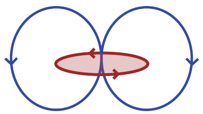
EM: Inductive Sources

DISC

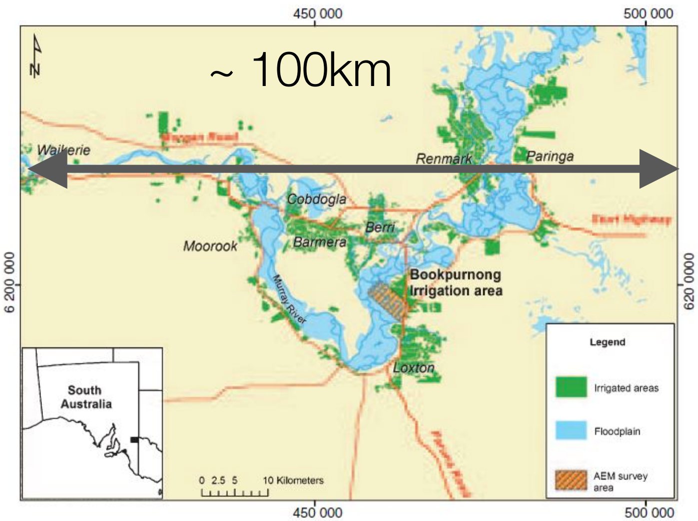
2017

Electromagnetics
Fundamentals and Applications

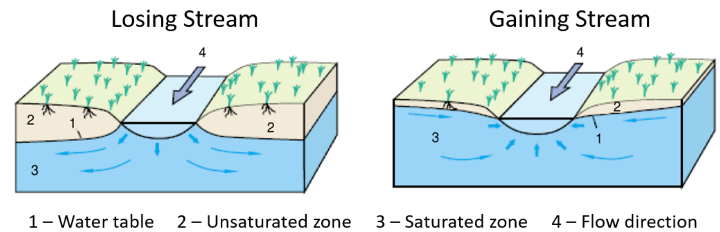


Motivation

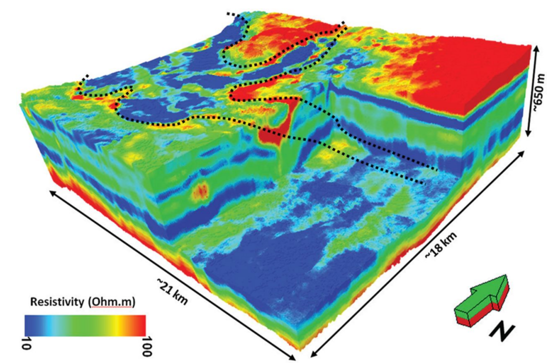
Large areas to be covered



Groundwater



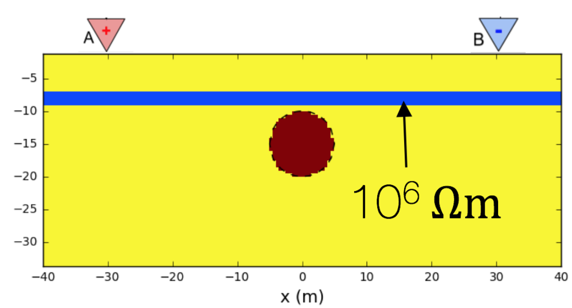
High resolution near surface



Rugged terrain

Minerals

Shielding problem



Outline

Setup

- Basic experiment
- Transmitters, Receivers

Time Domain EM

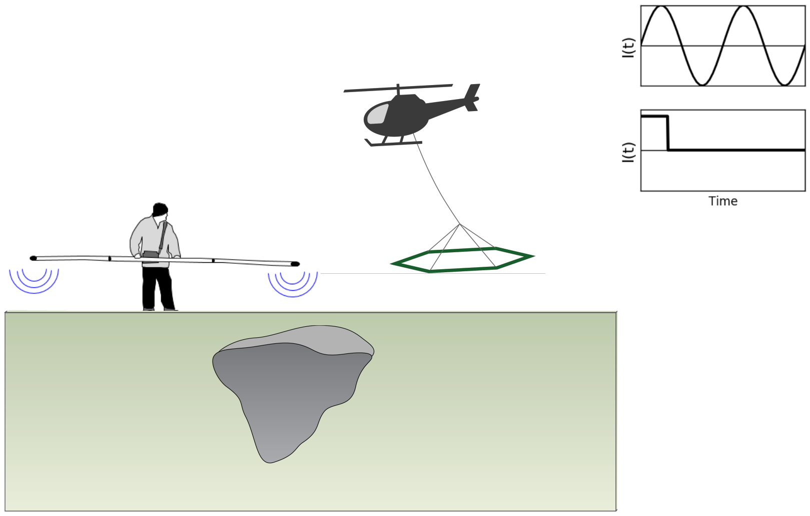
- Vertical Magnetic Dipole
- Propagation with Time
- Case History – Groundwater

Frequency Domain EM

- Vertical Magnetic Dipole
- Effects of Frequency
- Case History – Groundwater

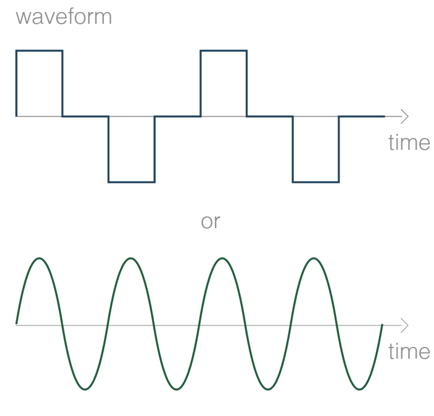
Important questions

- What is the target?
  - at the surface? At depth?. 1D, 2D, 3D?
- Transmitter
  - Location: surface? in the air?
  - Waveform: frequency or time?
  - “Size” and orientation?
- Exciting the target
  - Conductivity of the target and host
  - Geometry of the target (Coupling)
- Receiver and data
  - What fields to measure?
  - What instrument?

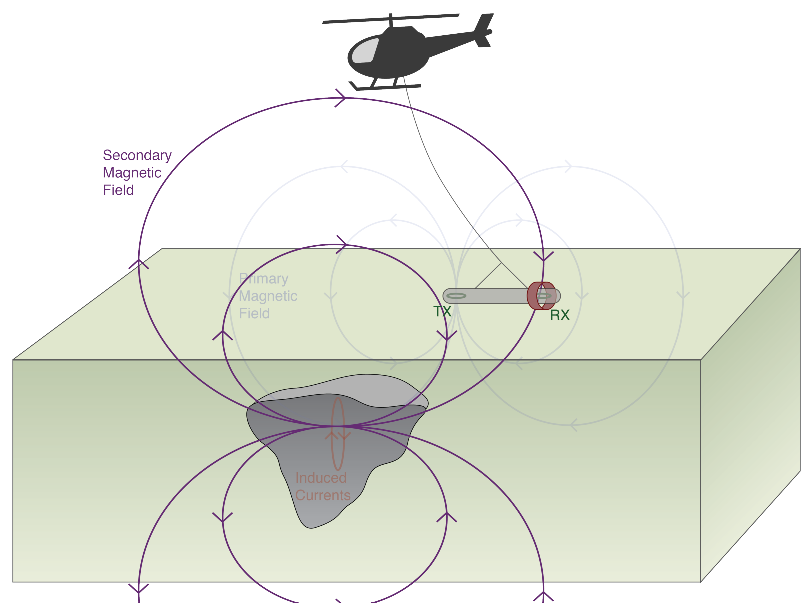


- Where to collect data? How many? How accurate?
- What is depth of investigation?
- What is the “footprint” of the transmitter”
  - These are questions of **SURVEY DESIGN**

Basic Experiment

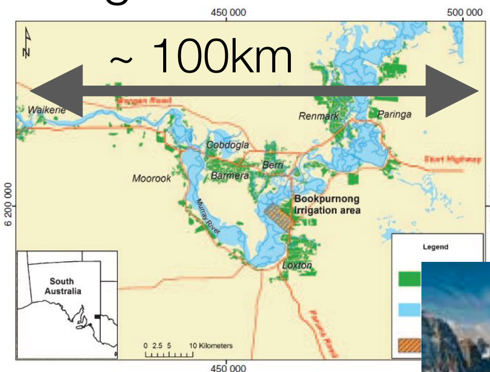


- **Transmitter:**
  - Produces a primary magnetic field
- **Exciting the target:**
  - Time varying magnetic fields generate electric fields everywhere
  - Producing currents in conductors
- **Receiver:**
  - Induced currents produce secondary magnetic fields



Transmitter

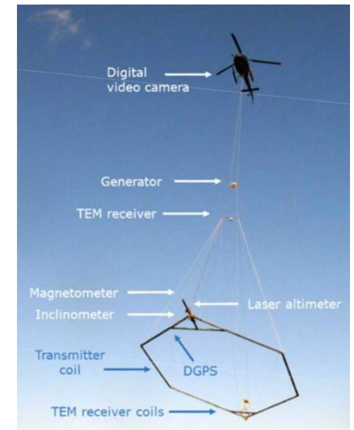
Large areas



Rugged terrain

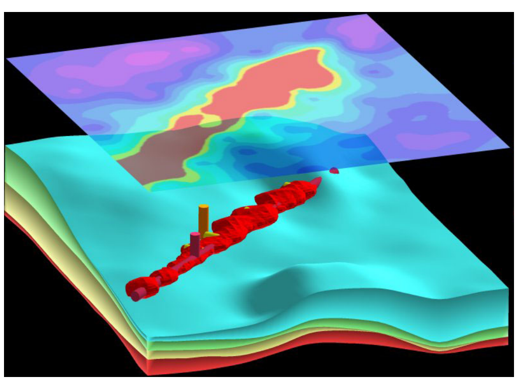
Airborne Survey

Resolve

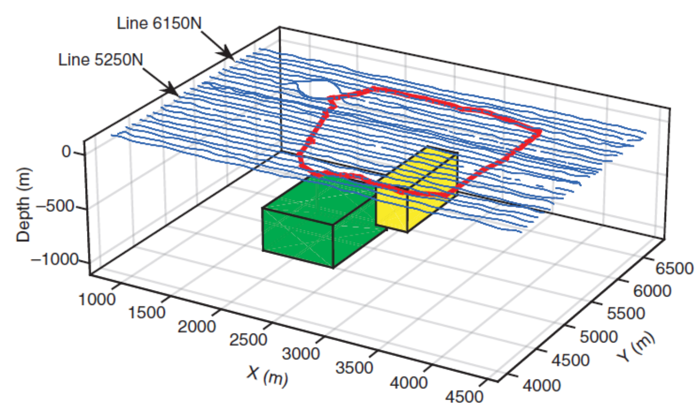


SkyTEM

Deep Targets

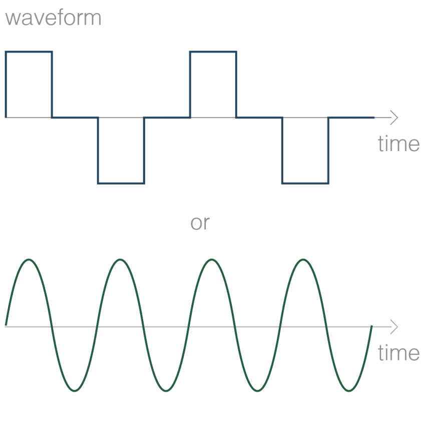


Large Loop



Transmitter

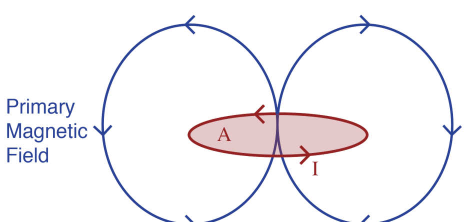
- Time or frequency?



- Key factor is moment

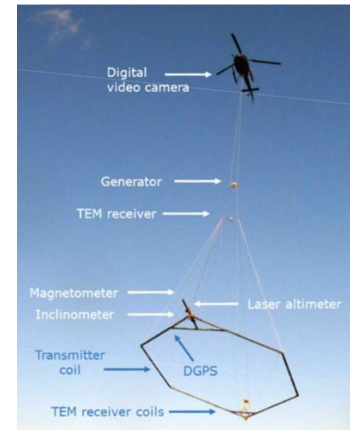
$$m = I\ (\text{current})\ A\ (\text{area})\ N\ (\#\text{ of turns})$$

$$\mathbf{B}(\mathbf{r}) = \frac{\mu_0}{4\pi}\left(\frac{3\mathbf{r}(\mathbf{m}\cdot\mathbf{r})}{|\mathbf{r}|^5} - \frac{\mathbf{m}}{|\mathbf{r}|^3}\right)$$



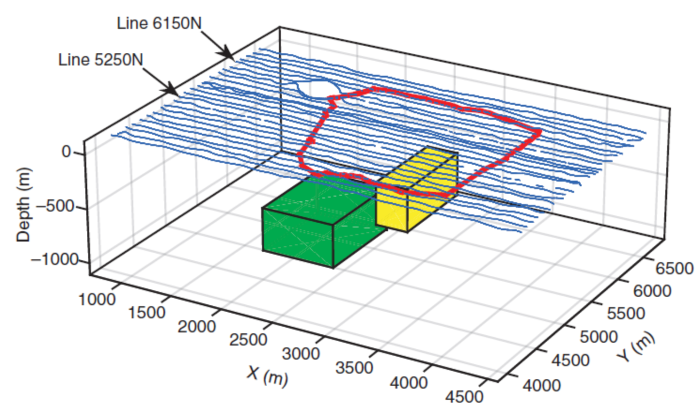
Airborne Survey

Resolve



SkyTEM

Large Loop

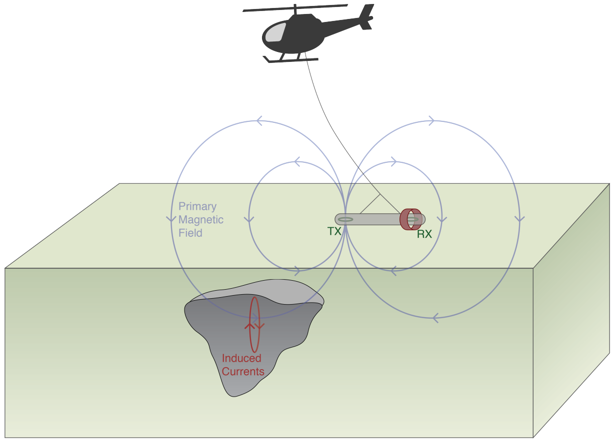


Exciting the target

- Primary field from a loop

$$\mathbf{B}(\mathbf{r}) = \frac{\mu_0}{4\pi}\left(\frac{3\mathbf{r}(\mathbf{m}\cdot\mathbf{r})}{|\mathbf{r}|^5} - \frac{\mathbf{m}}{|\mathbf{r}|^3}\right)$$

- Fields fall off
  - $1/r^3$ geometric decay
  - Attenuation
- Want to be as close as possible to target
  - Ground based systems
  - Helicopter
  - Fixed wing aircraft
- Always concerned about coupling

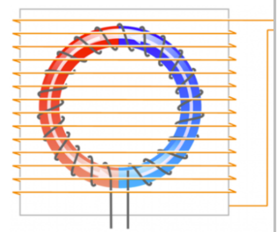


Receiver and Data

Magnetometer

- Measures:
  - Magnetic field
  - 3 components
- eg. 3-component fluxgate

$$\mathbf{b}(t)$$

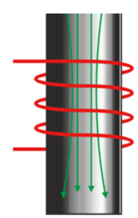


Fluxgate

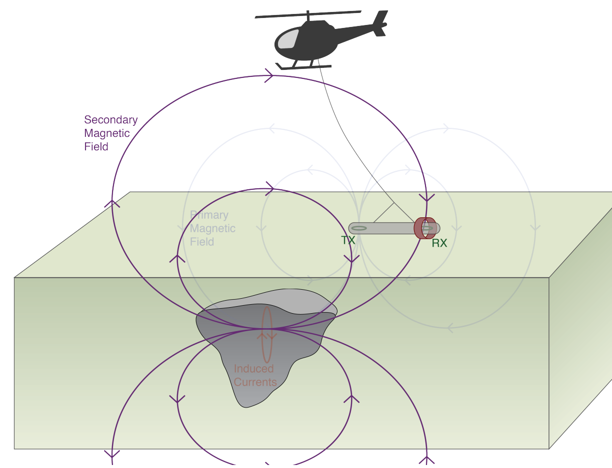
Coil

- Measures:
  - Voltage
  - Single component that depends on coil orientation
    - Coupling matters
- eg. airborne frequency domain.
  - ratio of Hs/Hp is the same as Vs/Vp

$$\frac{\partial b}{\partial t}$$



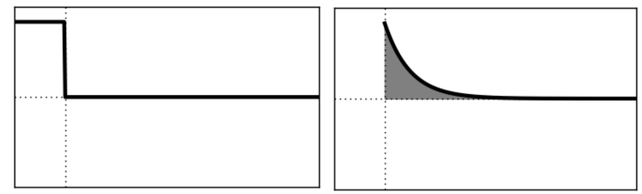
Coil



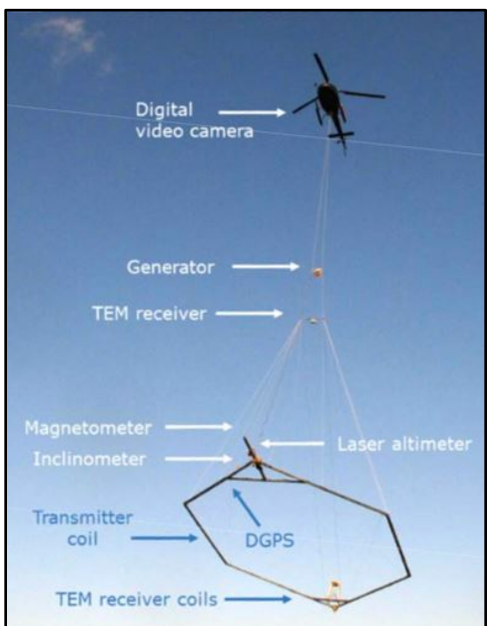
Receiver: Time Domain

- Primary field has off-time
- Measure secondary fields
- Receivers can be mounted on transmitter loop or above it

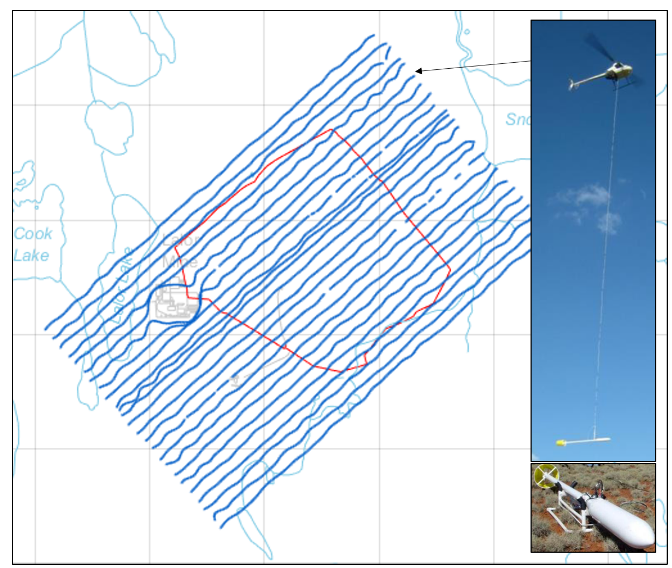
Current Response



SkyTEM

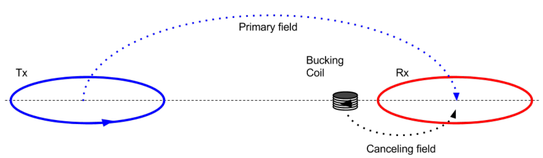


HeliSAM

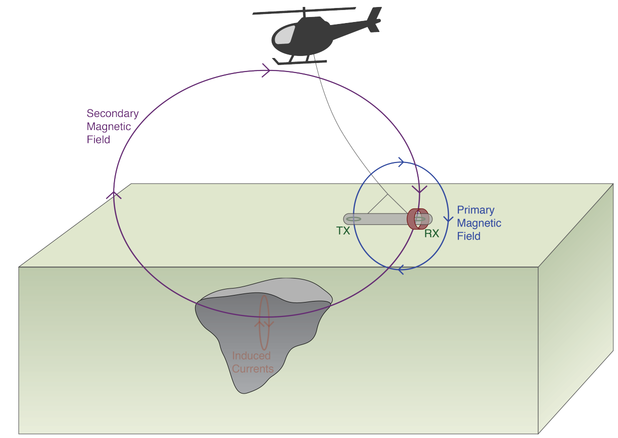


Receiver: Frequency Domain

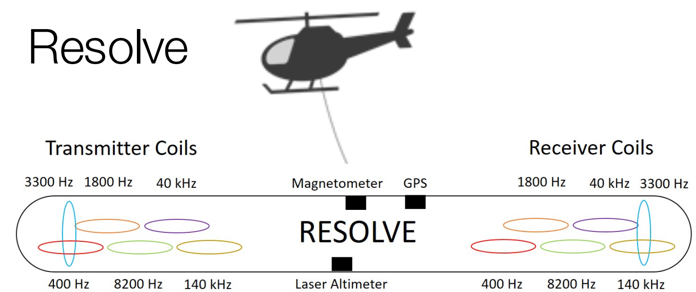
- Primary field
  - always “on”
  - large compared to secondary fields
- Primary removal
  - Compute and subtract
  - Bucking coil



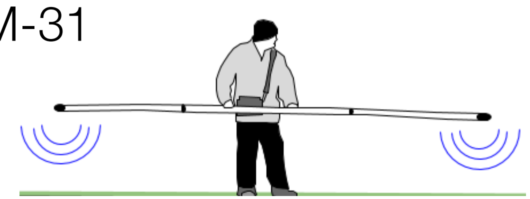
- Main requirement:
  - Know positions of Tx and Rx
  - Keep them in one unit



Resolve



EM-31



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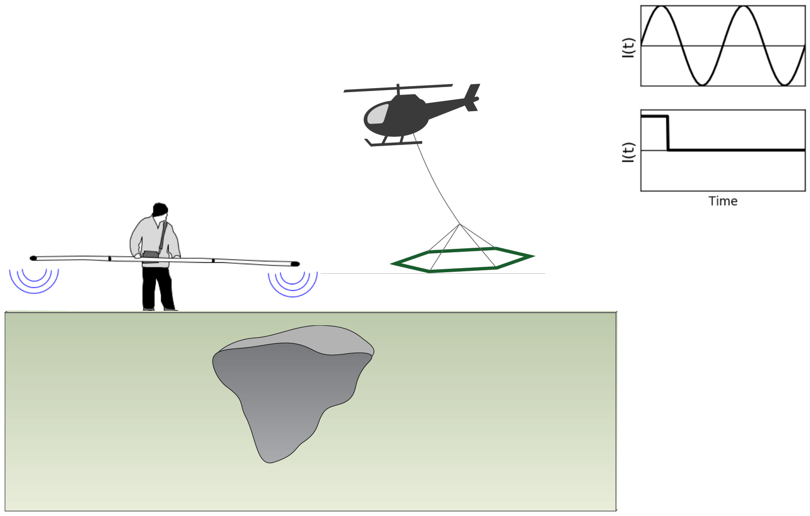

- Where to collect data? How many? How accurate?
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- What is the “footprint” of the transmitter”
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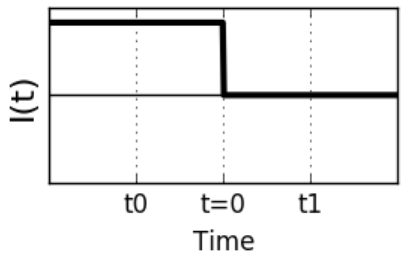
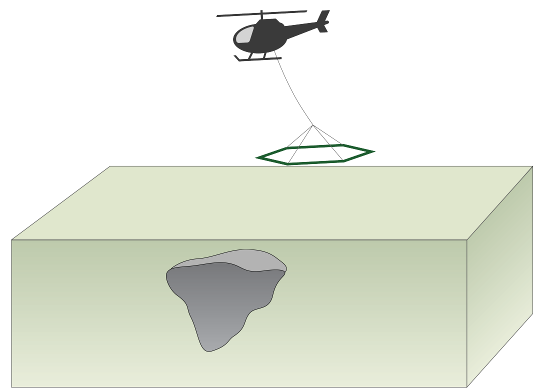
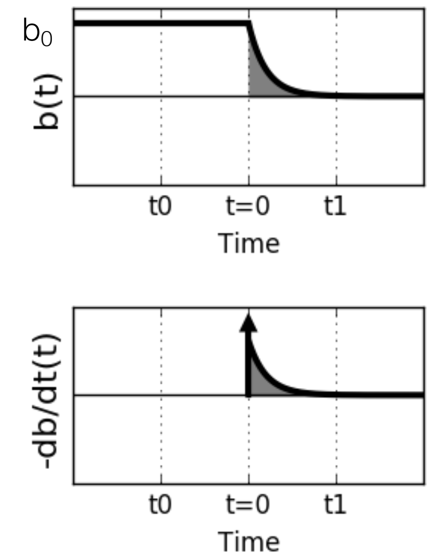
EM with Inductive Sources: Time Domain

