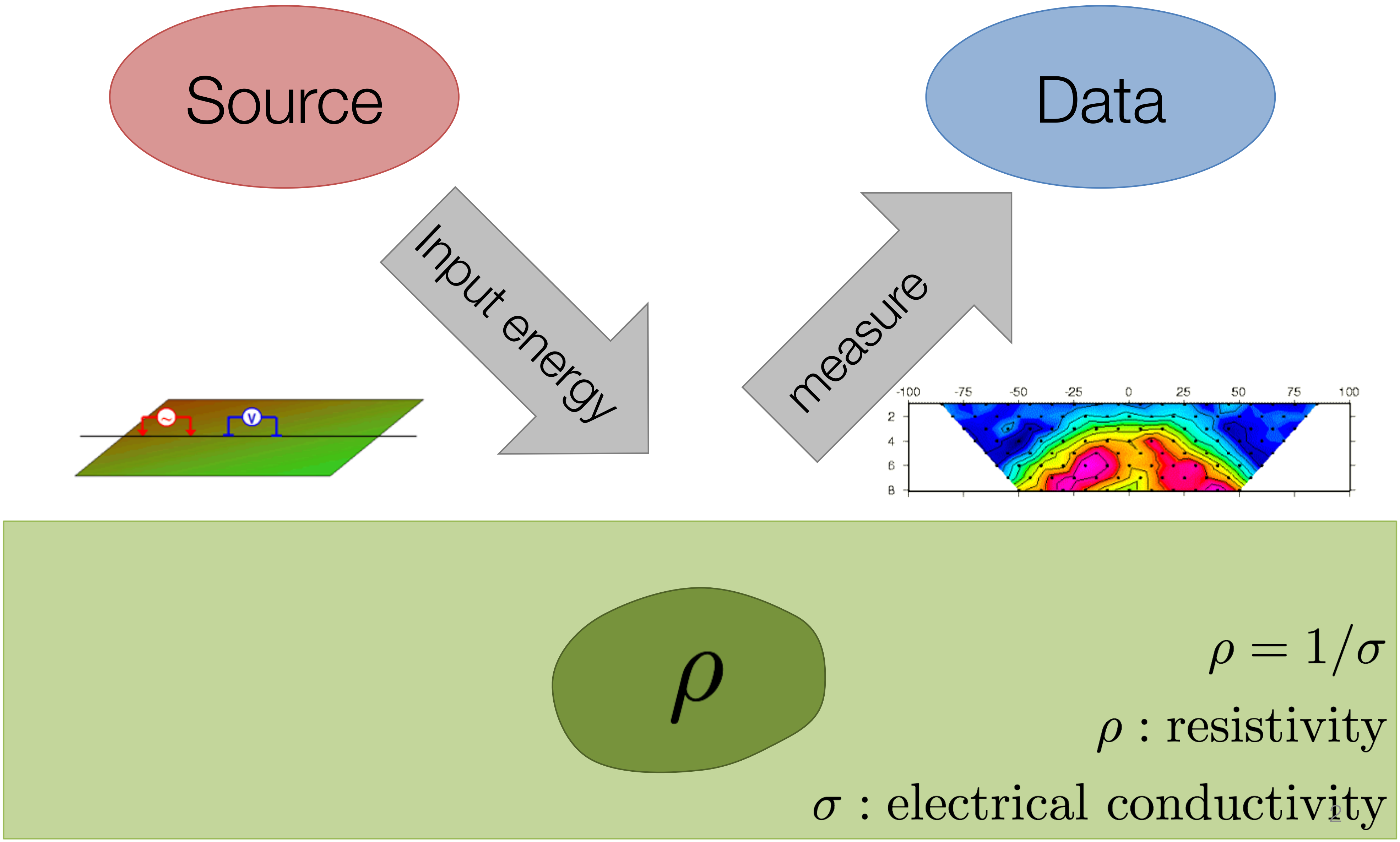
DC Resistivity

DISC

2017

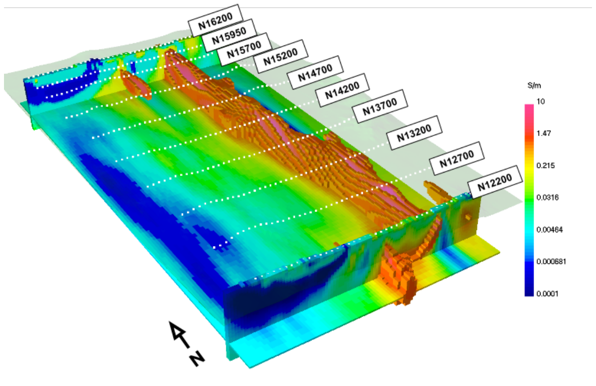
Electromagnetics
Fundamentals and Applications

DC Resistivity Survey



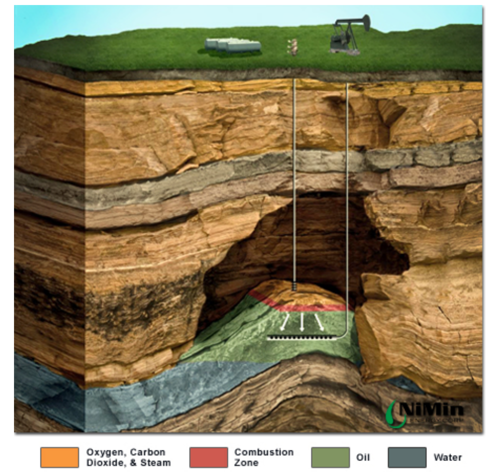
Motivation

Minerals

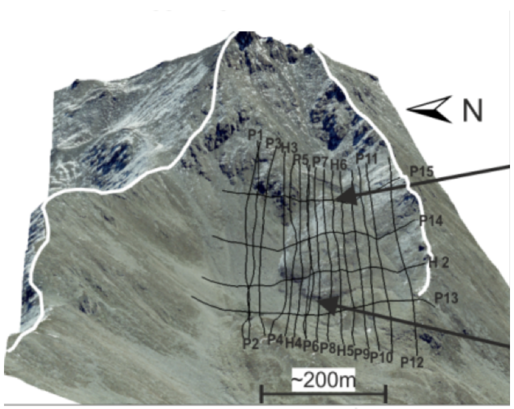


Water inflow in mine

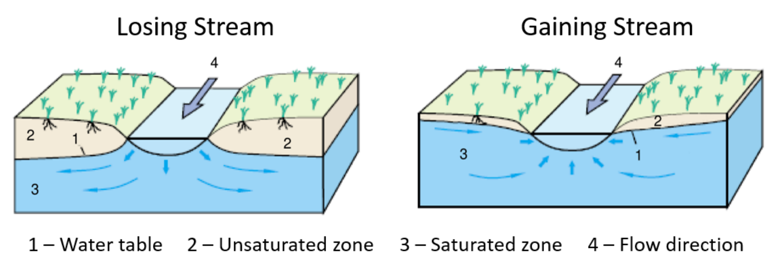
Oil and Gas



Geotechnical

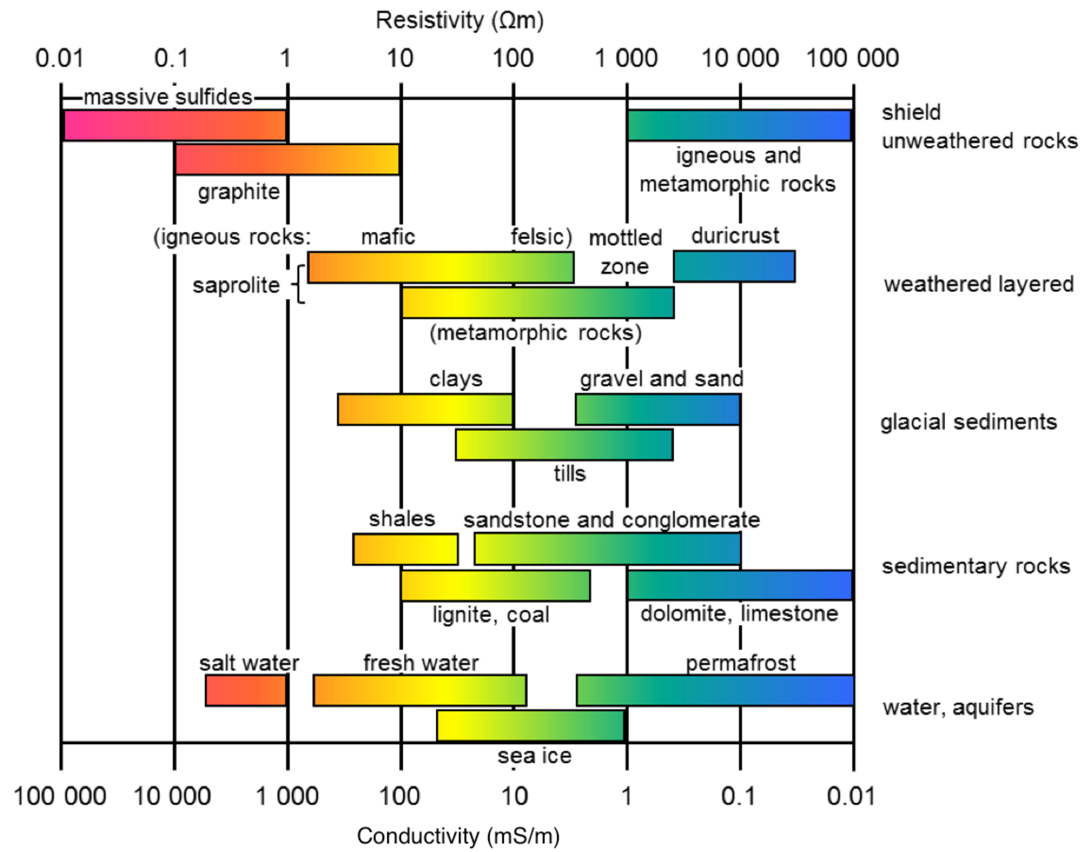


Groundwater



Electrical conductivity

- DC resistivity is sensitive to:
  - σ: Conductivity [S/m]
  - ρ: Resistivity [Ωm]
  - $\sigma = 1/\rho$
- Varies over many orders of magnitude
- Depends on many factors:
  - Rock type
  - Porosity
  - Connectivity of pores
  - Nature of the fluid
  - Metallic content of the solid matrix



Outline

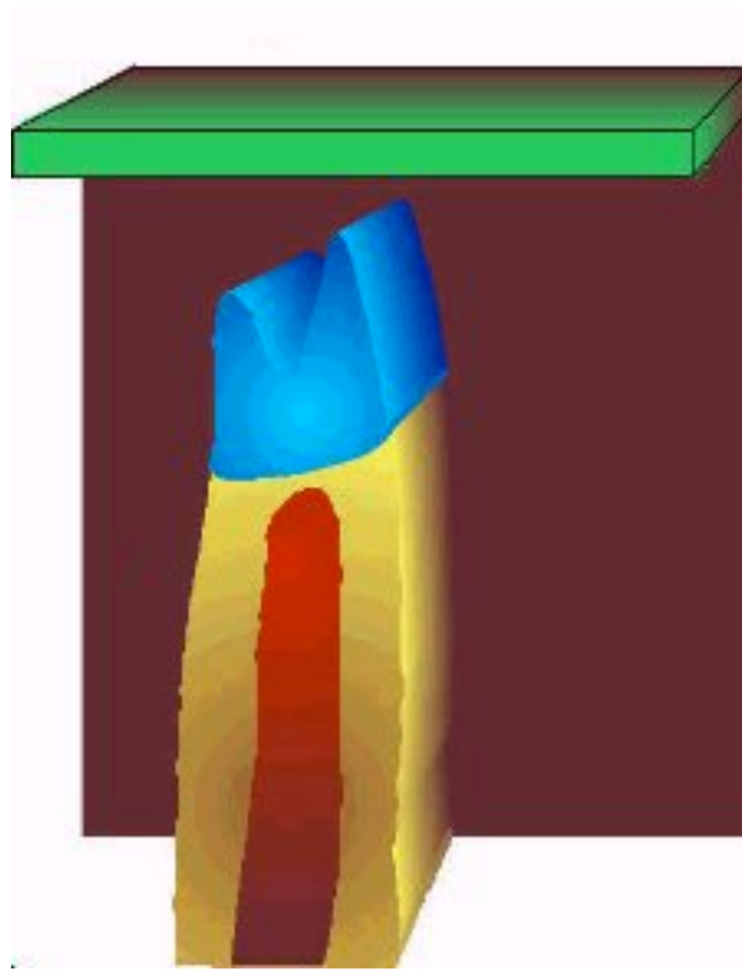
- Basic experiment
- Currents, charges, potentials and apparent resistivities
- Soundings, profiles and arrays
- Data, pseudosections and inversion
- Sensitivity
- Survey Design
- Case History – Mt Isa
- Effects of background resistivity

Basic Experiment

- **Target:**
  - Ore body. Mineralized regions less resistive than host

Elura Orebody Electrical resistivities

Rock Type	*Ohm-m*
Overburden	12
Host rocks	200
Gossan	420
Mineralization (pyritic)	0.6
Mineralization (pyrrhotite)	0.6



Basic Experiment

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- **Setup:**
  - Tx: Current electrodes
  - Rx: Potential electrodes

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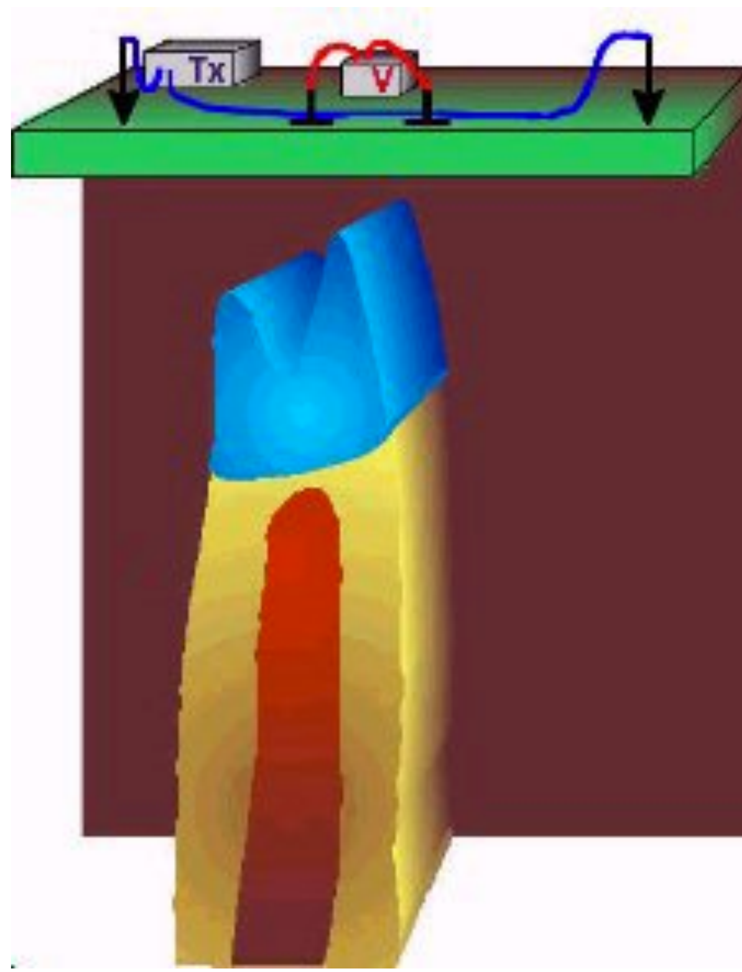

Basic Experiment

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  - Preferentially flow through conductors

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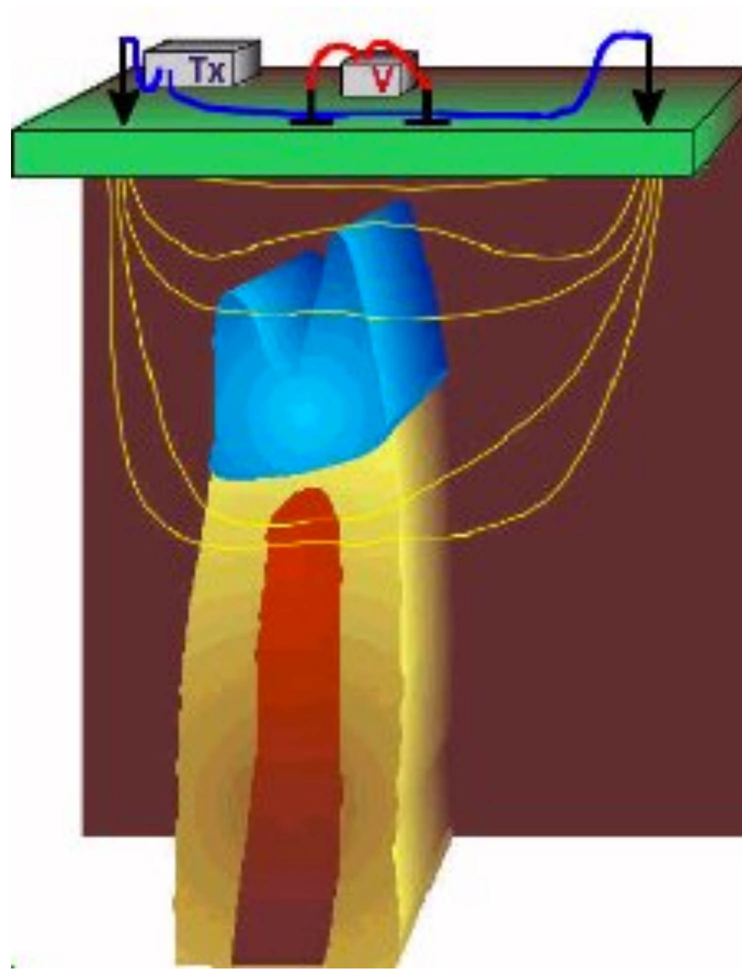

Basic Experiment

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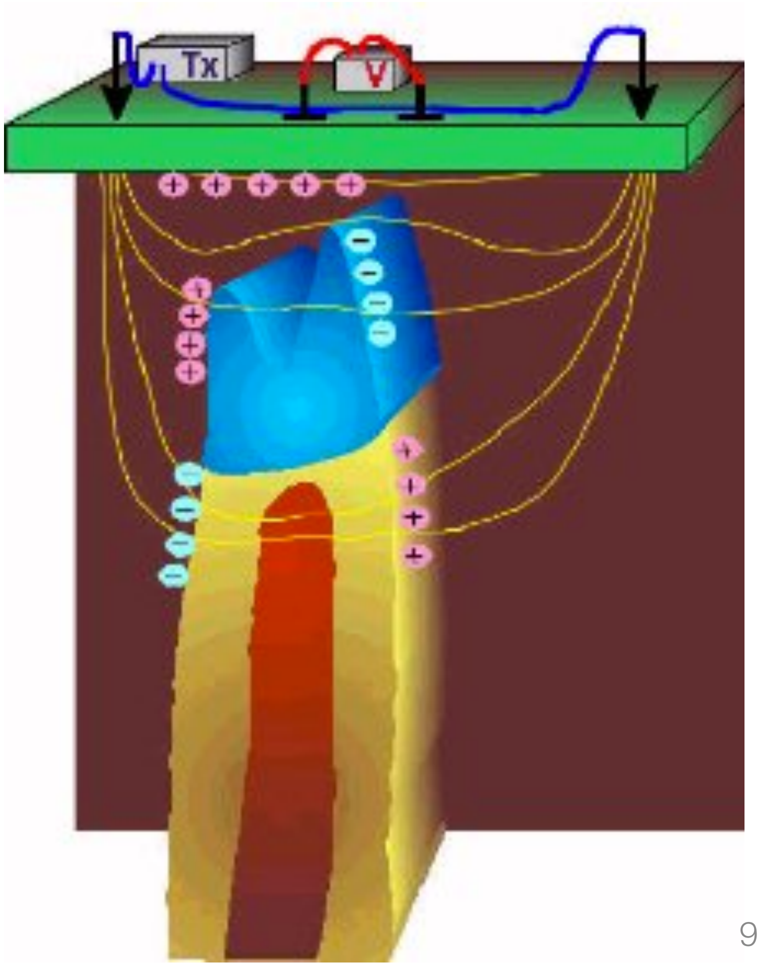
- **Charges:**
  - Build up at interfaces



Basic Experiment

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- **Setup:**
  - Tx: Current electrodes
  - Rx: Potential electrodes
- **Currents:**
  - Preferentially flow through conductors
- **Charges:**
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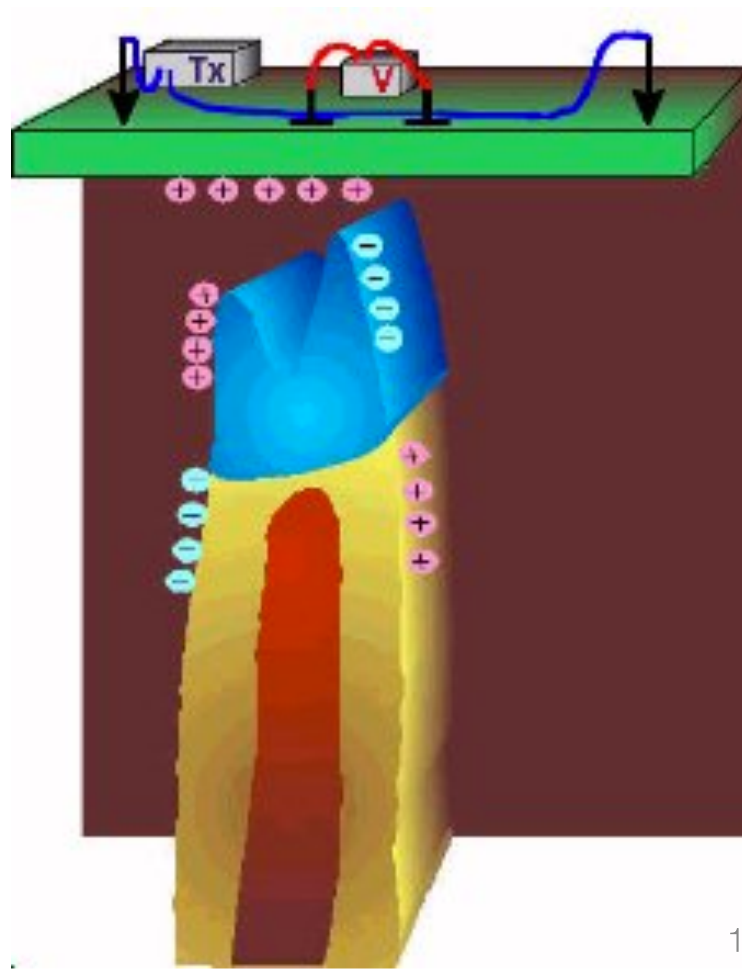
- **Potentials:**
  - Associated with the charges are measured at the surface



How do we obtain resistivity?

