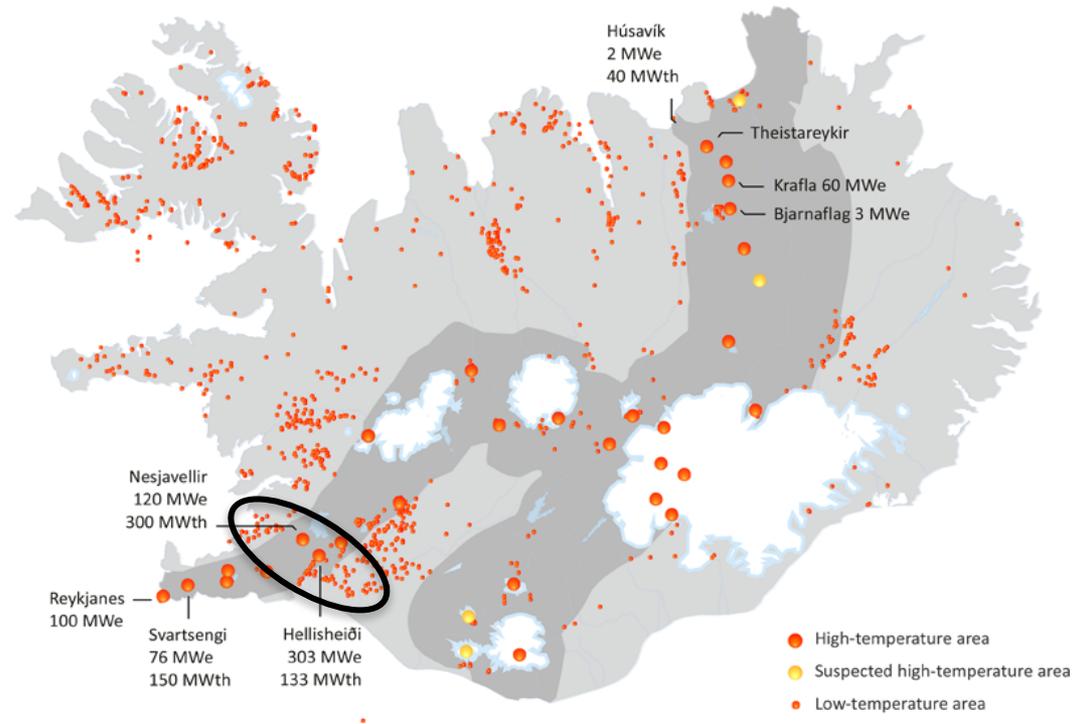
# MT case history

- Iceland



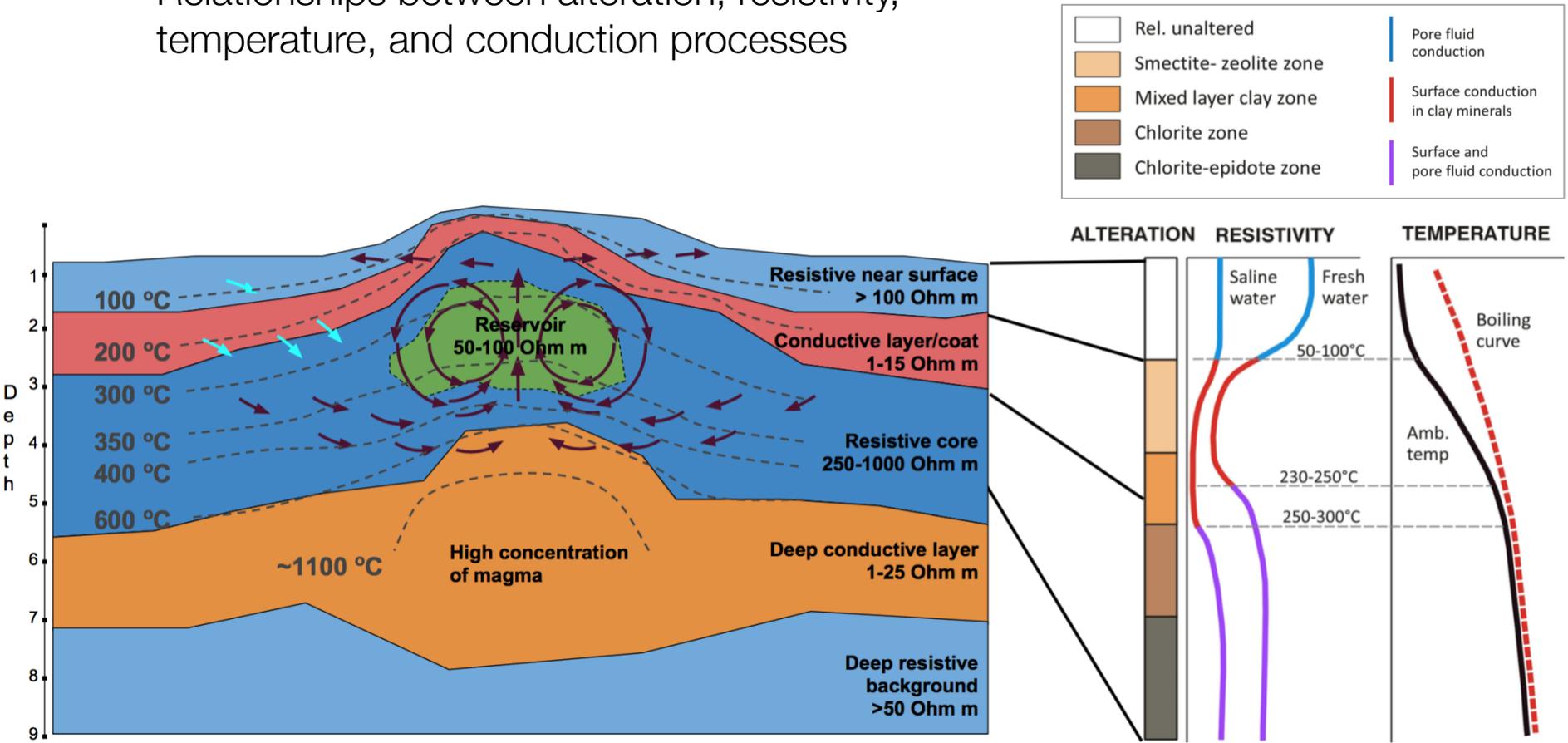
# Hengill geothermal region: setup

- Iceland: geothermal hot spot
  - On the mid-Atlantic ridge
  - Hosts multiple high temperature geothermal systems
- Hengill geothermal area
  - Supplies majority of hot water in Reykjavik
  - Contributes ~450 Mwe to National power grid