Transmitter current





Receiver



time	**b**	**db/dt**
$t < 0$	b_0	0
$t = 0$	b_0	$-b_0\delta(t)$
$t > 0$	secondary	secondary

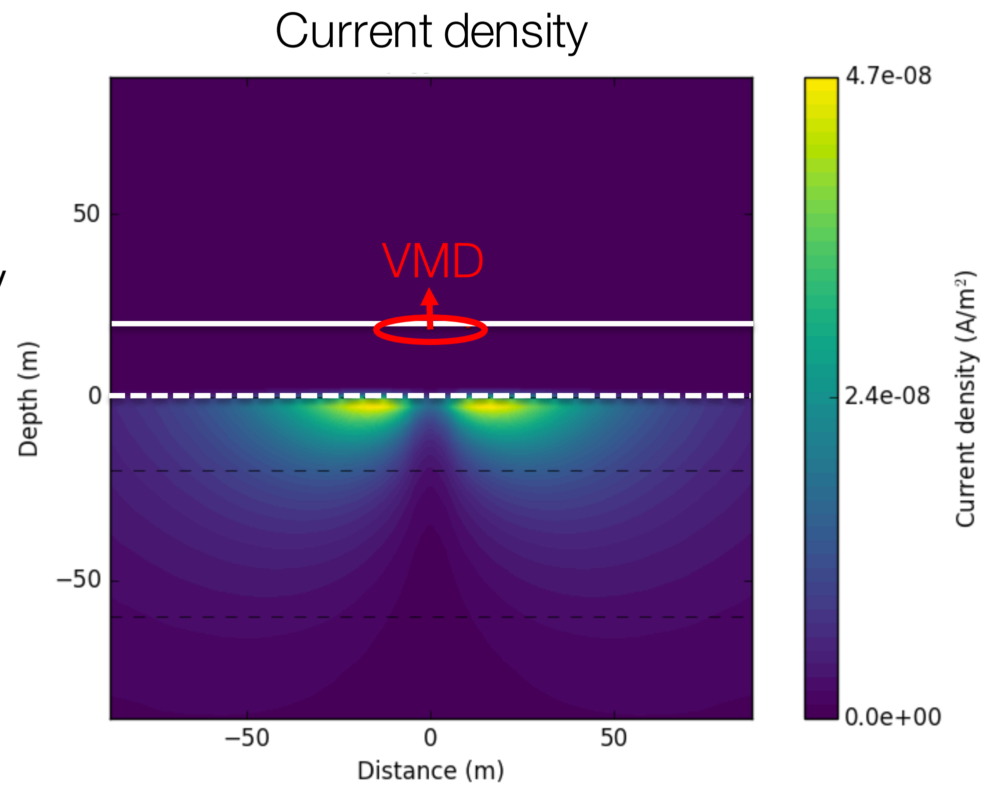
$\delta(t)$: Dirac-delta function

Footprint of Airborne EM system

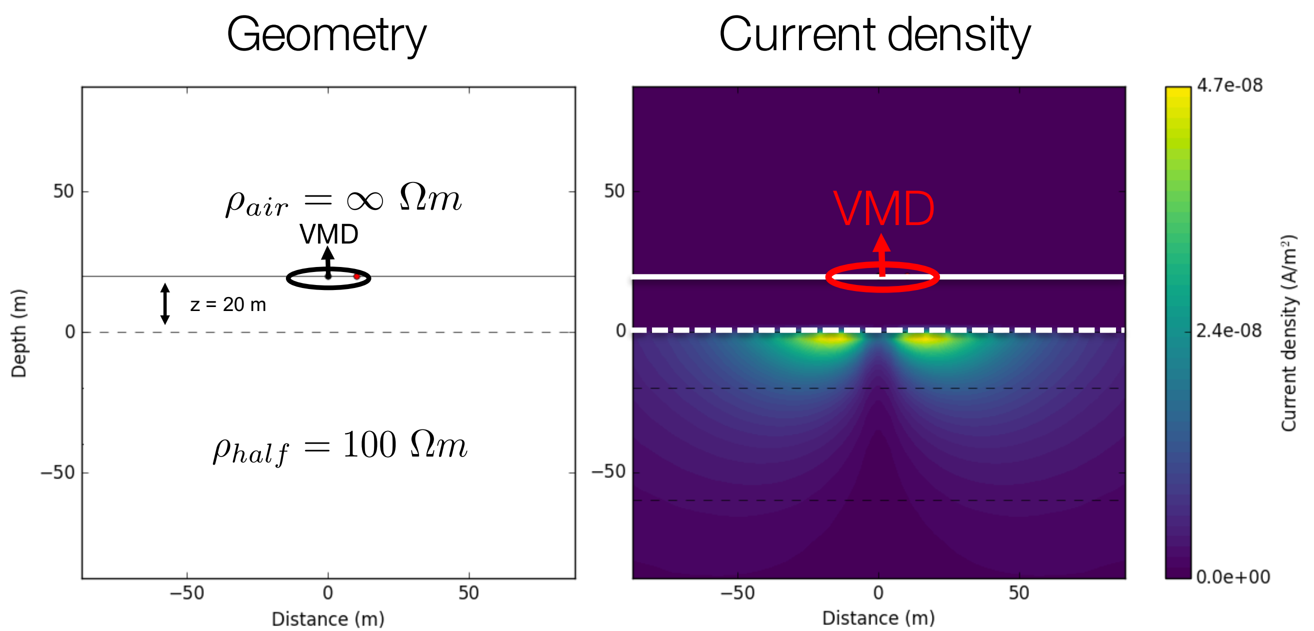
- What volume of earth is "seen" by the airborne system?
  - Where are the currents?

- Currents depend on
  - Transmitter
  - Waveform: time or frequency
  - Background conductivity

- Simple case: loop source over homogeneous earth

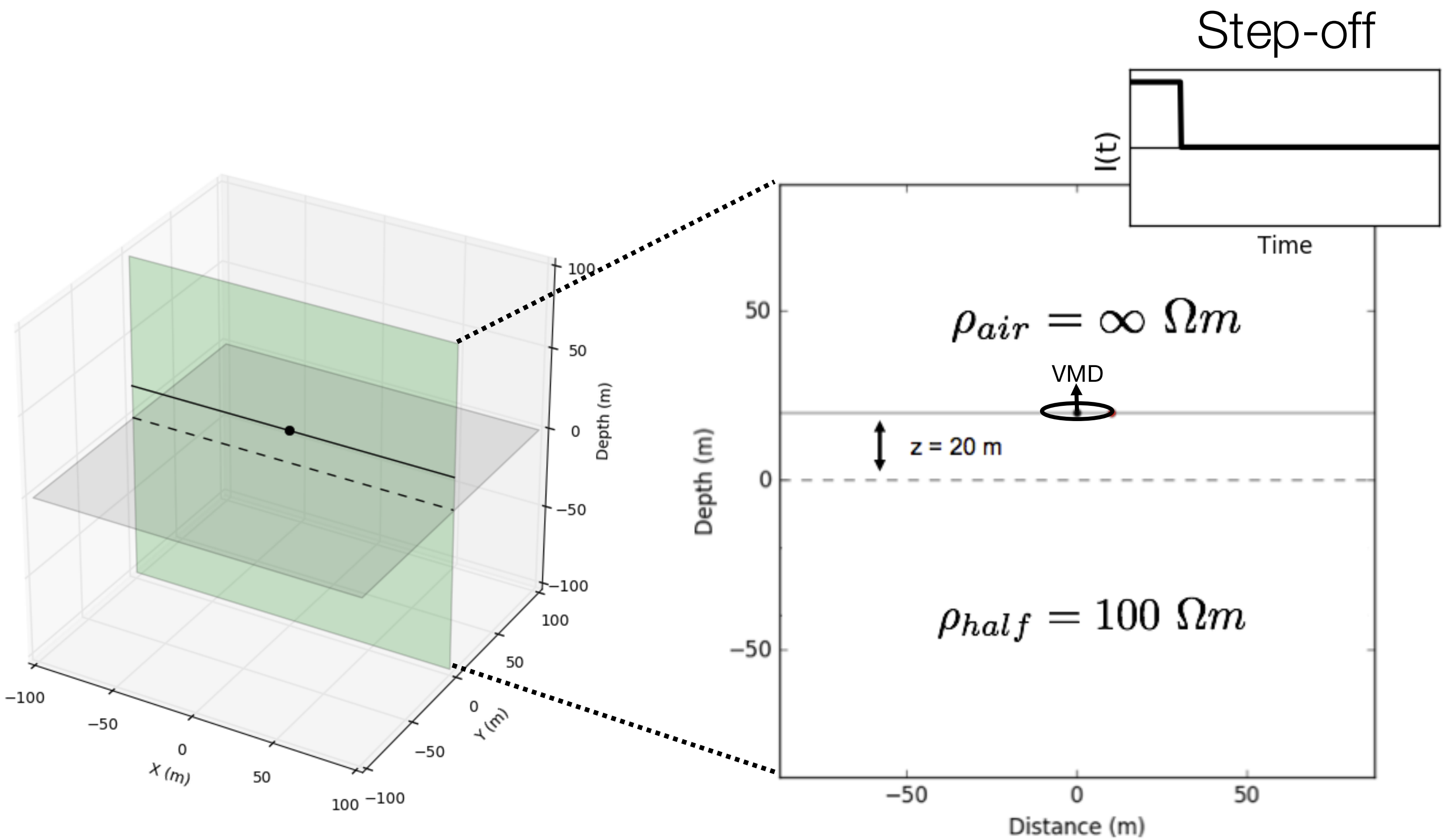


Vertical Magnetic Dipole (VMD)



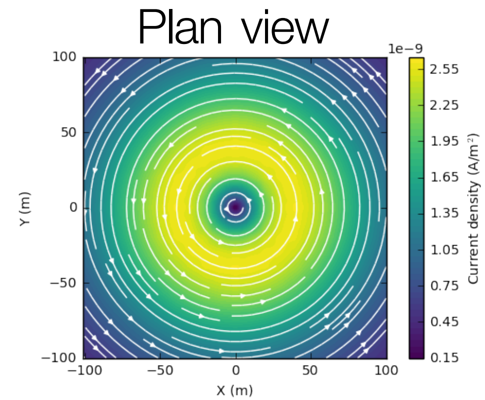
- Some questions
  - Where, and how strong, are the currents?
  - How do they depend upon the conductivity?
  - What do the resulting magnetic fields look like?

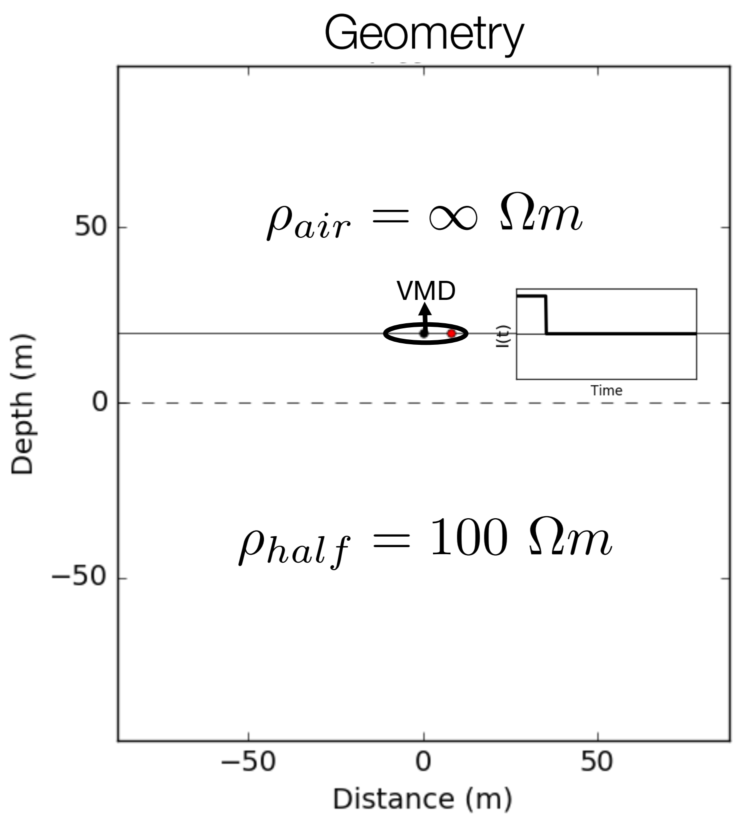
Vertical Magnetic Dipole over a halfspace (TDEM)