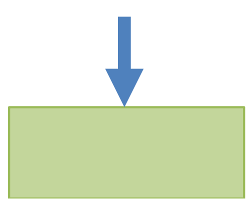
Steady State Maxwell equations

	Full	Steady State
Faraday	$\nabla \times \vec{e} = -\frac{\partial \vec{b}}{\partial t}$	$\nabla \times \vec{e} = 0 \qquad \vec{e} = -\nabla V$
Ampere	$\nabla \times \vec{h} = \vec{j} + \frac{\partial \vec{d}}{\partial t} + \vec{j}_s$	$\nabla \cdot \vec{j} = -\nabla \cdot \vec{j}_s$
Ohm's Law	$\vec{j} = \sigma \vec{e}$	

Put it together

$$\nabla \cdot \sigma \nabla V = I\delta(r)$$

Potential in a homogeneous halfspace

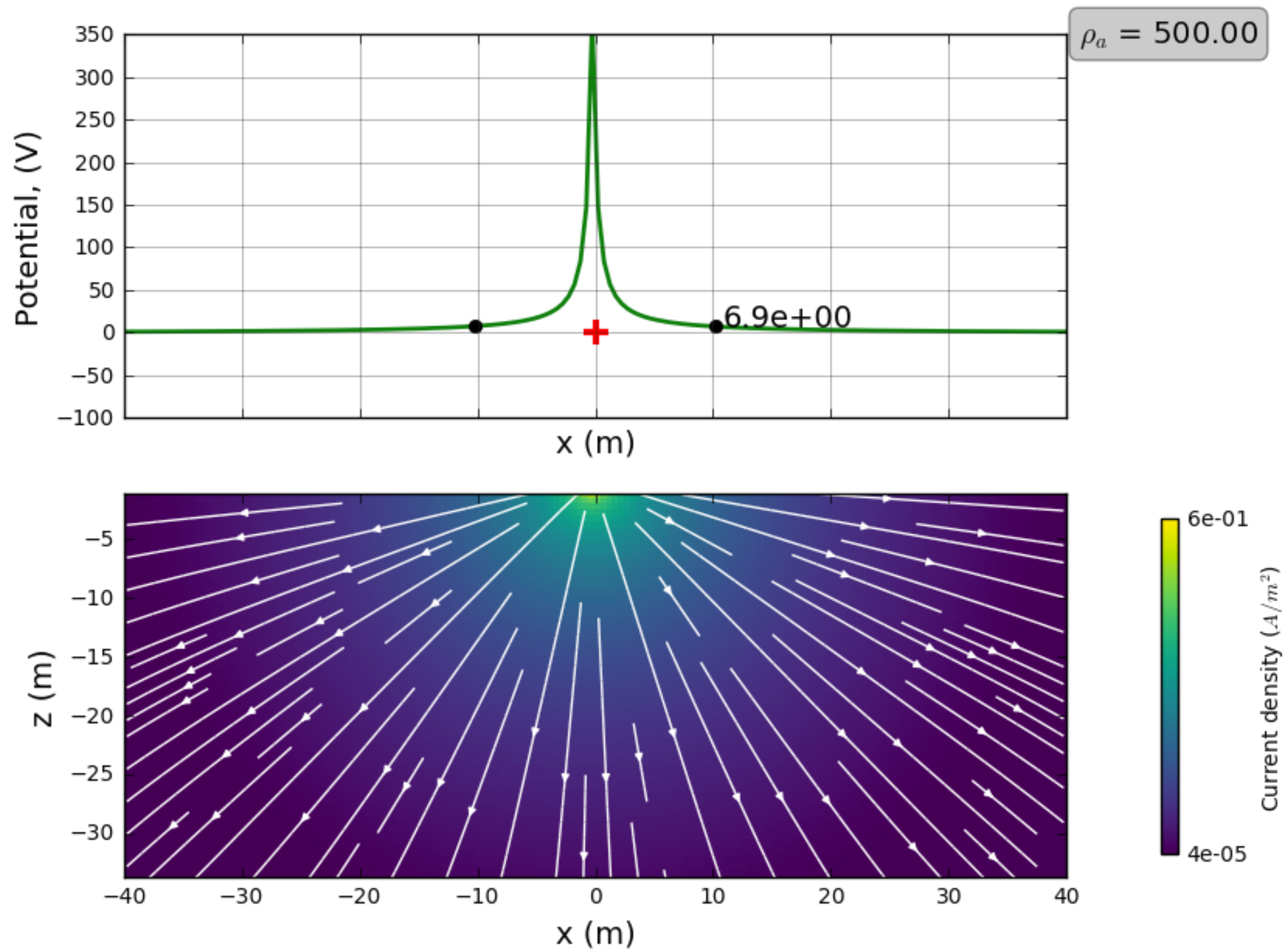


$$V = \frac{I}{2\pi\sigma}\frac{1}{r} \qquad V = \frac{\rho I}{2\pi r}$$

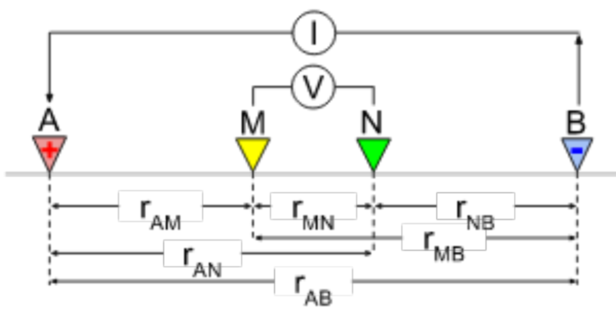
Currents and potentials: halfspace

$$V = \frac{\rho I}{2\pi r}$$

$$\rho = \frac{2\pi r V}{I}$$



Currents and potentials: 4-electrode array

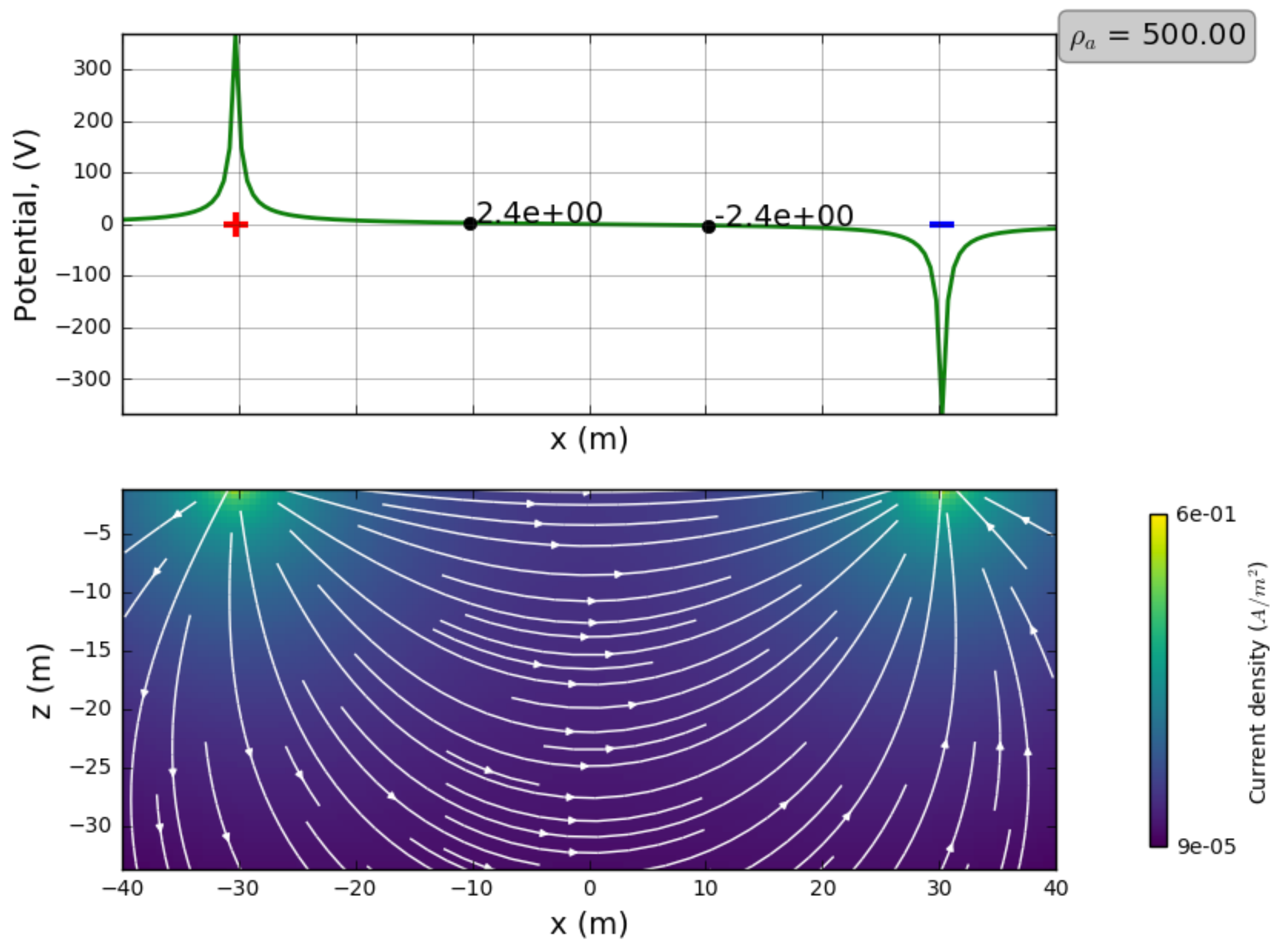


$$\Delta V_{MN} = \rho I \underbrace{\frac{1}{2\pi}\left[\frac{1}{AM} - \frac{1}{MB} - \frac{1}{AN} + \frac{1}{NB}\right]}_{G}$$

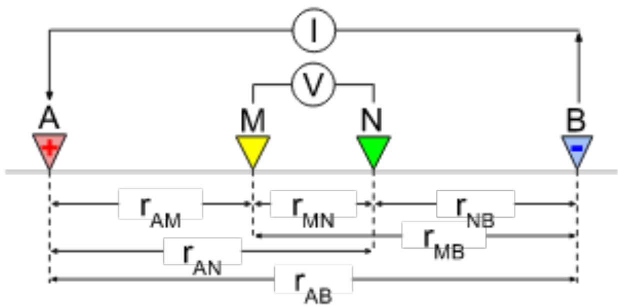
Resistivity

$$\rho = \frac{\Delta V_{MN}}{IG}$$

Halfspace (500 Ωm)



Currents and Apparent Resistivity

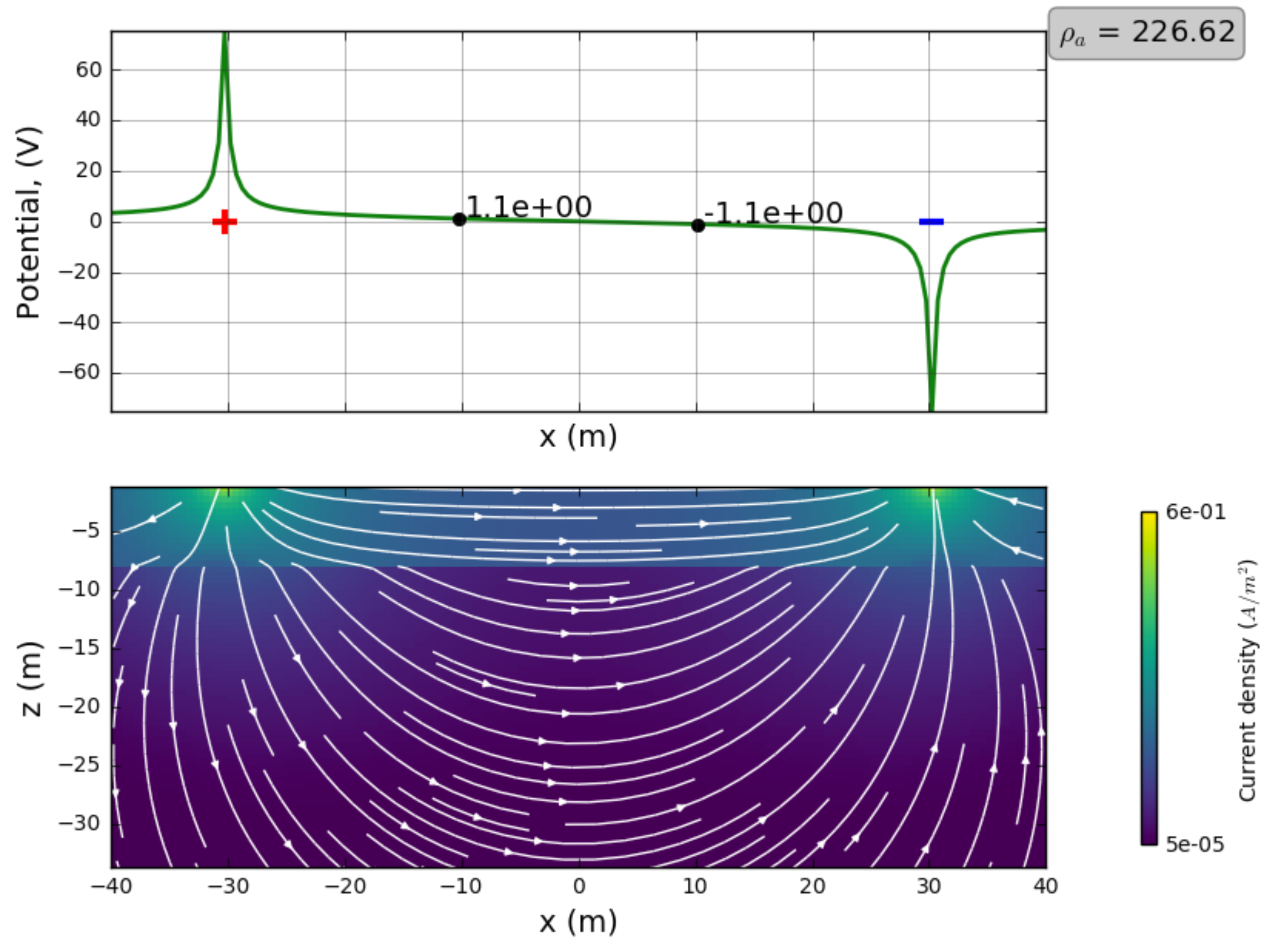


$$\Delta V_{MN} = \rho I \underbrace{\frac{1}{2\pi}\left[\frac{1}{AM} - \frac{1}{MB} - \frac{1}{AN} + \frac{1}{NB}\right]}_{G}$$

Apparent resistivity

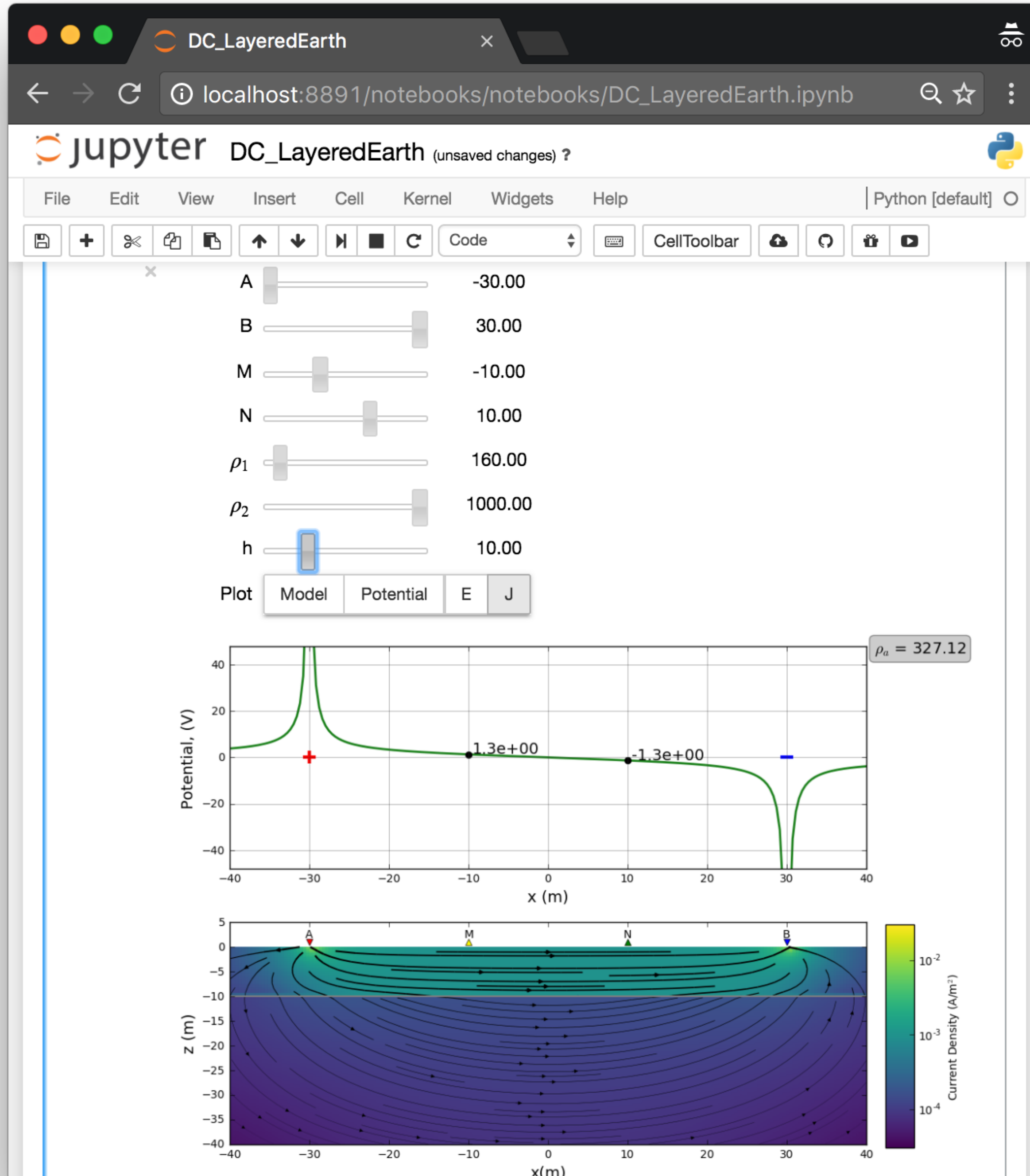
$$\rho_a = \frac{\Delta V_{MN}}{IG}$$

Conductive overburden (100 Ωm)



Why interactive apps?

- Visualization aids understanding
- Learn through interaction
  - ask questions and investigate
- Open source:
  - Free to use
  - Welcome contributions!