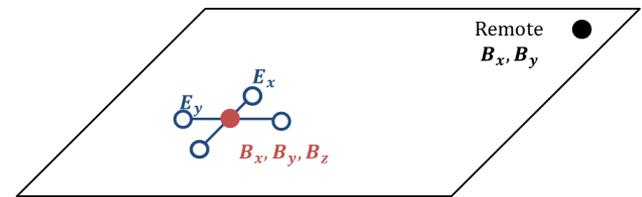


# Physical properties

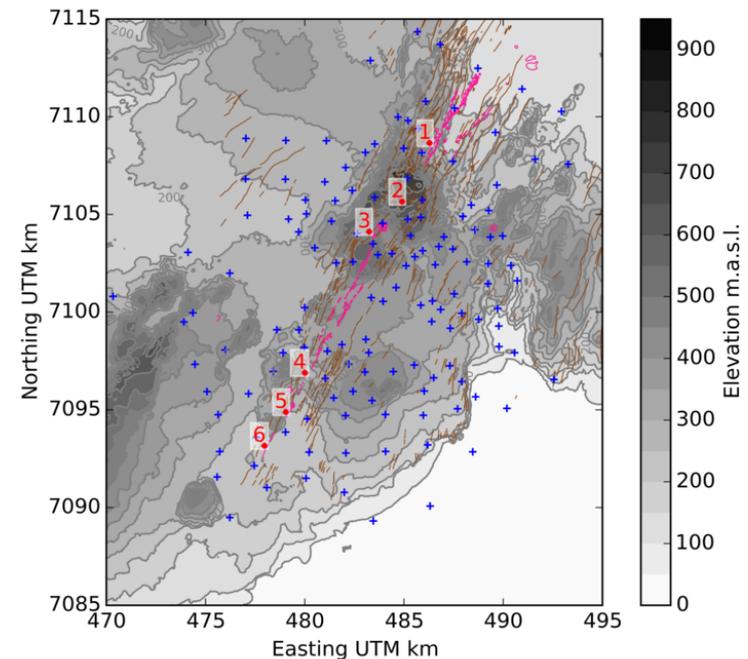
- Relationships between alteration, resistivity, temperature, and conduction processes



# Survey

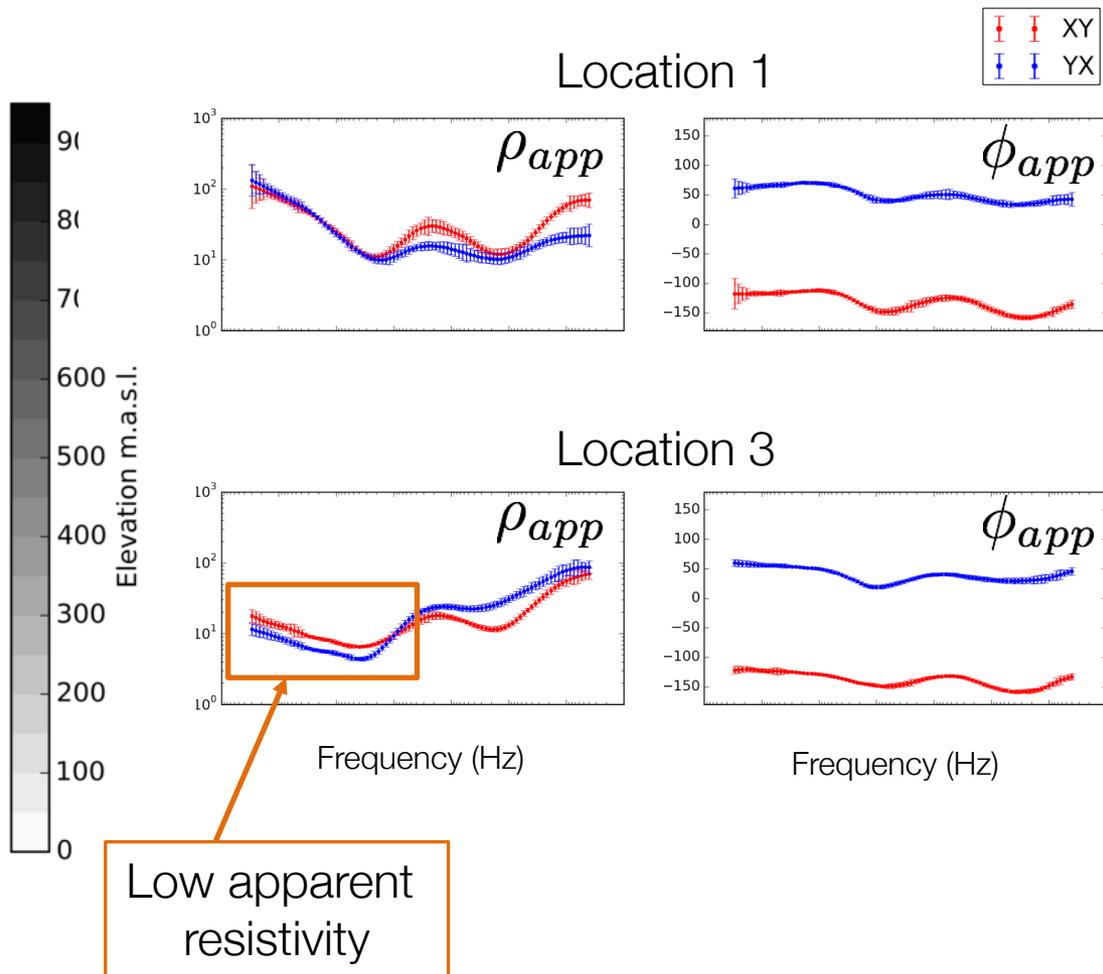
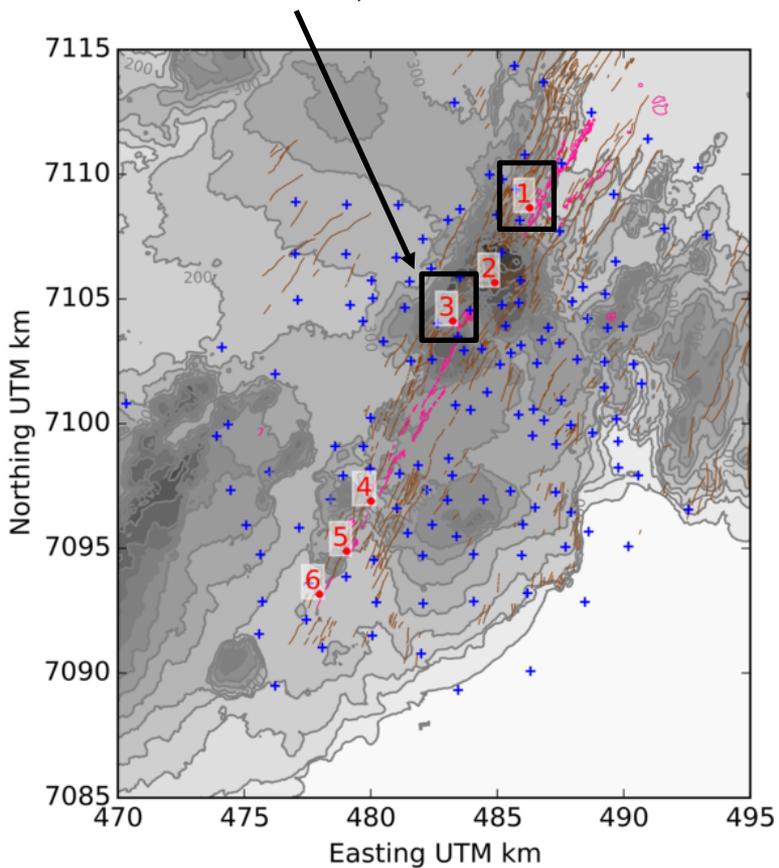


- MT instrumentation
  - Phoenix MTU5's
- Survey
  - 133 stations used
  - Combination of 2E and 2E+3H setup
  - Frequencies: 300 – 0.001 Hz
- Remote reference
  - About 40 km away
- Raw data processing using Phoenix software