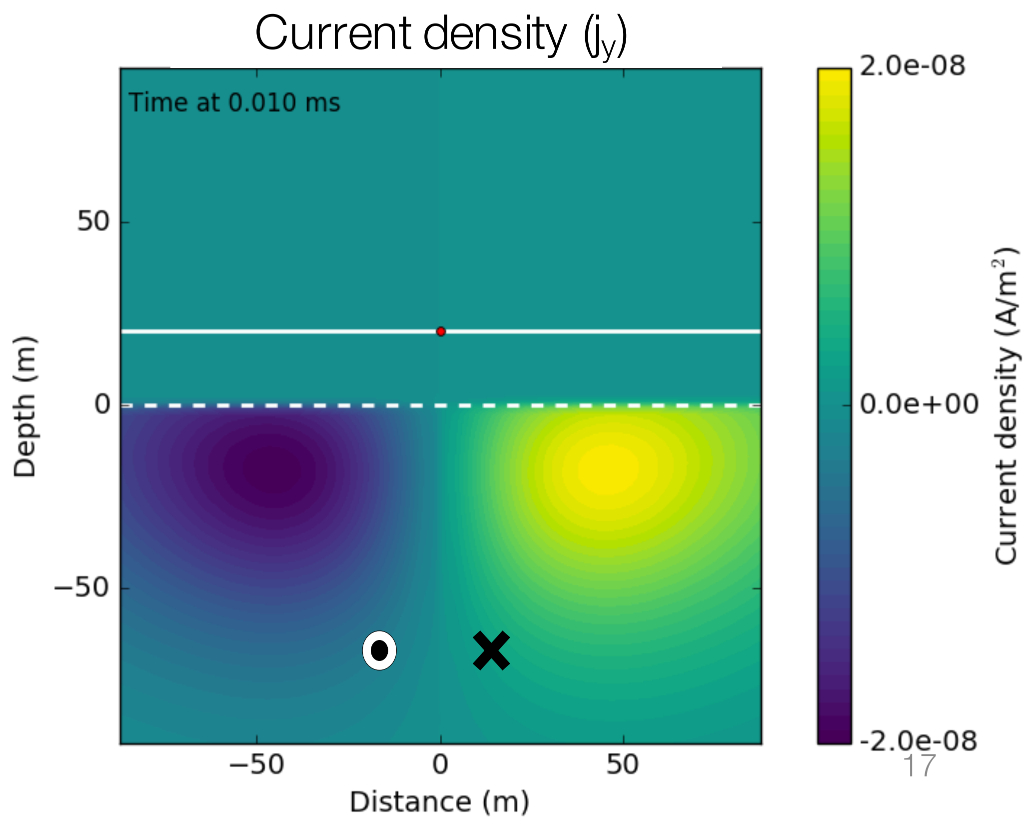


Current Density

- Time: 0.01ms

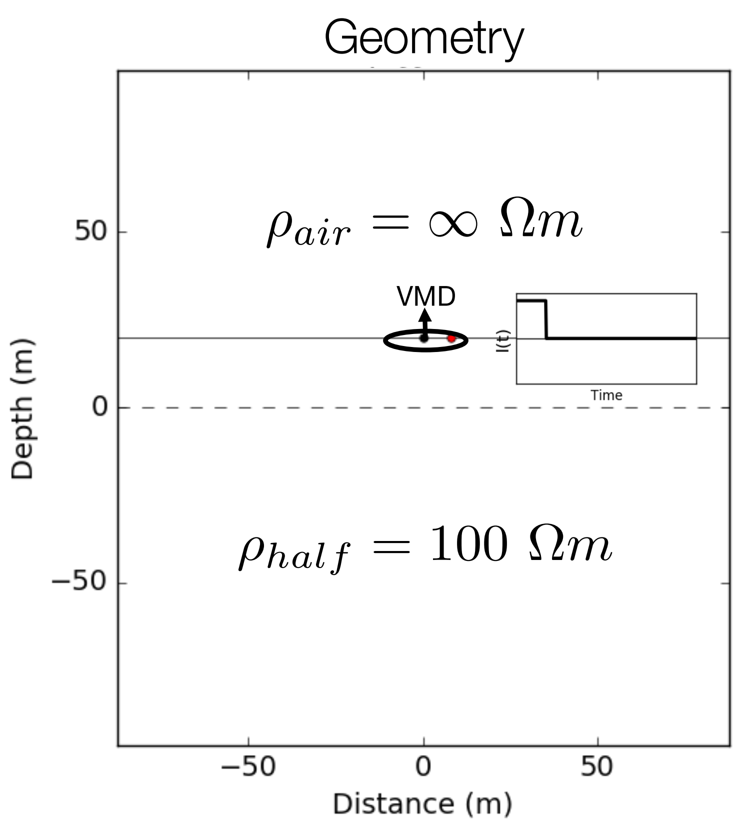


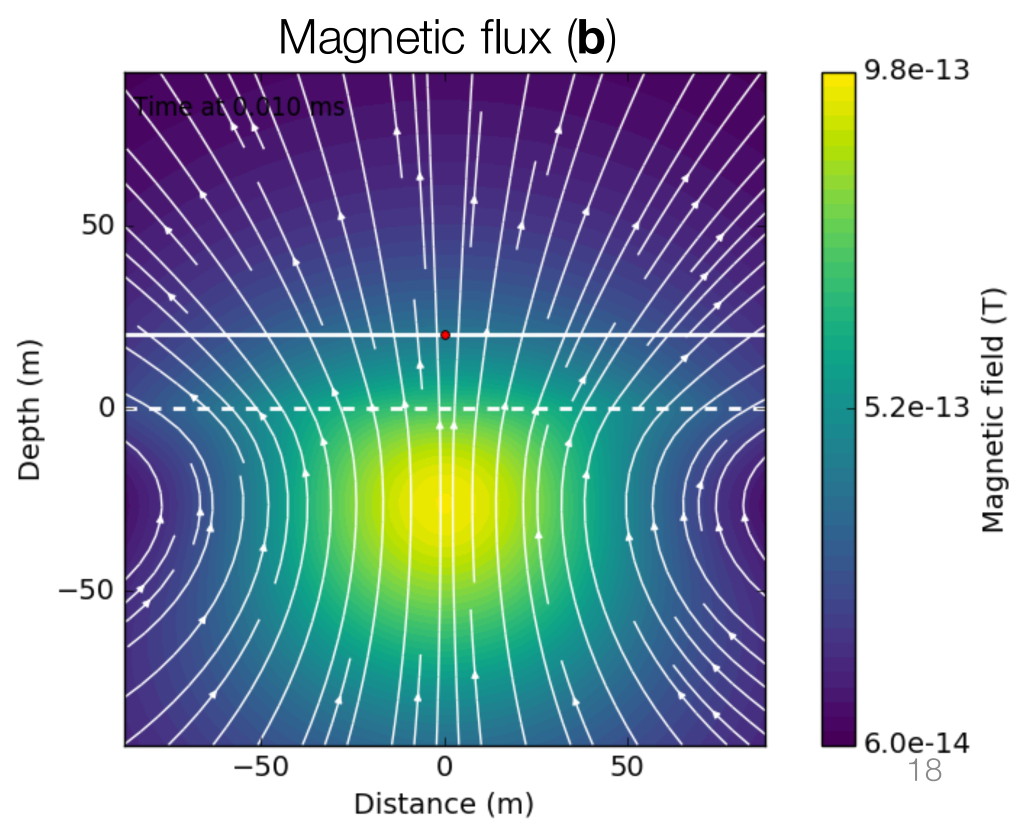




Magnetic flux density

- Time: 0.01ms

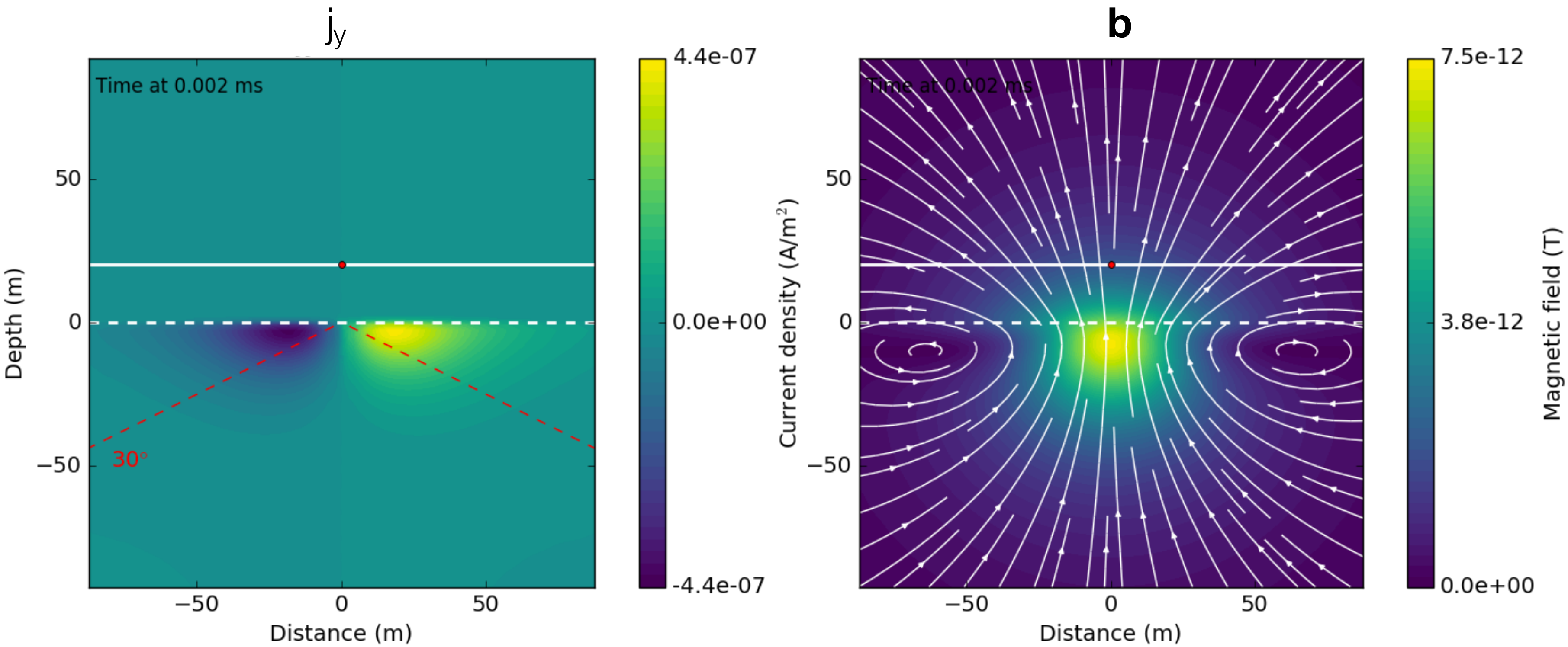




Propagation through time

- Time: 0.002ms
- diffusion distance = 18 m

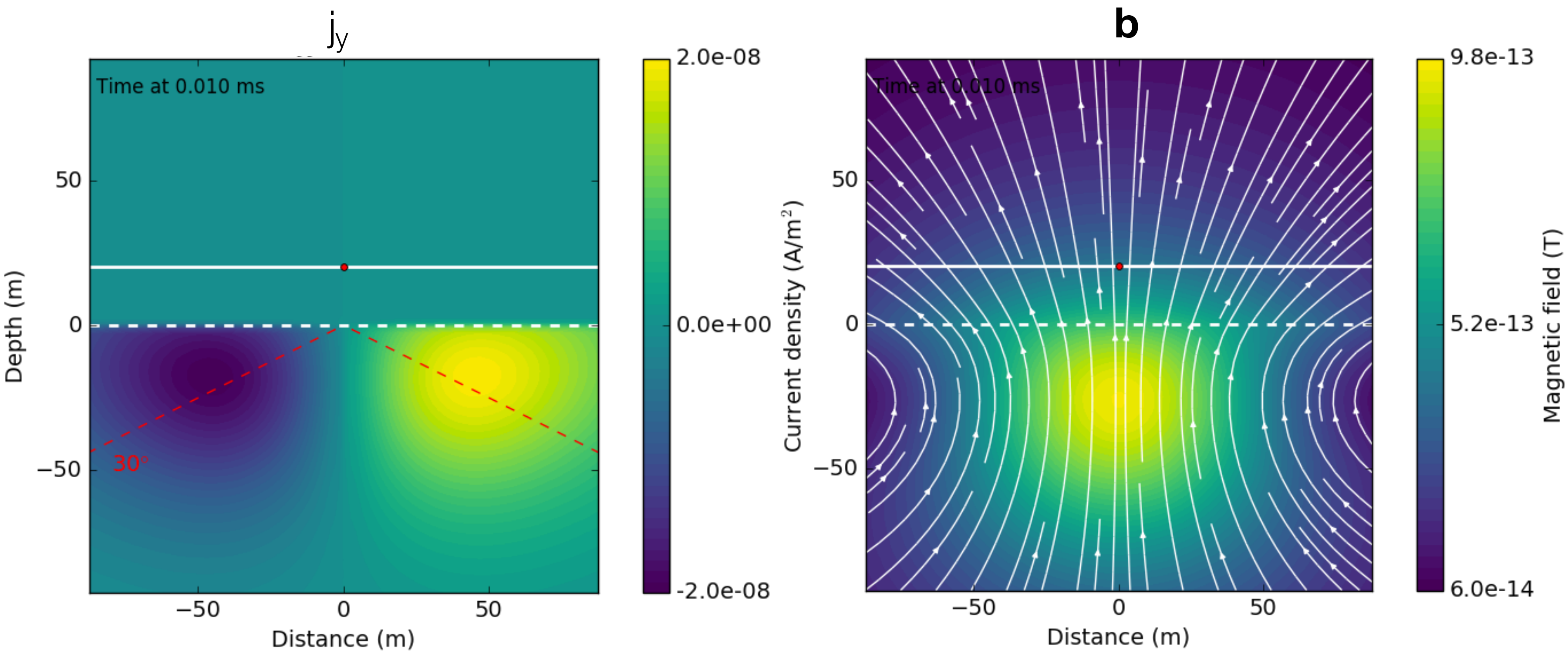
$$d = 1260\sqrt{t\rho}$$



Propagation through time

- Time: 0.01ms
- diffusion distance = 38 m

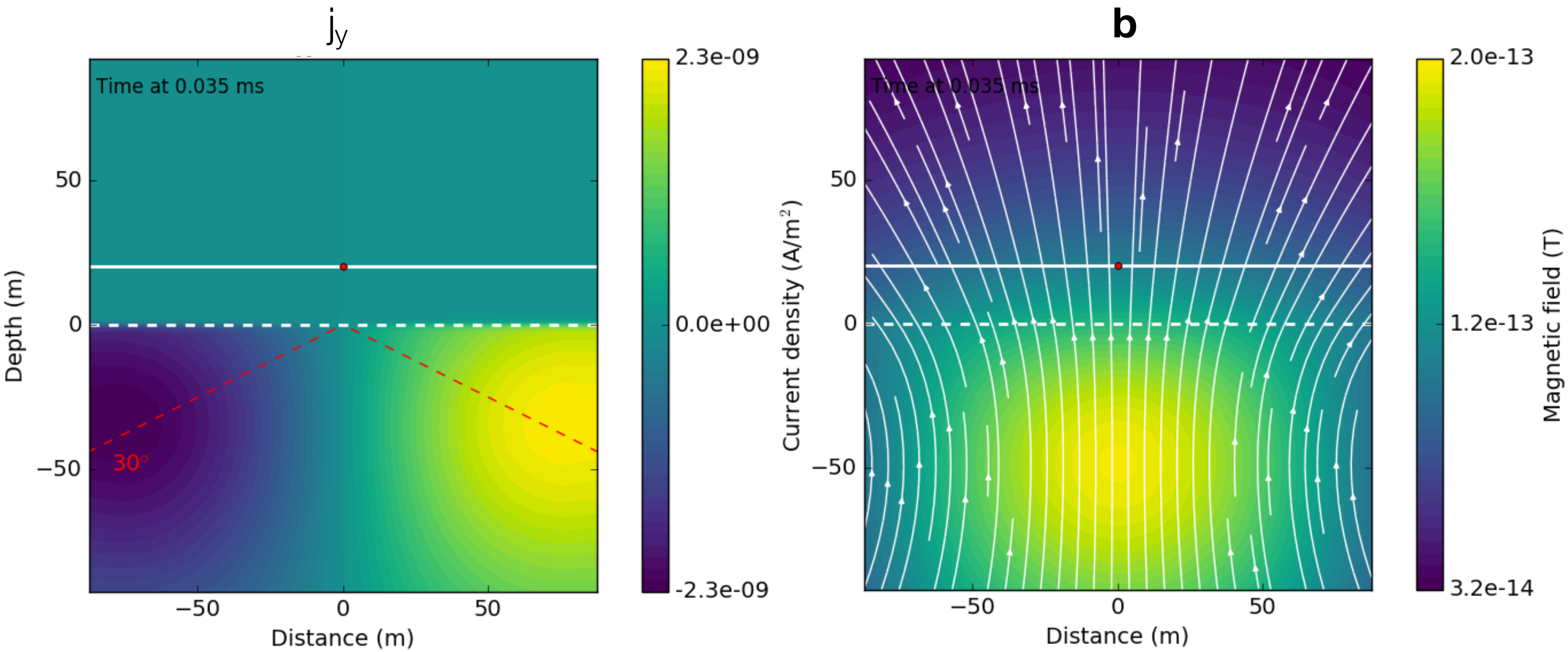
$$d = 1260\sqrt{t\rho}$$



Propagation through time

- Time: 0.035ms
- diffusion distance = 75 m

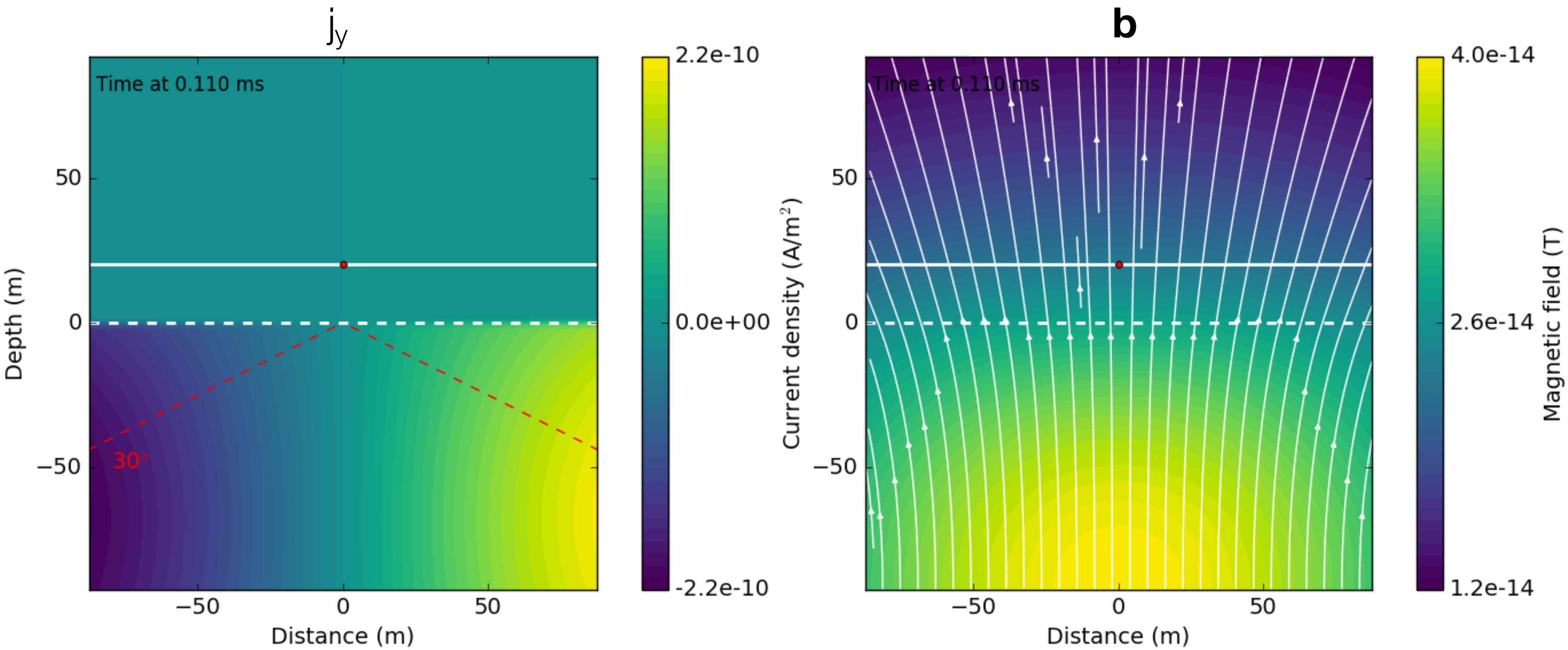
$$d = 1260\sqrt{t\rho}$$



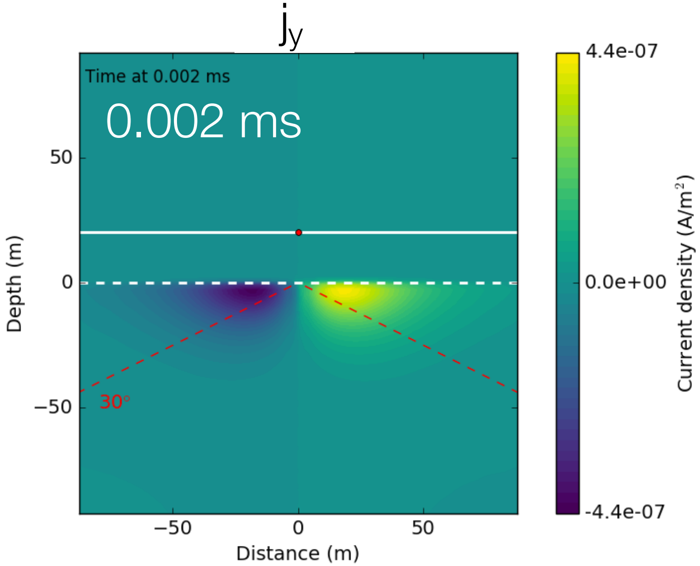
Propagation through time

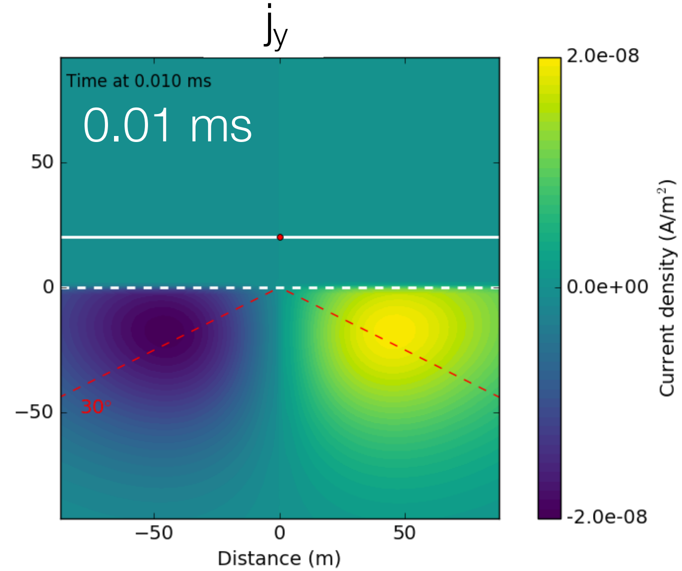
- Time: 0.110ms
- diffusion distance = 132 m

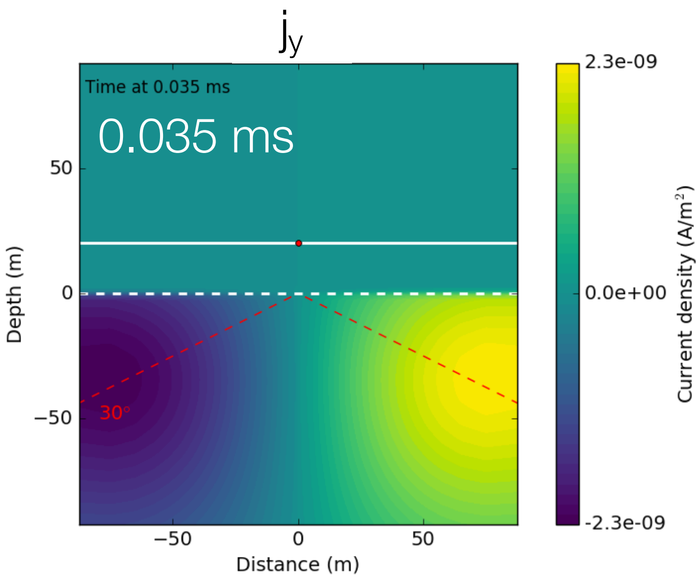
$$d = 1260\sqrt{t\rho}$$



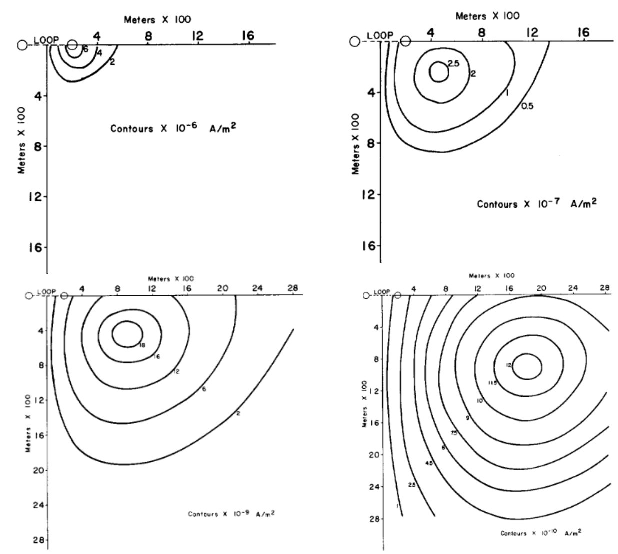
Summary: propagation through time







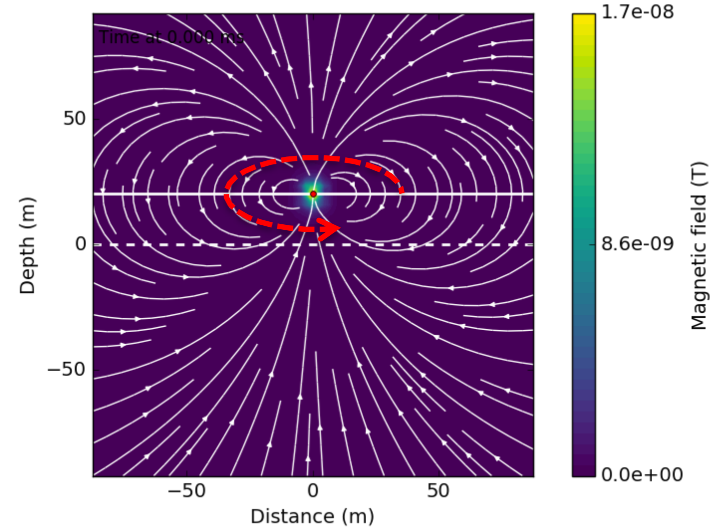
Nabighian (1979)



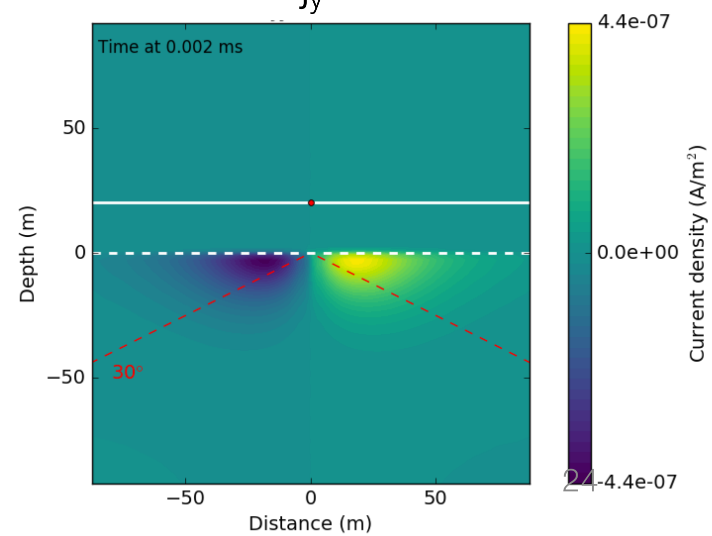
Important points

- Currents flow in same plane as transmitter currents
- Currents diffuse outward downward
- Each transmitter has a "footprint"
- Max resolution controlled by earliest time
- Depth of investigation controlled by latest time

magnetic field (on-time)



j_y



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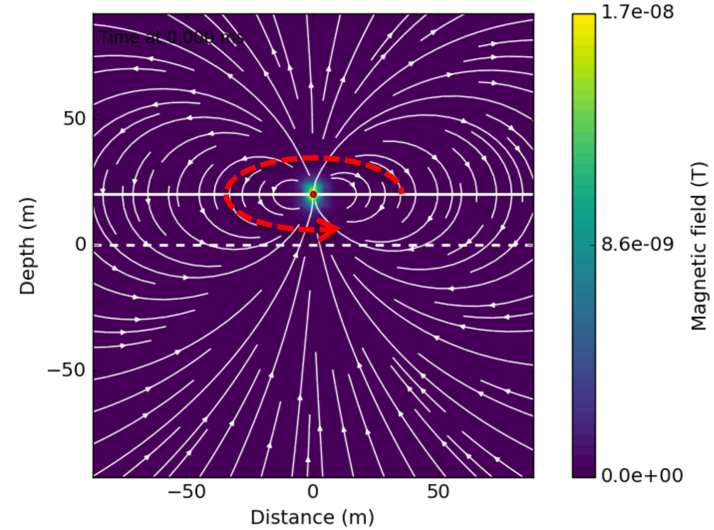
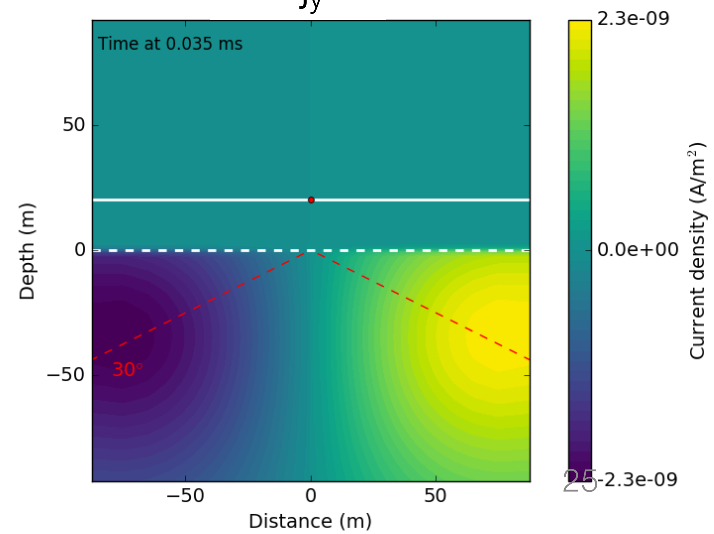
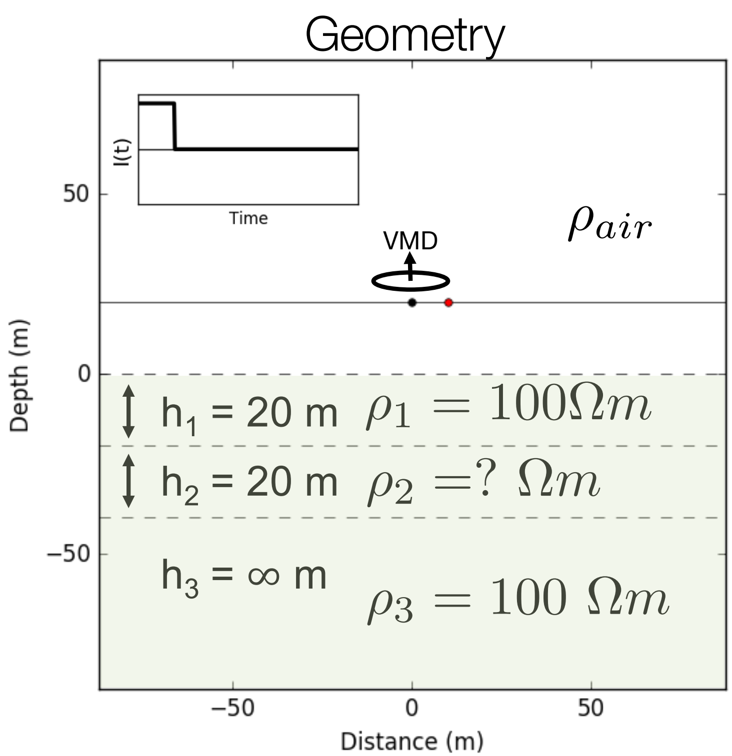


j_y



Layered earth

- 3 layers + air,
- ρ_2 varies

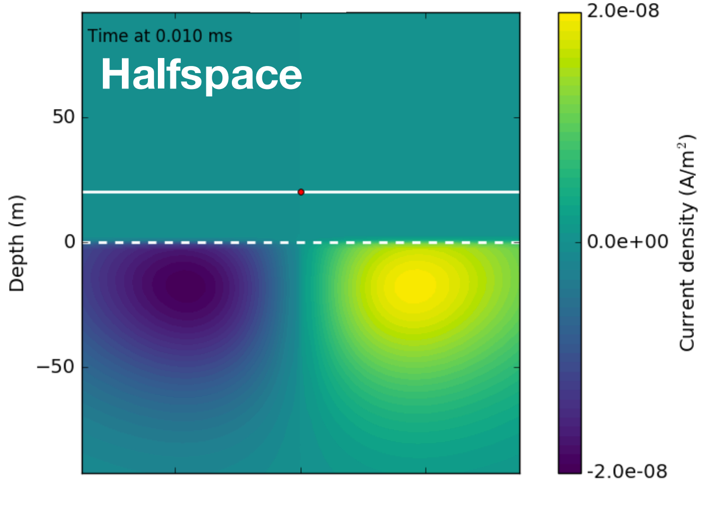


- Four different cases:
  - Halfspace
    $\rho_2 = 100 \ \Omega\mathrm{m}$
  - Resistive
    $\rho_2 = 1000 \ \Omega\mathrm{m}$
  - Conductive
    $\rho_2 = 10 \ \Omega\mathrm{m}$
  - Very conductive
    $\rho_2 = 1 \ \Omega\mathrm{m}$