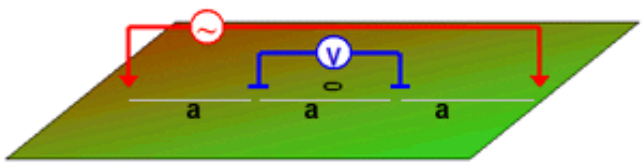


http://em.geosci.xyz/apps.html

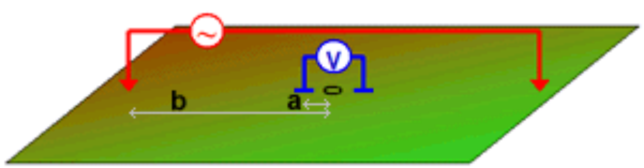
Soundings and Arrays

Geometry

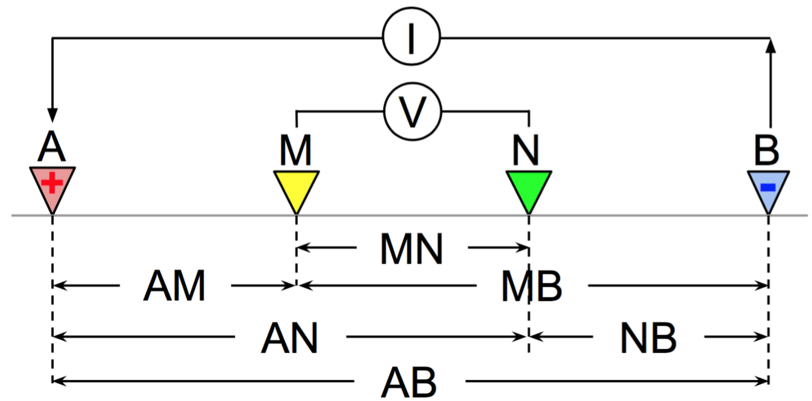
Wenner



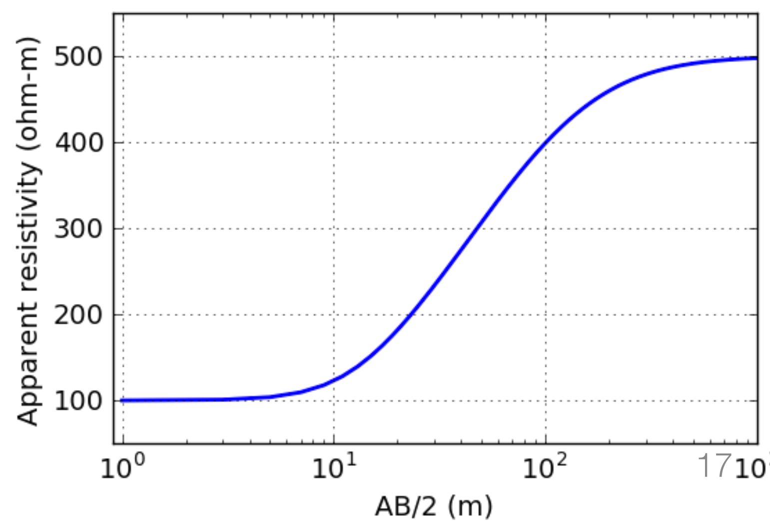
Schlumberger



4 electrode Array

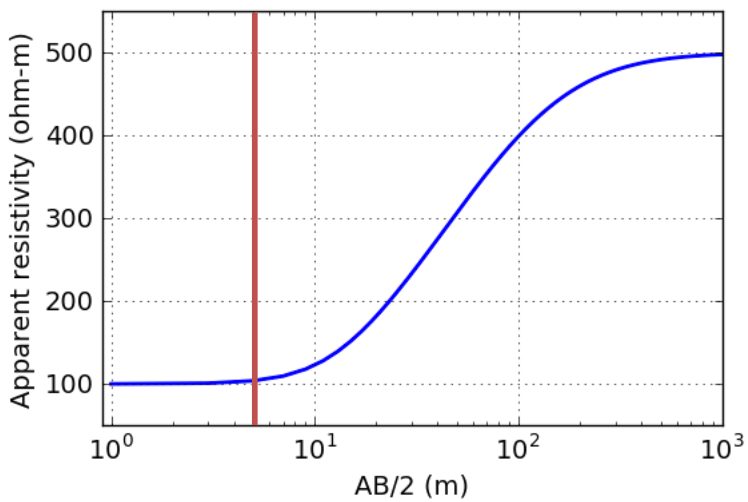


Sounding

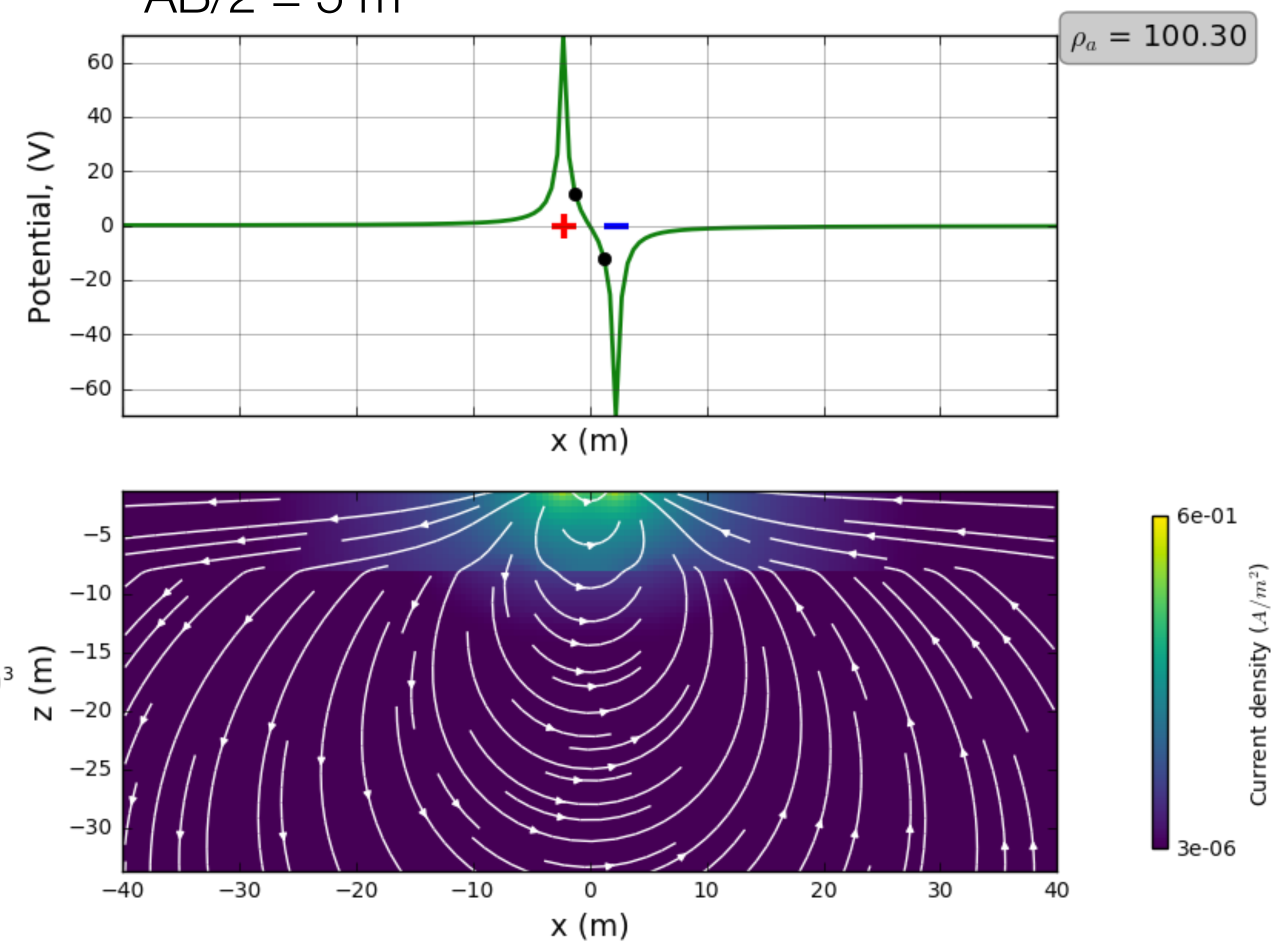


Soundings

DC Sounding curve



AB/2 = 5 m



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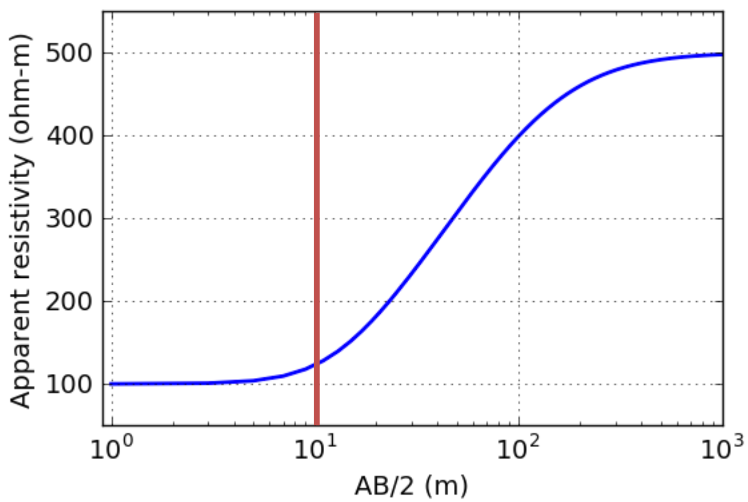
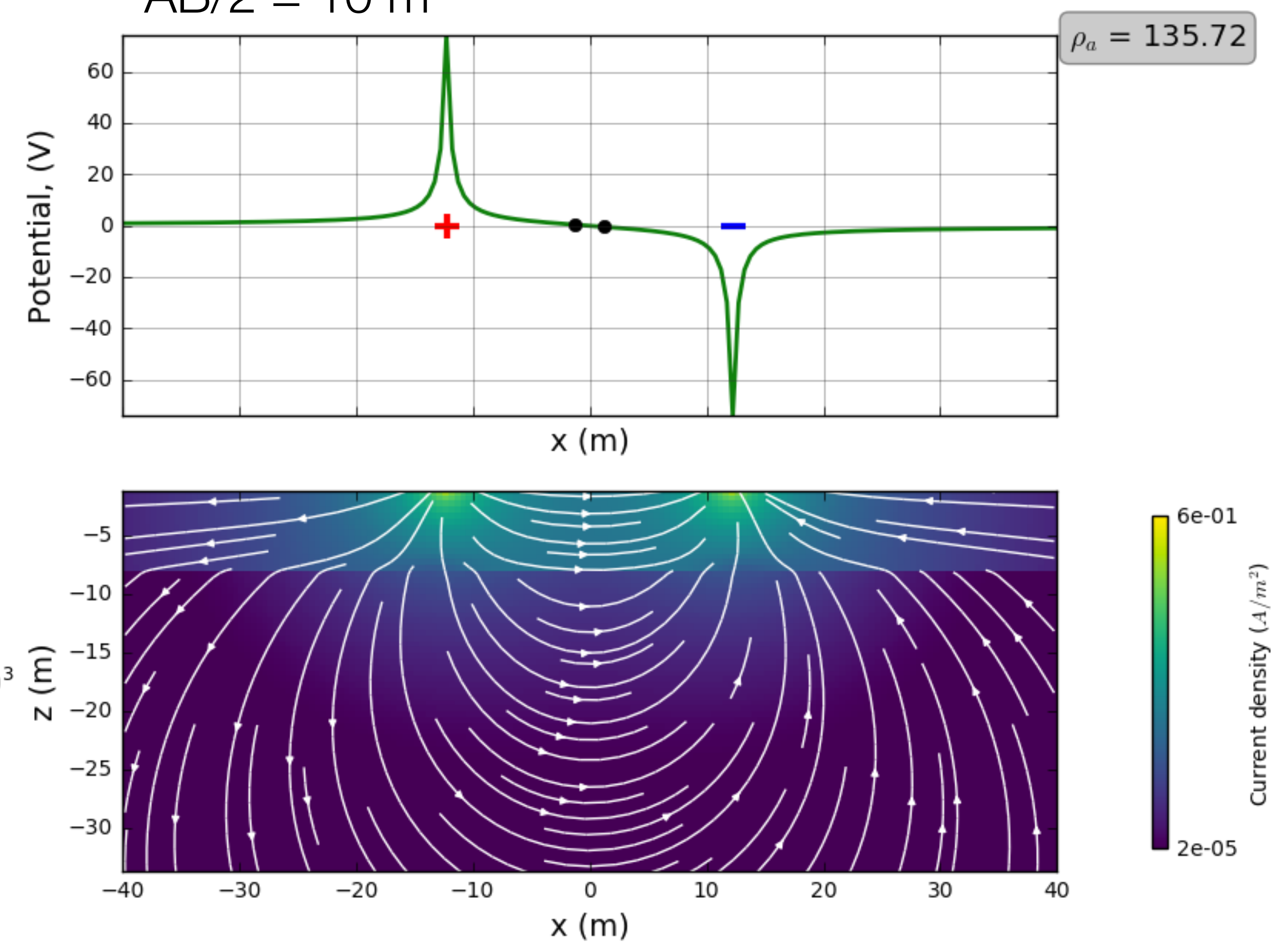
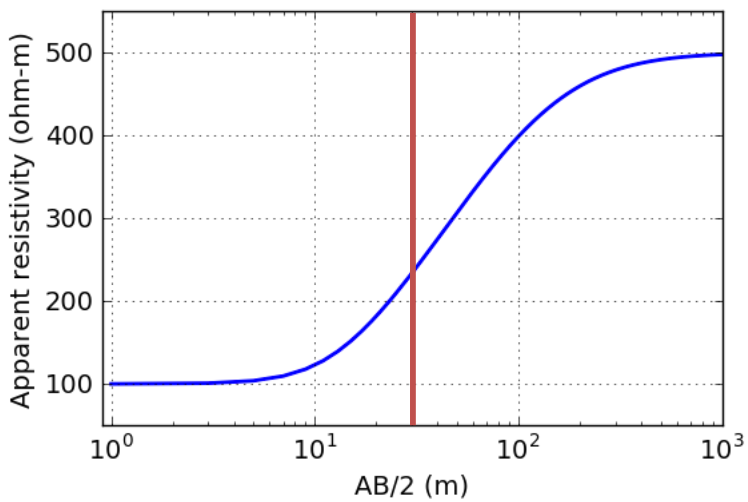


AB/2 = 10 m

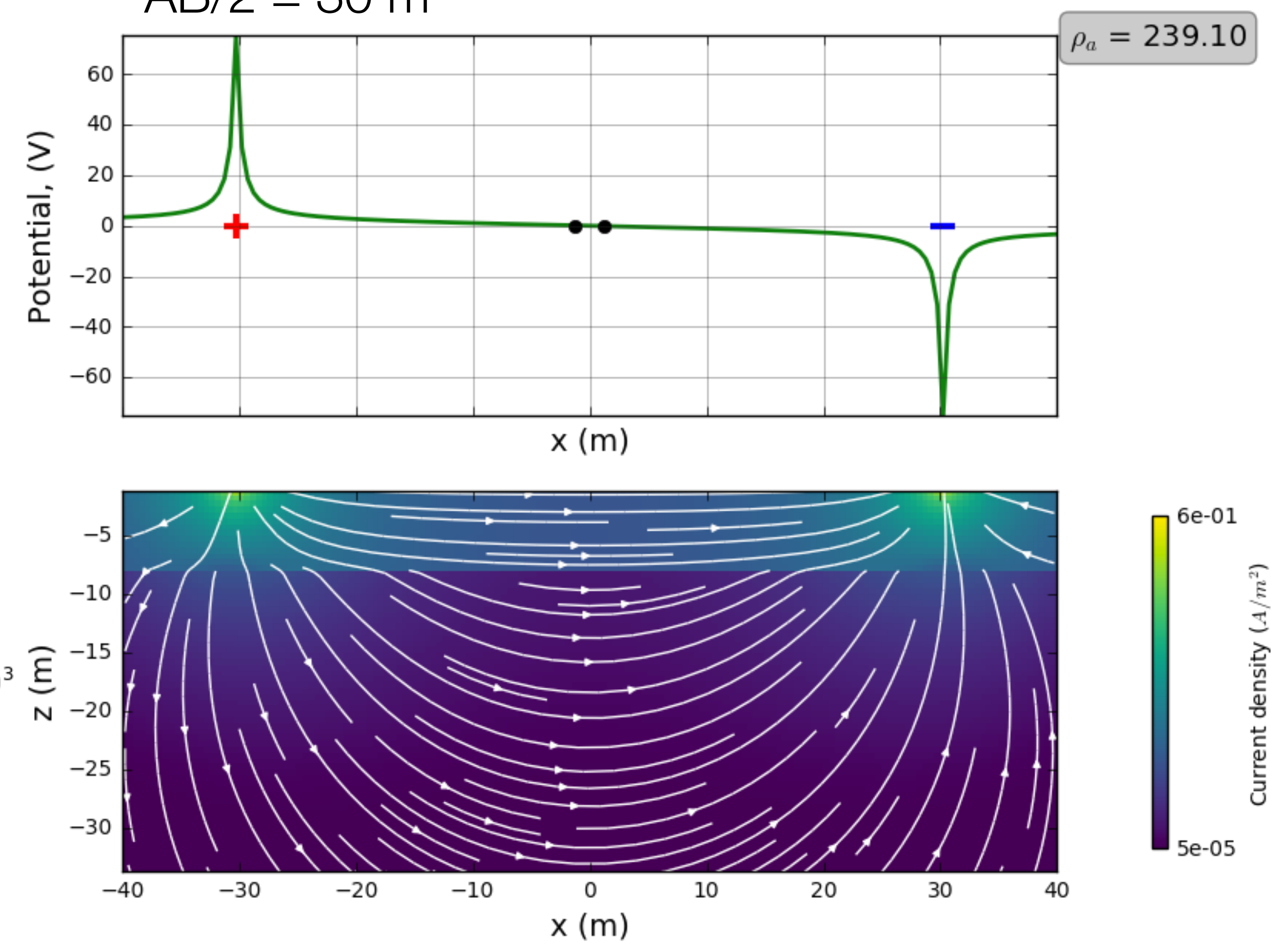


Soundings

DC Sounding curve

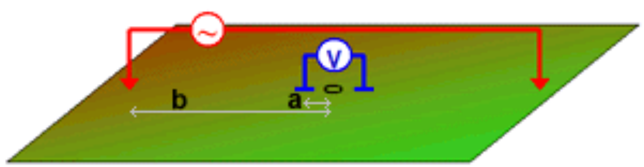


AB/2 = 30 m

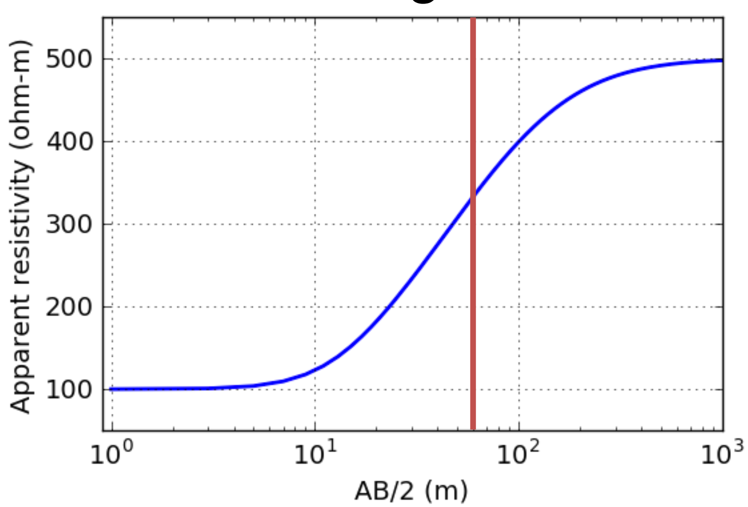


Summary: soundings

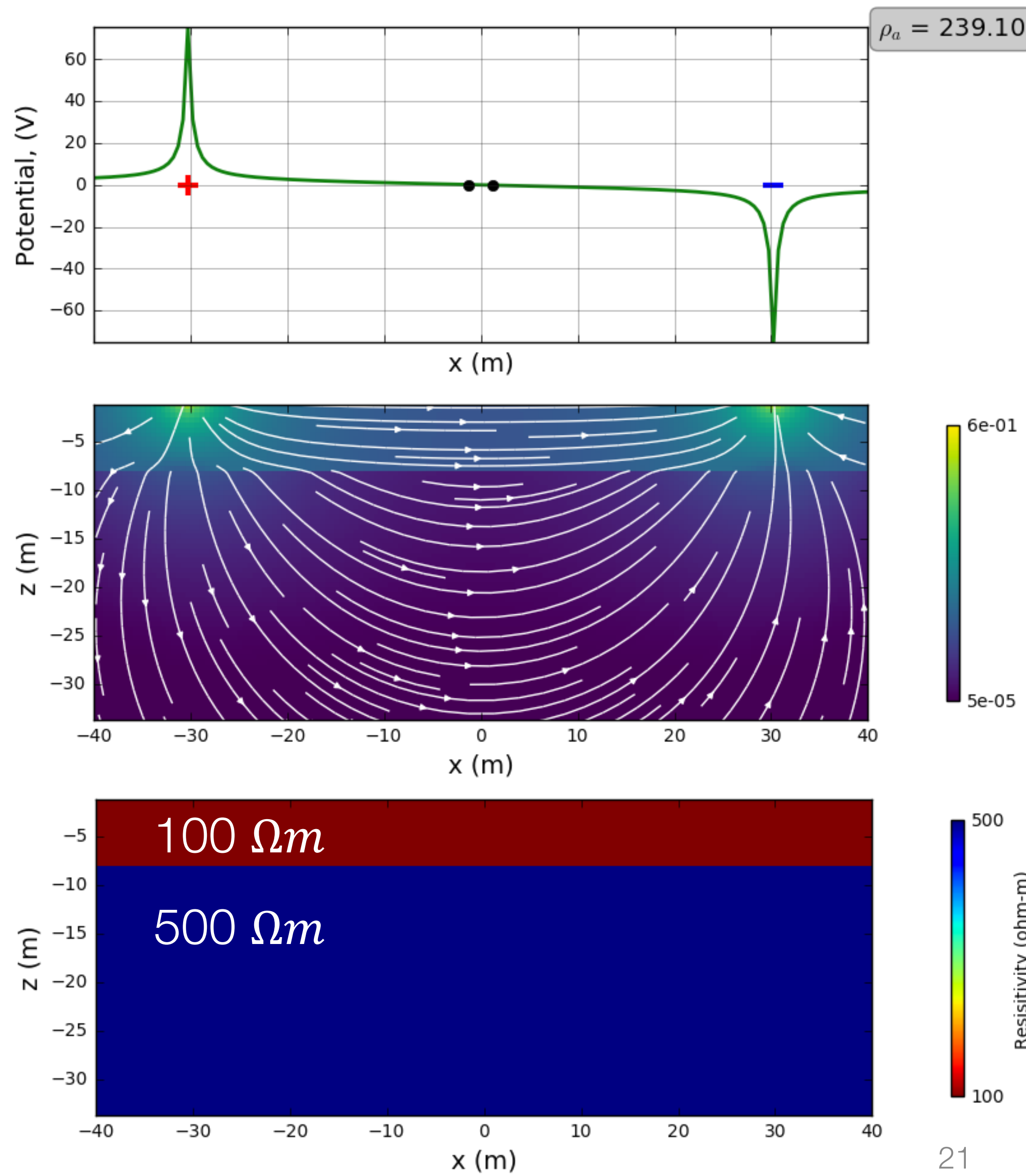
Schlumberger array



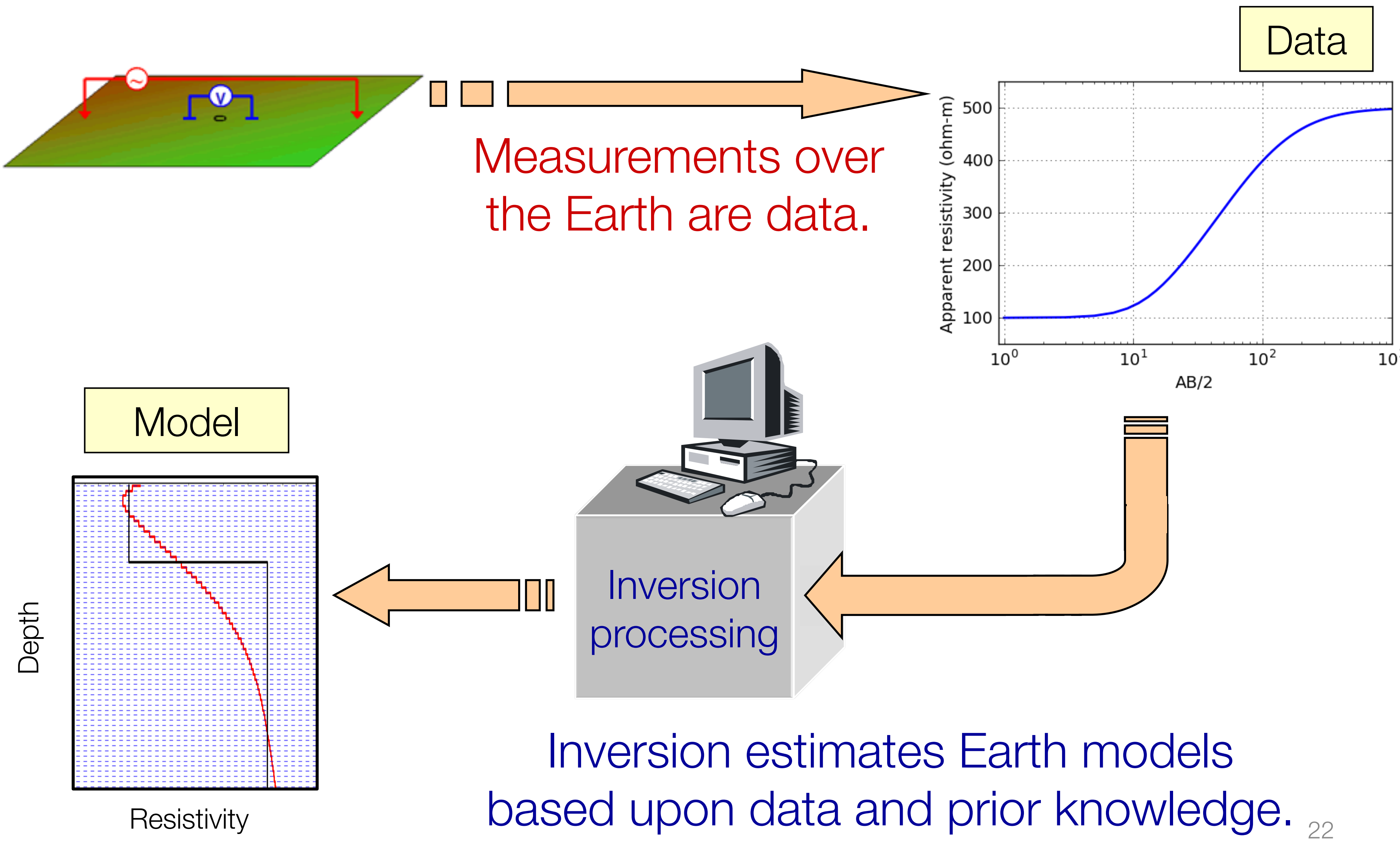
DC **Sounding** curve



Scale length of array must be large to see deep



Inversion


Data
Measurements over the Earth are data.
Apparent resistivity (ohm-m)
100
200
300
400
500
10^0
10^1
10^2
10^3
AB/2
Model
Depth
Resistivity
Inversion processing
Inversion estimates Earth models based upon data and prior knowledge.

DCR for a confined body

- Useful to formally bring in the concept of charges

Normal component of current density is continuous

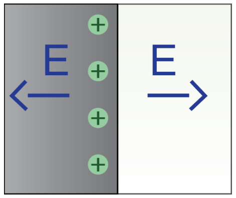
$$J_{1n} = J_{2n}$$

$$\sigma_1 E_{1n} = \sigma_2 E_{2n}$$

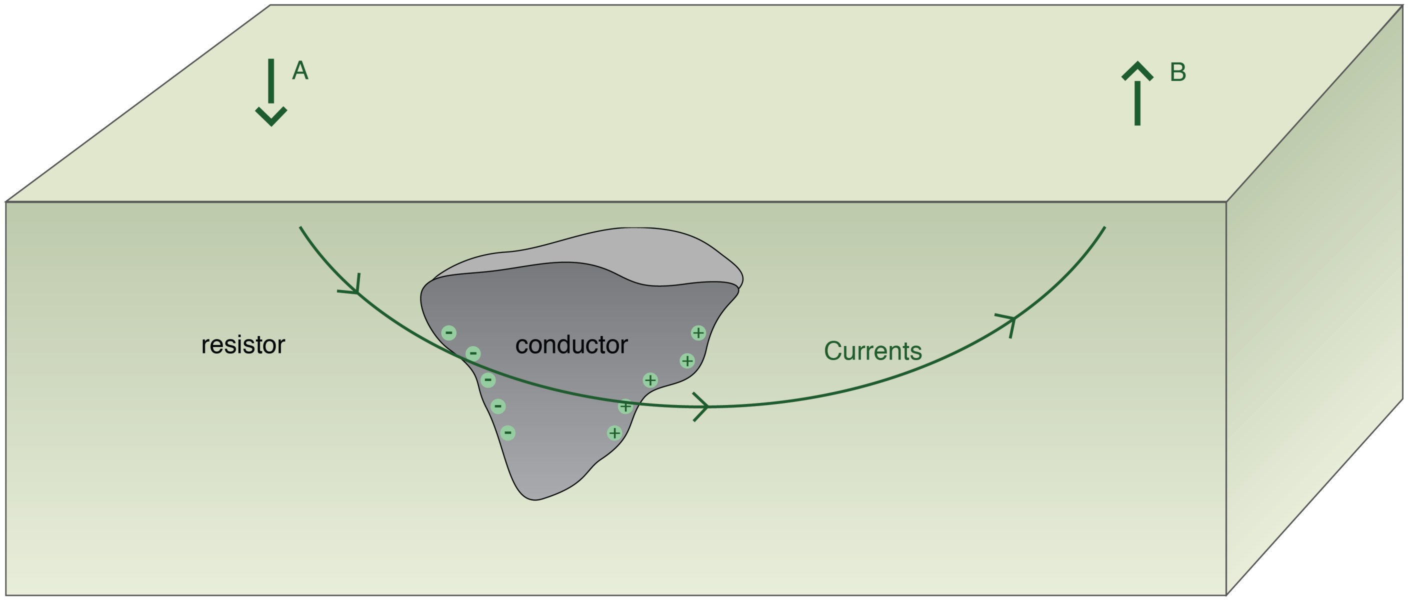
Conductivity contrast

$$\sigma_1 \neq \sigma_2$$

- Electric field discontinuous
- Charge build-up

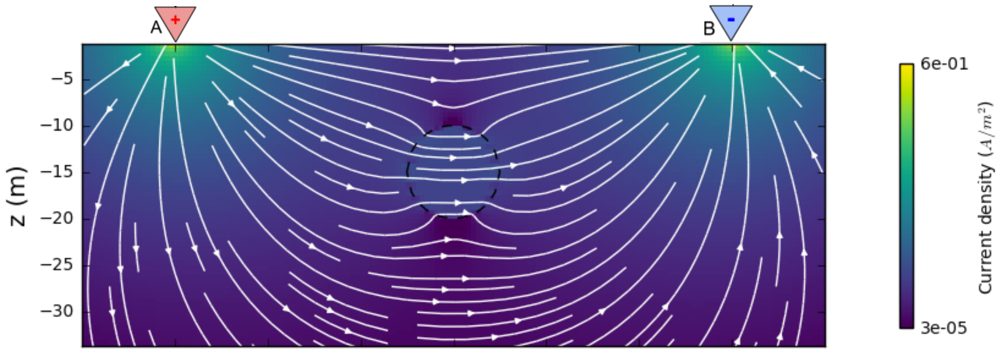


$$\mathbf{E} = \frac{Q}{4\pi\varepsilon_0 |\mathbf{r} - \mathbf{r}'|^2}\hat{\mathbf{r}}$$