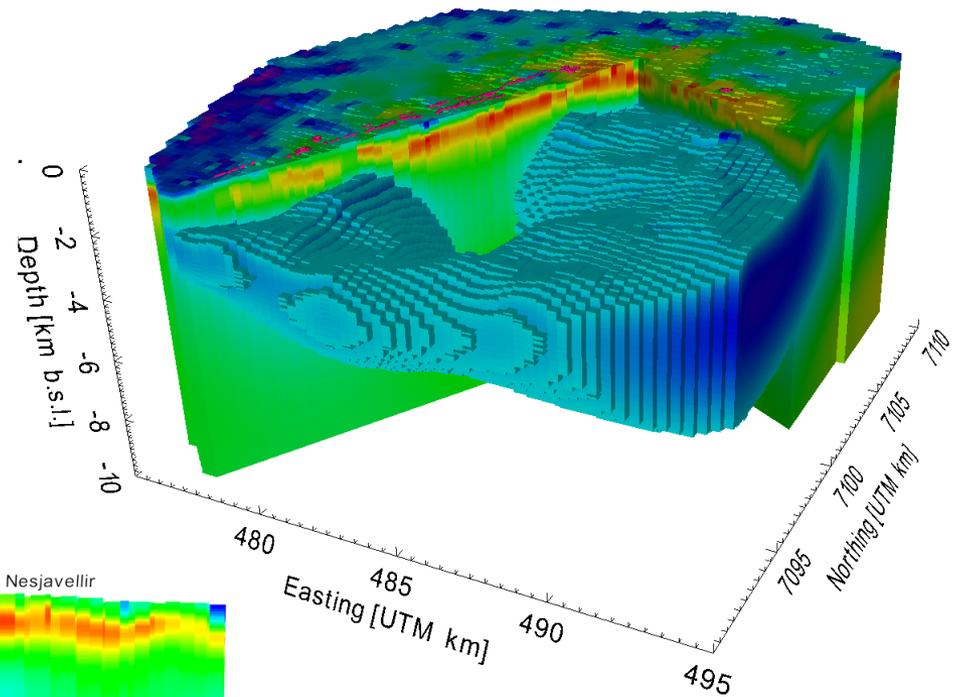
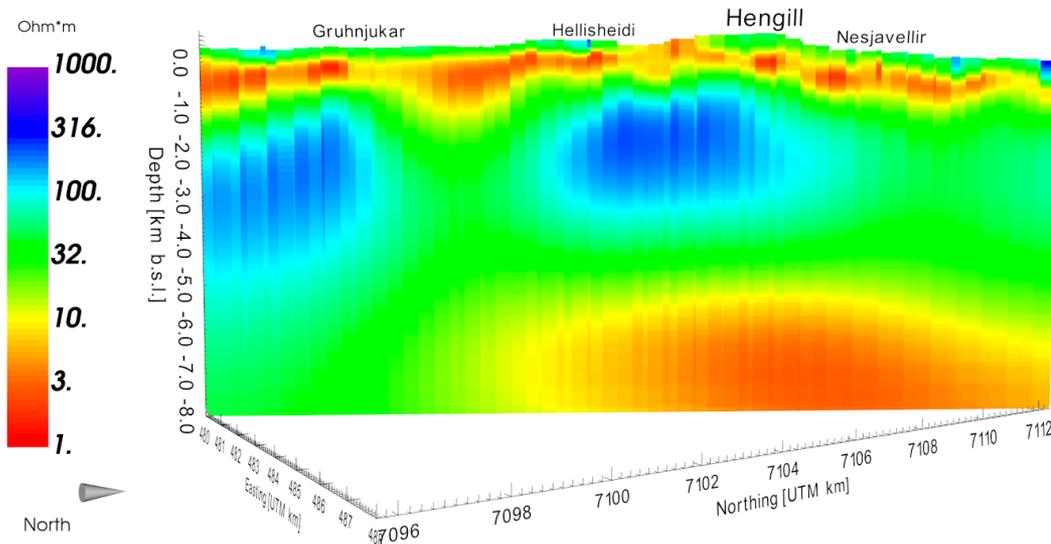
# Data

Surface alteration,  
hot water, fumaroles



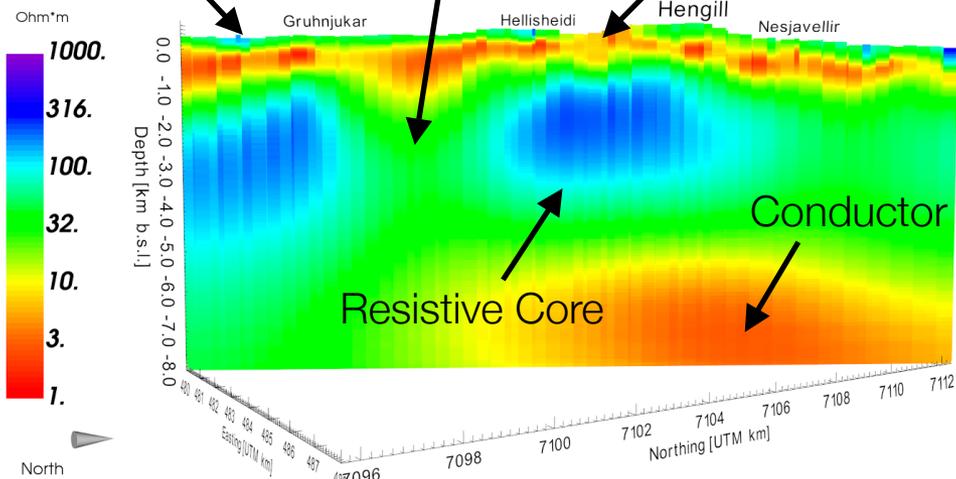
# 3D inversion

- Off-diagonal impedance ( $Z_{xy}$  and  $Z_{yx}$ ) used
- Combined multi-frequency inversion (300 Hz – 0.001 Hz)

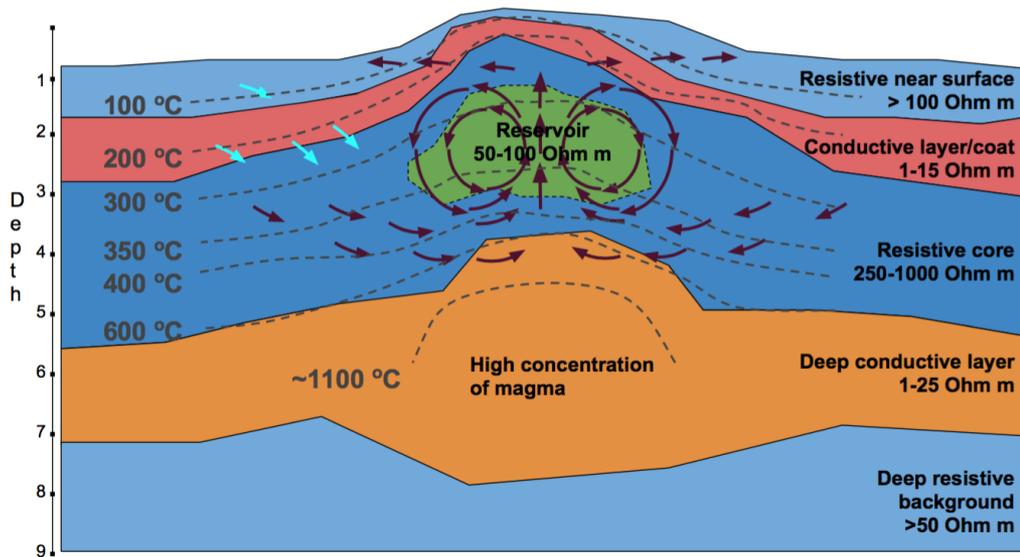
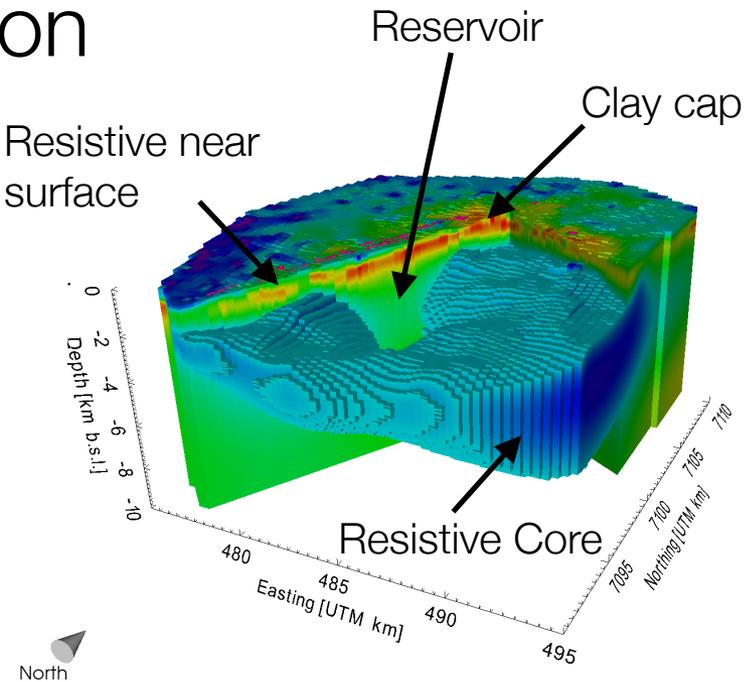