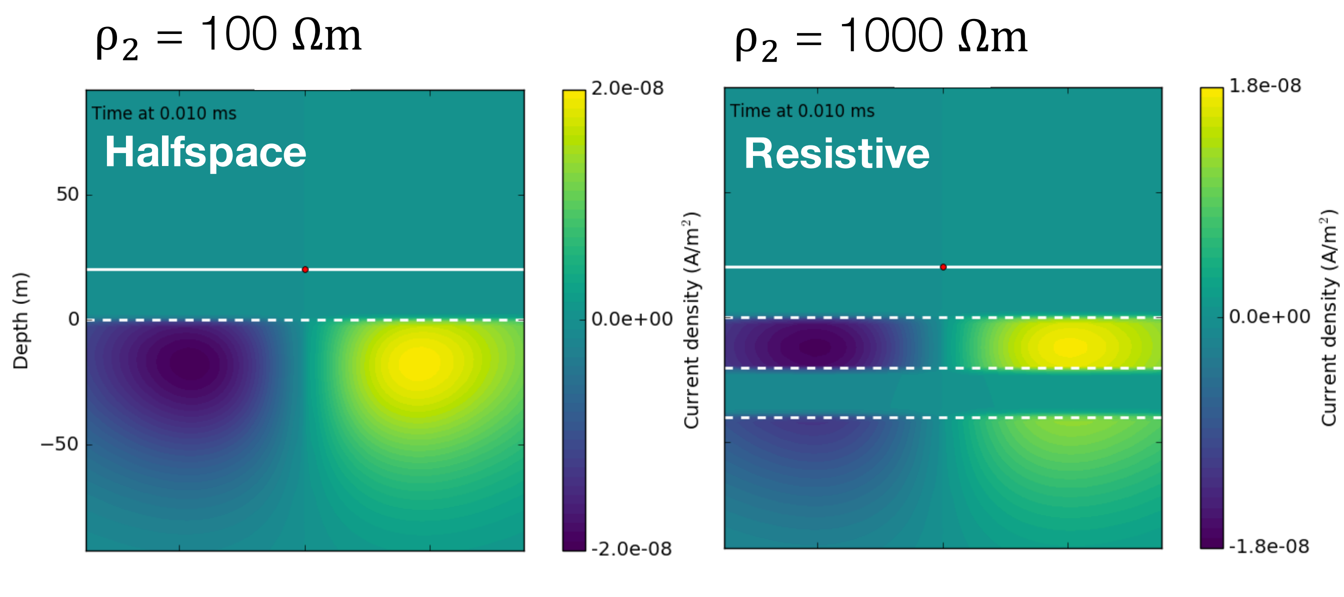
- Fields
  - j_y off-time
  - **b** off-time

Layered earth currents (j_y)

$\rho_2 = 100\ \Omega m$

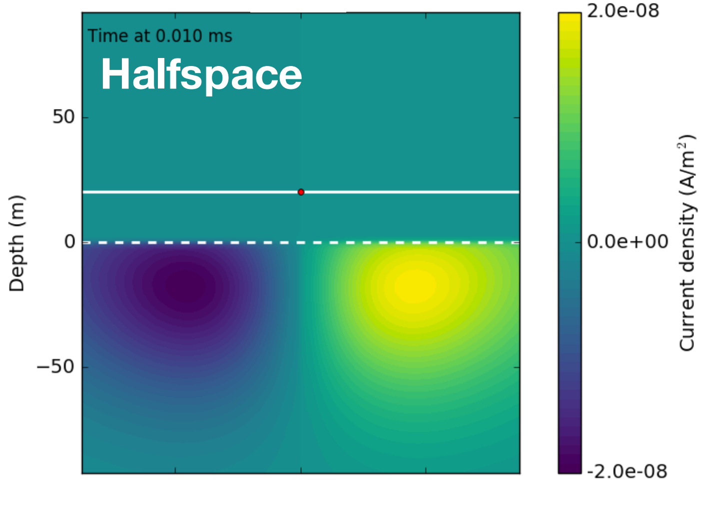


Layered earth currents (j_y)

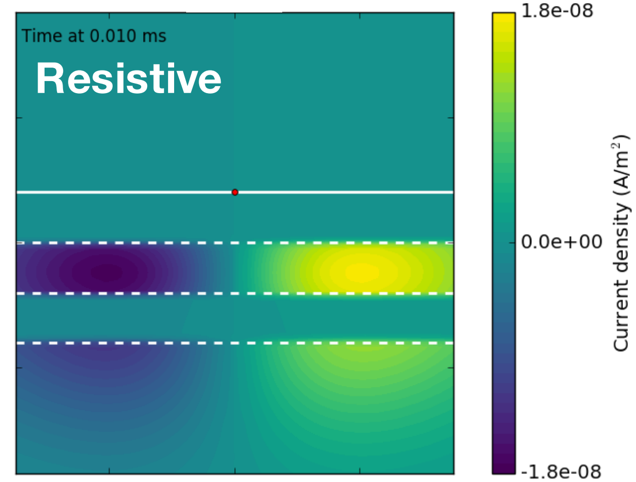


Layered earth currents (j_y)

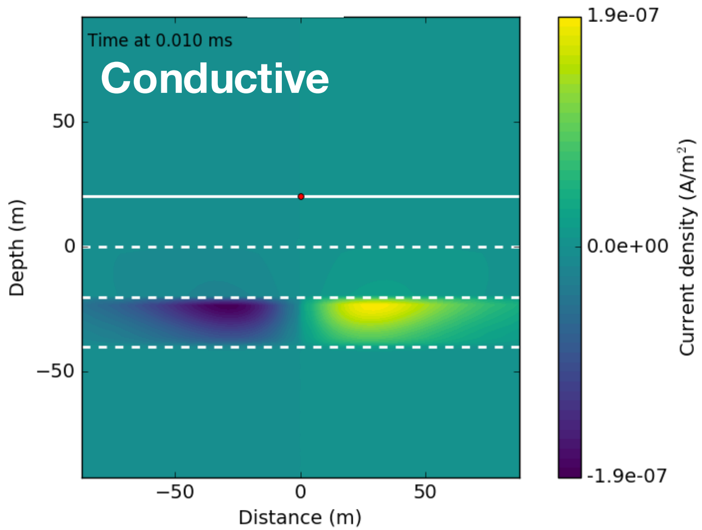
$\rho_2 = 100\ \Omega m$



$\rho_2 = 1000\ \Omega m$



$\rho_2 = 10\ \Omega m$



Layered earth currents (j_y)

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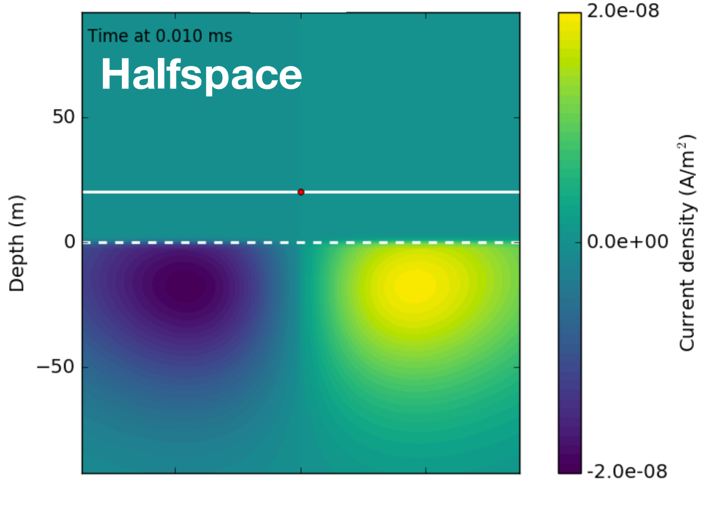
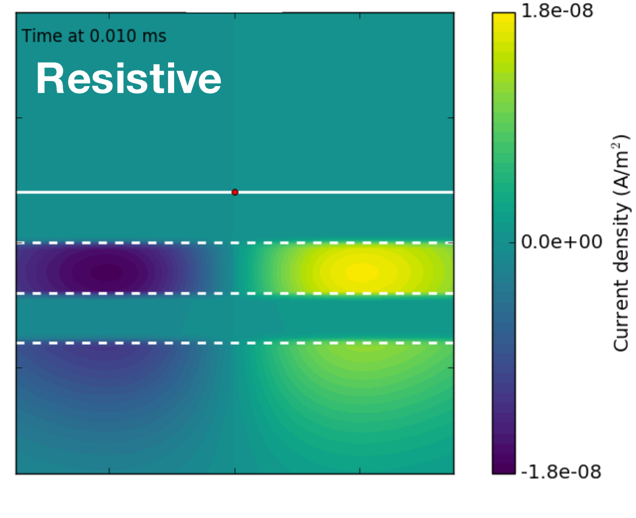
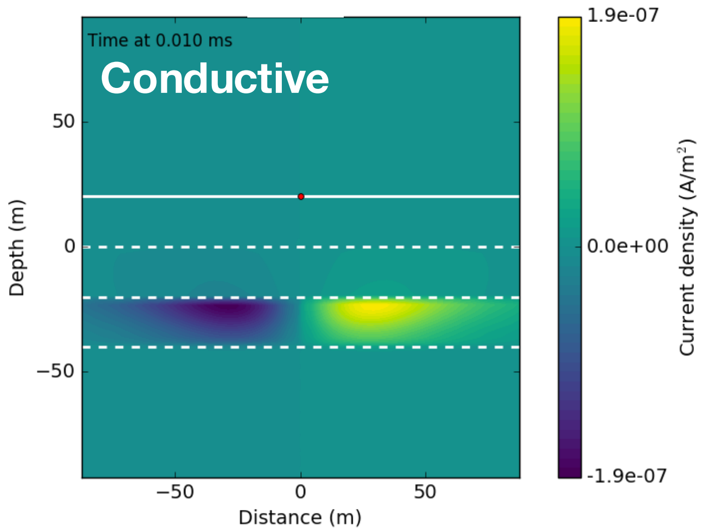


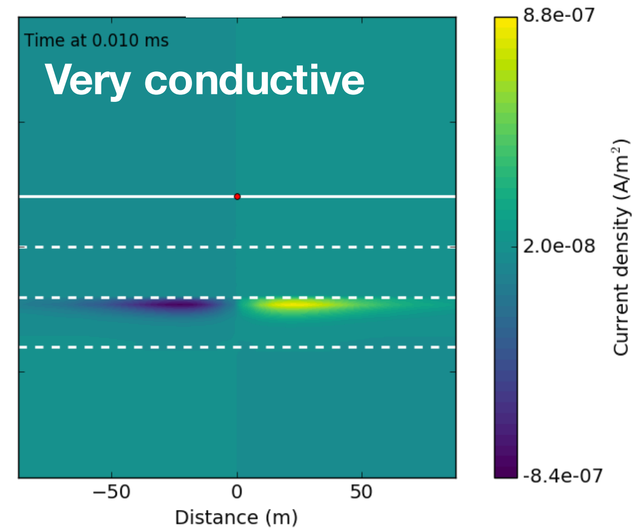
$\rho_2 = 1000\ \Omega m$



$\rho_2 = 10\ \Omega m$

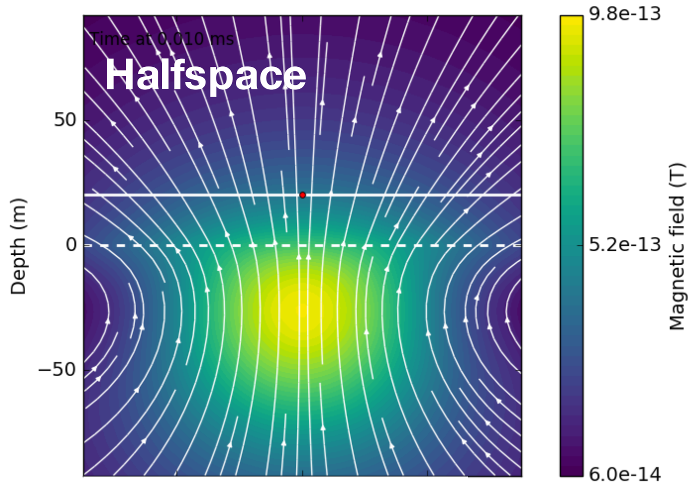


$\rho_2 = 1\ \Omega m$

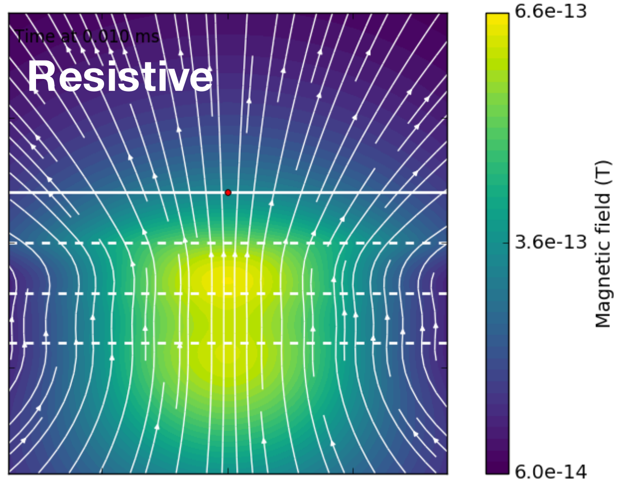


Layered earth mag. fields (**b**)