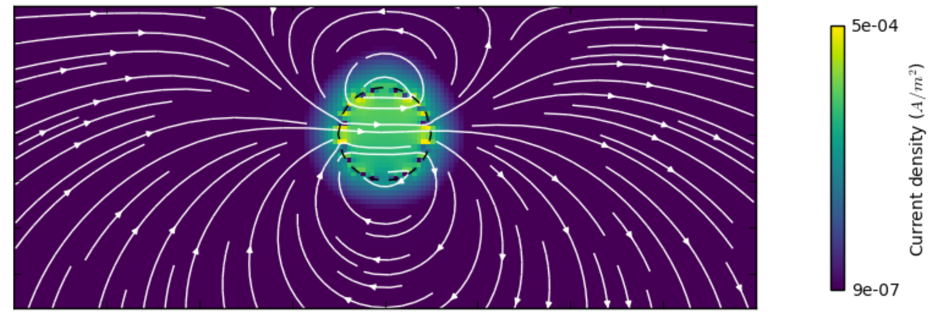


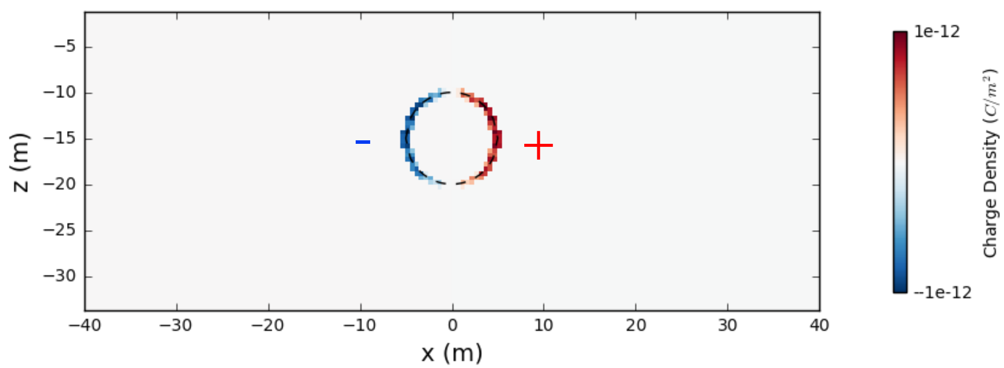
Currents, charges, and potentials

Total currents: J

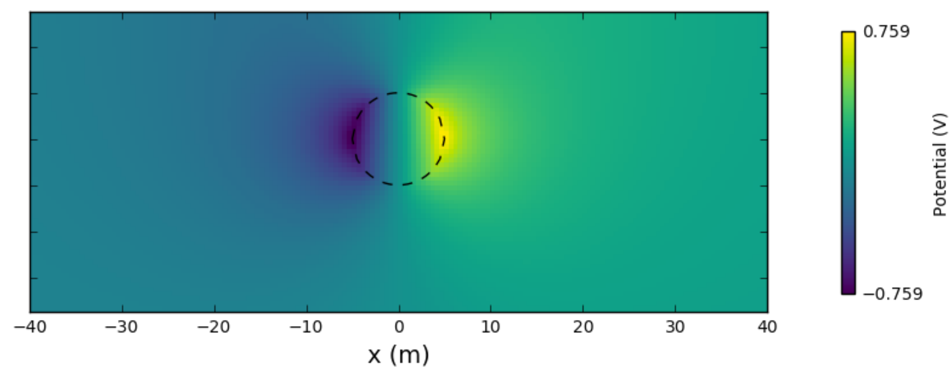


Secondary currents: J_s





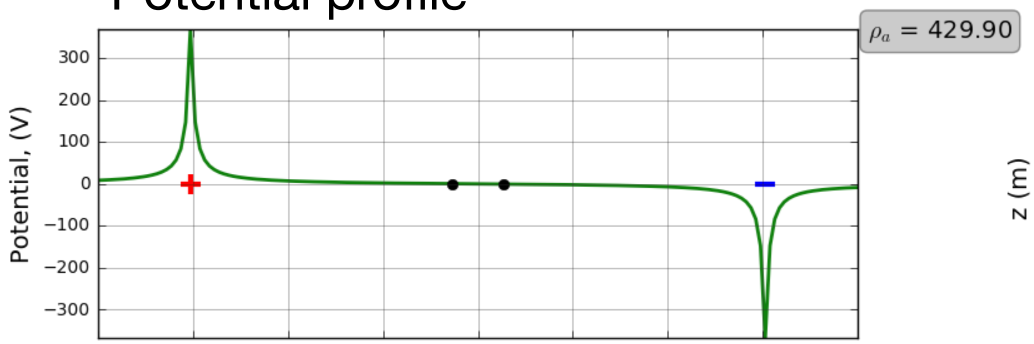
Secondary charges: Q_s



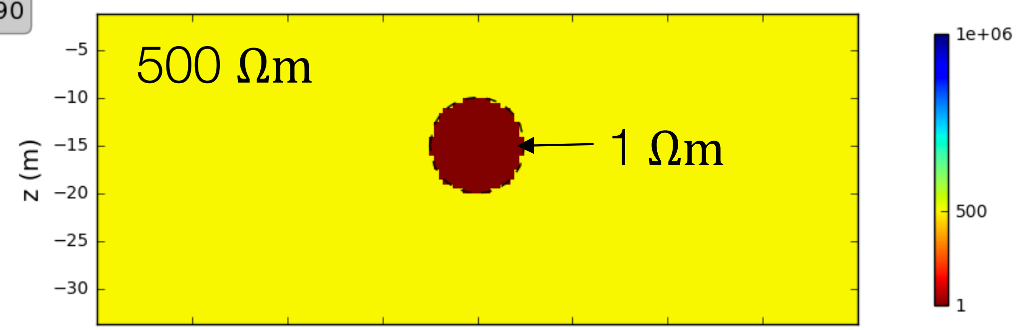
Secondary potential: ϕ_s

Measurements of DC data: gradient array

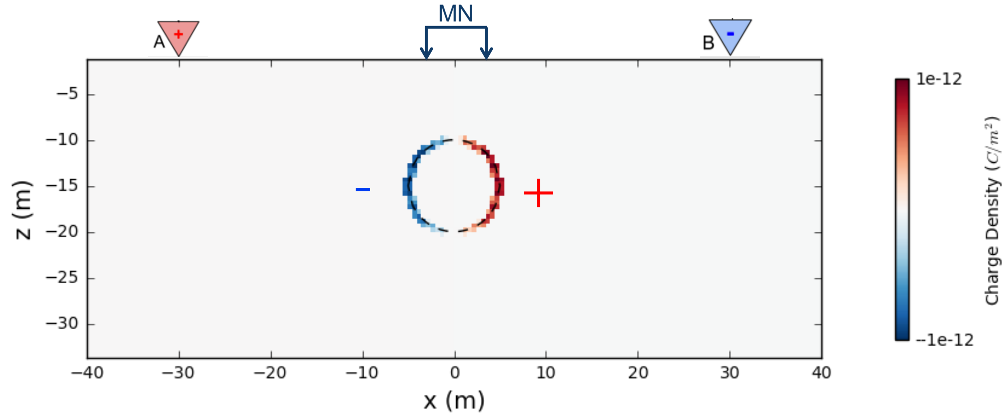
Potential profile



Resistivity model



ρ_a = 430



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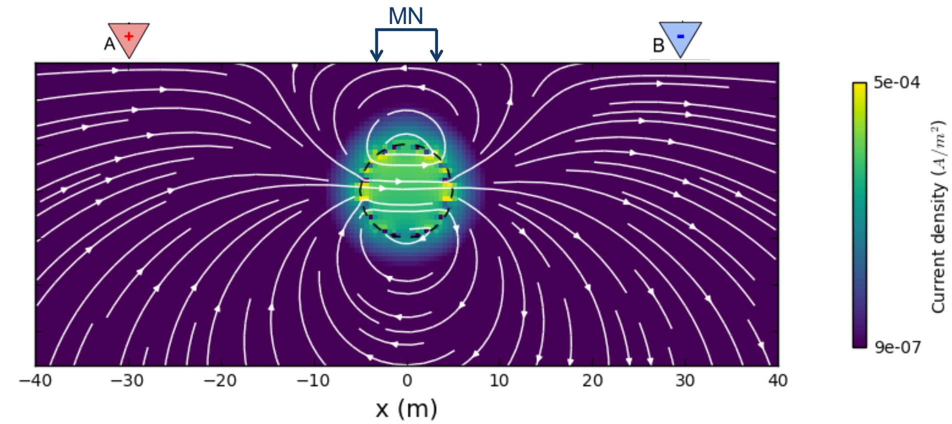
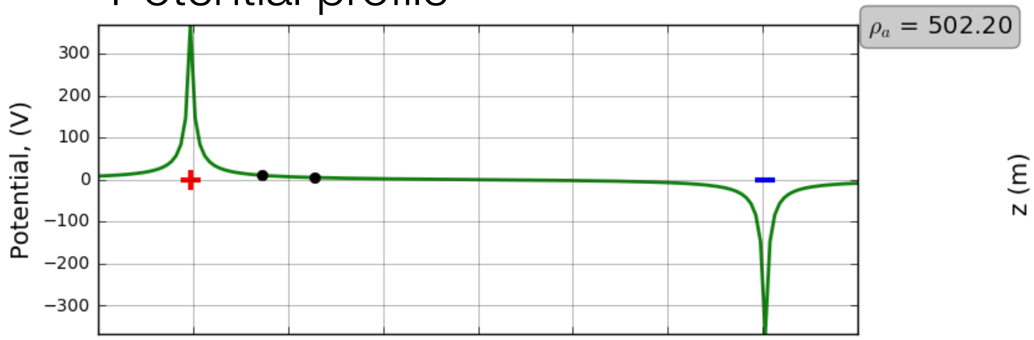


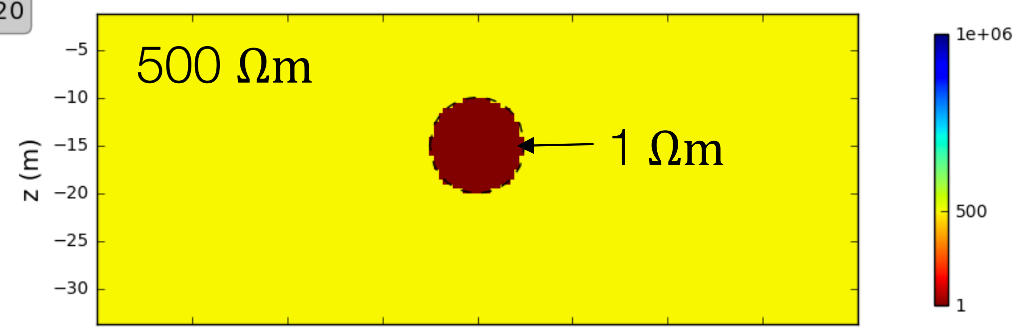
Secondary currents: J_s

Measurements of DC data: gradient array

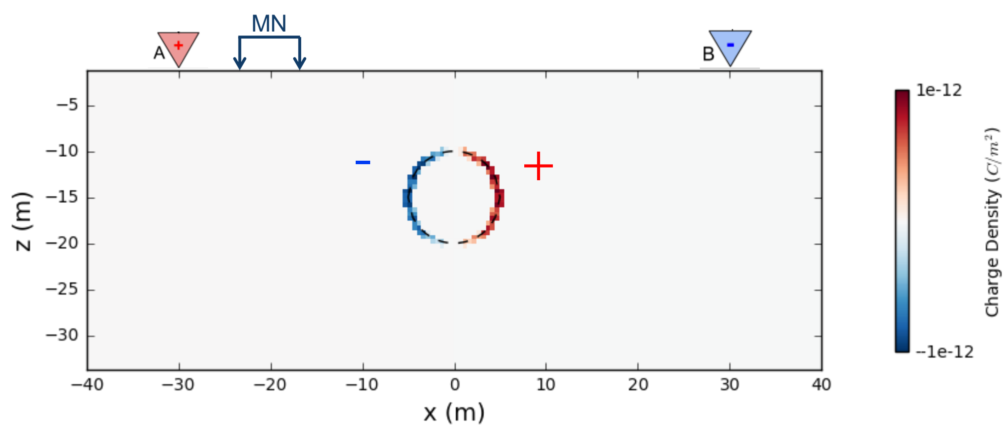
Potential profile



Resistivity model



$\rho_a = 502$



Secondary charges: Q_s

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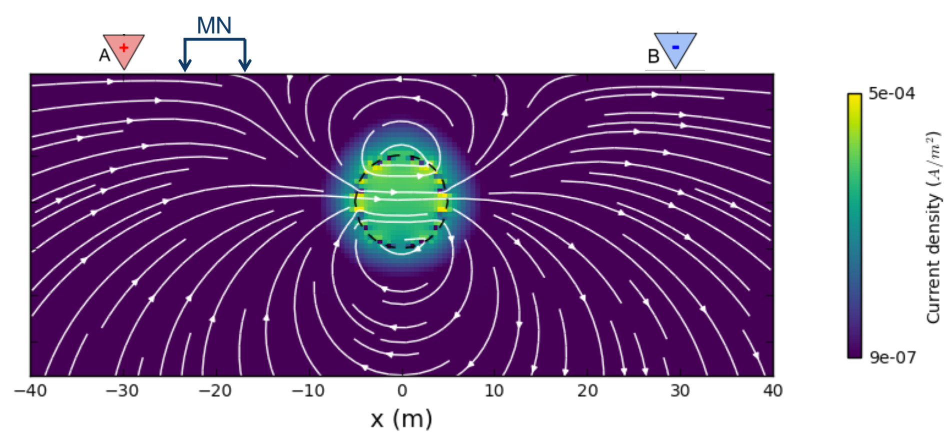
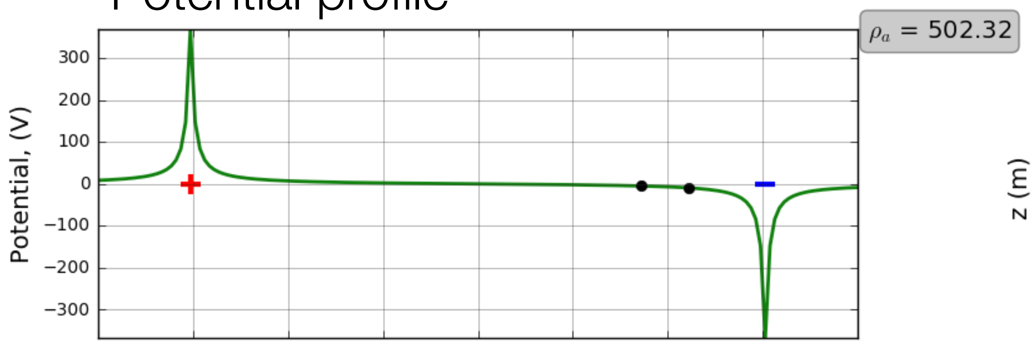


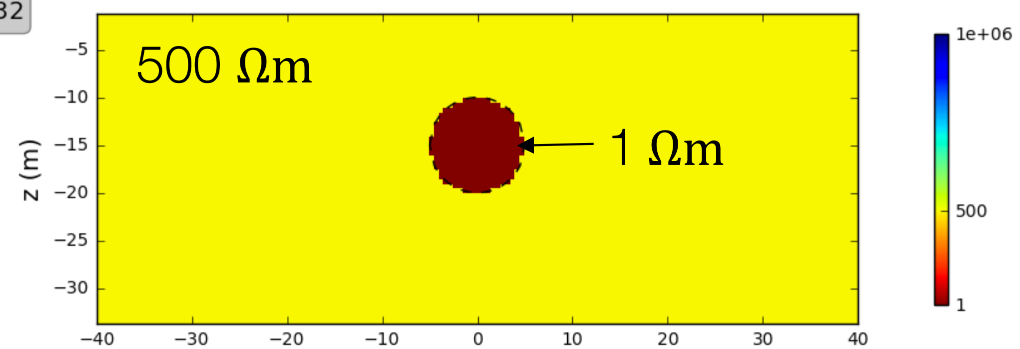
Secondary currents: J_s

Measurements of DC data: gradient array

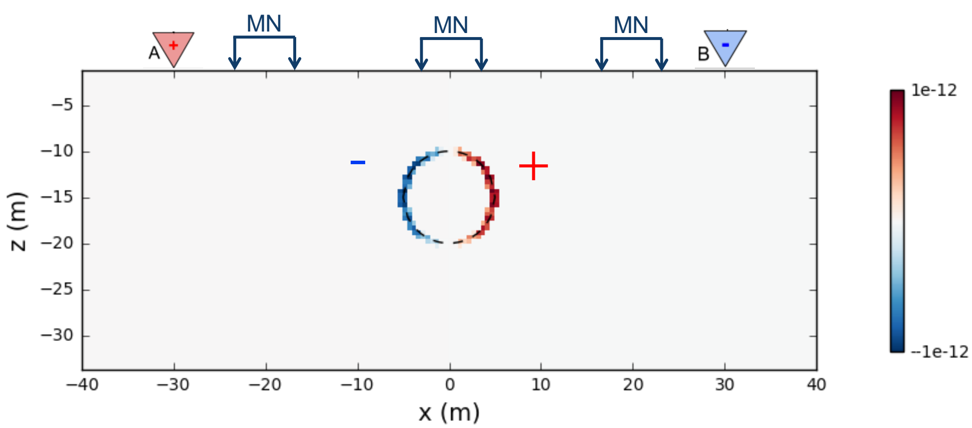
Potential profile



Resistivity model



$\rho_a = 502$ $\rho_a = 430$ $\rho_a = 502$



Secondary charges: Q_s

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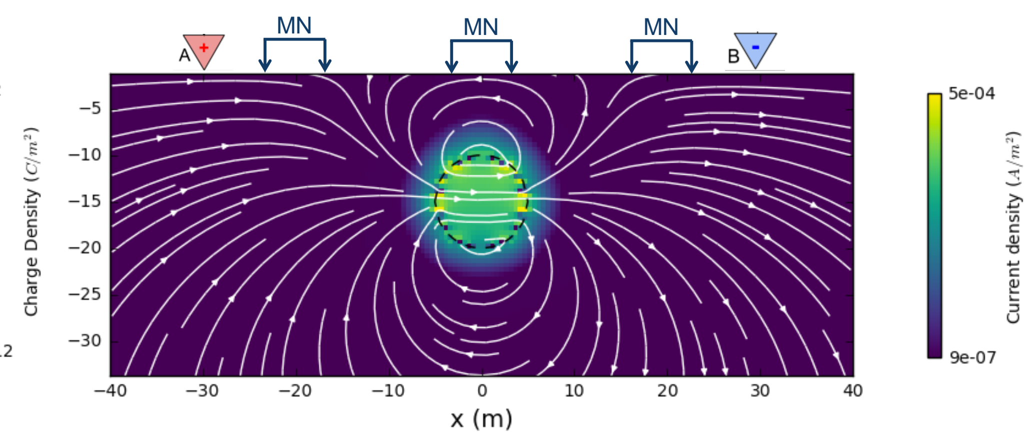