# Interpretation

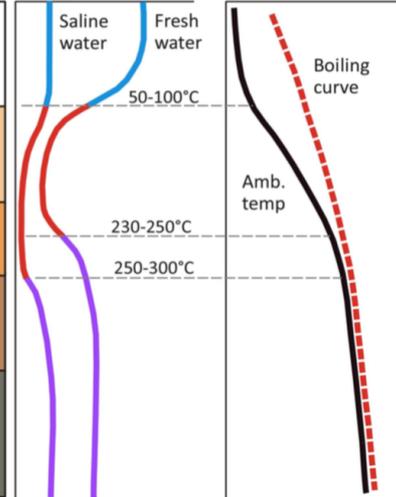
Resistive near surface



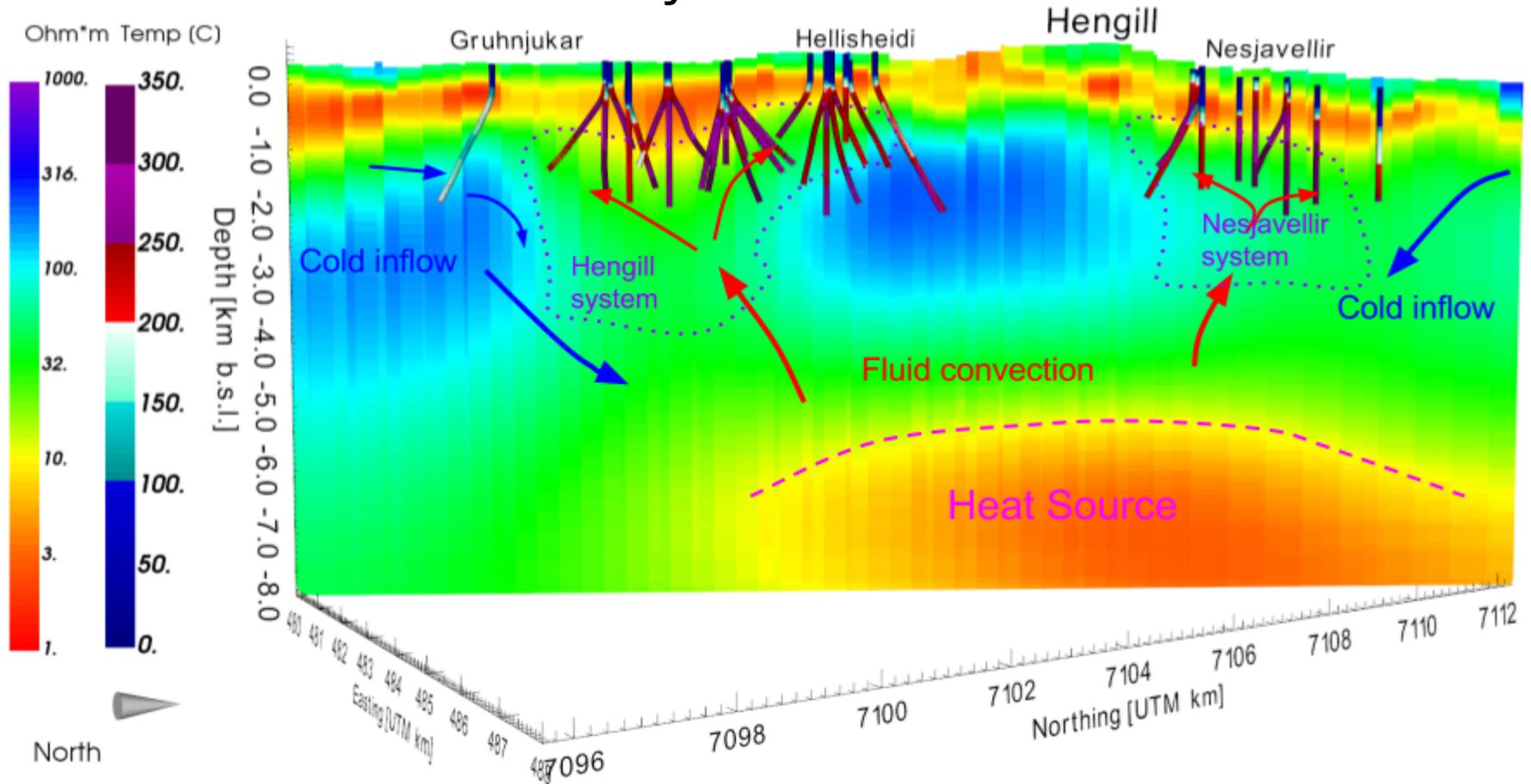
Resistive near surface



ALTERATION RESISTIVITY TEMPERATURE



# Synthesis

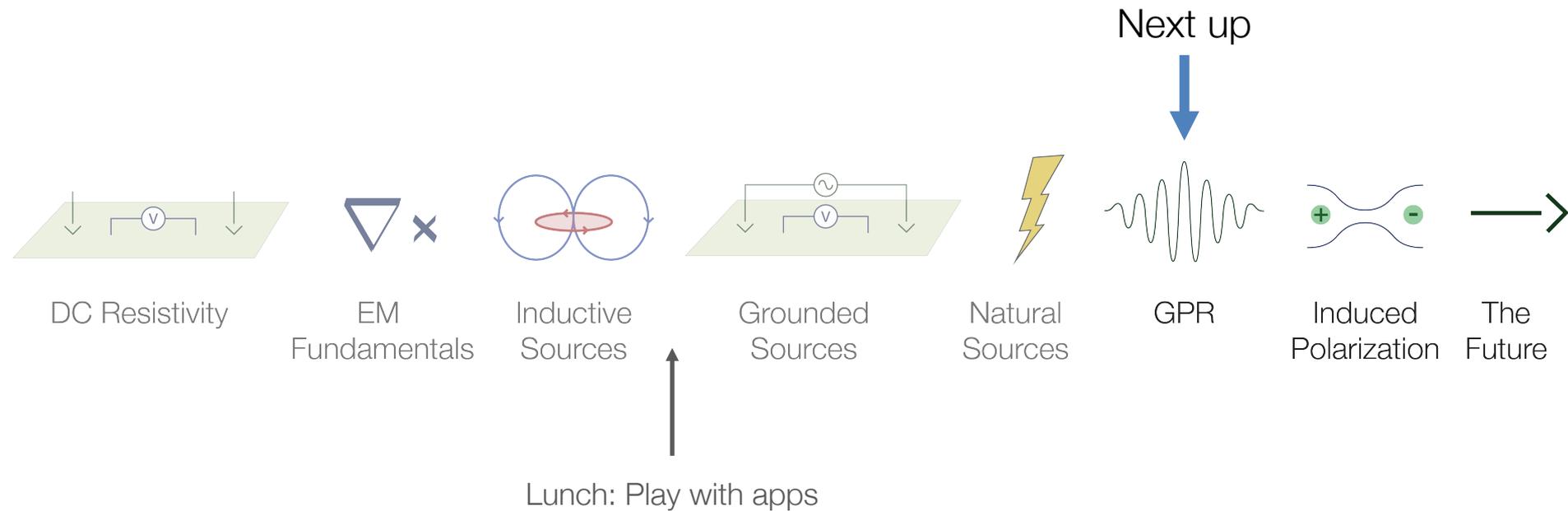


- Conductive layer corresponds with formation temperature
- Two main production fields: Hengill and Nesjavellir
- Deep conductive heat source