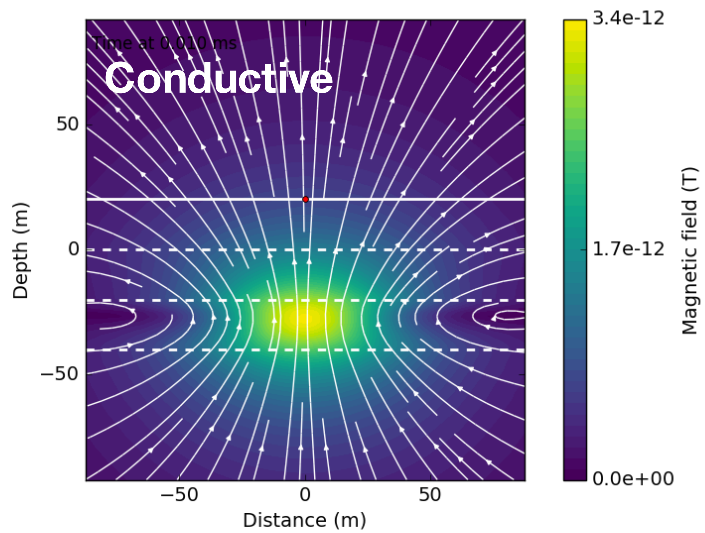
$\rho_2 = 100\ \Omega m$



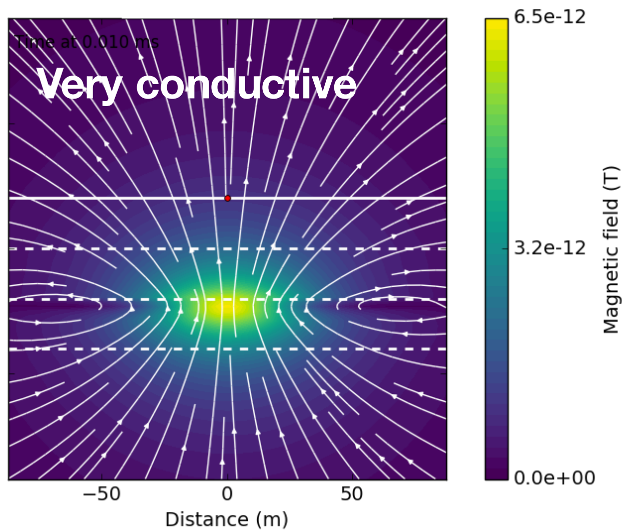
$\rho_2 = 1000\ \Omega m$



$\rho_2 = 10\ \Omega m$

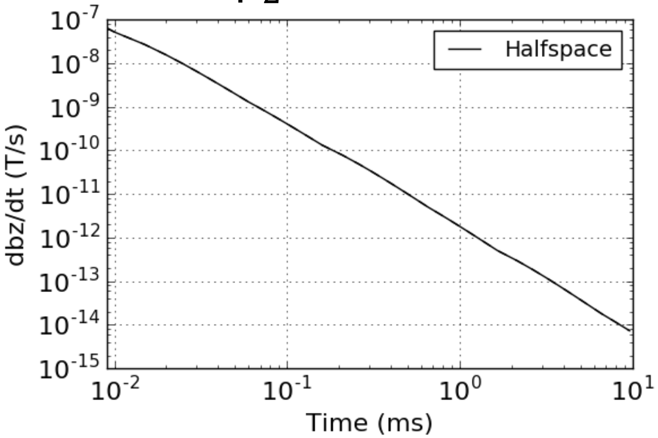


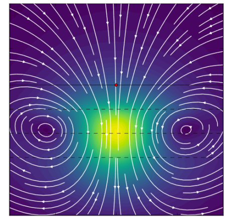
$\rho_2 = 1\ \Omega m$



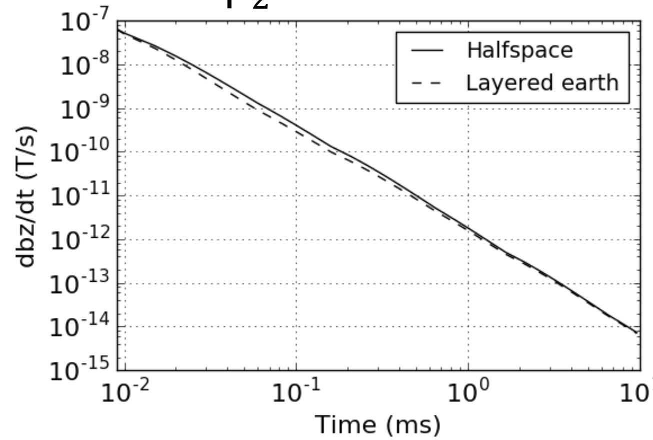
db_z/dt sounding curves

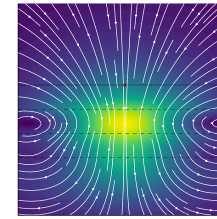
$\rho_2 = 100\ \Omega m$



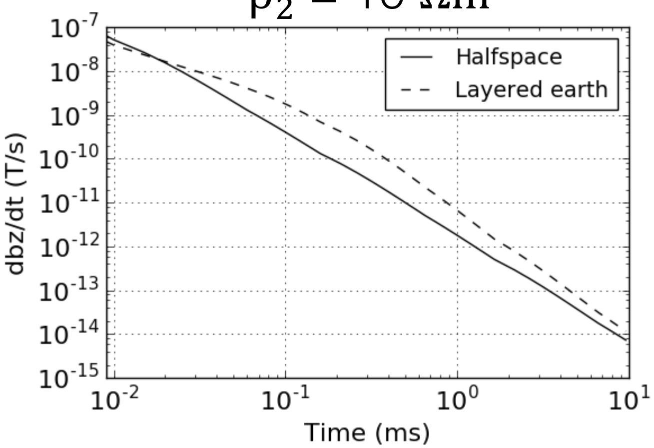


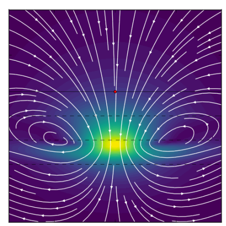
$\rho_2 = 1000\ \Omega m$



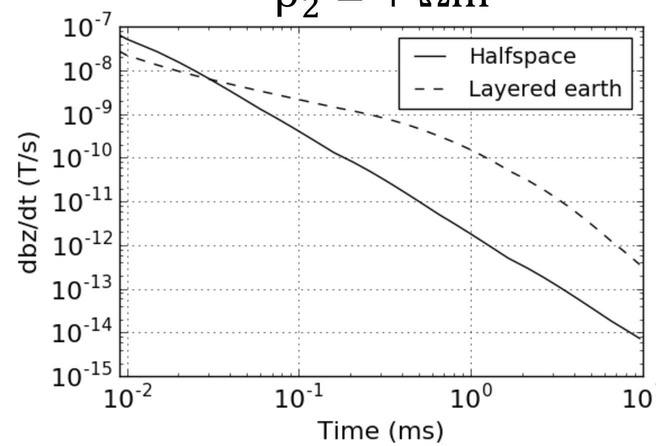


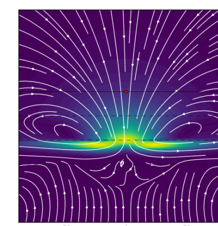
$\rho_2 = 10\ \Omega m$





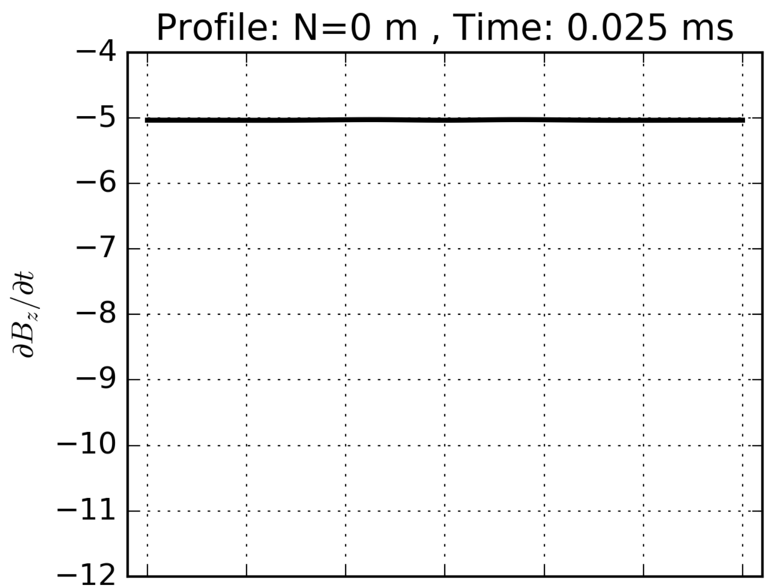
$\rho_2 = 1\ \Omega m$



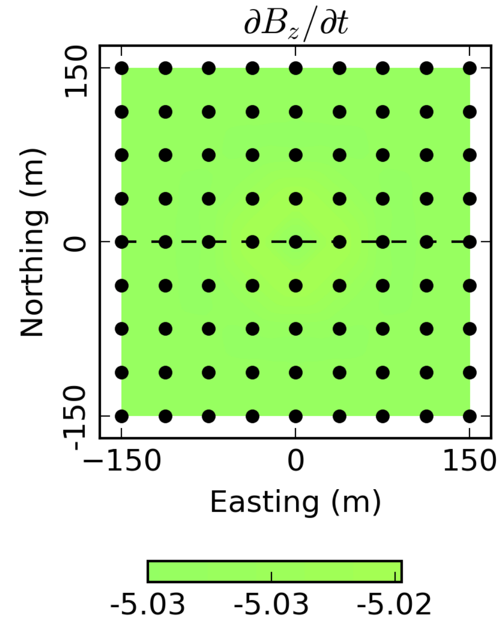


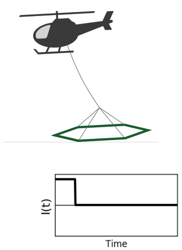
Airborne example: conductive sphere

Data profile

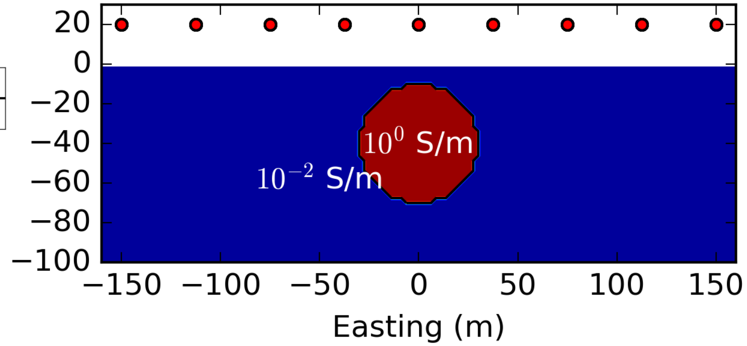


Data map



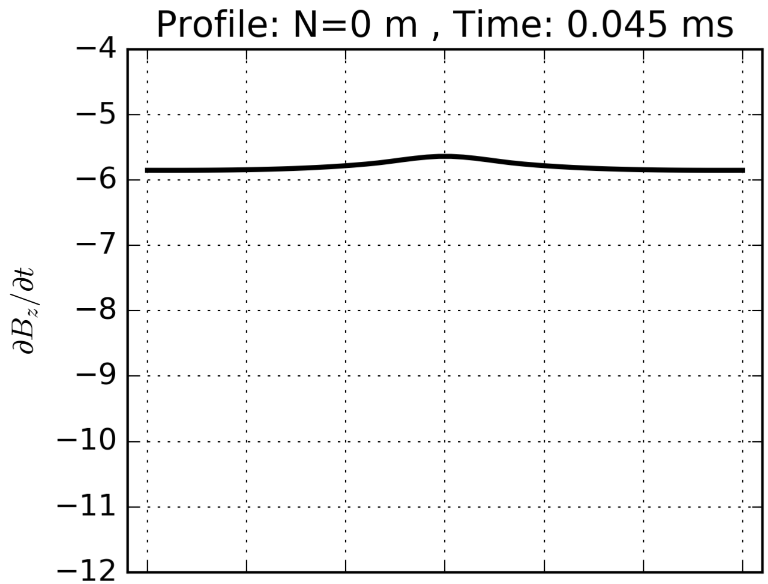


Conductivity

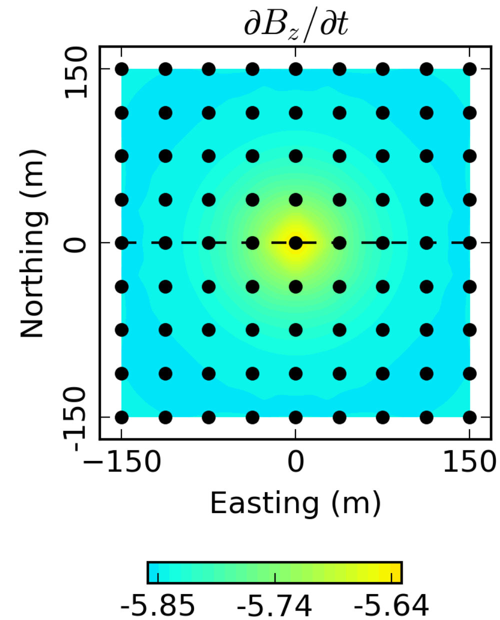


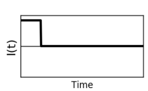
Airborne example: conductive sphere

Data profile

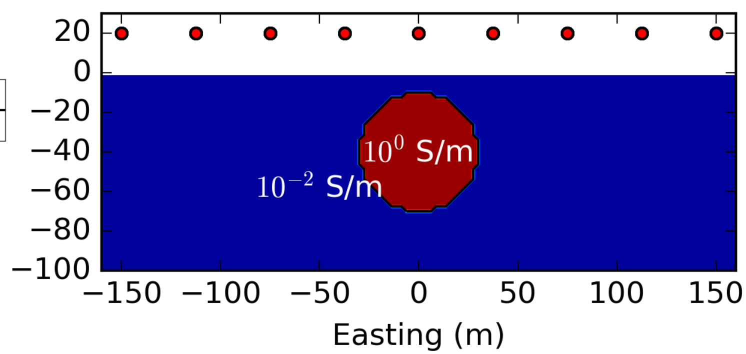


Data map



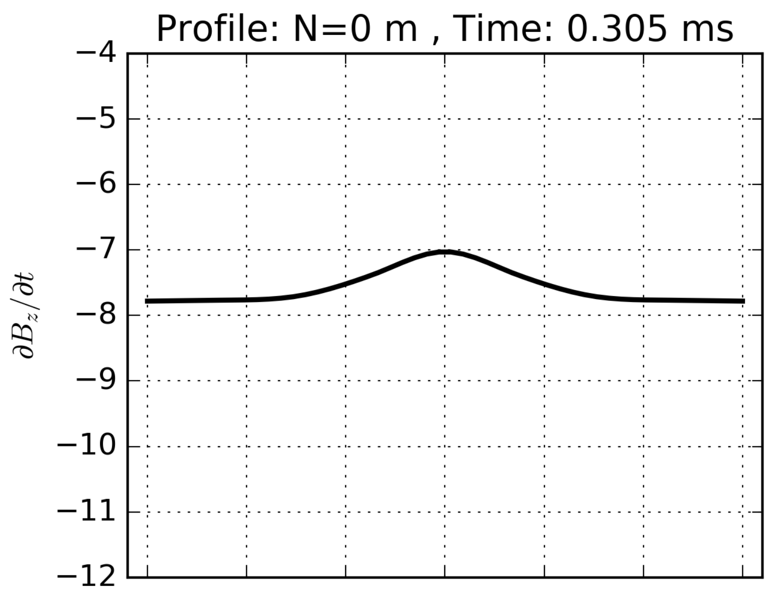


Conductivity

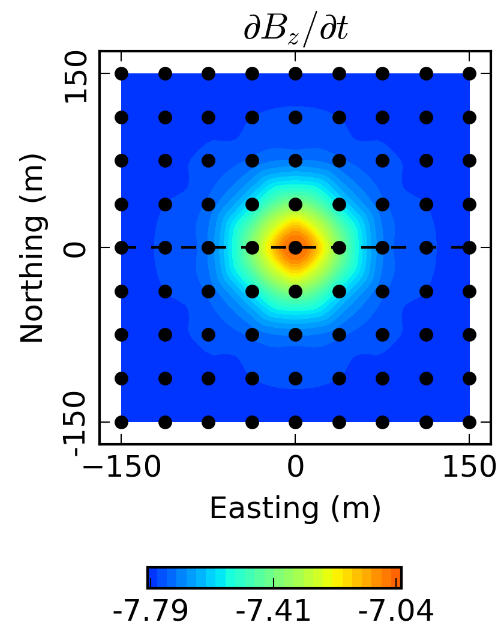


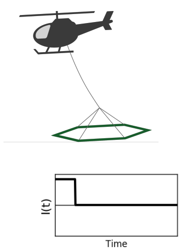
Airborne example: conductive sphere

Data profile

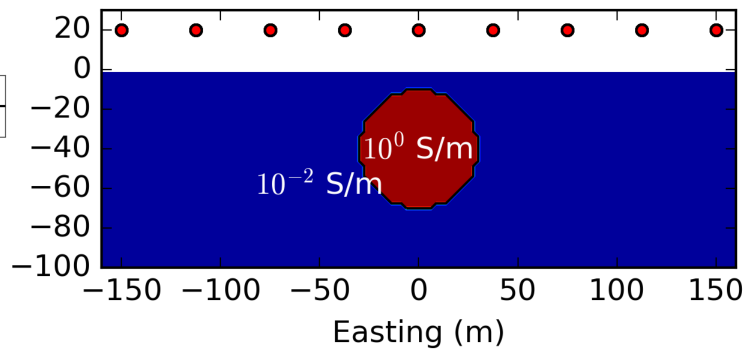


Data map



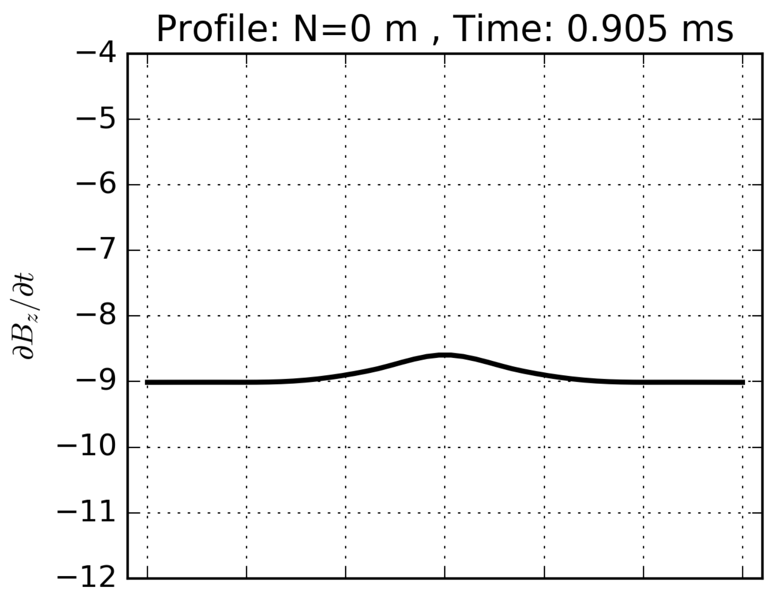


Conductivity

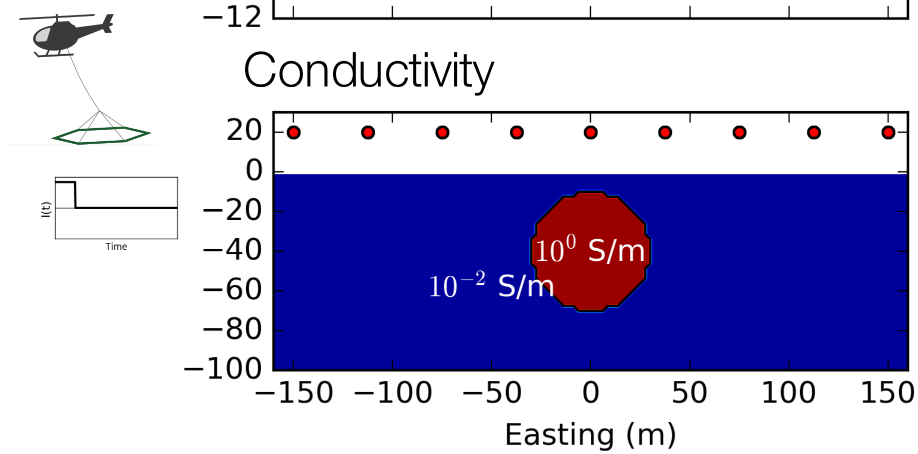


Airborne example: conductive sphere

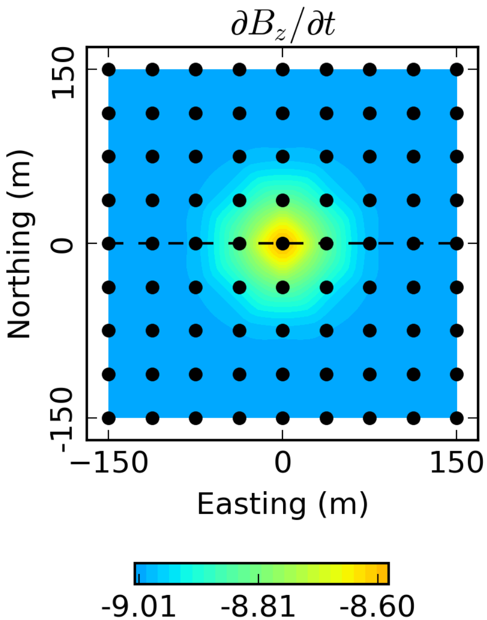
Data profile



Conductivity

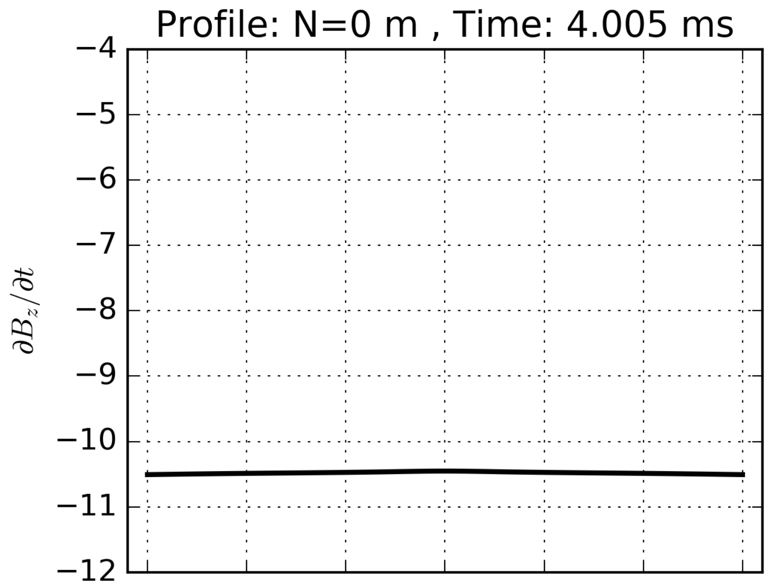


Data map

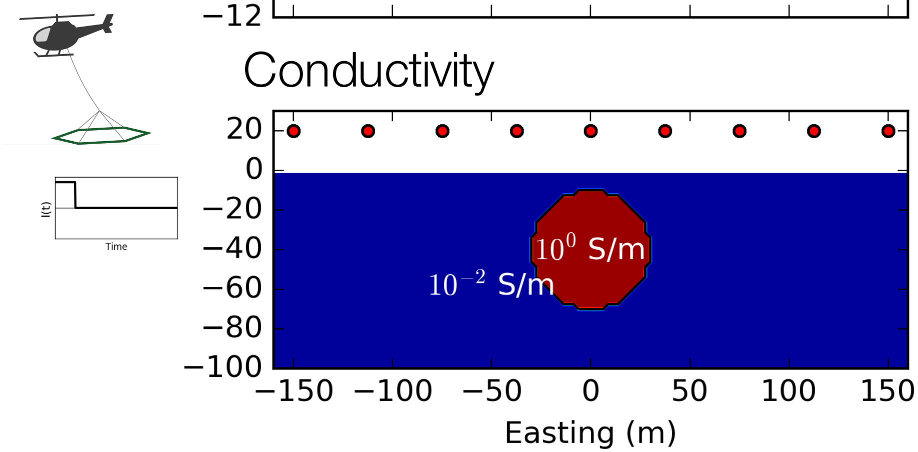


Airborne example: conductive sphere

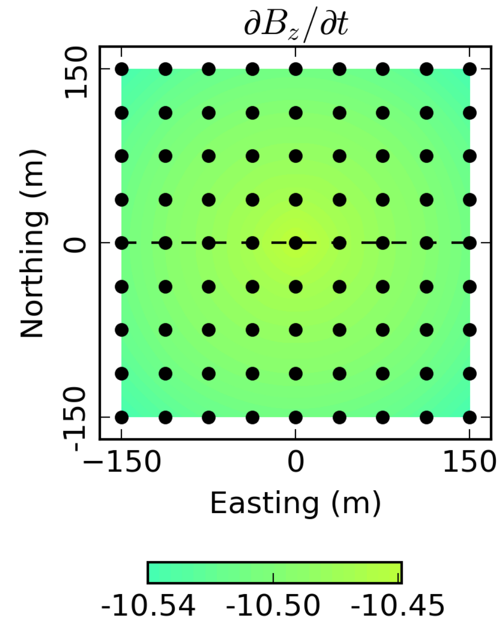
Data profile



Conductivity

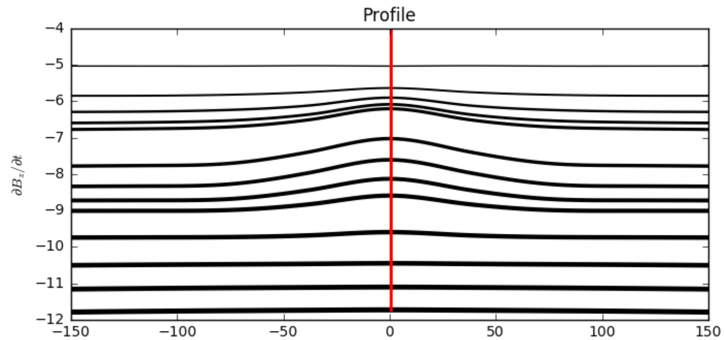


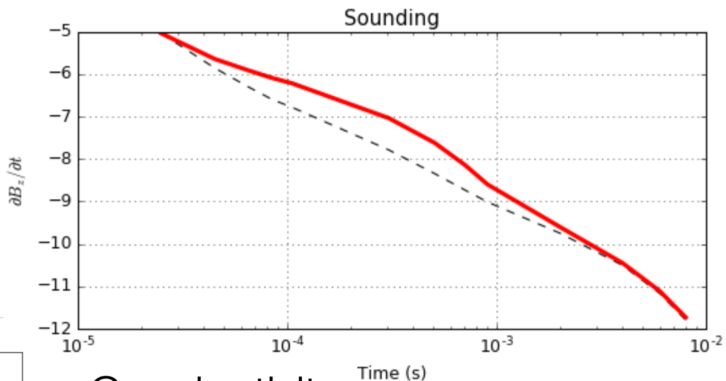
Data map

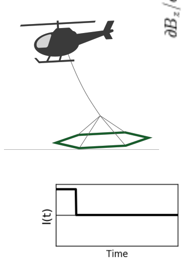


Summary: airborne example

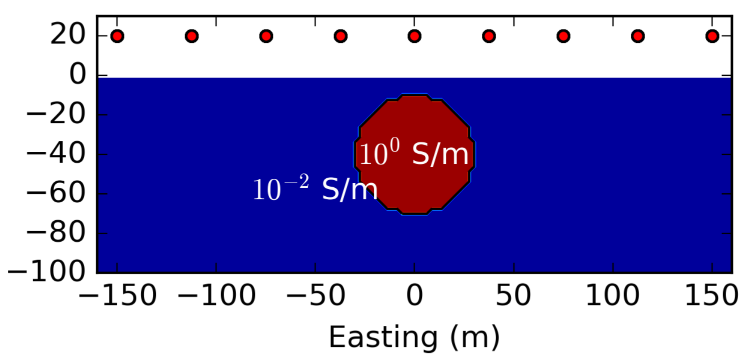
Data profile



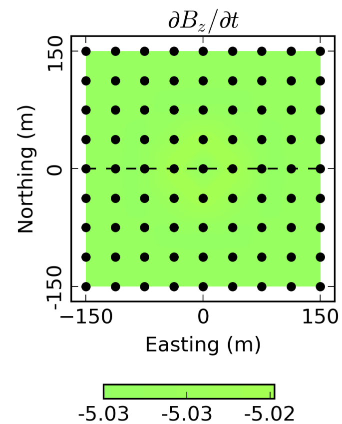




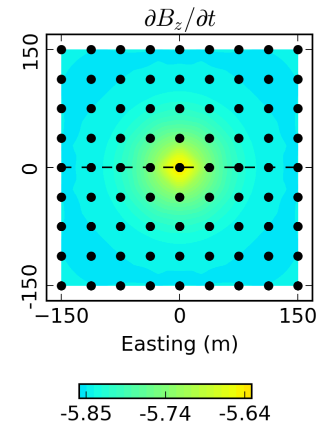
Conductivity



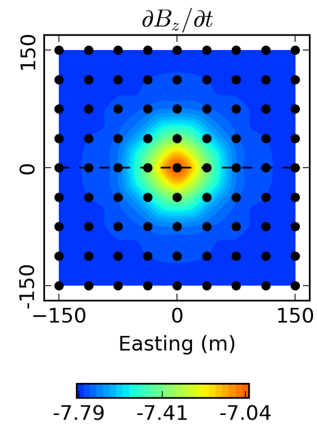
0.025 ms



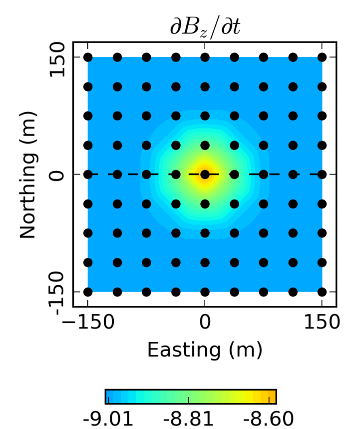
0.045 ms



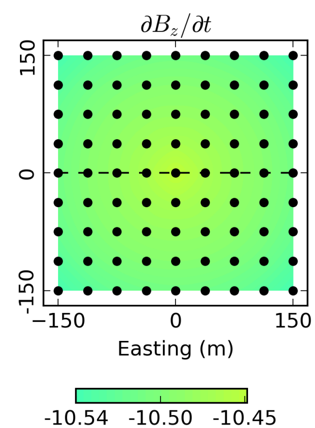
0.305 ms



0.905 ms

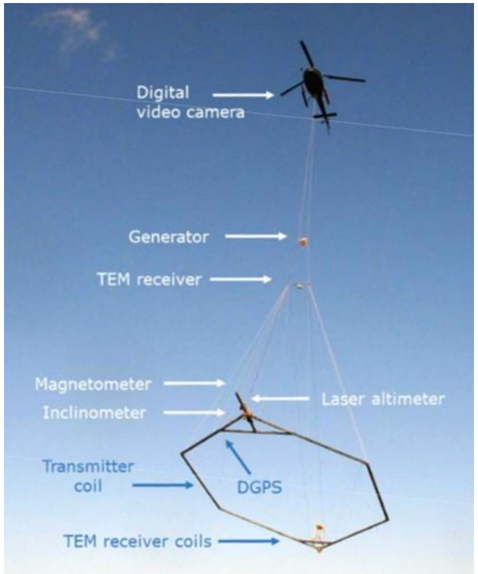


4.005 ms



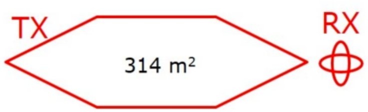
Some Airborne TDEM Systems

SkyTEM (2006)

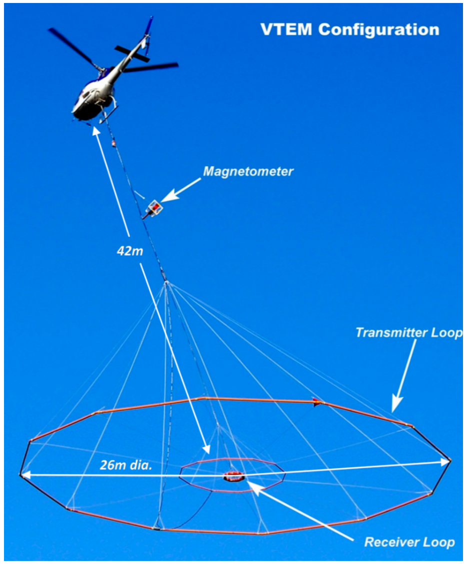


Area = 314 m^2

Peak dipole moment:
- HM: 113040 NIA
- LM: 12560 NIA

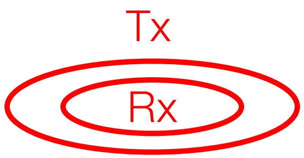


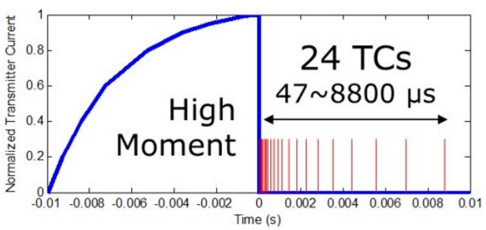
VTEM (2007)



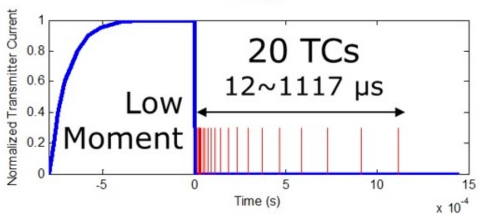
Area = 535 m^2

Peak dipole moment:
- 503,100 NIA

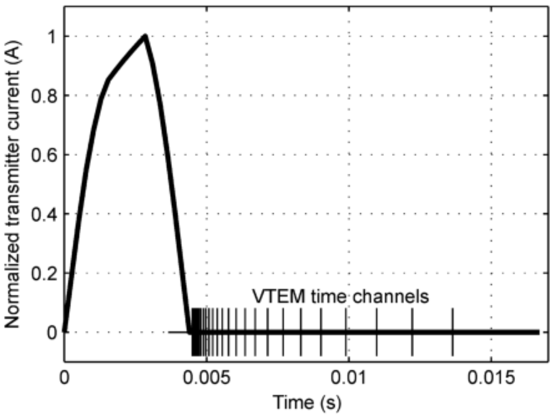




Peak current: 90 A
Turns: 4
On-time: 10 ms
Off-time: 10 ms



Peak current: 40 A
Turns: 1
On-time: 0.8 ms
Off-time: 1.45 ms



Peak current: 235 A
Turns: 4
On-time: 4.5 ms
Off-time: 9.1 ms

Outline

Setup

Time Domain EM

- Vertical Magnetic Dipole
- Propagation with Time
- Effects of Background Conductivity
- Transmitters and receivers
- Decay Curves
- Case History: Ground water