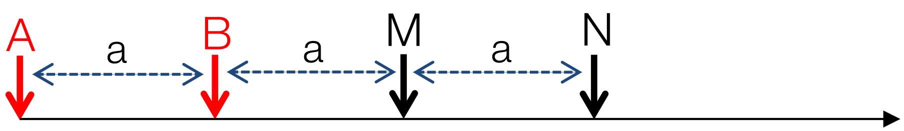
Secondary currents: J_s

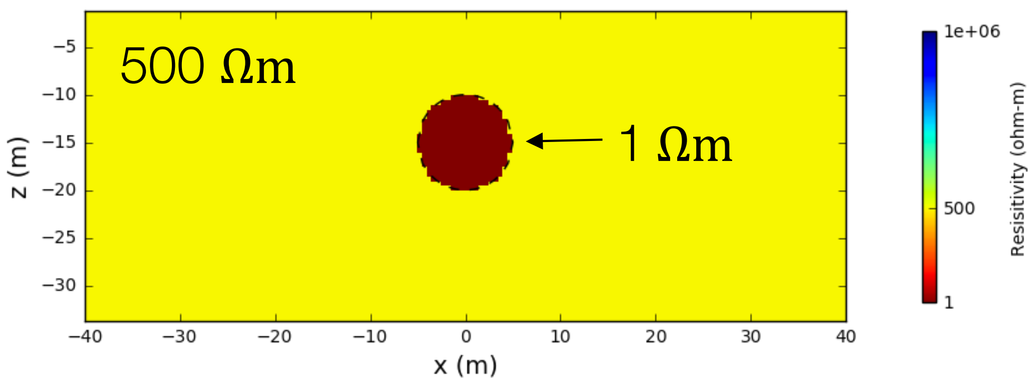
Profiling

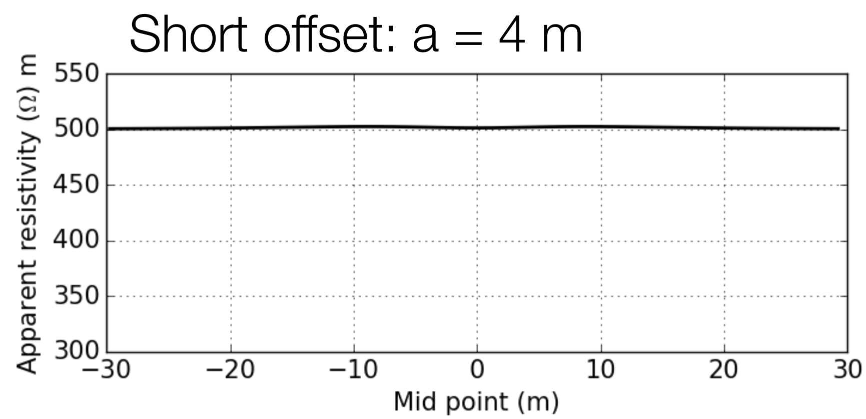
Fixed geometry: Move laterally

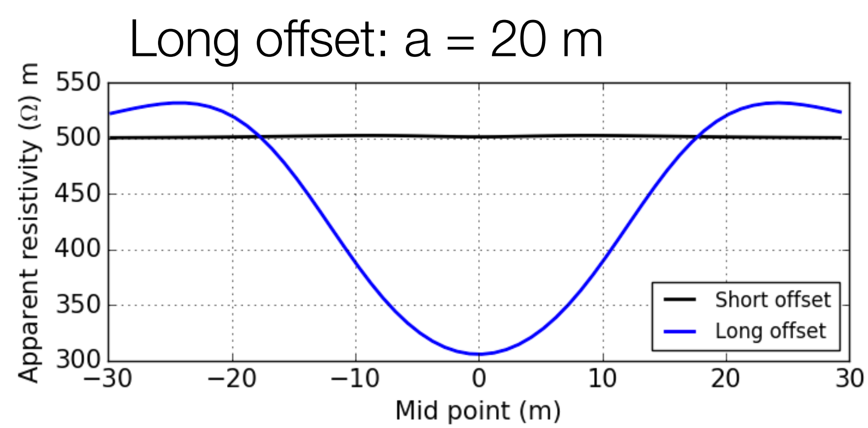
Short offset, a=4m

Long offset, a=20m







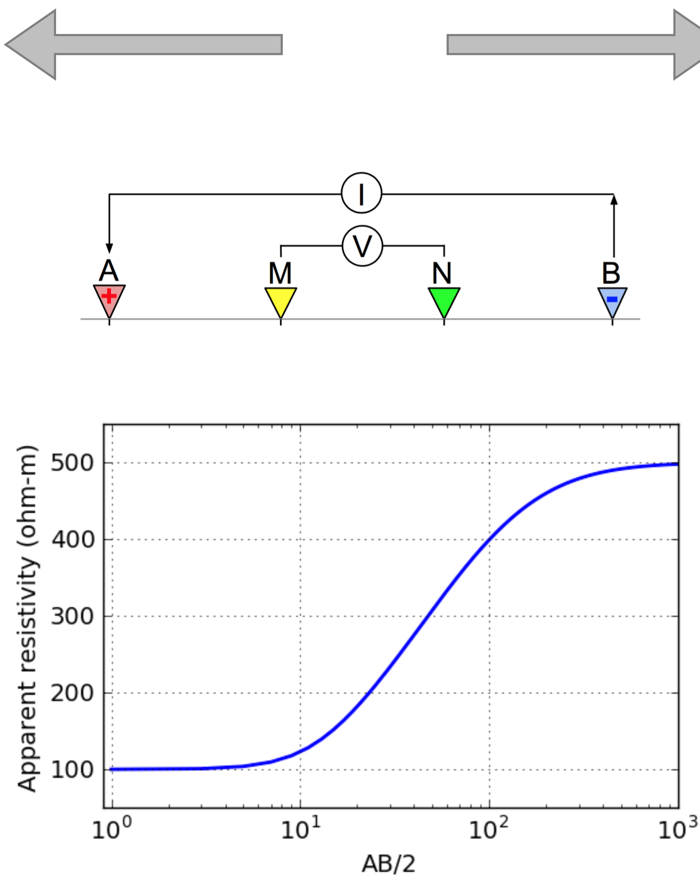


Depth of investigation depends upon offset or array length

Summary: Soundings and Profiles

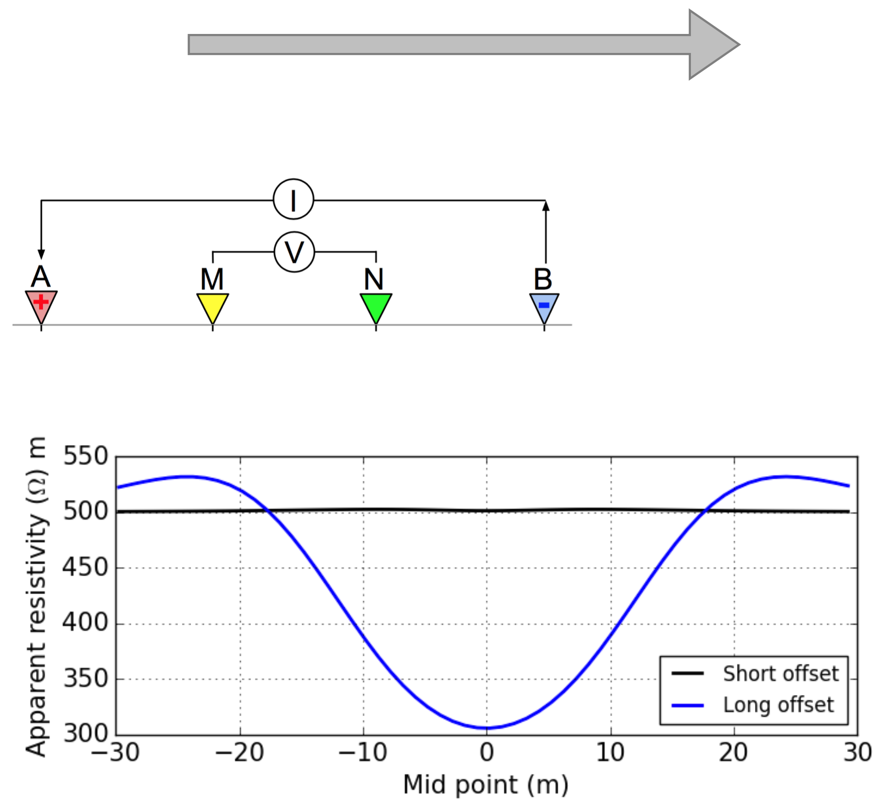
Sounding

Expand



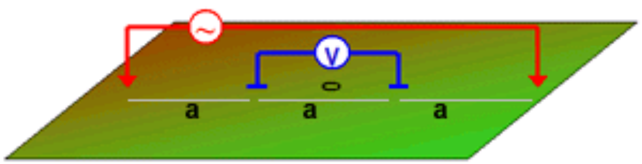
Profiling

Translate

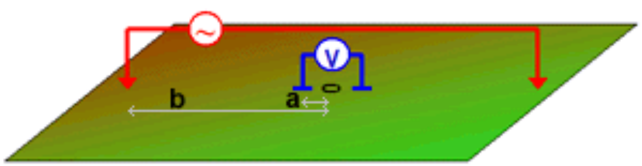


Basic Survey Setups

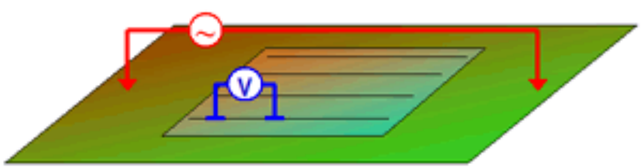
Wenner



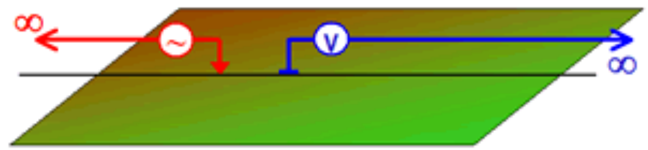
Schulmberger



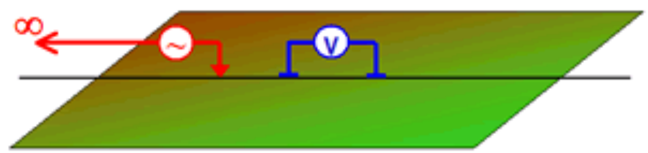
Gradient



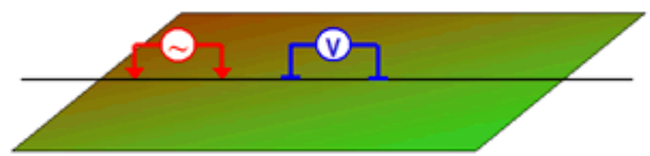
Pole-Pole



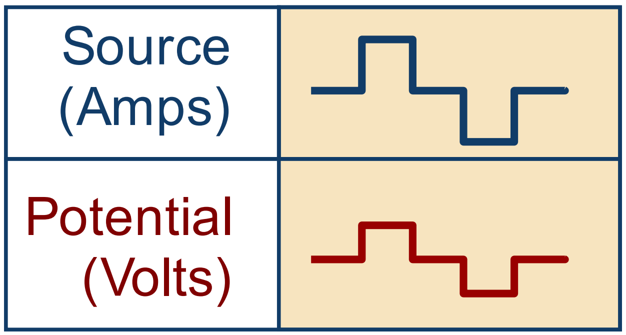
Pole-Dipole

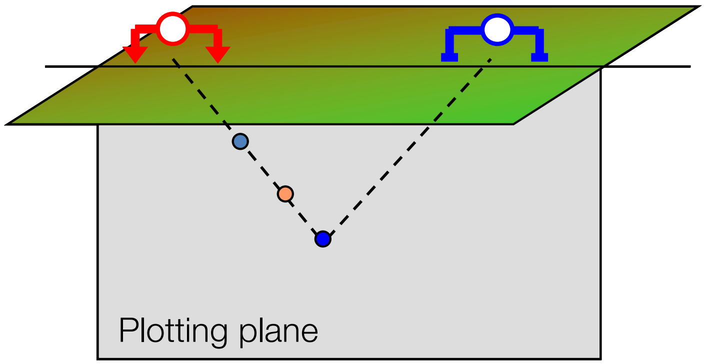


Dipole-Dipole

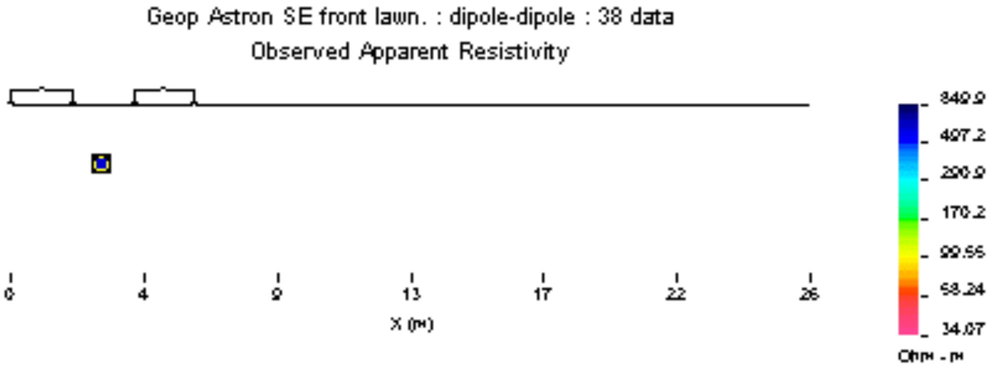


DC resistivity data



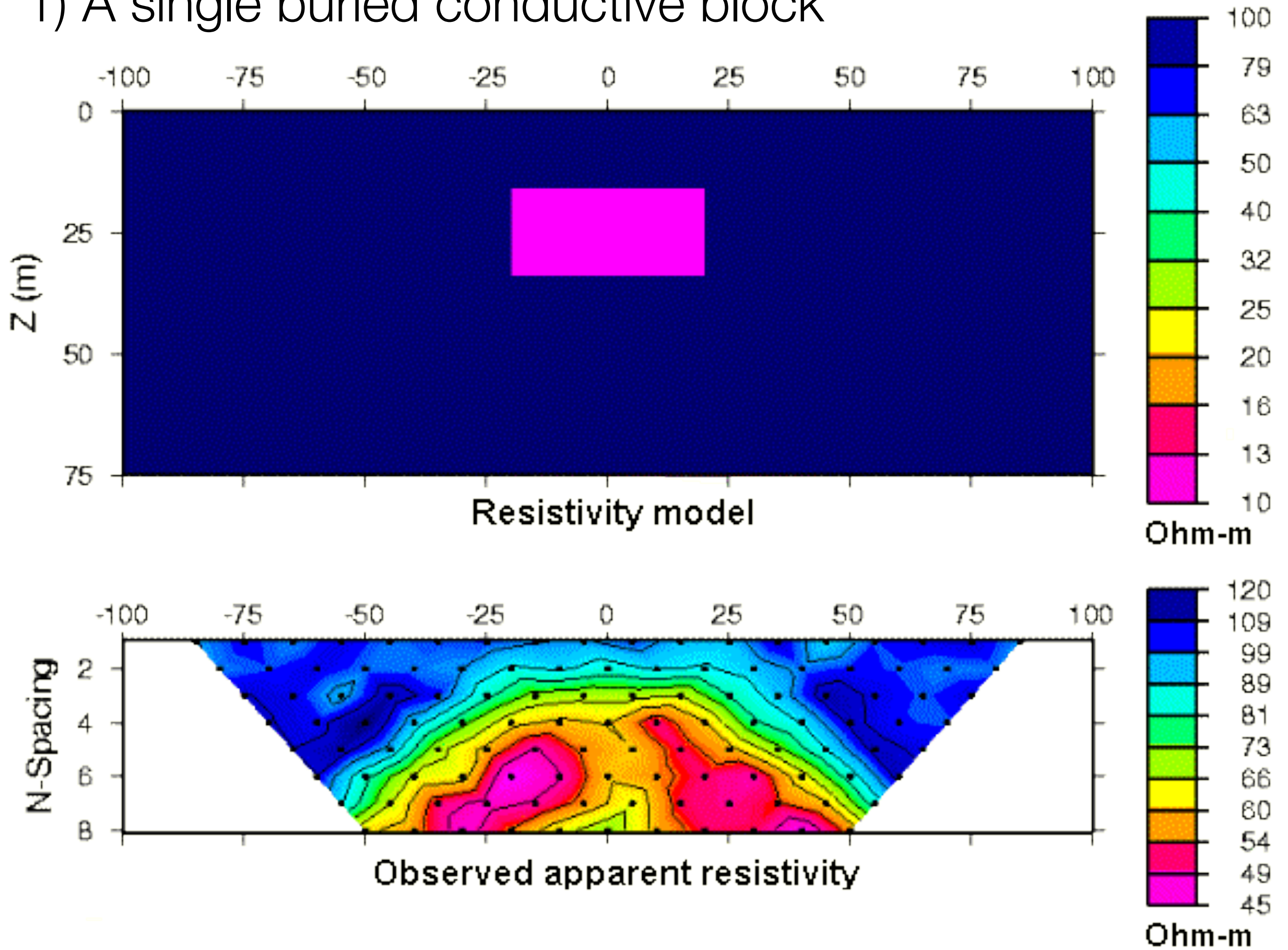


Each data point is an apparent resistivity: $$\rho_a = \frac{2\pi \Delta V}{IG}$$

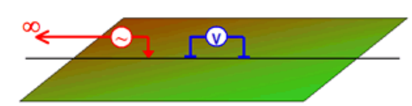


Example pseudosections

1) A single buried conductive block



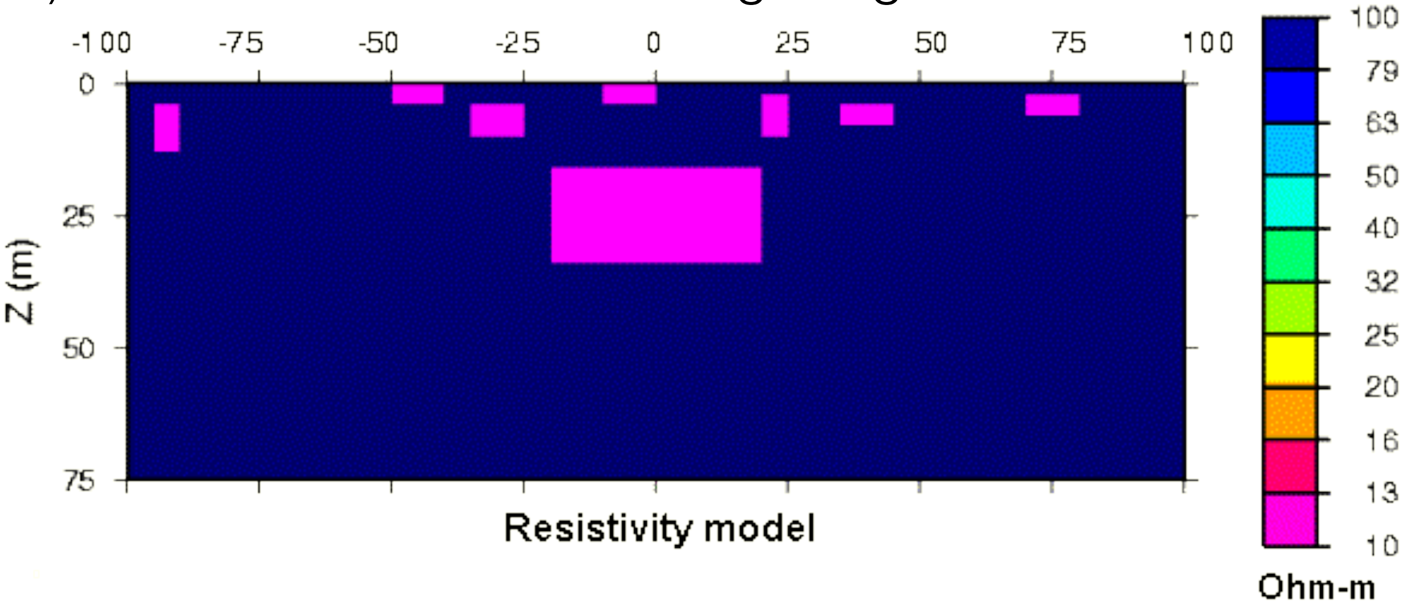
Pole-Dipole

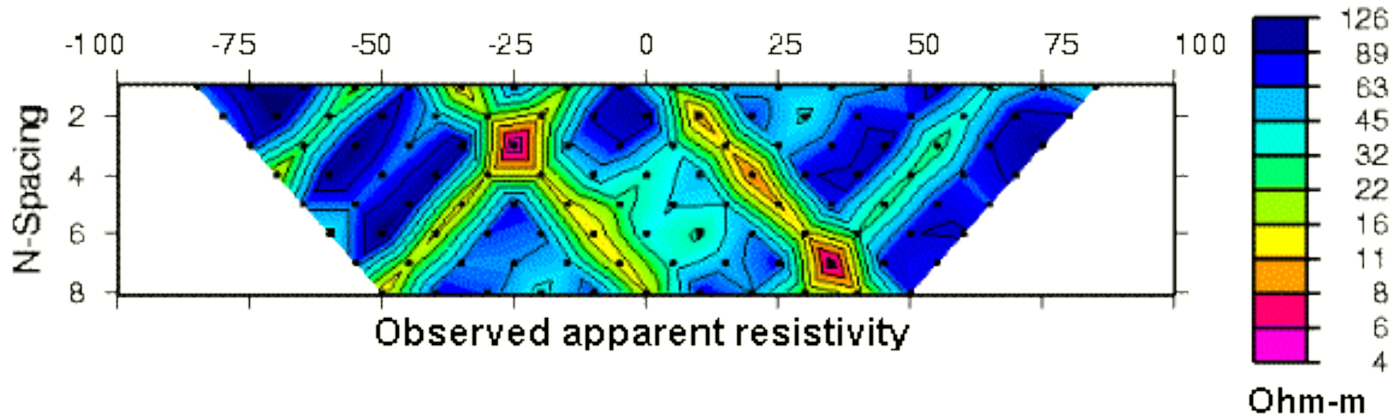


- Pole-dipole; n=1,8; a=10m; N=316

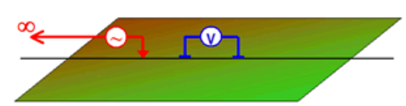
Example pseudosections

2) The conductive block with geologic noise.





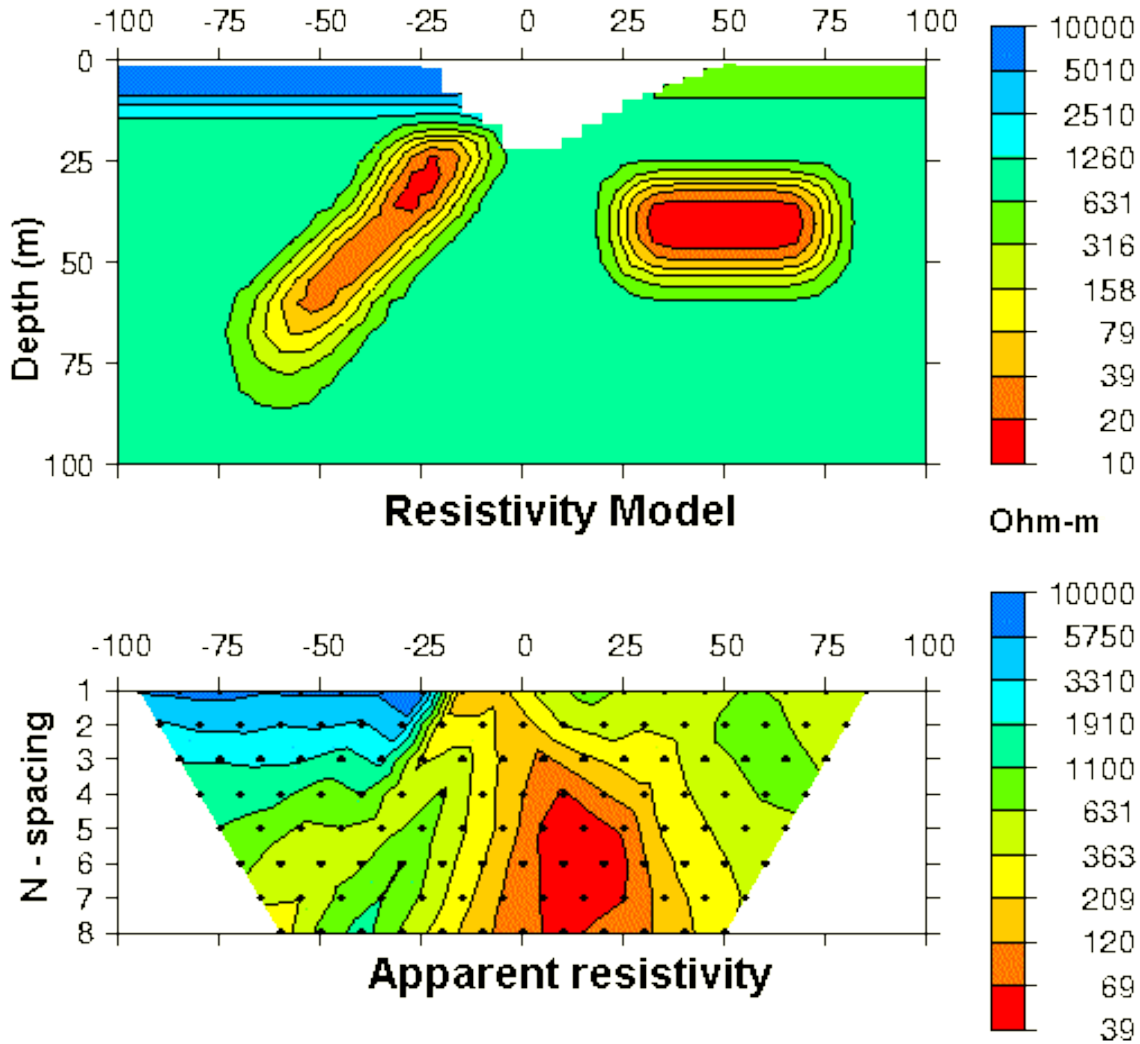
Pole-Dipole



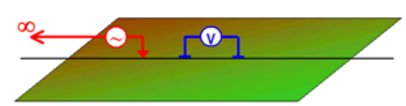
- Pole-dipole; n=1,8; a=10m; N=316

Example pseudosections

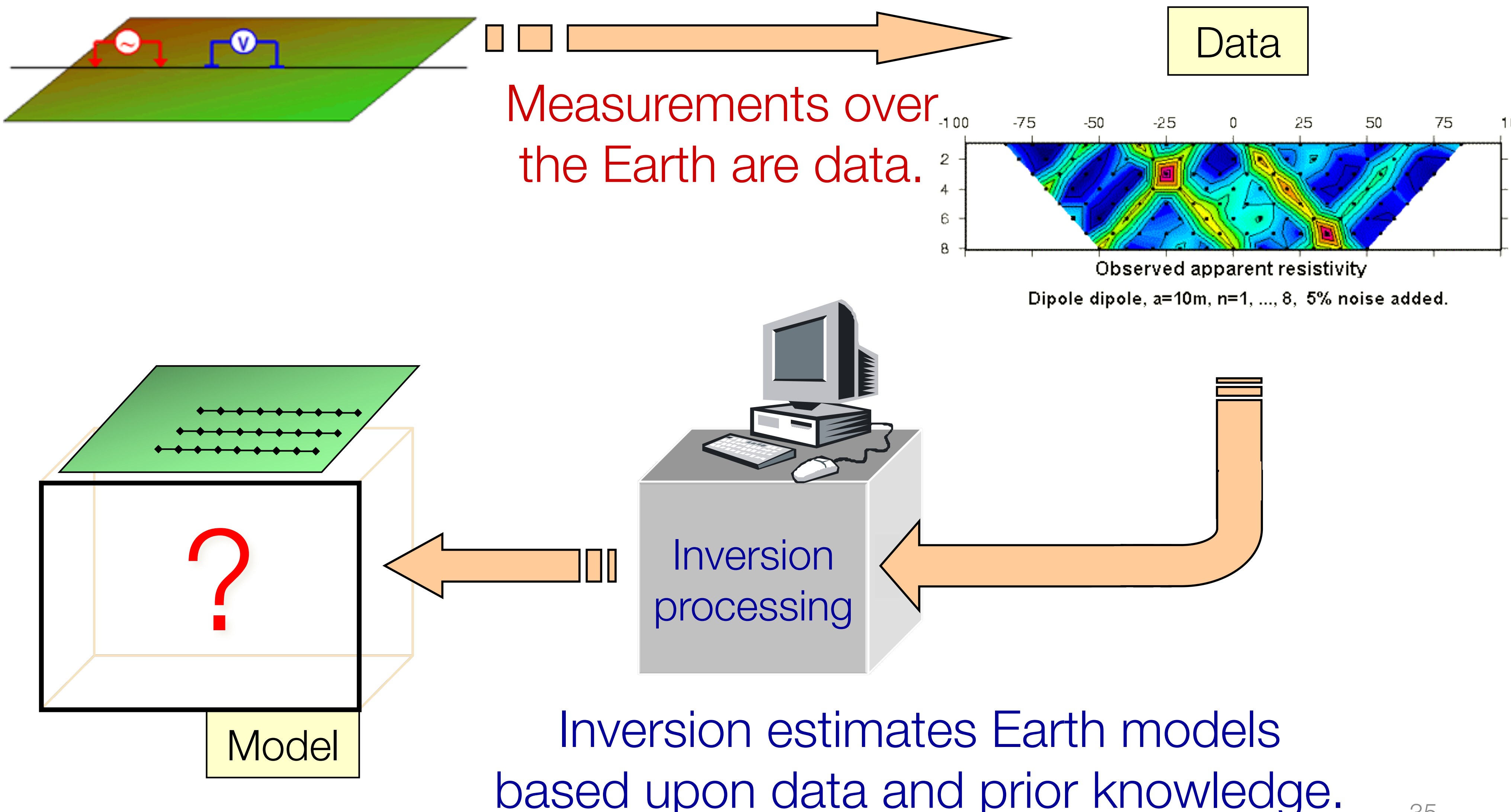
3) The "UBC-GIF model"