# Summary

- Background on natural source EM methods
- Magnetotellurics
- MT case history
- Z-axis tipper electromagnetics
- ZTEM case history

# End of Natural Sources





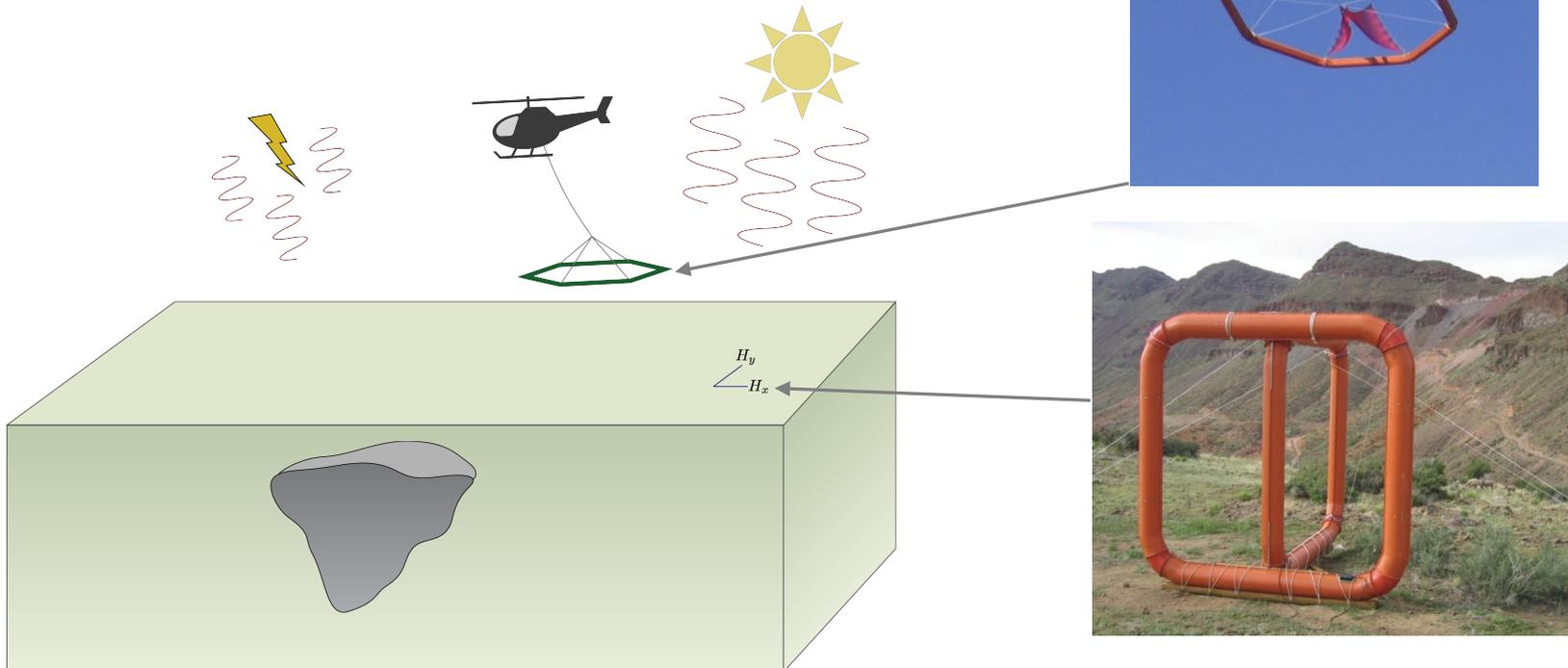
# Tipper data (ZTEM)

- Magnetic transfer function

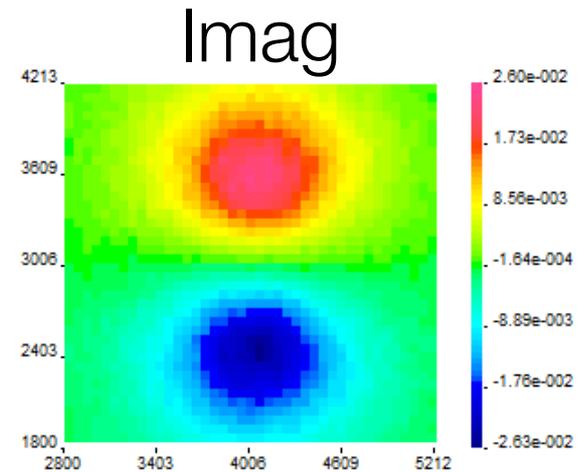
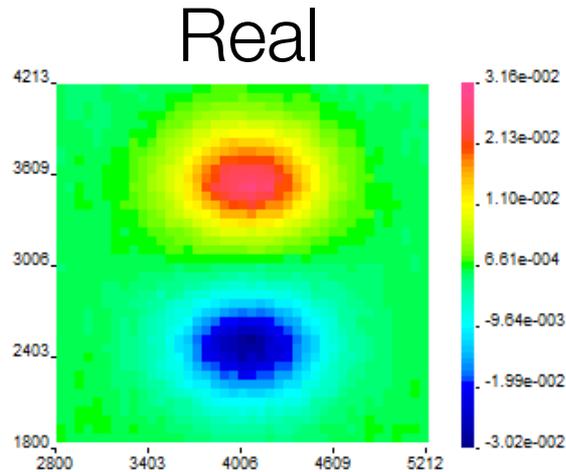
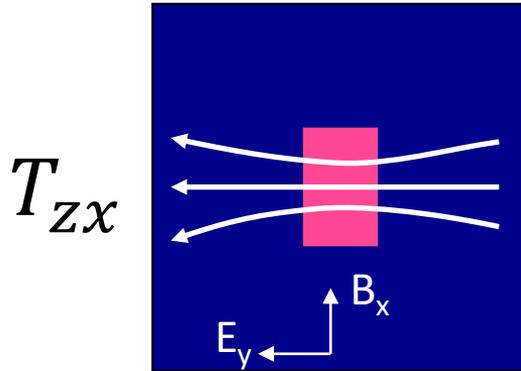
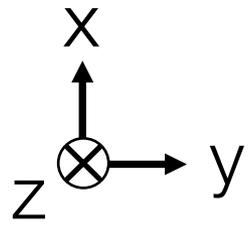
$$H_z = \mathbf{T}\mathbf{H}$$