Frequency Domain EM

Questions

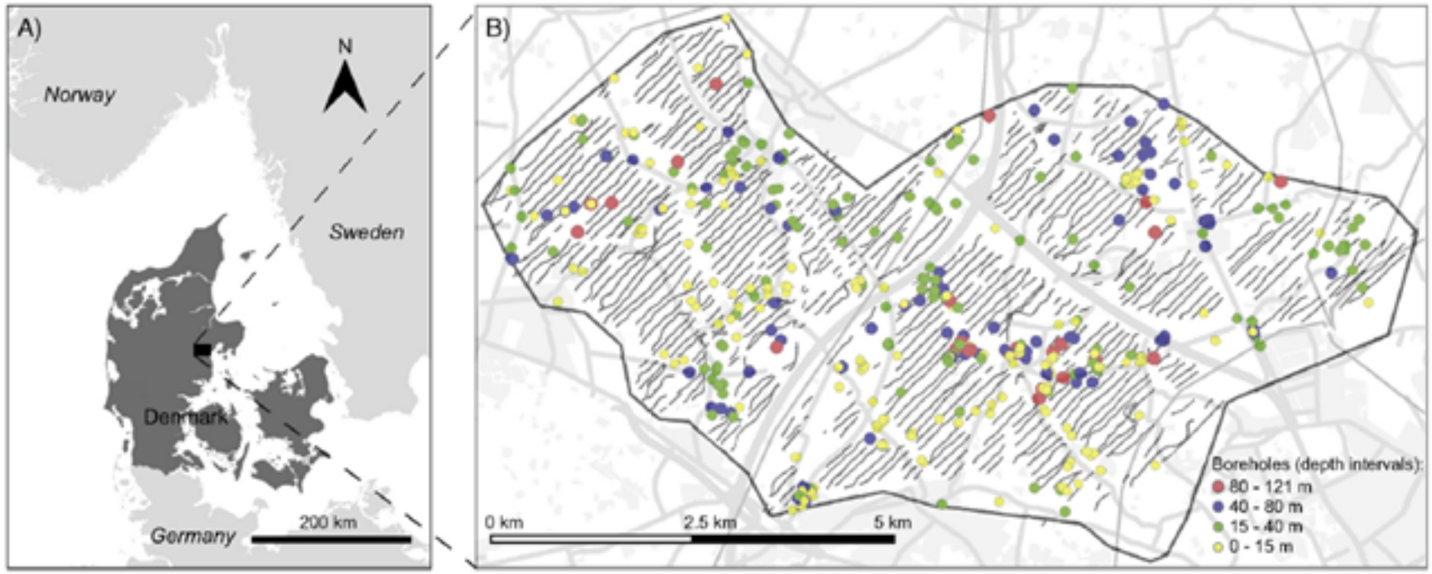
Case History: Kasted

Vilhelmsen et al. (2016)

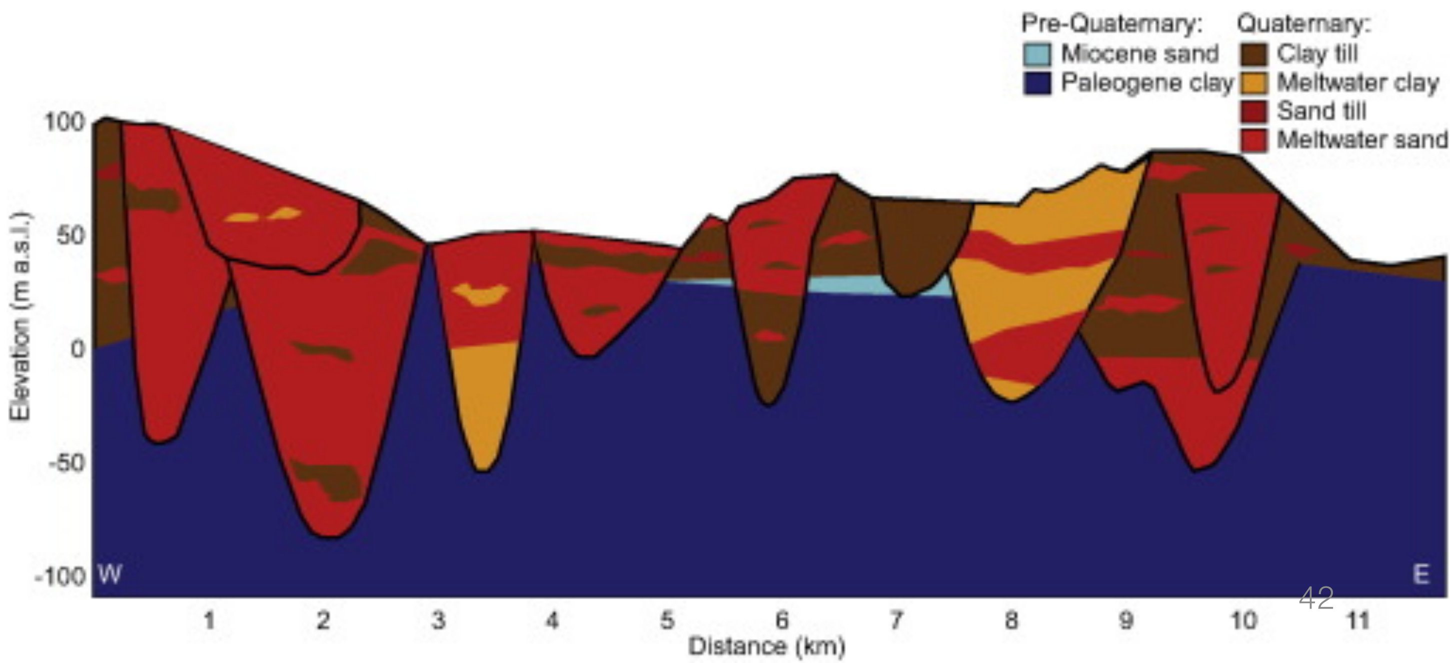
Setup

A) Survey Area: Kasted, Demark

B) Borehole locations

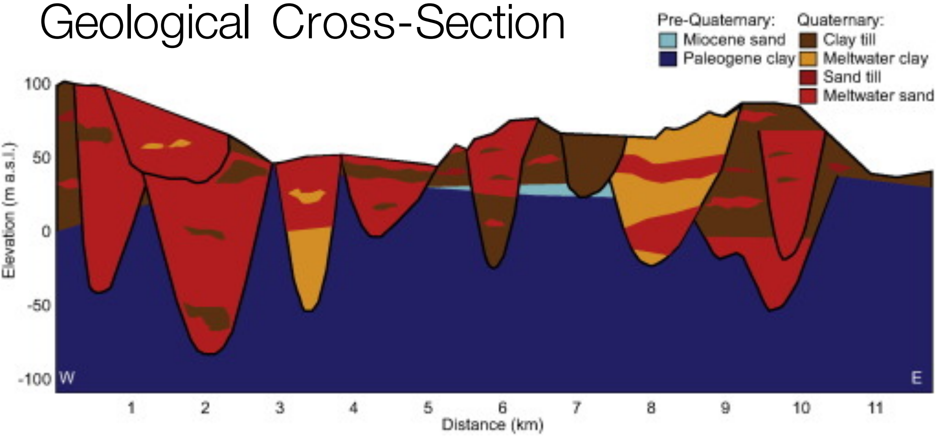


Local Geology: W-E cross-section



Properties

Geological Cross-Section

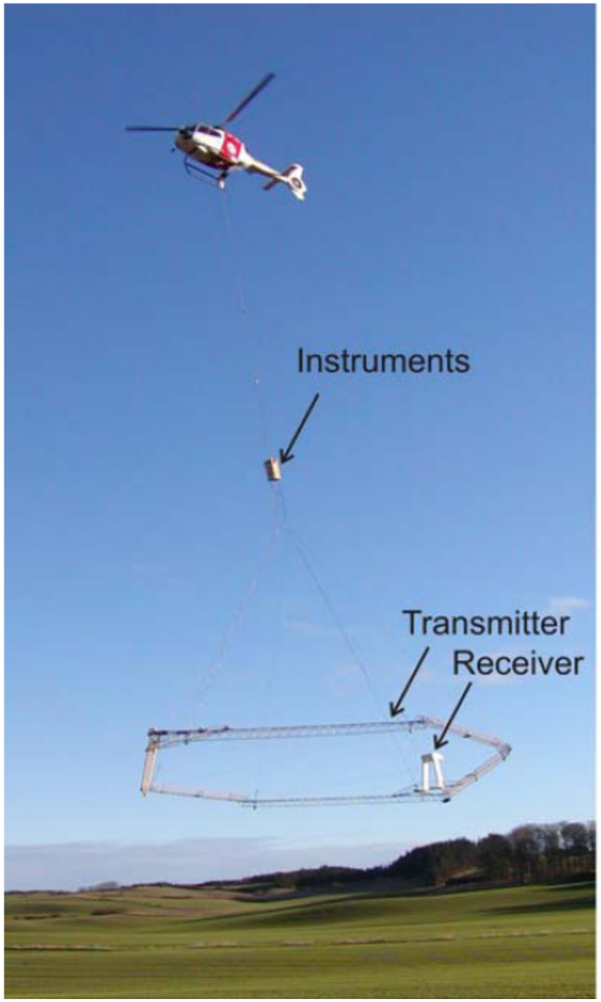


Geological Units	Resistivity (Ωm)
Palaeogene Clay	1-10
Clay Till	25-60
Sand Till	>50
Meltwater Sand and Gravel	>60
Glaciolacustrine Clay	10-40
Miocene Silt and Sand	>40
Miocene Clay	10-40
Sand	>40
Clay	1-60

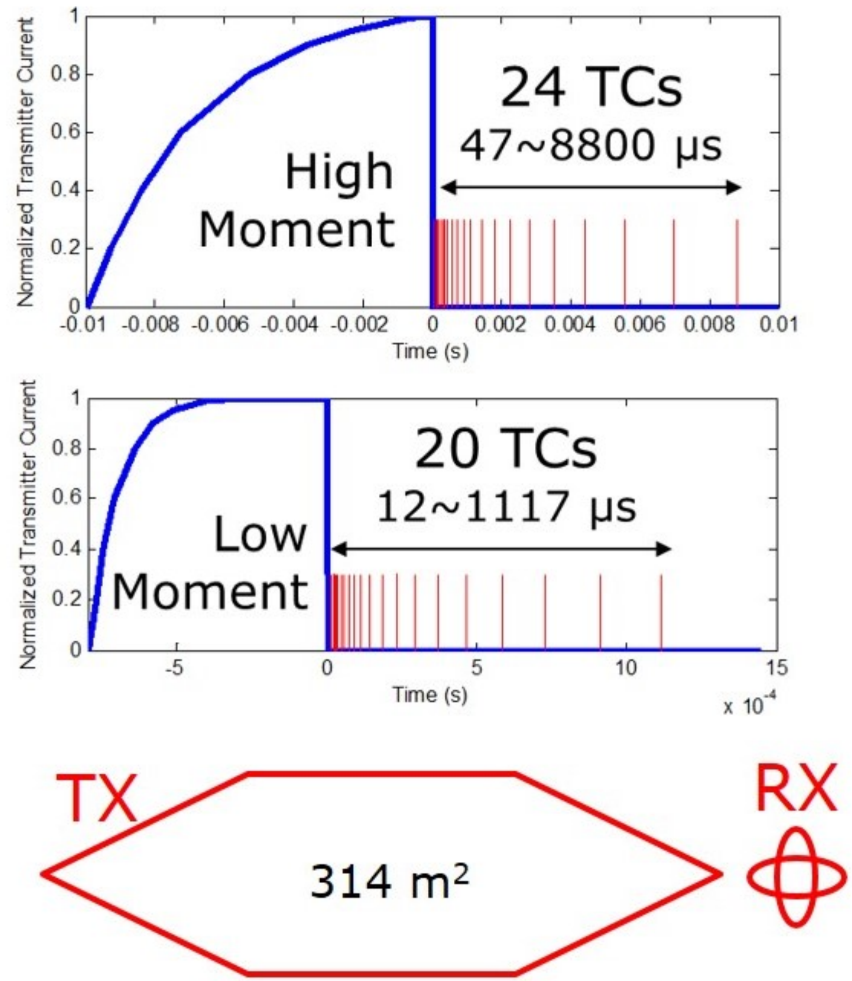
- Buried valleys with clays beneath
- Infill (water-bearing): coarse sand and gravel
- Clays are conductive (1-40 Ωm)
- Water-bearing sands and gravels are more resistive (>40 Ωm)

Survey

SkyTEM System



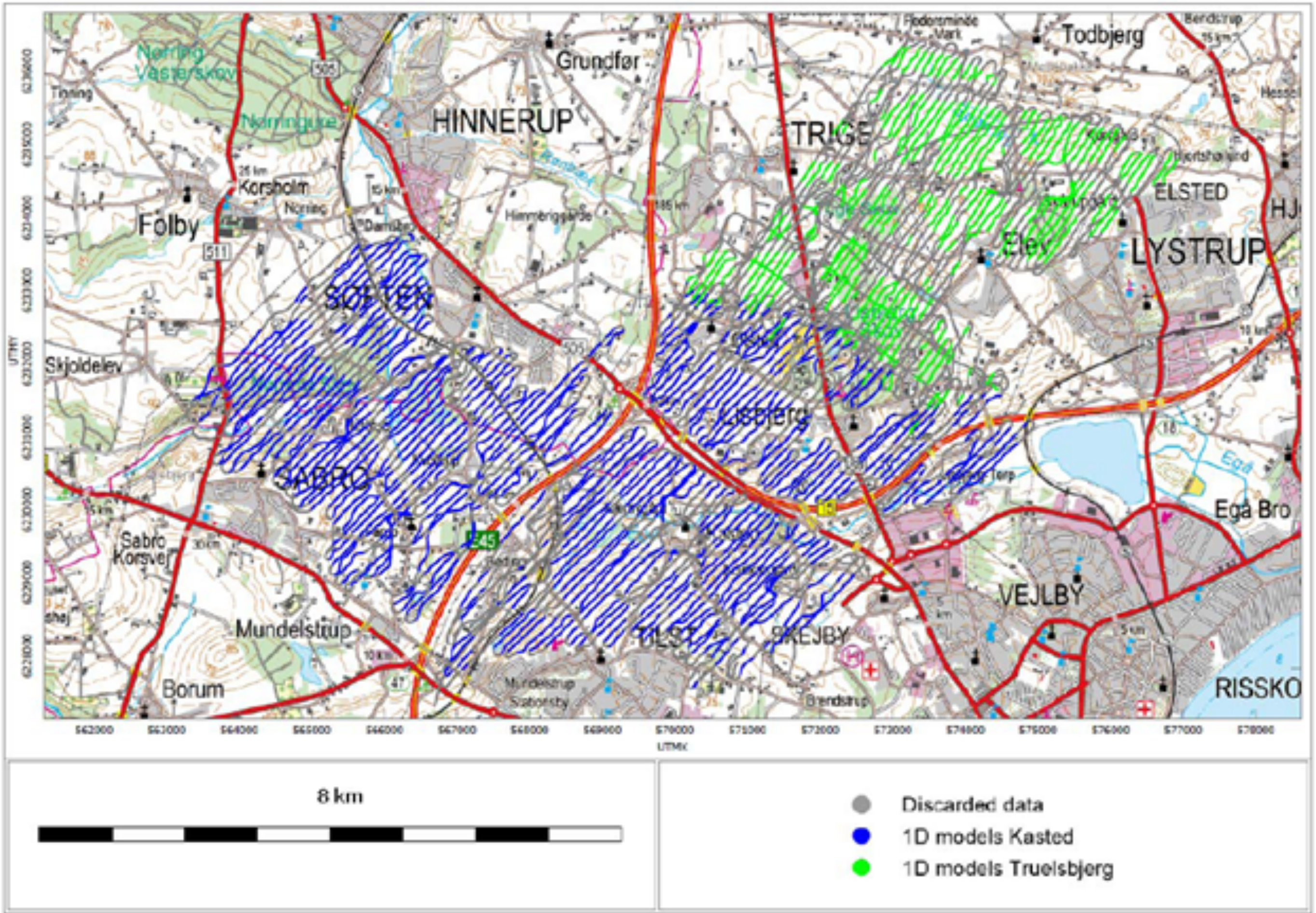
System Configuration



- Low moment (LM) used to image near surface structures
- High moment (HM) used to image deeper structures

Data

Blue: data used for Kasted study

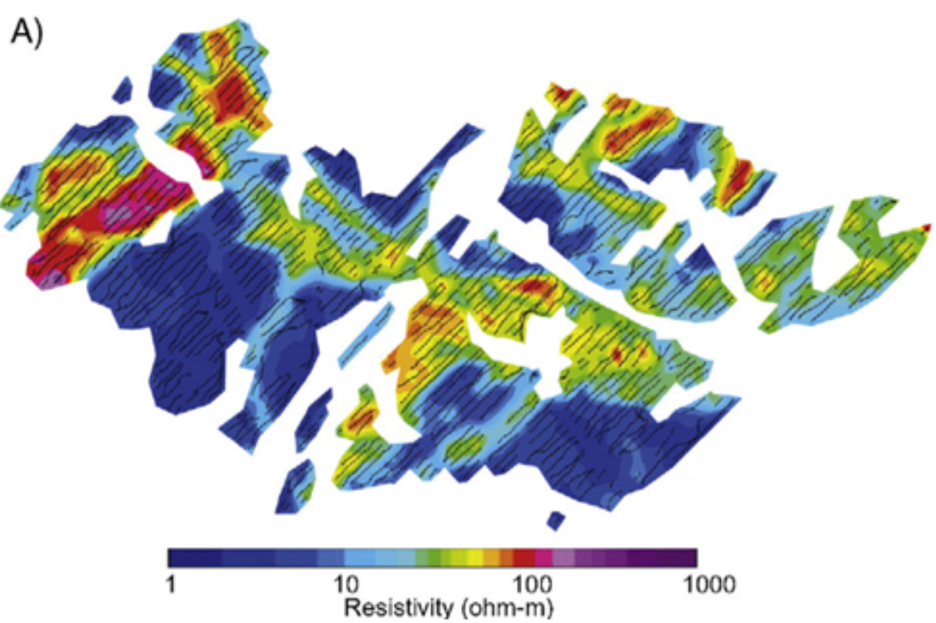


- 333 line km of data, 100 m line-spacing
- Data points with strong coupling to cultural noise were removed (~30%)

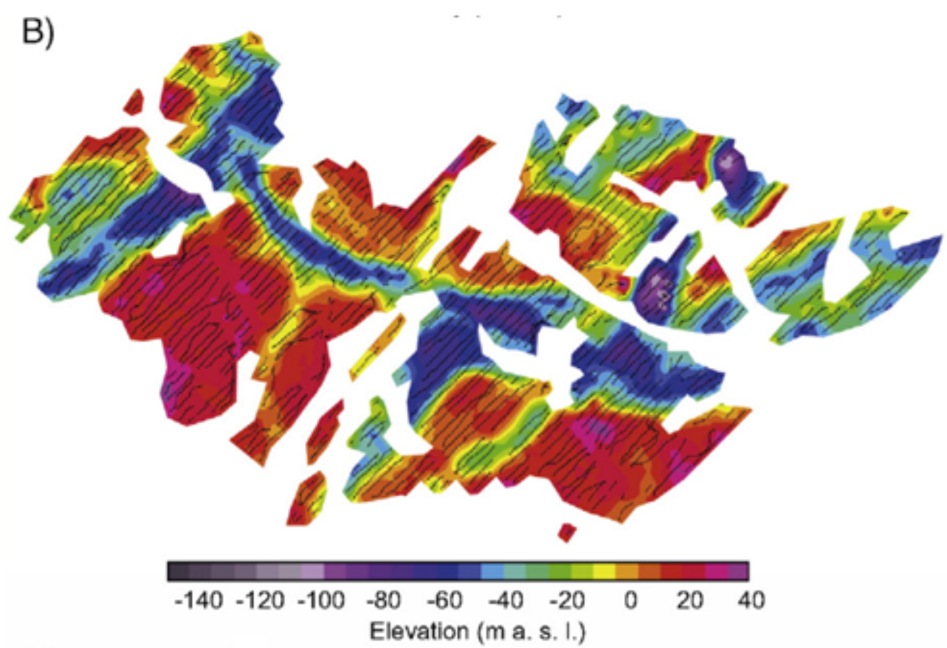
Processing (inversion)

- Spatially constrained 1D inversion → quasi-3D approach
- 9,500 soundings were inverted using 25 layers

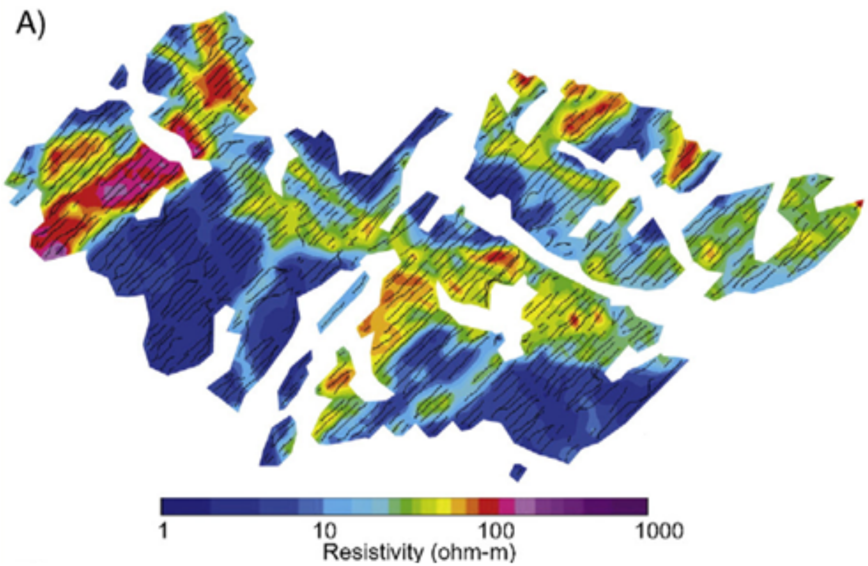
Depth slice 5 m above sea-level

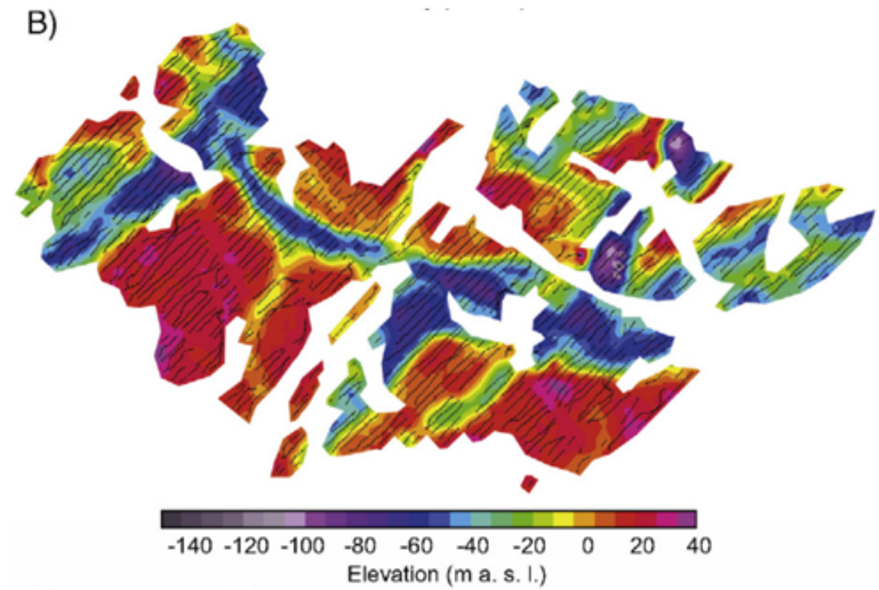


Approximate depth to the top of Paleogene clay layer

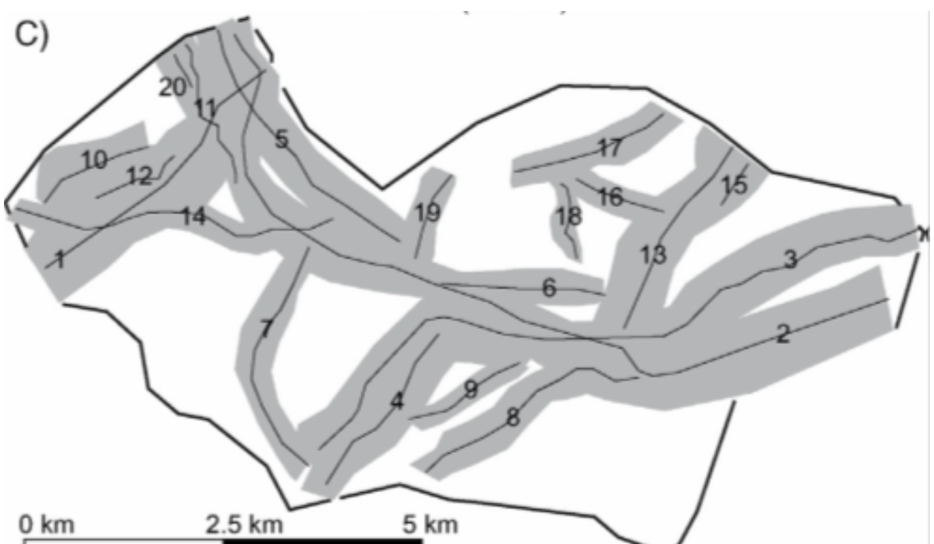


Interpretation



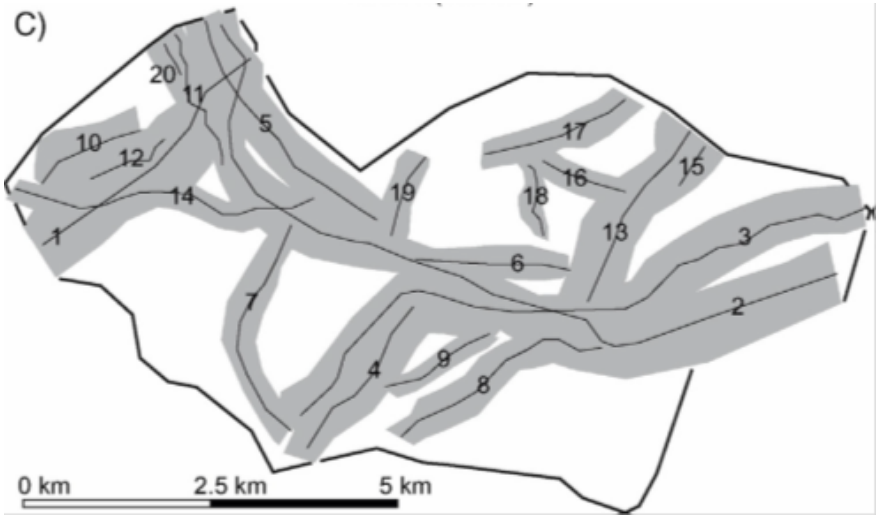


Delineation of valley structures

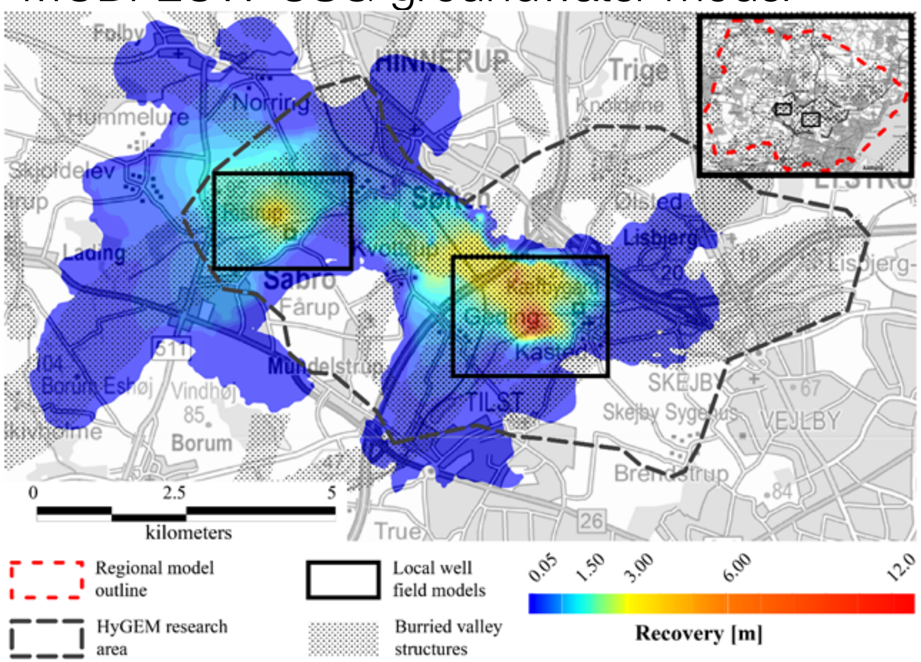


- Inversion results used to construct geological model.
- Delineated 20 buried and cross-cutting valley structures.