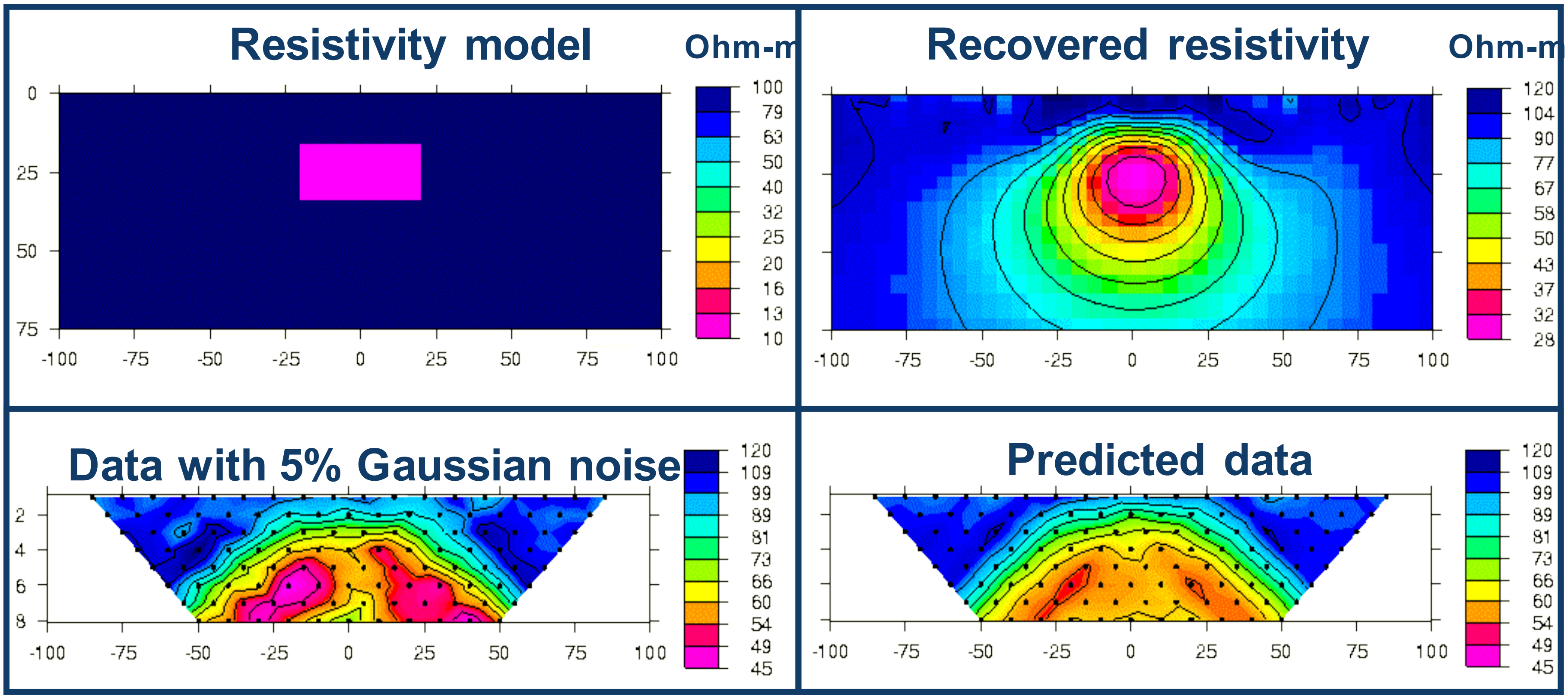
Pole-Dipole



Inversion

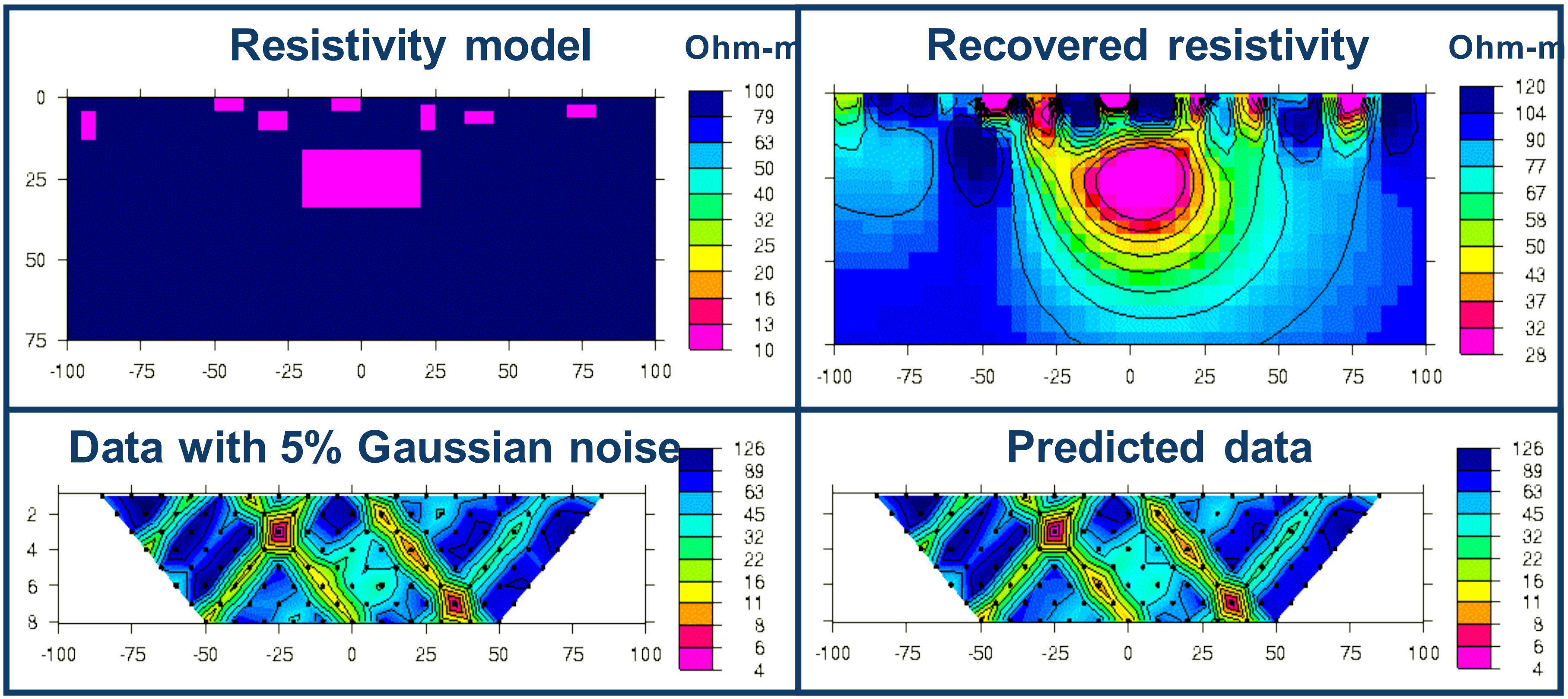

Data
Measurements over the Earth are data.
-100 -75 -50 -25 0 25 50 75 100
2
4
6
8
Observed apparent resistivity
Dipole dipole, a=10m, n=1, ..., 8, 5% noise added.
Inversion processing
?
Model
Inversion estimates Earth models based upon data and prior knowledge.

Example 1: buried prism



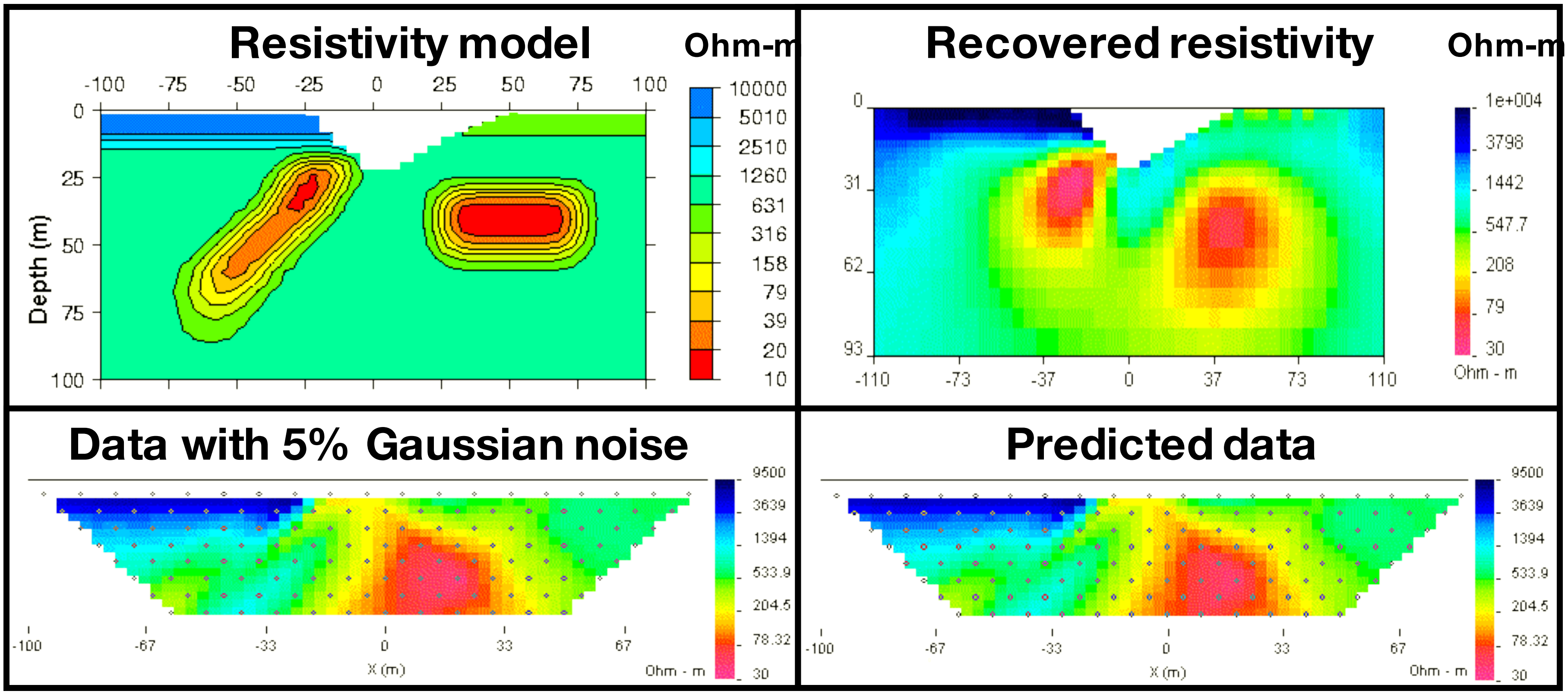
- Pole-dipole; n=1,8; a=10m; N=316; $(\alpha_s, \alpha_x, \alpha_z)$=(.001, 1.0, 1.0)

Example 2: prism with geologic noise



- Pole-dipole; n=1,8; a=10m; N=316; $(\alpha_s, \alpha_x, \alpha_z)$=(.001, 1.0, 1.0)

Example 3: UBC-GIF model



- Pole-dipole; n=1,8; a=10m

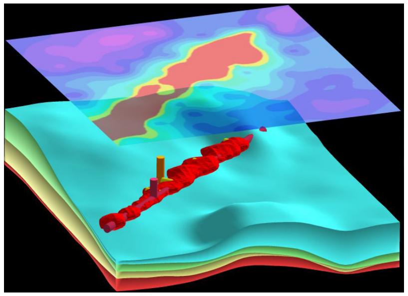
The world is 3D

- Target
  - Size, shape, depth
- Background
  - Variable resistivity
- Questions
  - Where to put currents? 2D acquisition? 3D?
  - Where to make measurements?
  - Which measurements?
  - Effects of topography?
- These are survey design questions
- Crucial element is the "sensitivity"

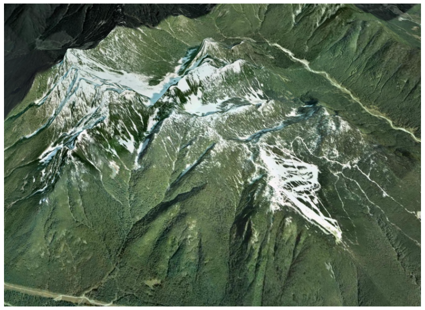
Host

Water underground

Ore body

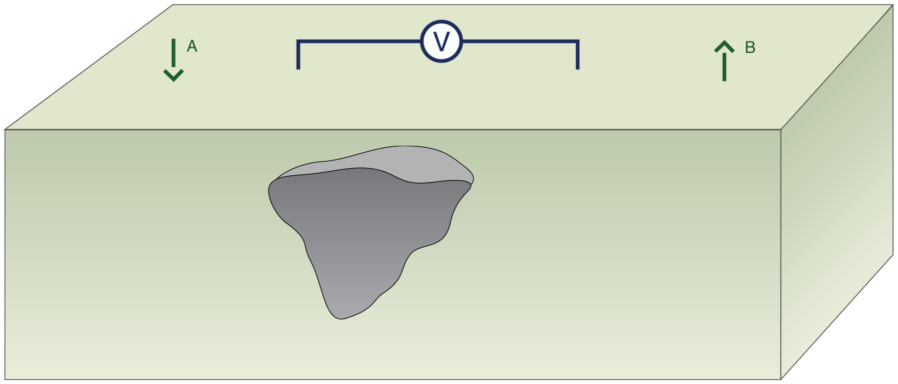


Topography



Sensitivity

Sensitivity Function



Is the measured potential *sensitive* to the target?

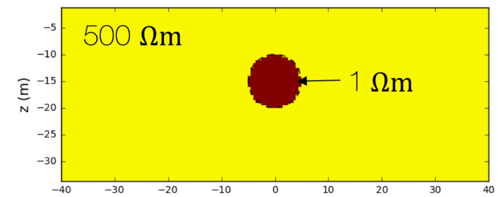
Quantified by the sensitivity

$$G = \frac{\Delta d}{\Delta p} = \frac{\text{change in data}}{\text{change in model}}$$

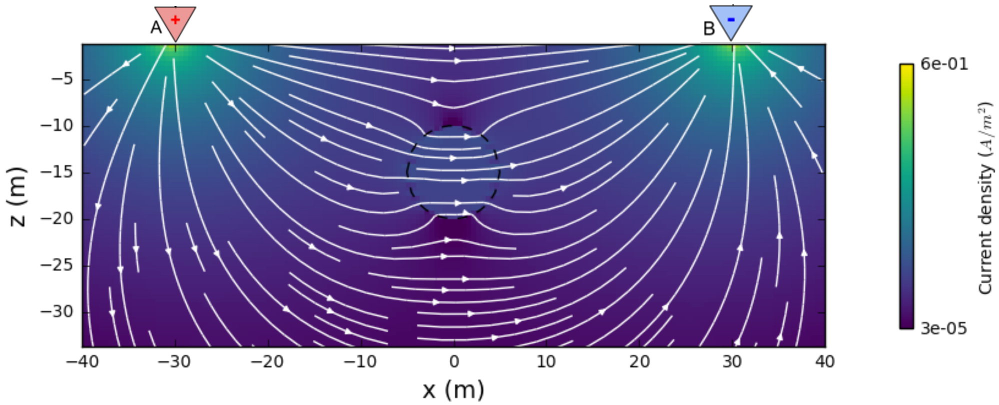
- Collect the data that are "sensitive" to the target
  - Need to "excite" the target
  - Need to have sensor "close" to to the target

Exciting the target

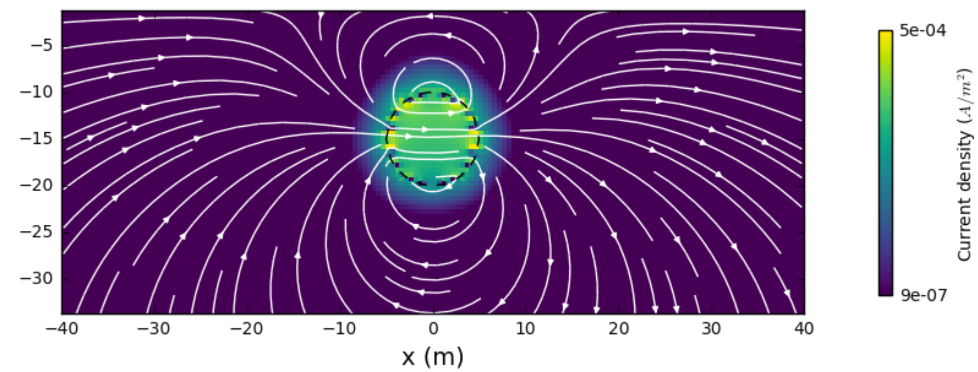
Resistivity model



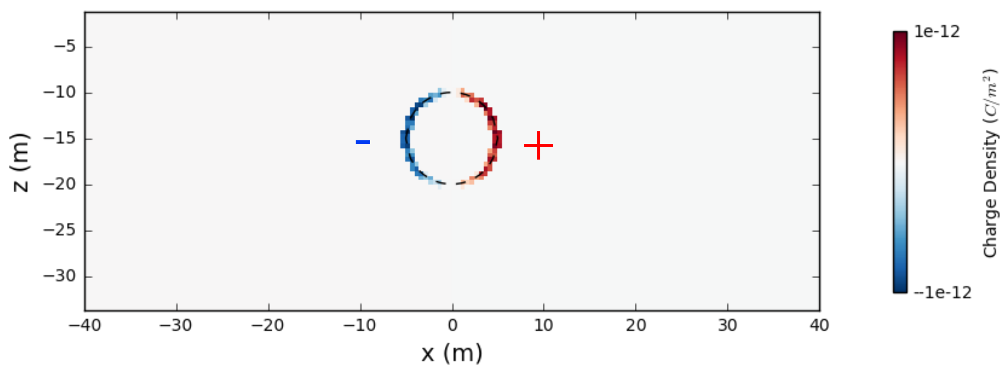
Total currents: **J**



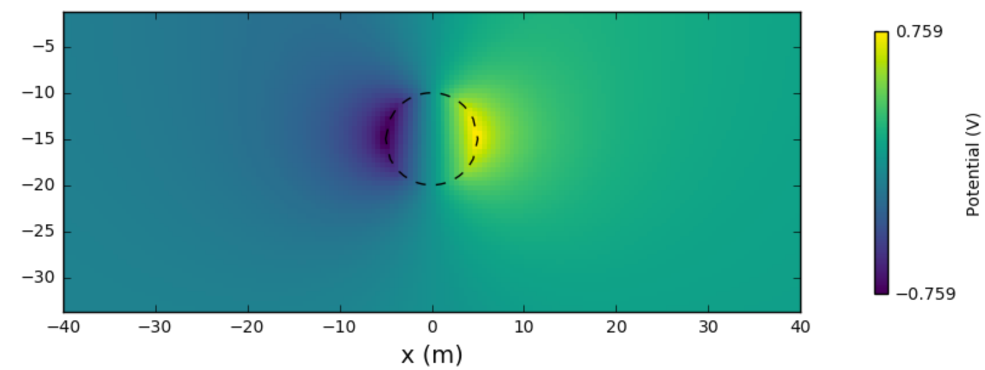
Secondary currents: $\mathbf{J}_s$



Secondary charges: Q_s

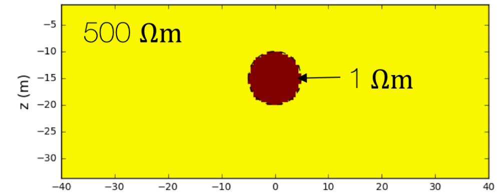


Secondary potential: ϕ_s

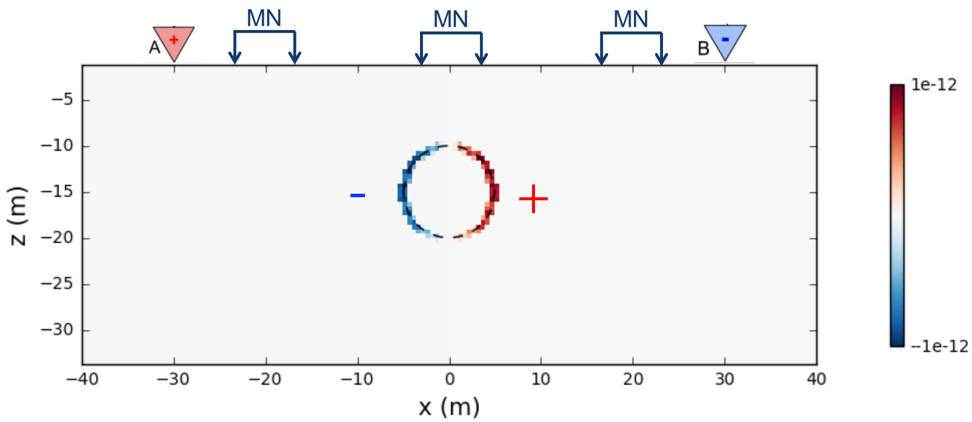


Measurements

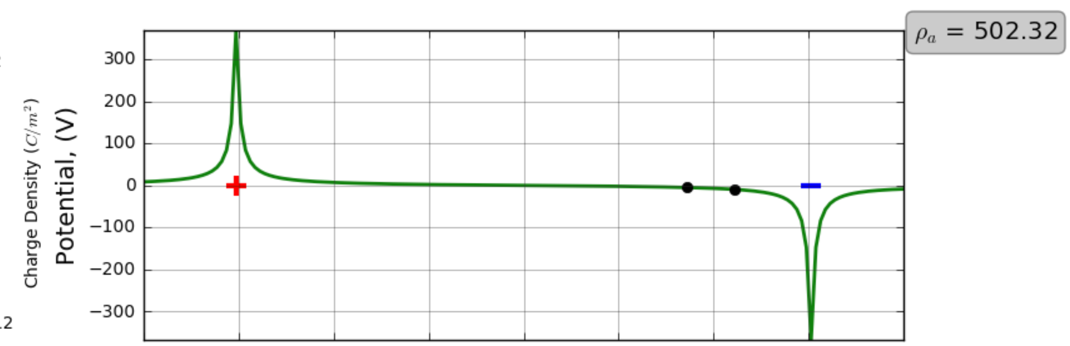
Resistivity model



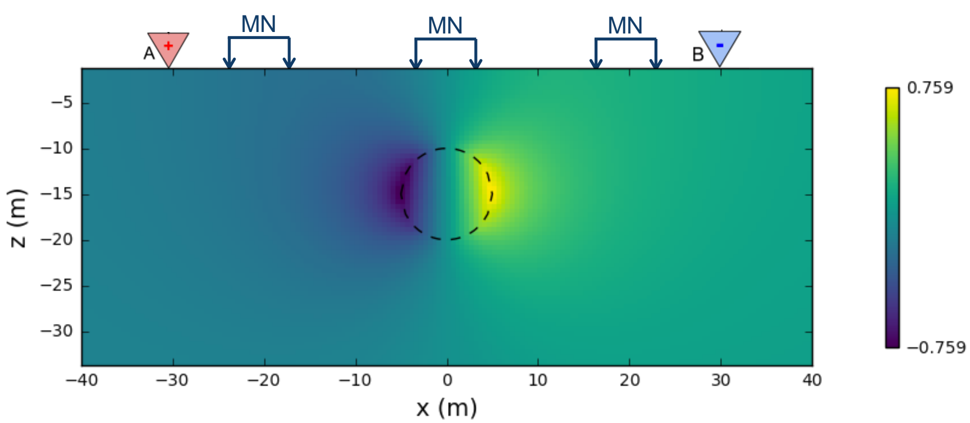
Secondary charges: Q_s



Potential profile

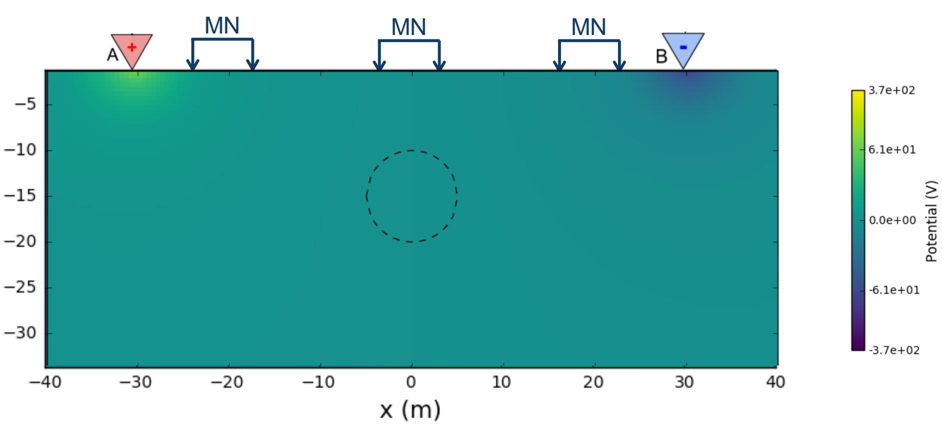


Secondary potential: ϕ_s



Total potential: ϕ

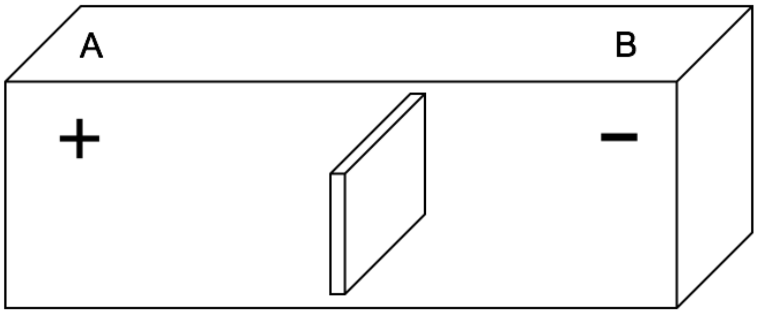
$\rho_a = 502$ $\rho_a = 430$ $\rho_a = 502$