$$H_z(r) = T_{zx}H_x(r_0) + T_{zy}H_y(r_0)$$

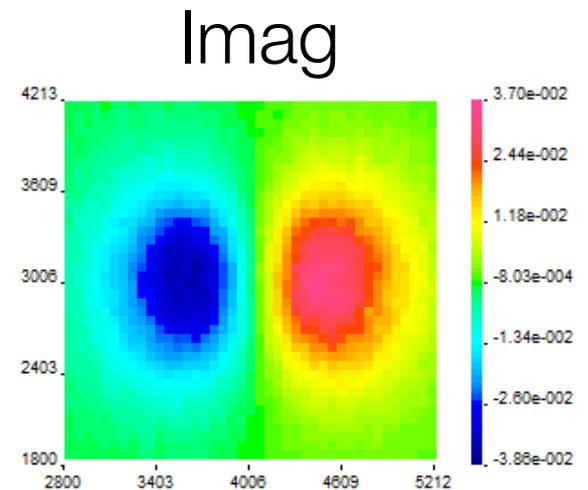
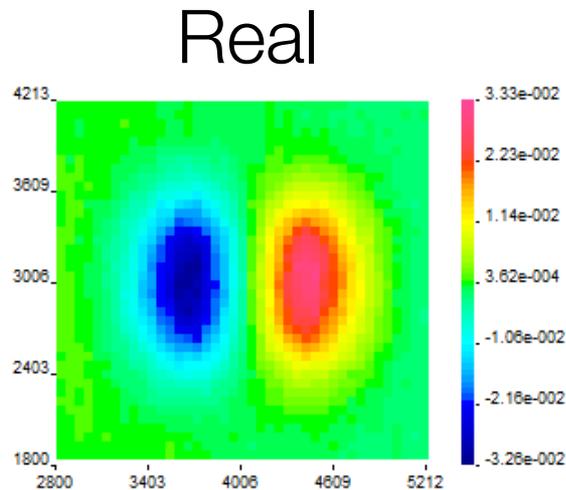
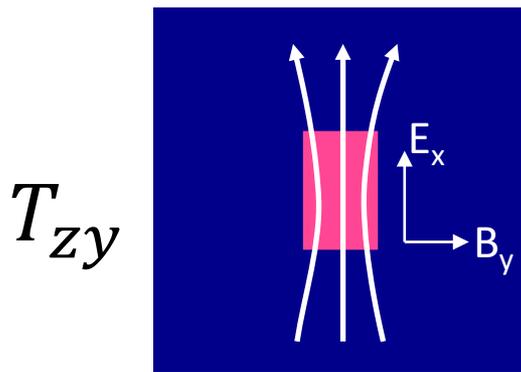
- Frequencies 30Hz – 720 Hz



# Synthetic example

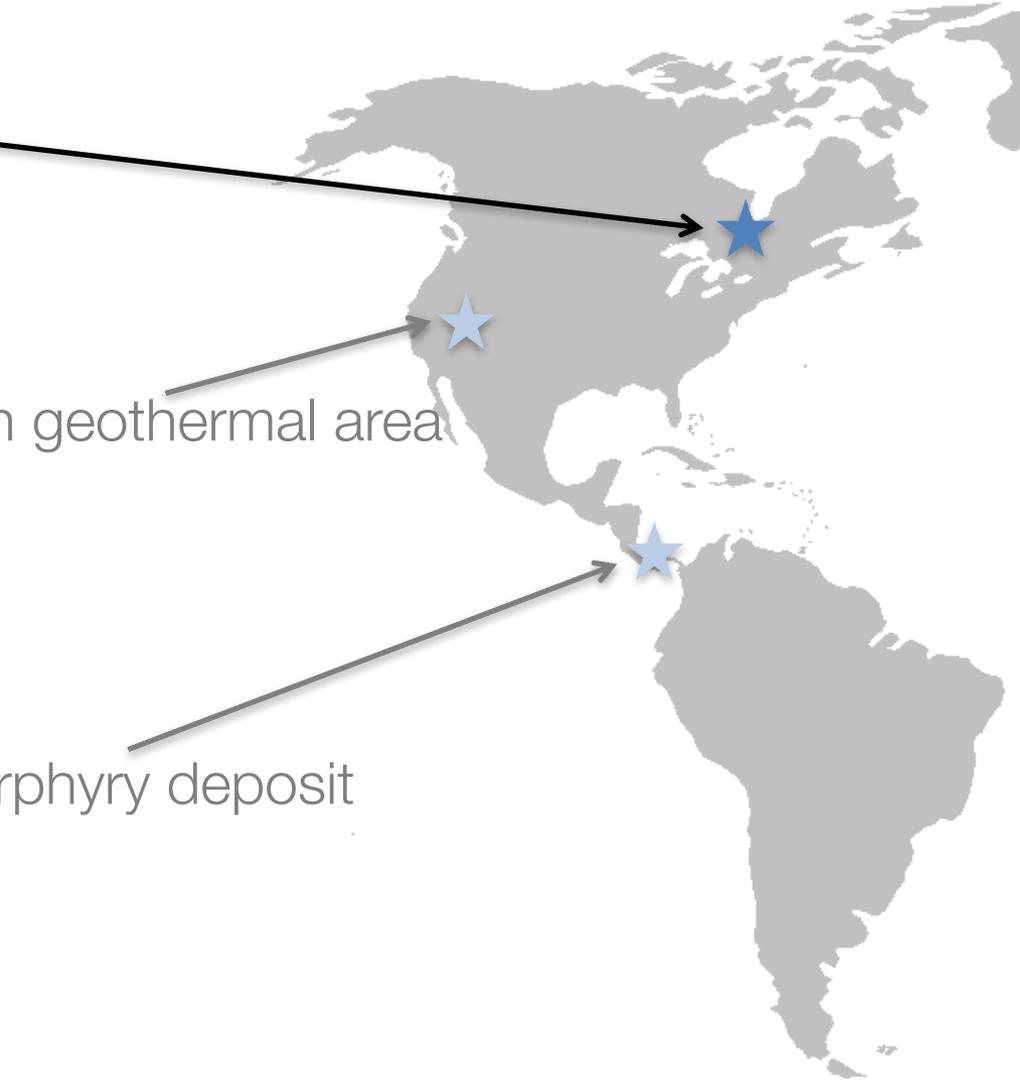


Conductor



# ZTEM case histories

- **Noranda district**

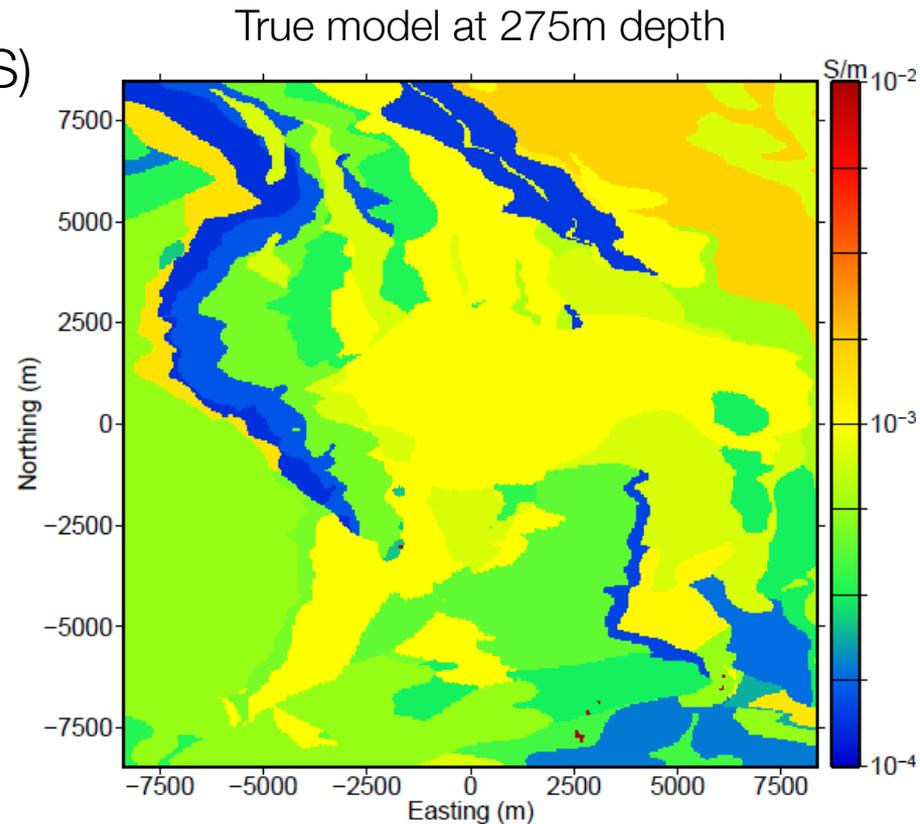


- Elevenmile Canyon geothermal area

- Balboa copper porphyry deposit

# Noranda district, Canada

- Hosts many deposits:
  - 20 economic volcanogenic massive sulphide deposits (VMS)
  - 19 orogenic gold deposits
  - Several intrusion-hosted Cu-Mo deposits
- Physical properties
  - Synthetic example from geologic model
  - 38 geologic units converted into expected conductivities