Synthesis



MODFLOW-USG groundwater model



- 3D geologic model incorporated into MODFLOW-USG groundwater modeling tool
- Extracted water from 2 wells.
- Downdraw between the two wells correlated with the resistive valley structures

Outline

Setup

- Basic experiment
- Transmitters, Receivers

Time Domain EM

- Vertical Magnetic Dipole
- Propagation with Time
- Case History

Frequency Domain EM

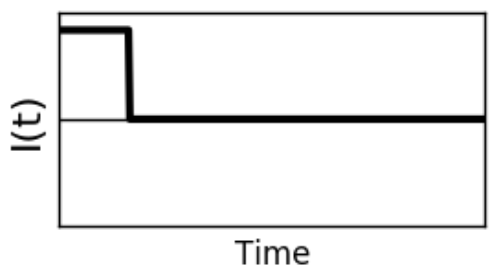
- Vertical Magnetic Dipole
- Effects of Frequency
- Case History – Ground water

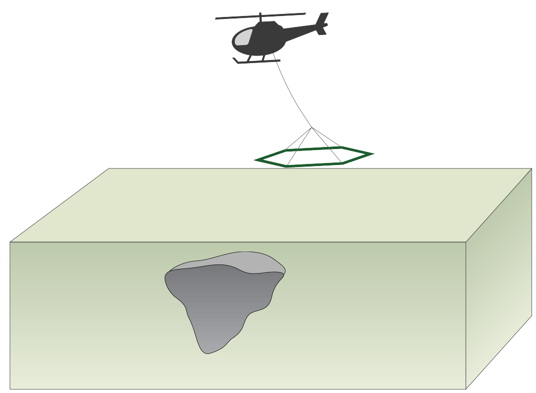
EM with Inductive Sources

- Induction principles are the same for
  - TDEM: Time domain EM
  - FDEM: Frequency domain EM

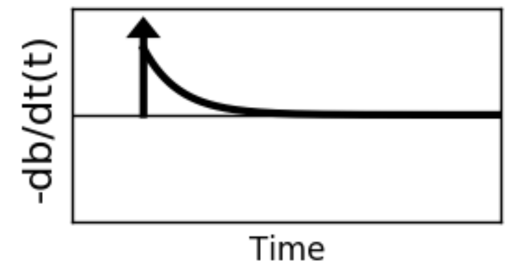
Transmitter current

TDEM

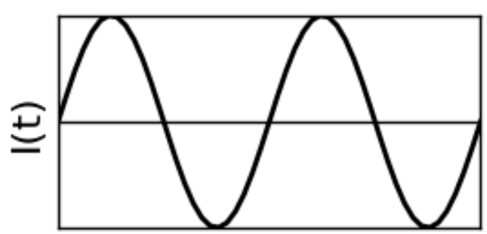


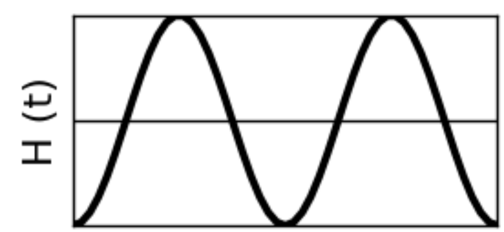


Receiver

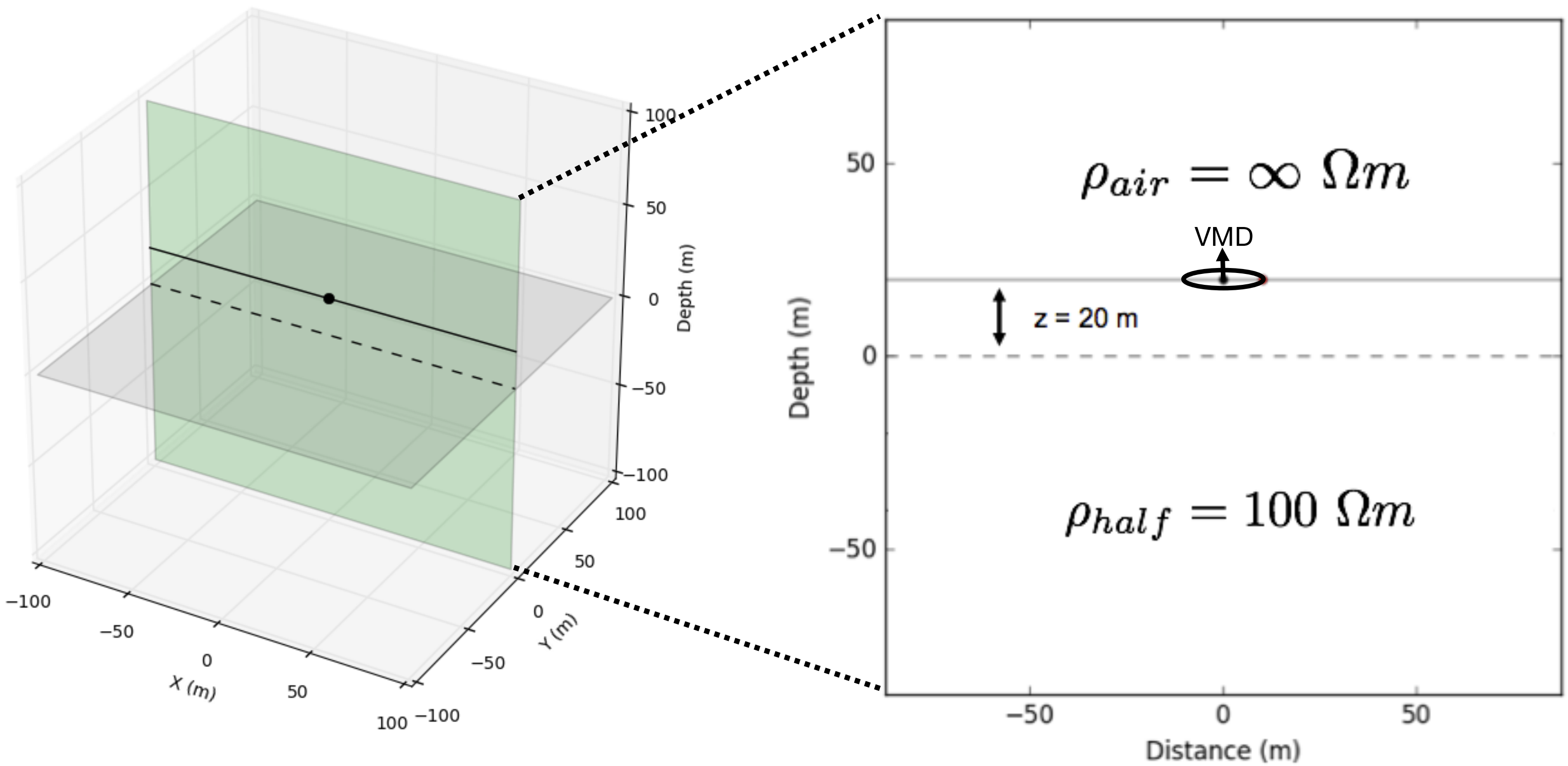


FDEM



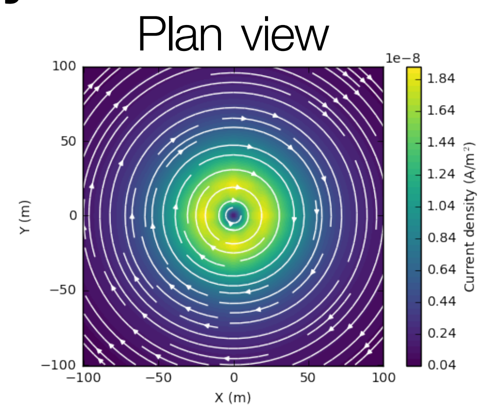


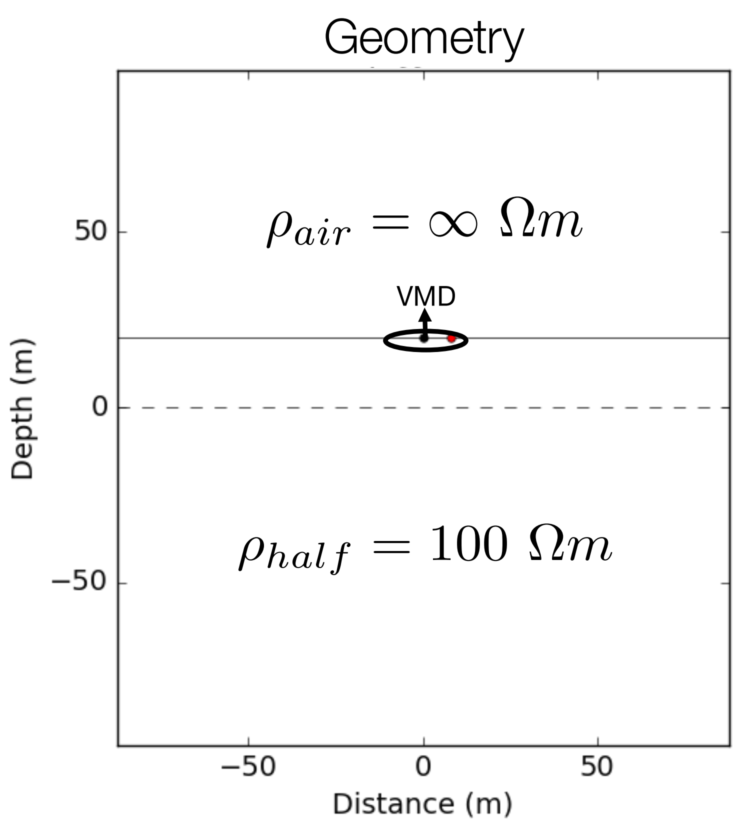
Vertical Magnetic Dipole over a halfspace (FDEM)

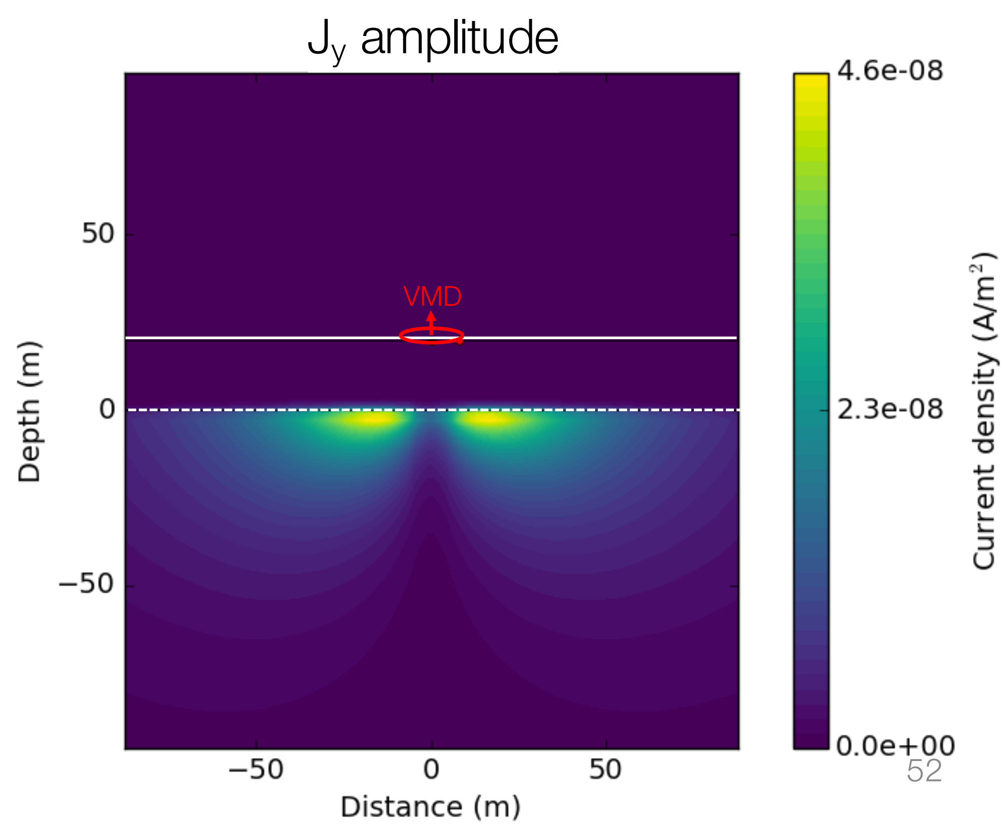


Current Density

- Frequency = 10 kHz
- Currents in the earth flow in planes parallel to the Tx

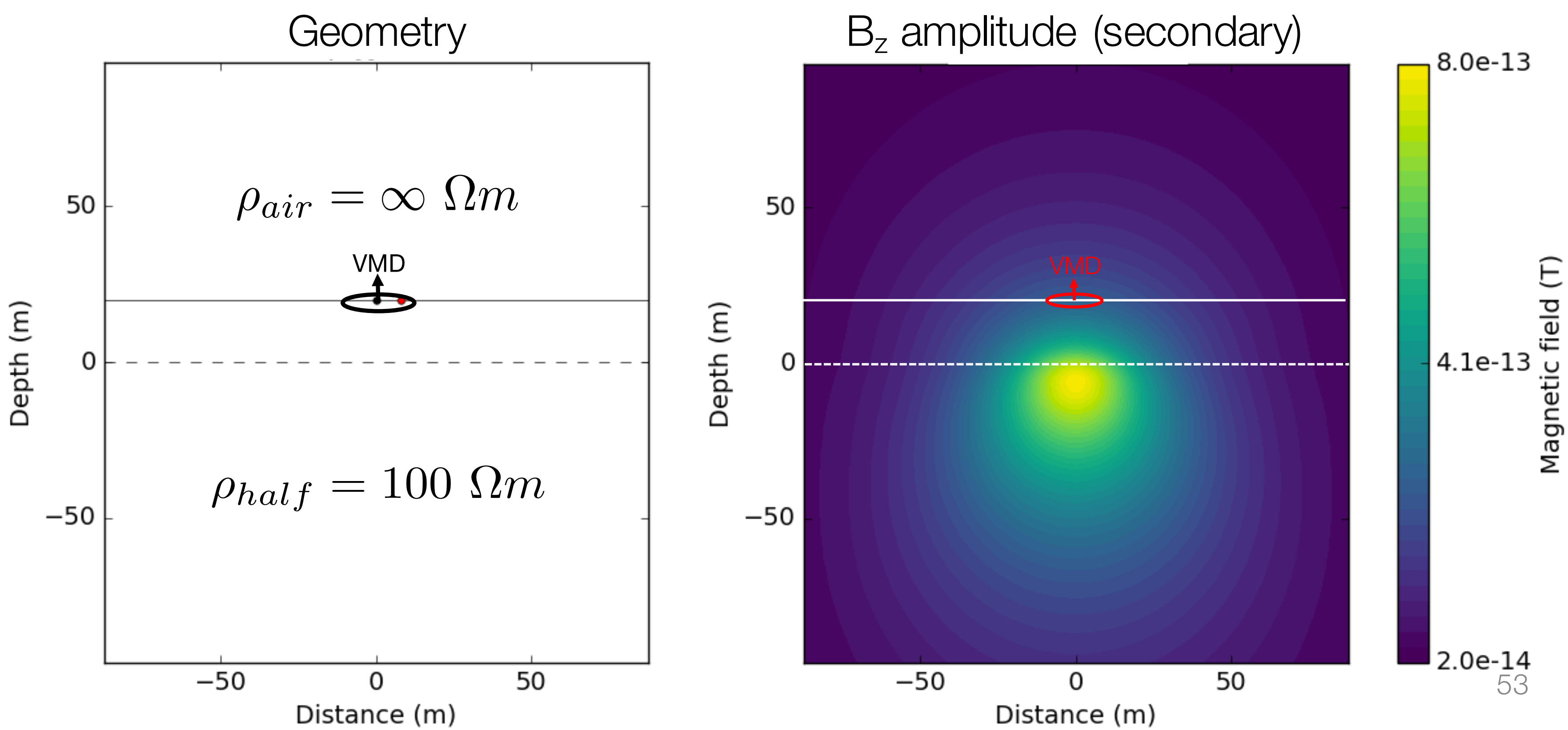






Secondary Magnetic Flux Density

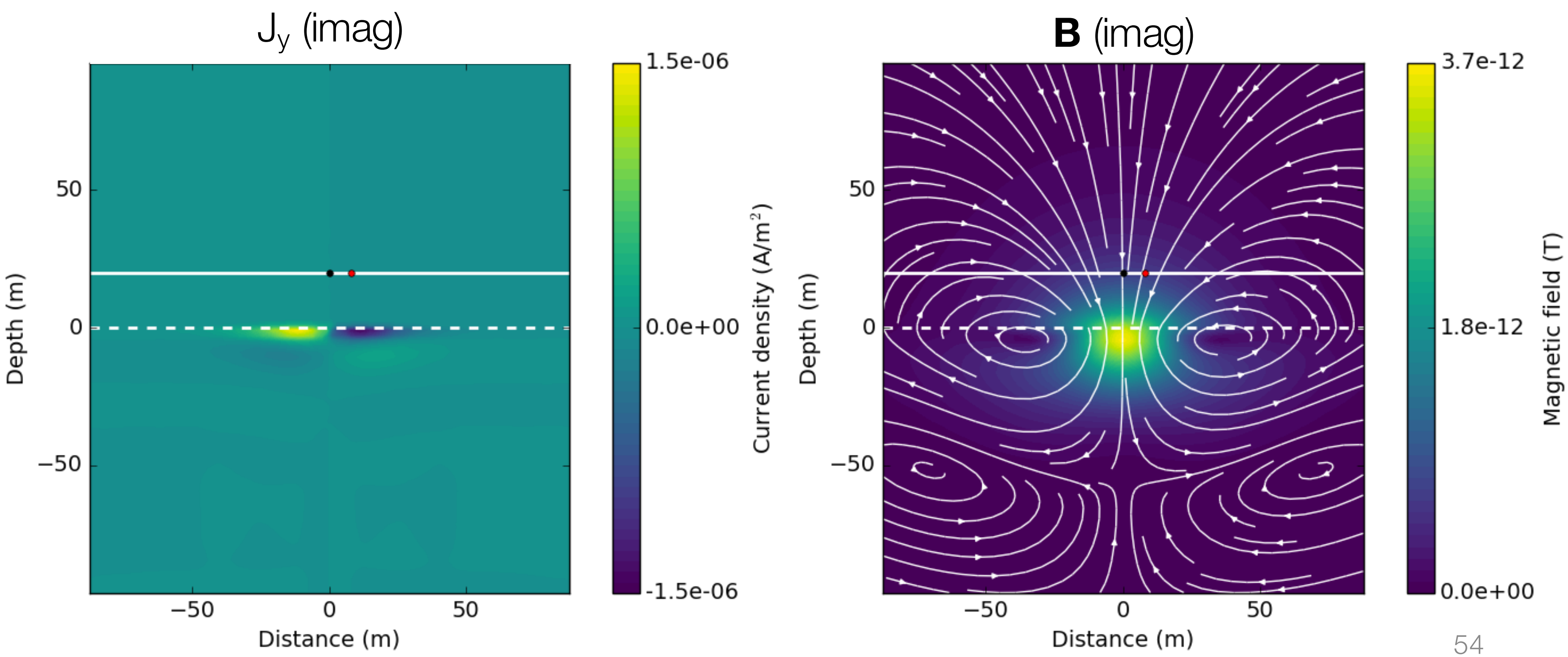
- Frequency = 10 kHz



Effects of Frequency

- Frequency at 100 kHz
- Skin depth = 16 m
- Currents are concentrated at surface