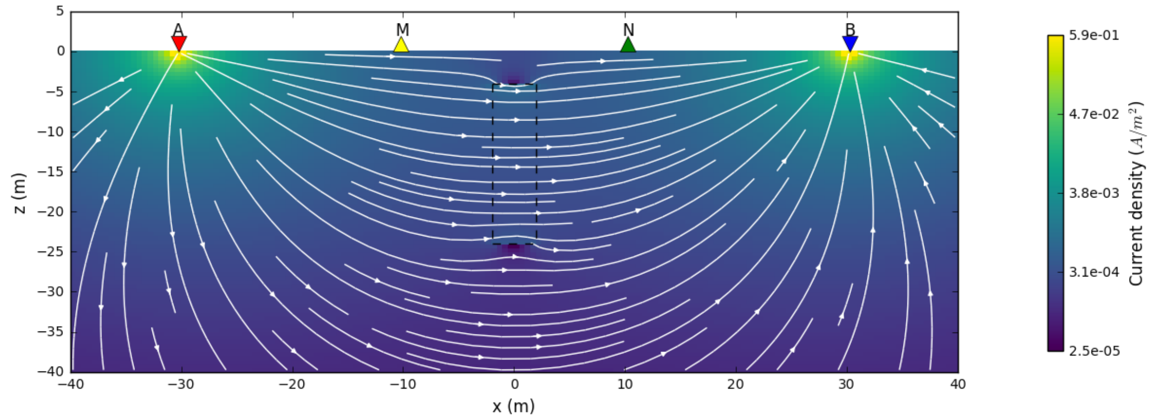


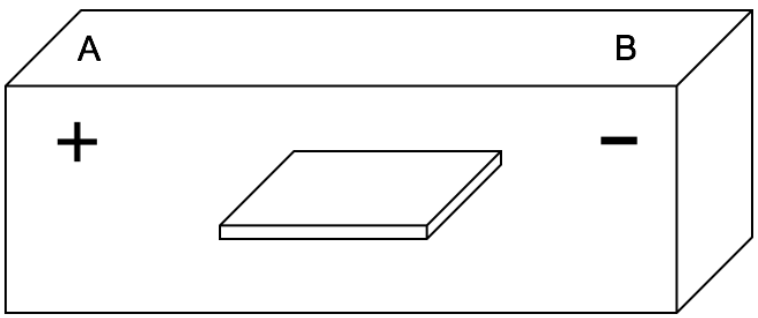
Coupling

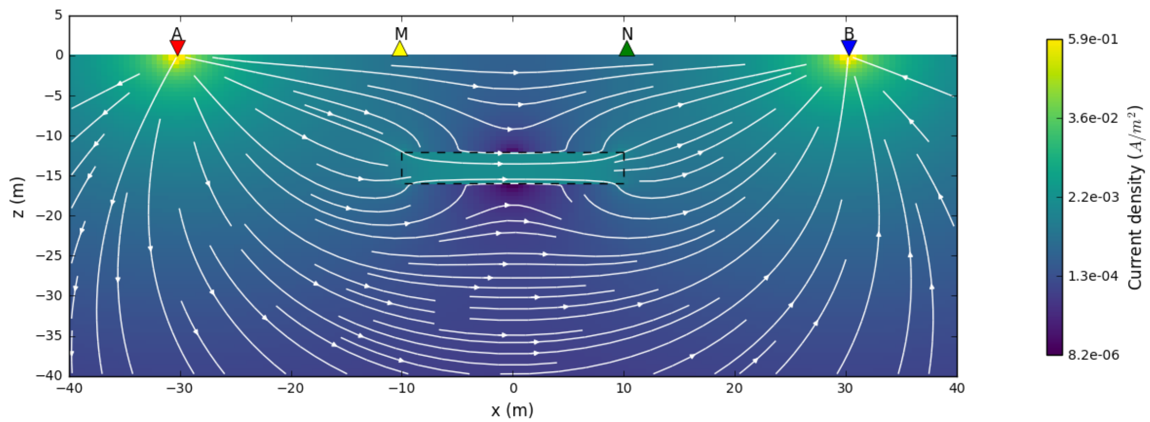
- Thin plate – different orientations
  → different data

Total currents: **J**





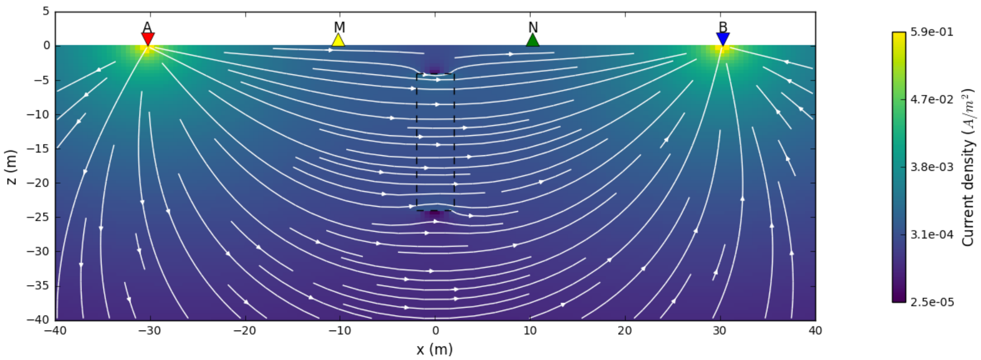




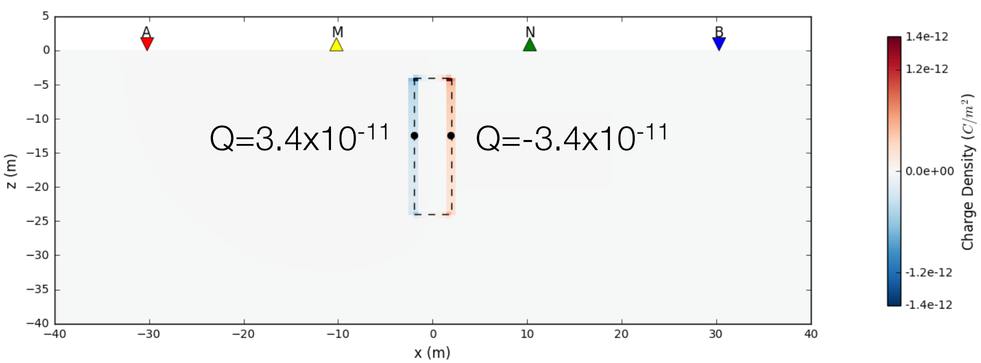
Conductive vs. Resistive Target

Conductive Target

Total currents: **J**

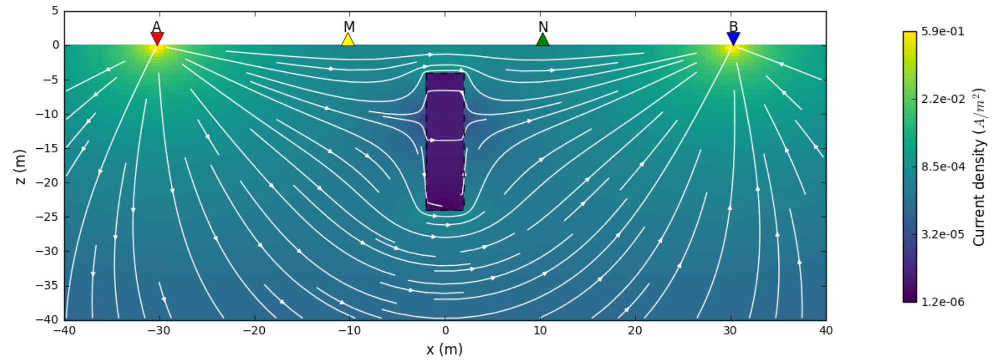


Secondary charges: Q_s

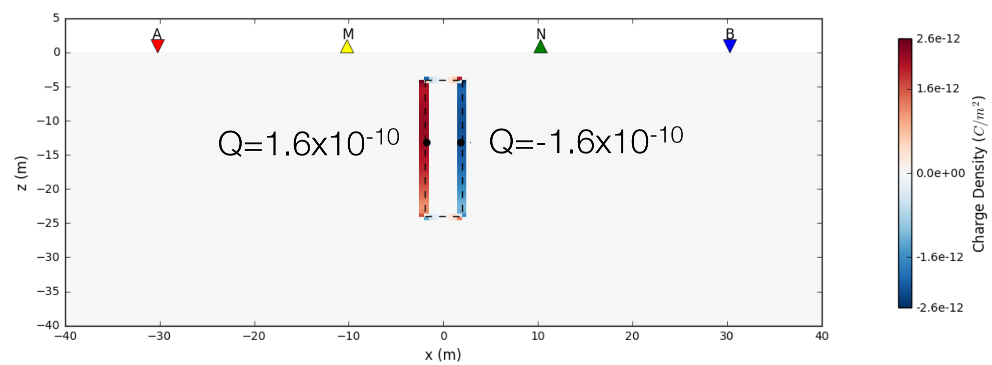


Resistive Target

Total currents: **J**



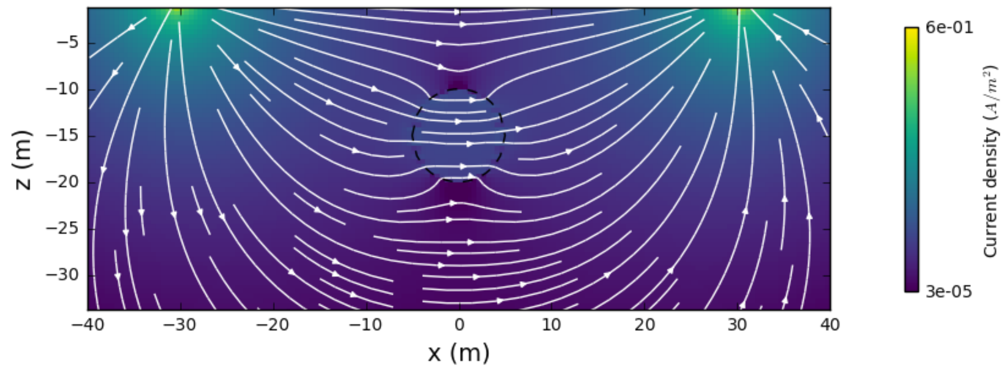
Secondary charges: Q_s



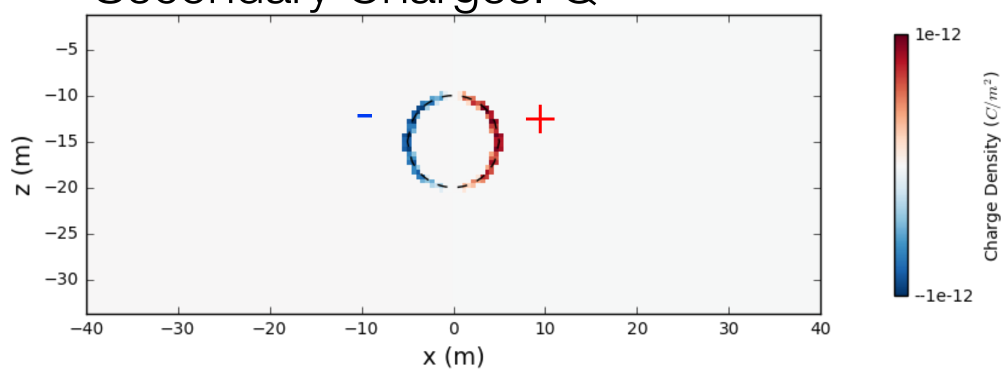
Summary: Sensitivity

- "Excite" the target
  - Drive currents to target
  - Need good coupling with target
- Measuring a datum
  - Proximity to target
  - Electrode orientation and separation
- Background resistivity is important

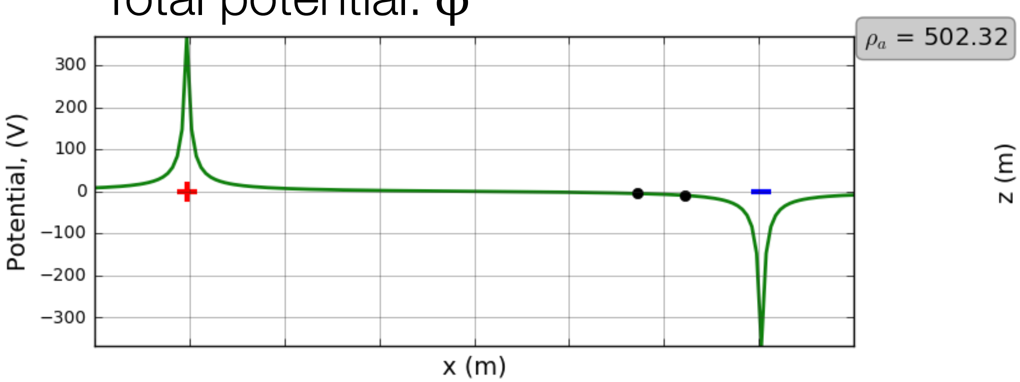
Total currents: **J**



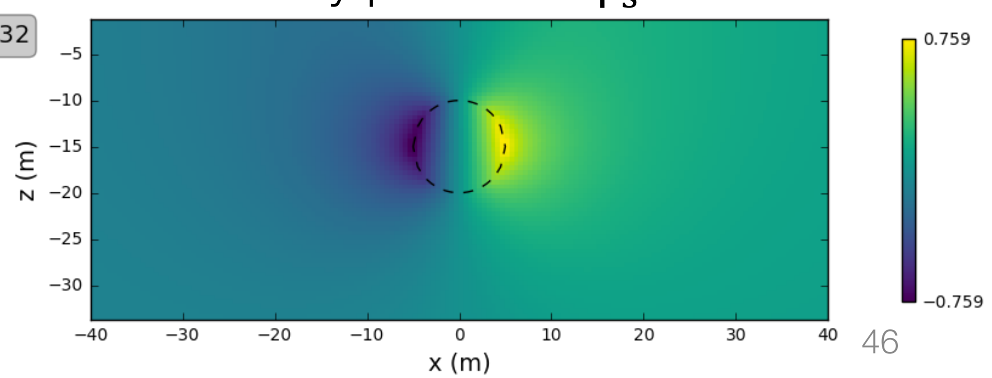
Secondary Charges: Q



Total potential: ϕ



Secondary potential: ϕ_s

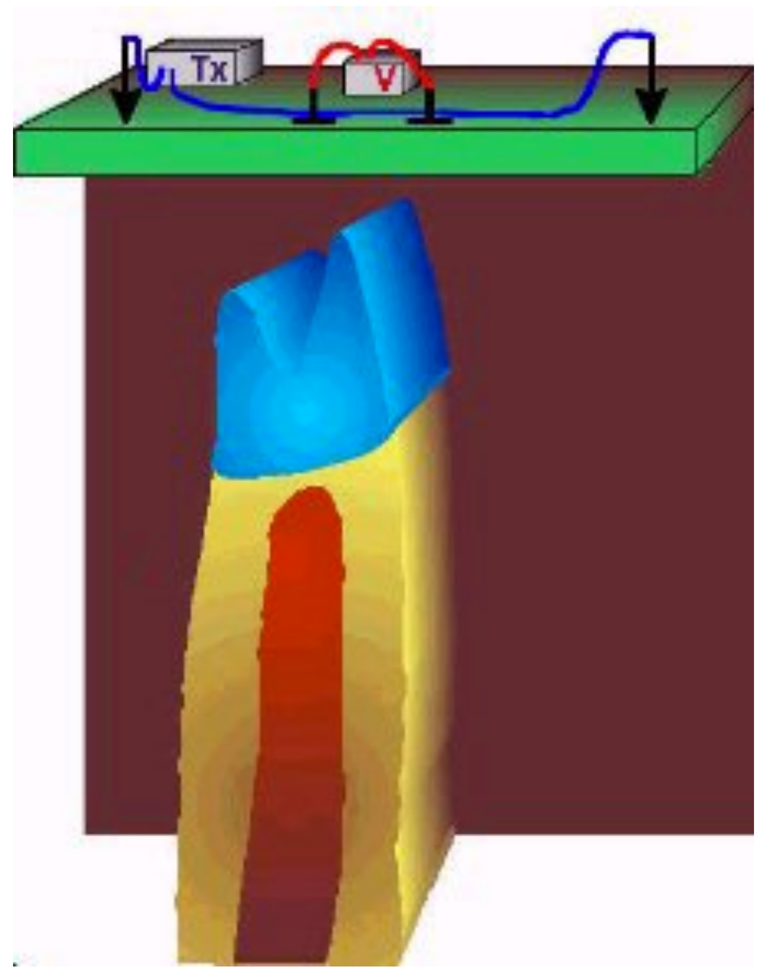


Survey Design: Questions

- What is objective?
  - Layered earth (1D)
    → do a sounding
  - Target body (2D)
    → profile, sounding perpendicular to geology
  - Target body (3d)
    → need 3D coverage

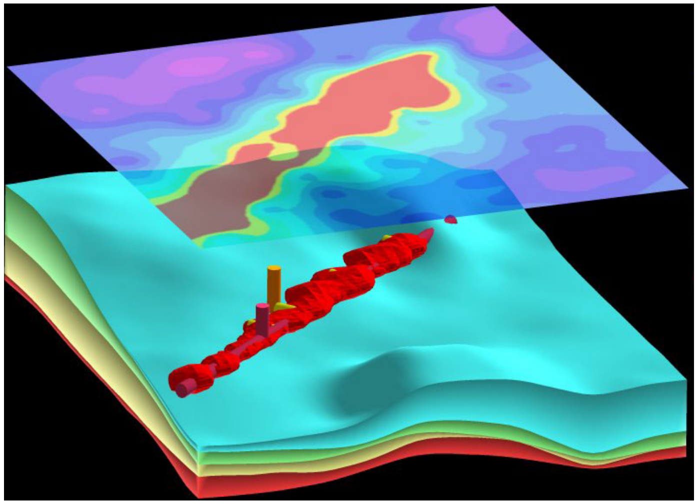
- What is the background resistivity?

- What are the noise sources? fences, power lines, …



Survey Design: in general

- Numerical simulation – can we **see** the target?
  - Secondary signals must be large enough
- Steps:
  - Define a representative geologic model
  - Assign physical properties
  - Select a survey (or surveys)
  - Simulate with and without target
  - Assess secondary signals



Absolute	Relative
ΔV	$\frac{V_s}{V_p}$

- Best practice
  - Assign uncertainties to the simulated data
  - Carry out an inversion with the code you will use to invert the field data

Outline

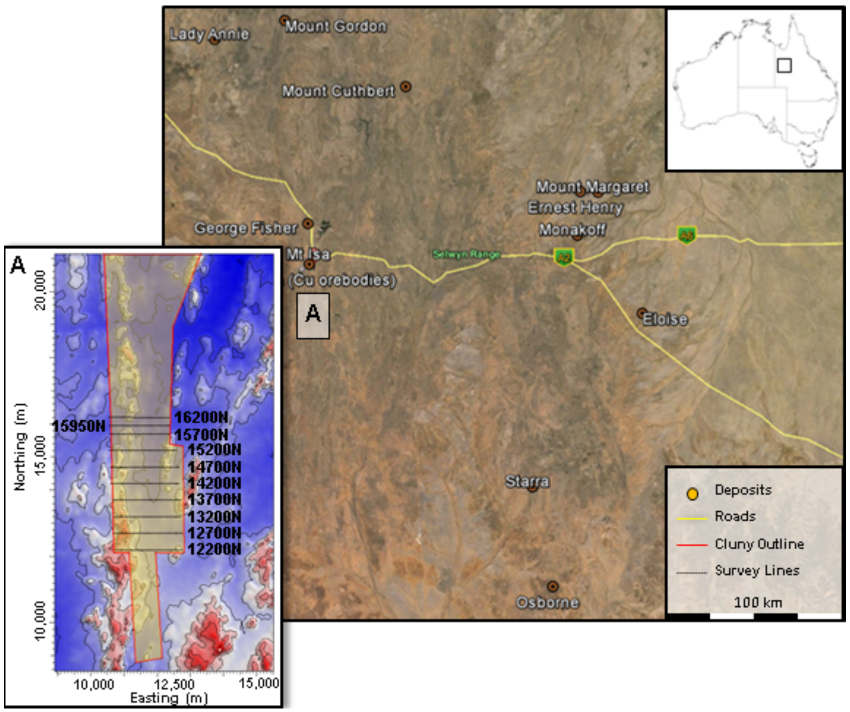
- Basic experiment
- Currents, charges, potentials and apparent resistivities
- Soundings, profiles and arrays
- Data, pseudosections and inversion
- Sensitivity
- Survey Design

- Questions

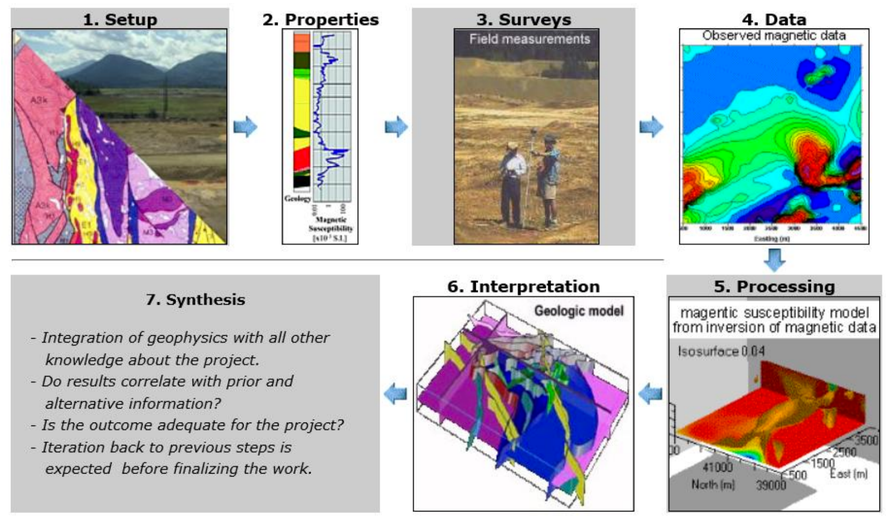
- Case History – Mt Isa
- Effects of background resistivity

Mt. Isa

Mt. Isa (Cluny prospect)

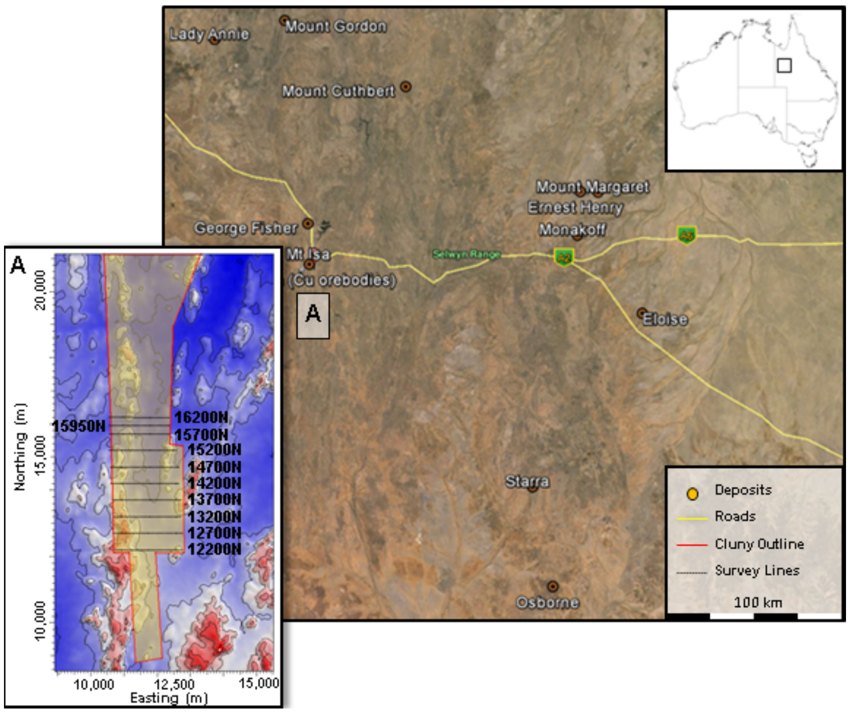


Seven Steps

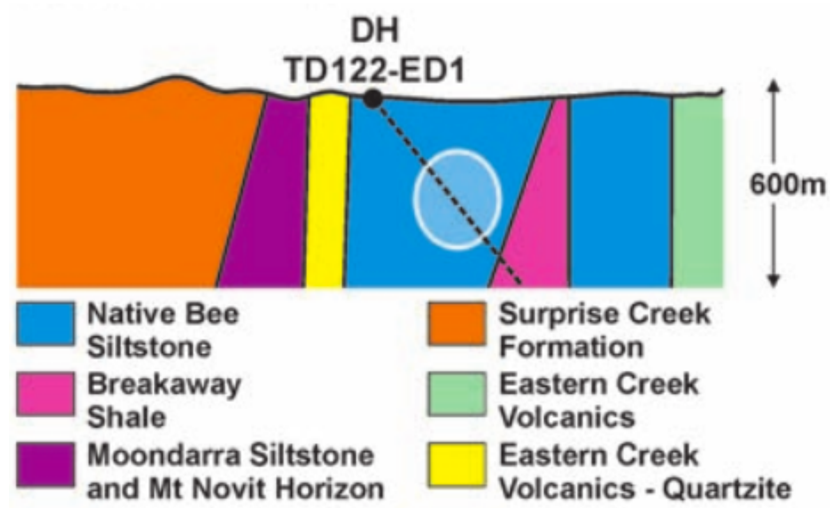


Setup

Mt. Isa (Cluny prospect)



Geologic model

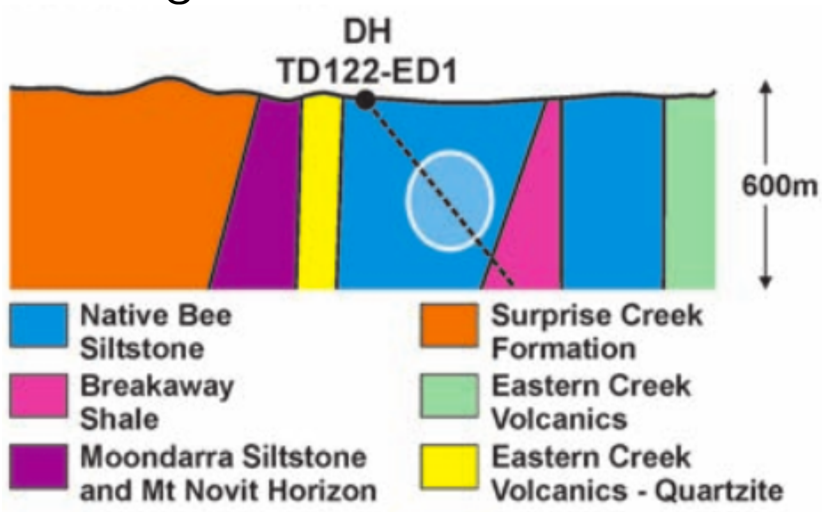


Question

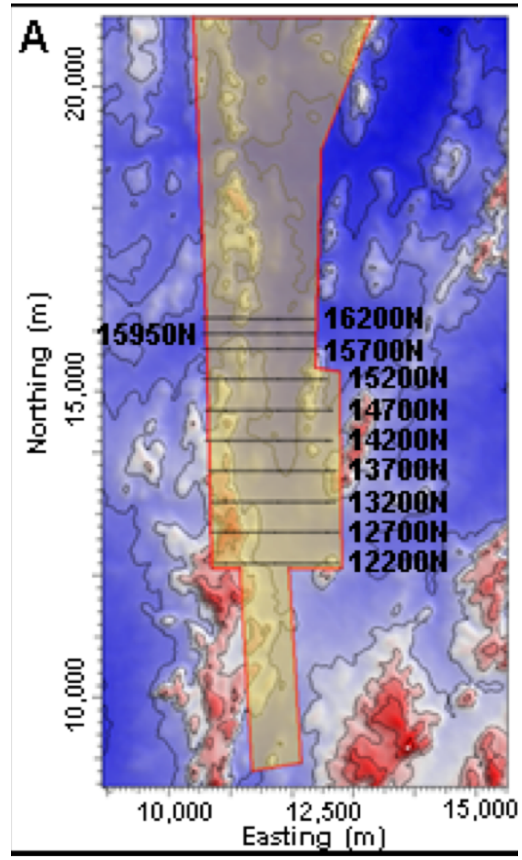
- Can conductive units, which would be potential targets within the siltstones, be identified with DC data?