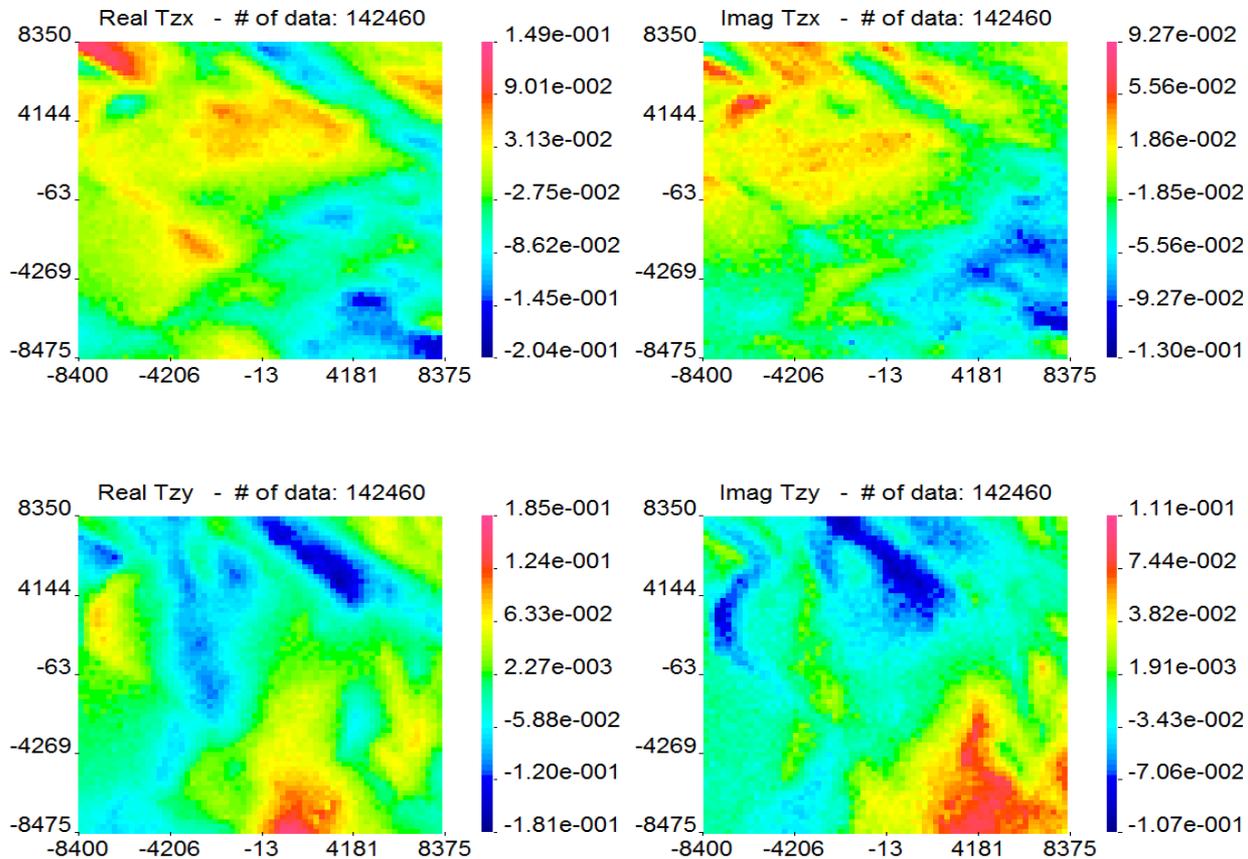
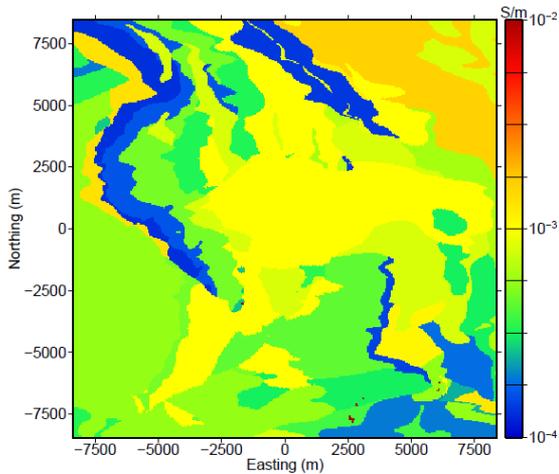


# Data

- Forward model data at 6 frequencies
  - 30, 45, 90, 180, 360, and 720 Hz
- Need to invert data

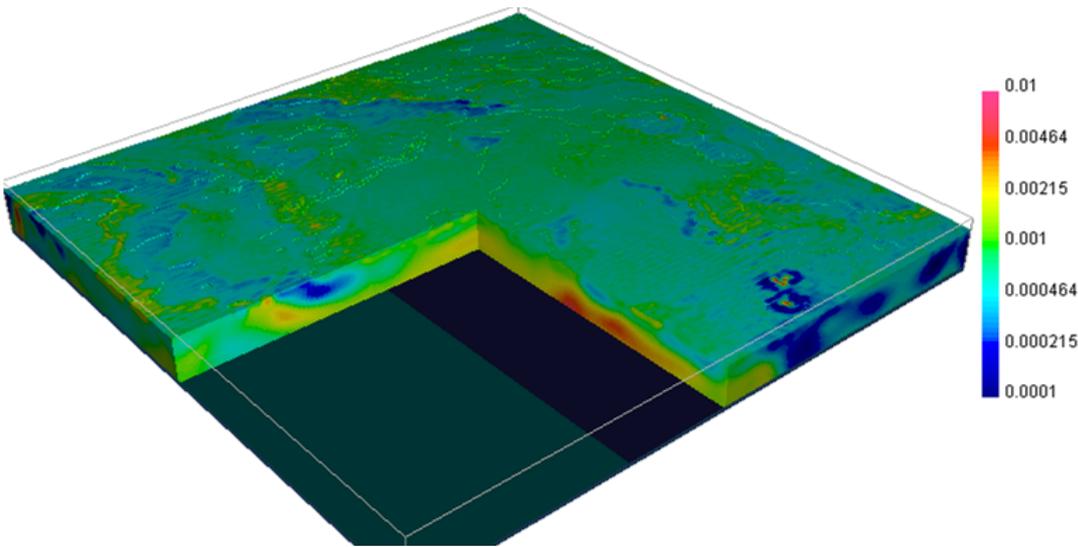
Observed (90 Hz)

True model at 275m depth

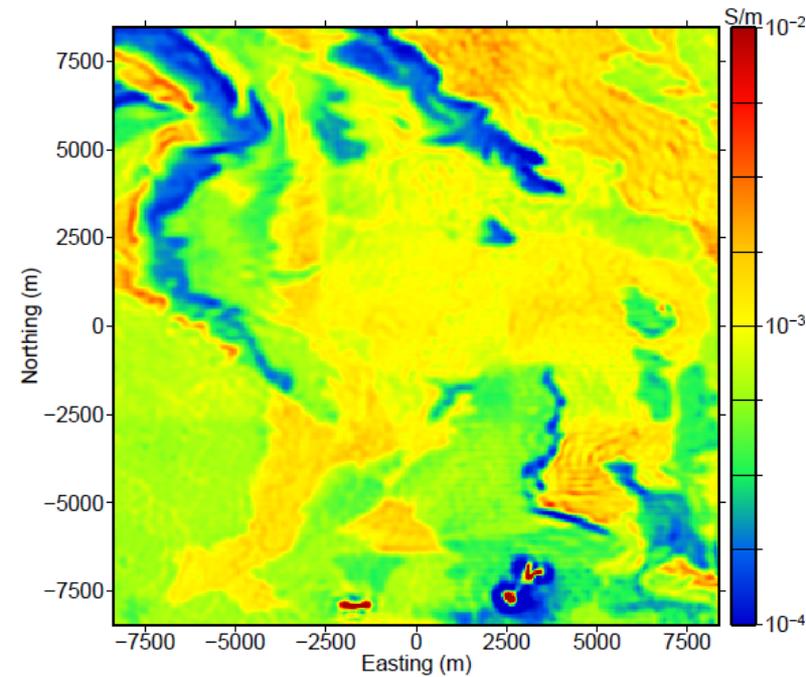


# Interpretation

## Recovered Model



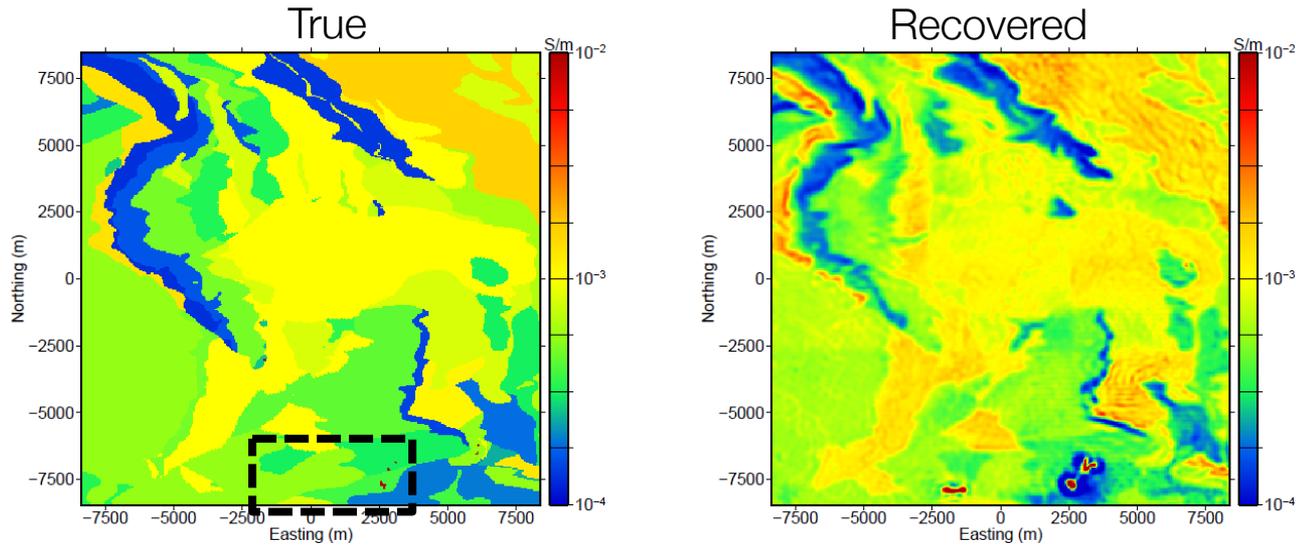
## Model at 275m depth



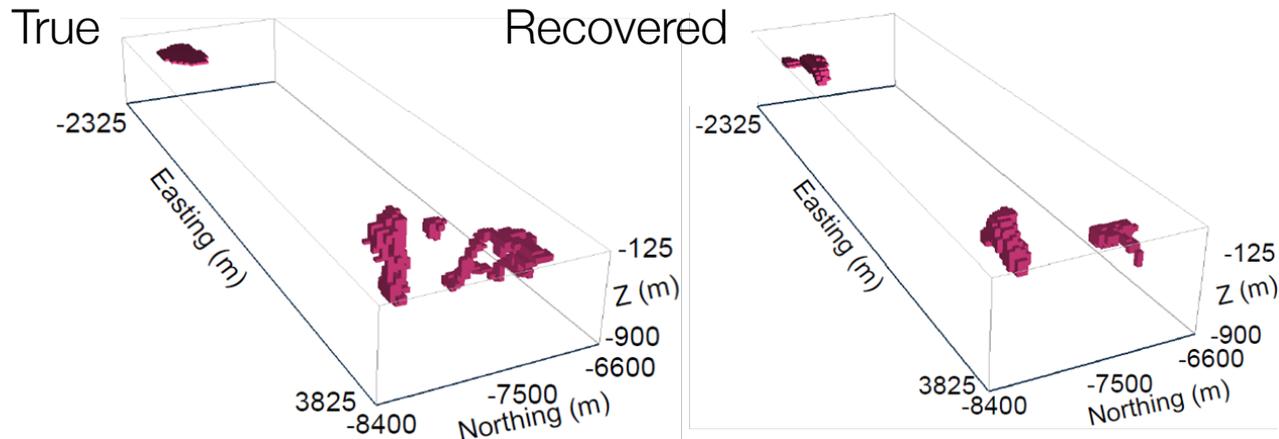
- Geologic units are well mapped
- Some mineralized bodies are located

# Synthesis

- Recovered model represents the regional geology

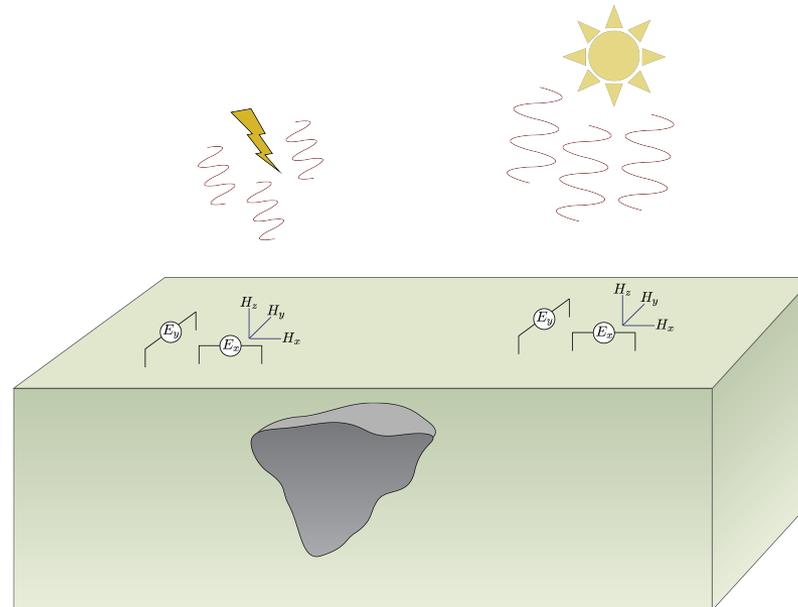


- Mineralized zones are recovered



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# End of Natural Sources