Properties

Geologic model



Surface topography

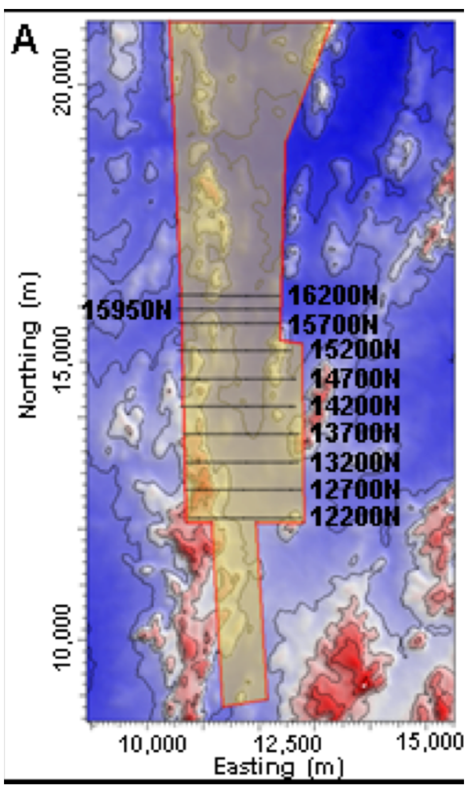


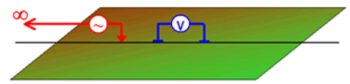
Rock Unit	Conductivity	Resistivity ($\Omega \cdot m$)
Native Bee Siltstone	Moderate	Moderate (~10)
Moondarra Siltstone	Moderate	Moderate (~10)
Breakaway Shale	Very High	Very Low (~0.1)
Mt Novit Horizon	High	Low (~1)
Surprise Creek Formation	Low	High (~1000)
Eastern Creek Volcanics	Low	High (~1000)

Survey and Data

- Eight survey lines
- Two survey configurations.

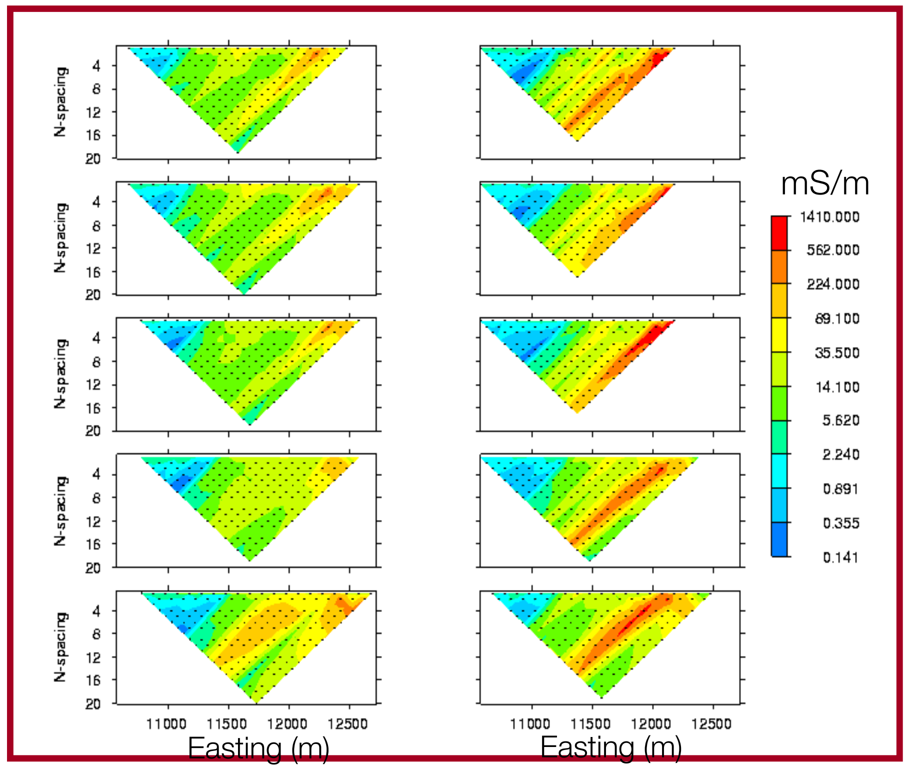
Surface topography





Data set #1:

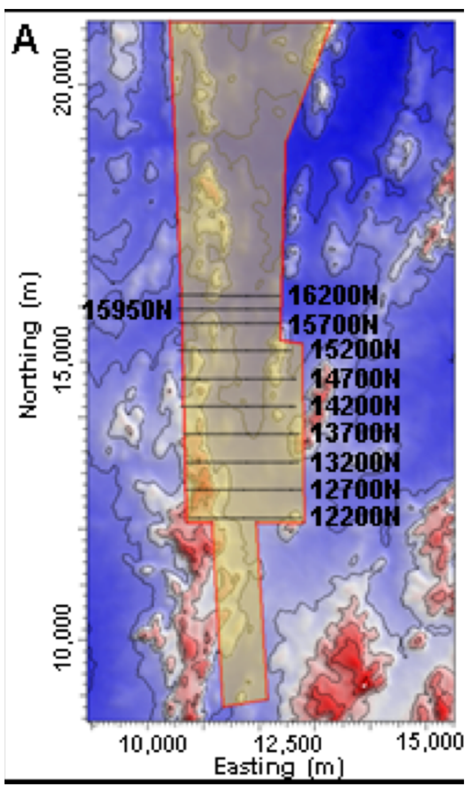
Apparent resistivity,
pole - dipole.

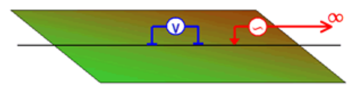


Survey and Data

- Eight survey lines
- Two survey configurations.

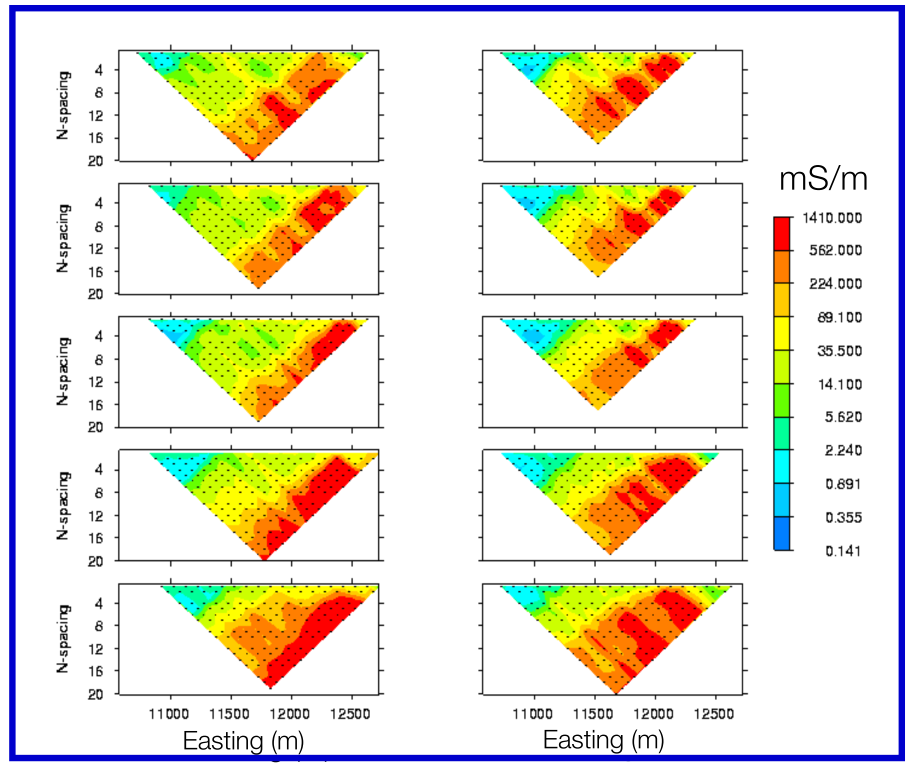
Surface topography





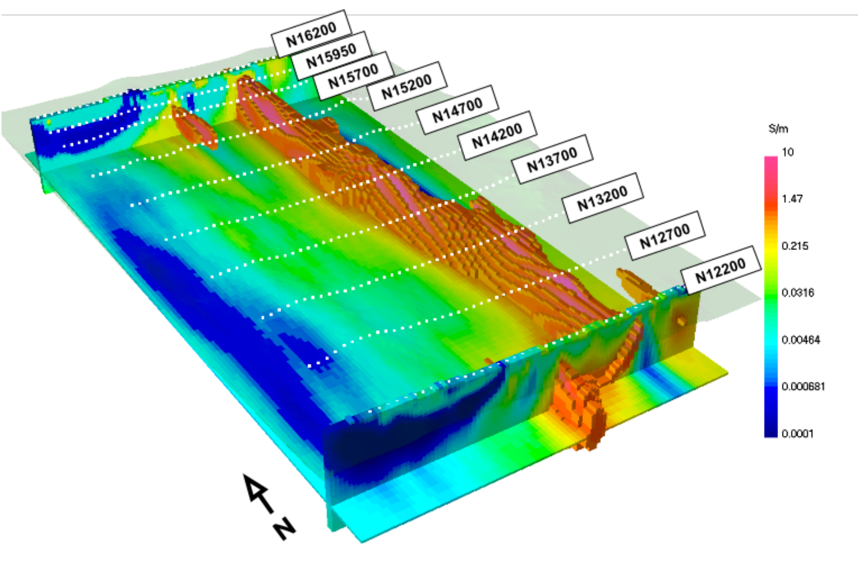
Data set #2:

Apparent resistivity, dipole - pole

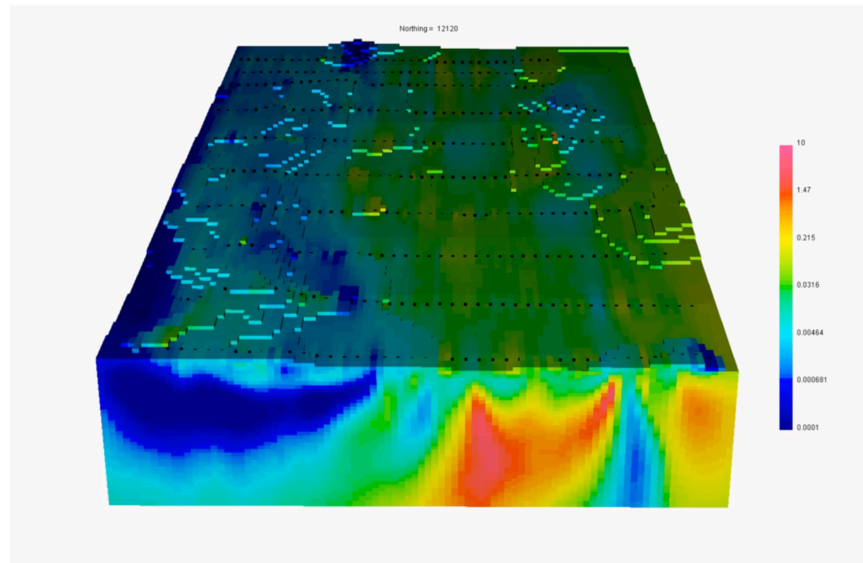


Processing and interpretation

3D resistivity model



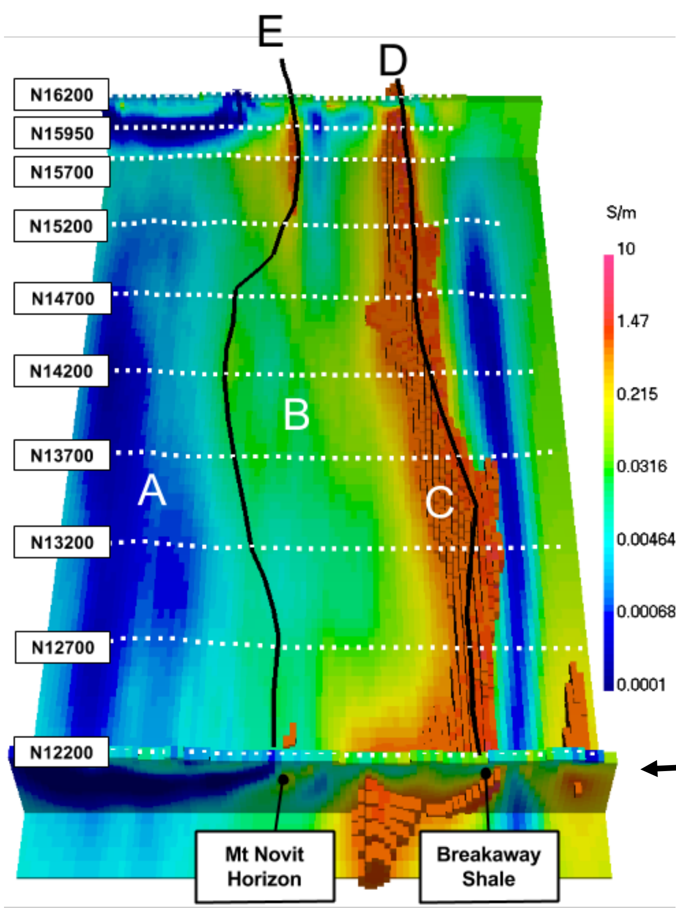
Animation



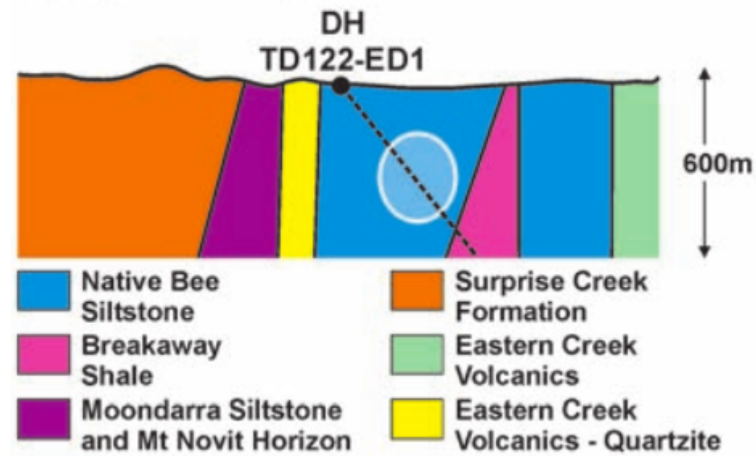
Synthesis

- Identified a major conductor → black shale unit
- Some indication of a moderate conductor

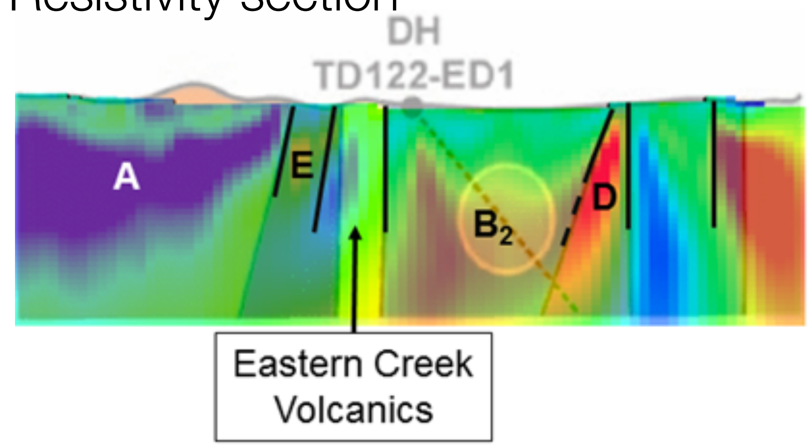
3D resistivity model



Geologic section



Resistivity section



Outline

- Basic experiment
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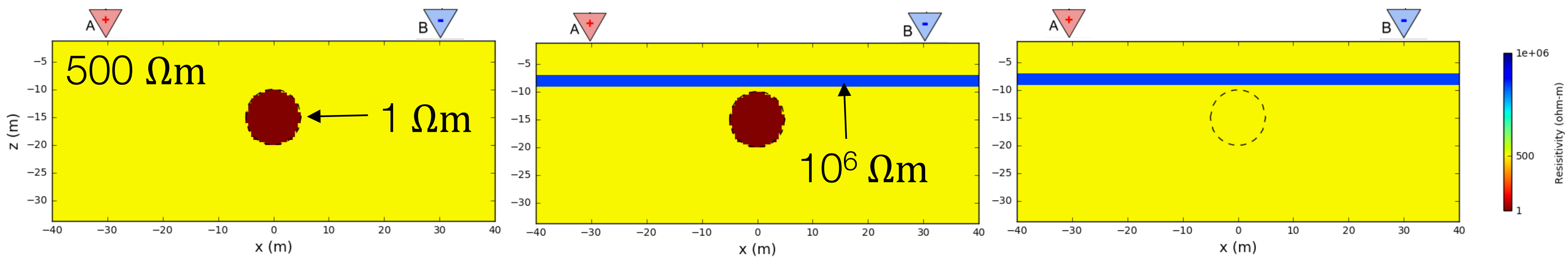
- Soundings, profiles and arrays
- Data, pseudosections and inversion
- Sensitivity
- Survey Design
- Case History – Mt Isa

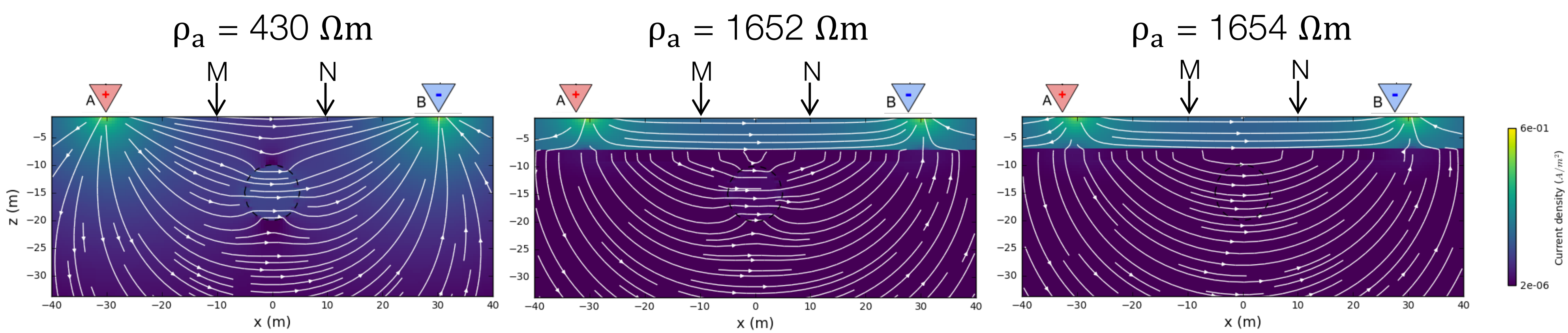
- Effects of background resistivity

Effects of background resistivity

Resistivity models (thin resistive layer)

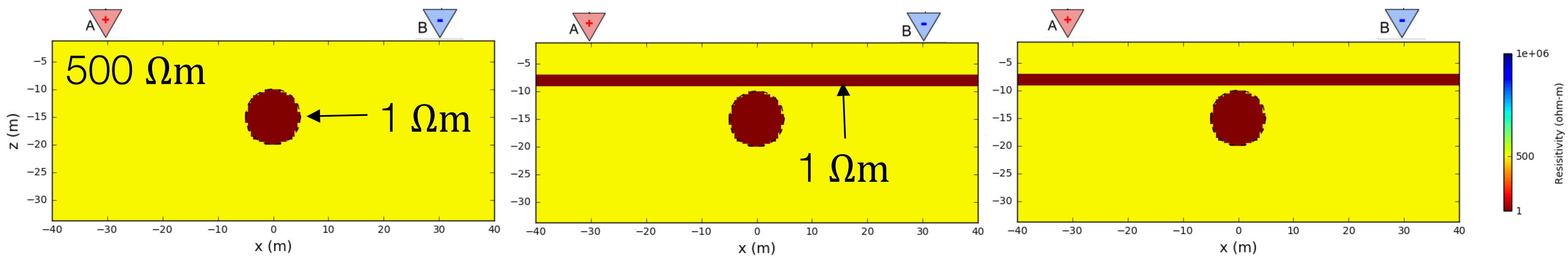


Currents and measured data at MN

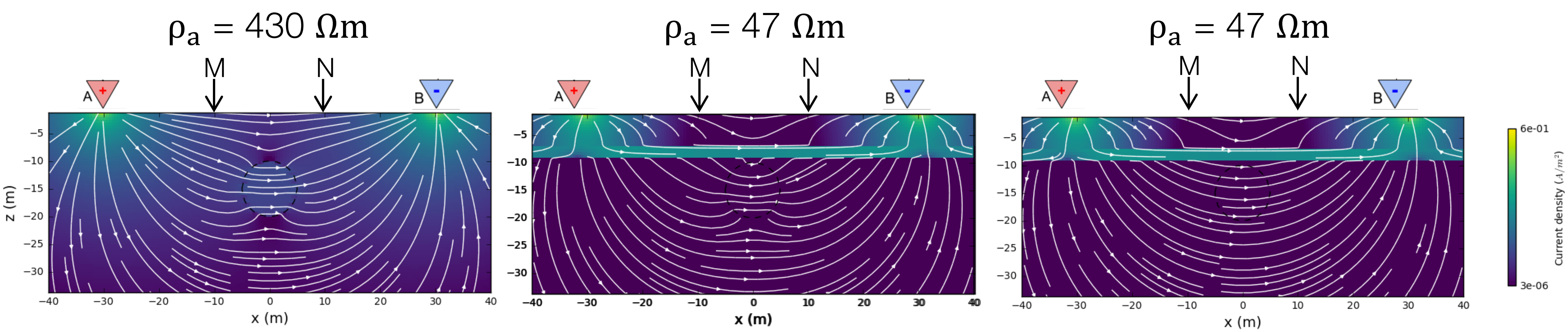


Effects of background resistivity

Resistivity models (thin conductive layer)



Currents and measured data at MN



End of DCR

- Introduction to EM
- DCR
- EM Fundamentals ← Next up
- Inductive sources
    - Lunch: Play with apps
- Grounded sources
- Natural sources
- GPR
- Induced polarization
- The Future