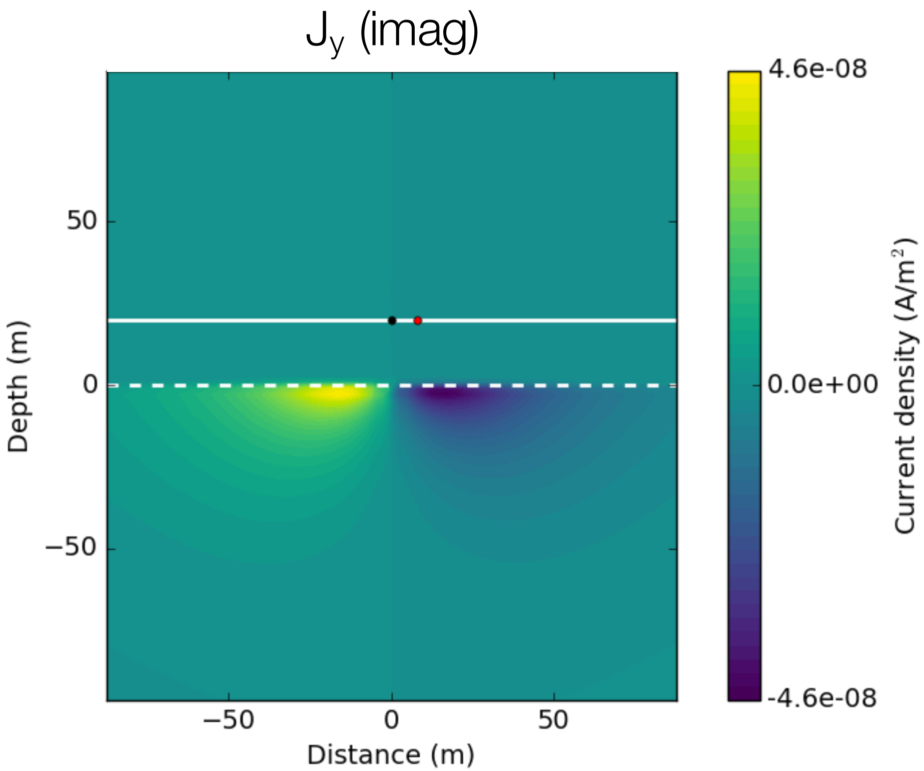
$$\delta = 503\sqrt{\frac{\rho}{f}}$$

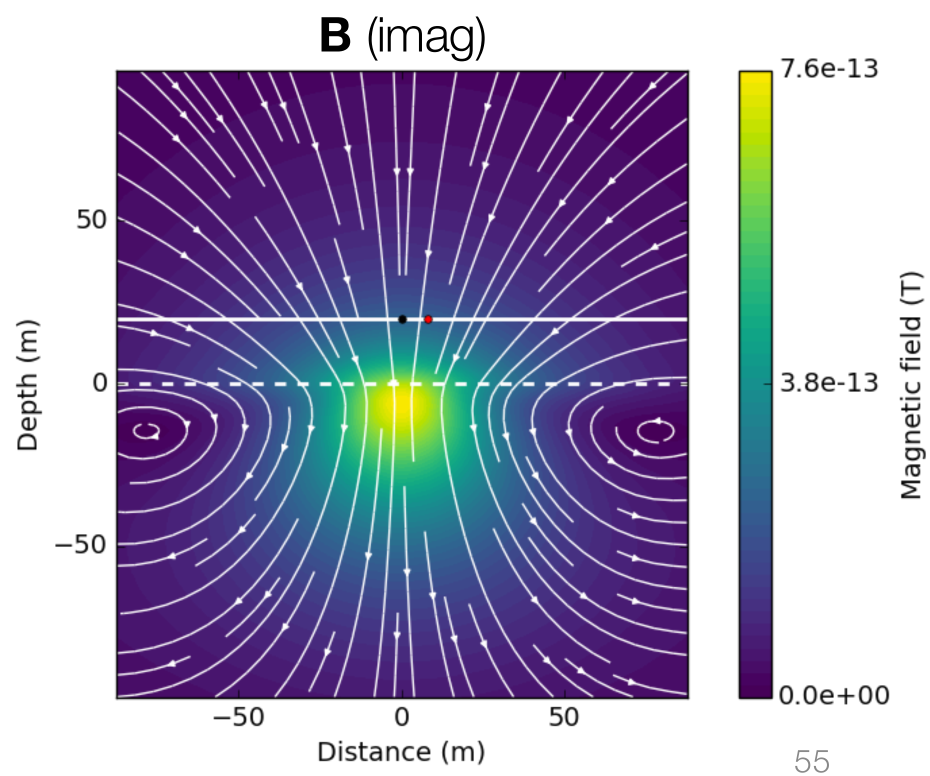


Effects of Frequency

- Frequency at 10 kHz
- Skin depth = 50 m
- Currents diffusing downward and outward

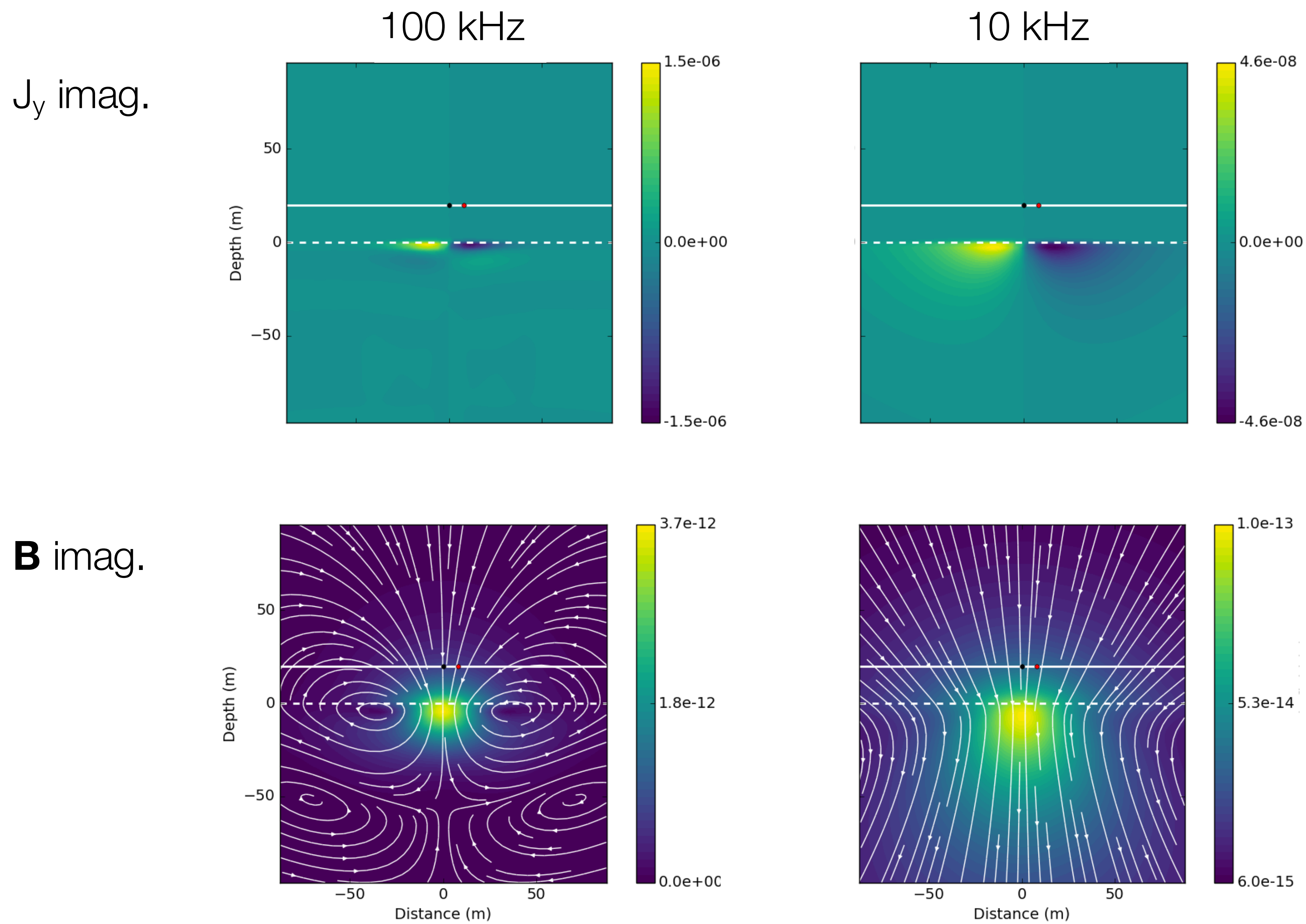
$$\delta = 503\sqrt{\frac{\rho}{f}}$$





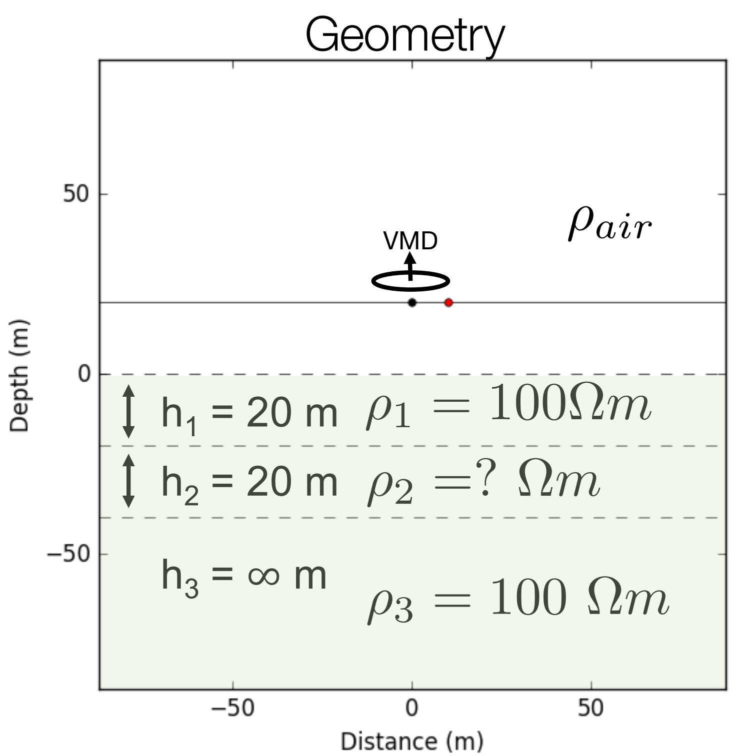
Summary: Effects of Frequency

$$\delta = 503\sqrt{\frac{\rho}{f}}$$



Layered earth

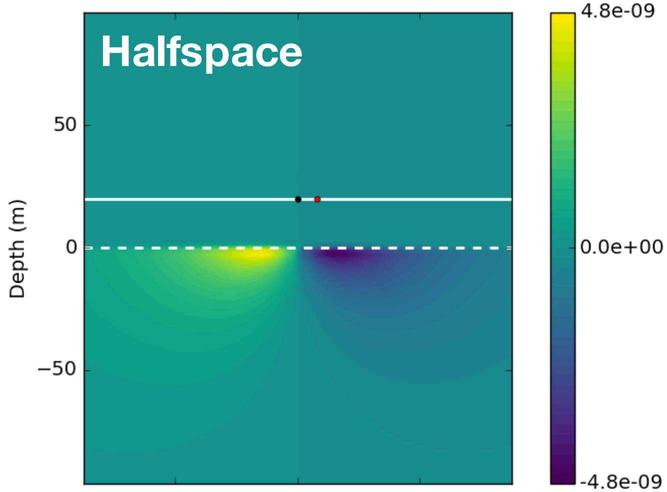
- 3 layers + air,
- ρ_2 varies



- Four different cases:
  - Halfspace
    $\rho_2 = 100\ \Omega m$
  - Resistive
    $\rho_2 = 1000\ \Omega m$
  - Conductive
    $\rho_2 = 10\ \Omega m$
  - Very conductive
    $\rho_2 = 1\ \Omega m$
- Fields
  - J_y imag
  - Secondary **B** imag

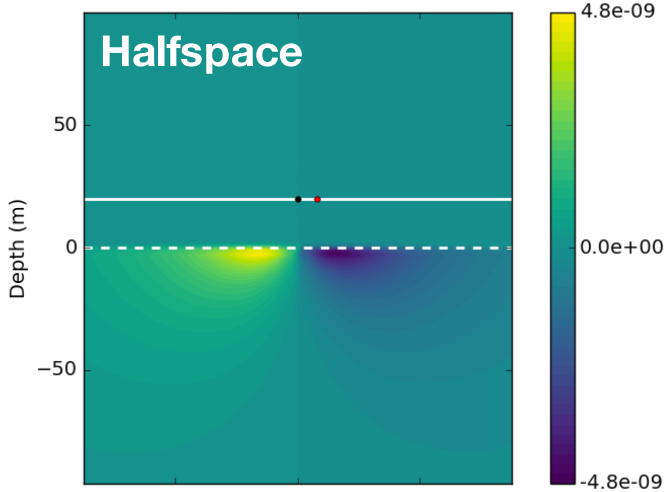
Current density (J_y imag)

$\rho_2 = 100\ \Omega m$

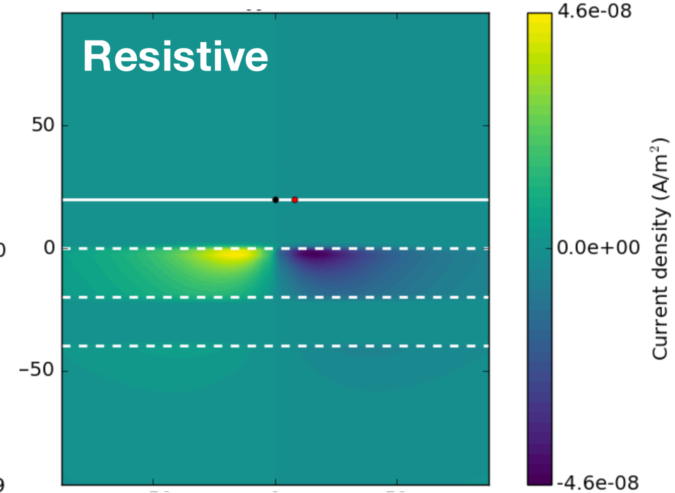


Current density (J_y imag)

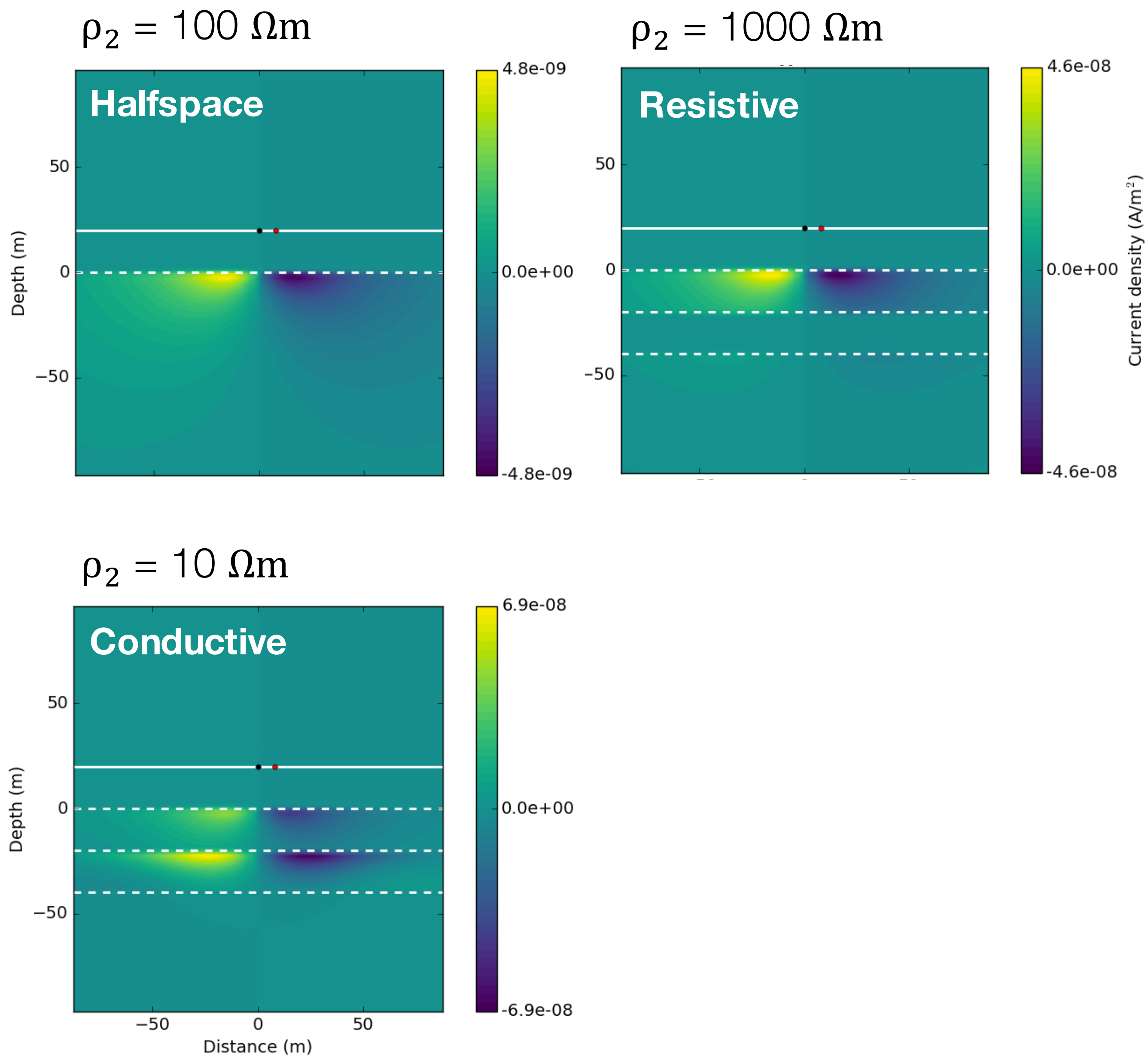
ρ_2 = 100 Ωm



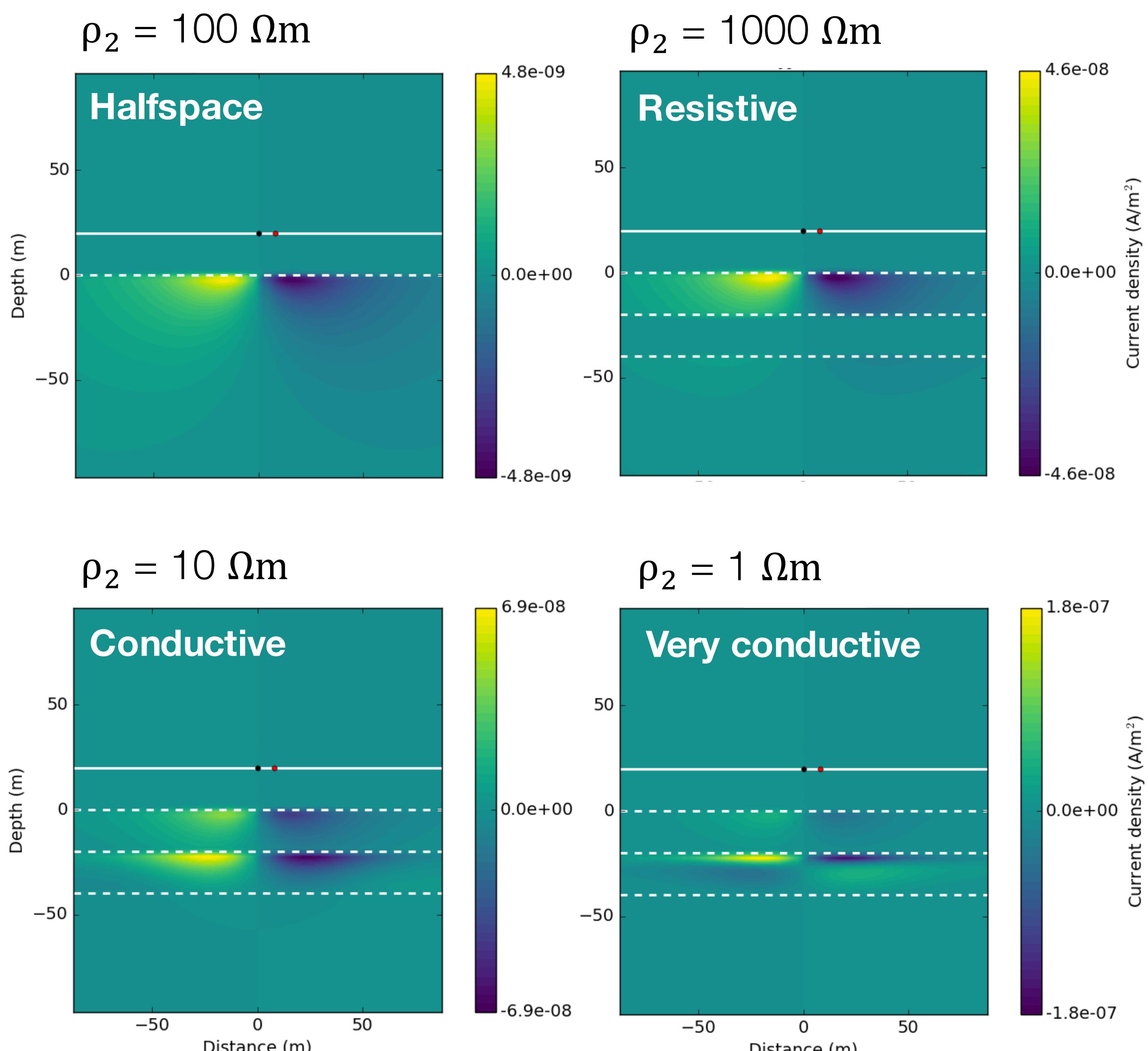
ρ_2 = 1000 Ωm



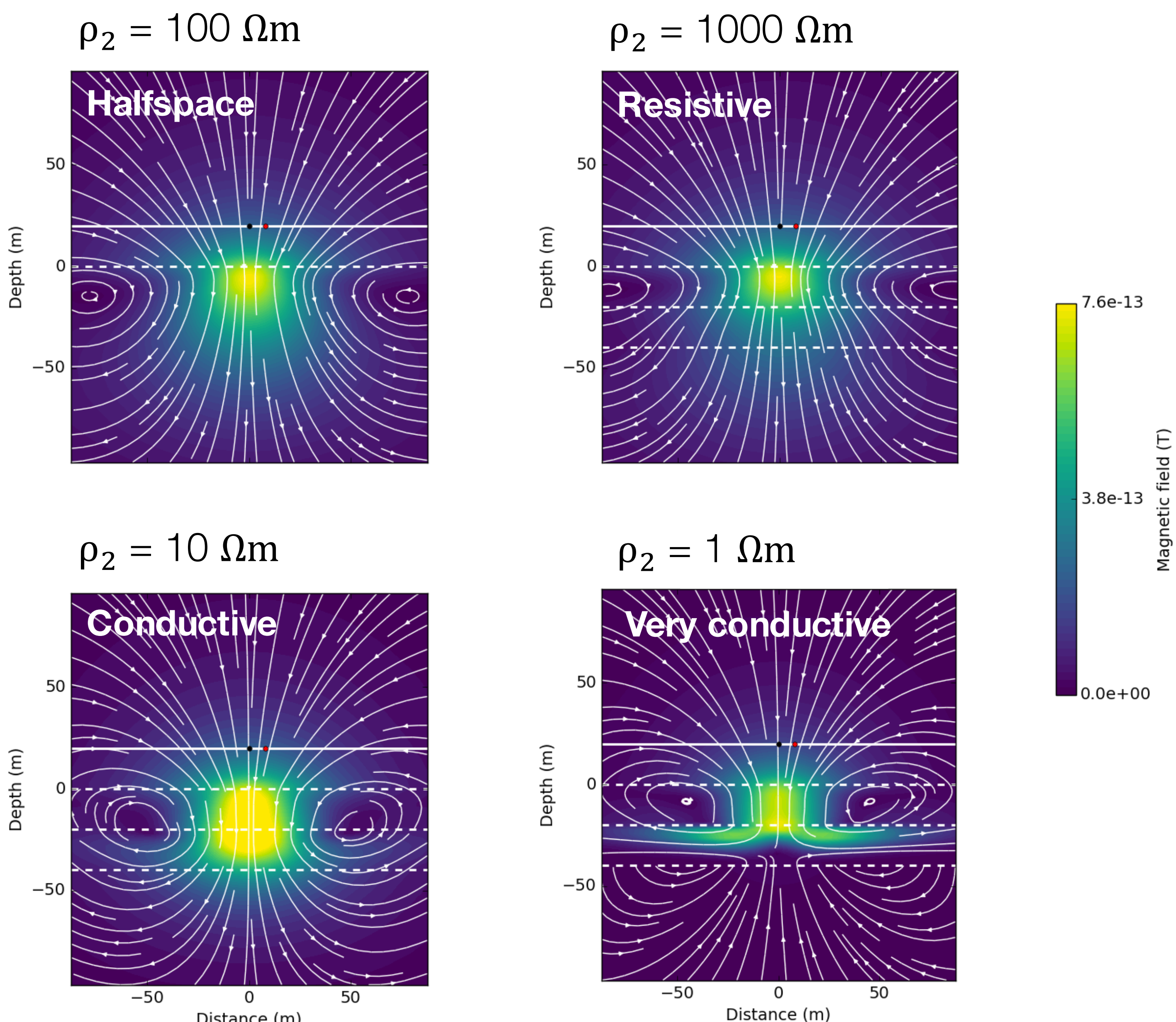
Current density (J_y imag)



Current density (J_y imag)

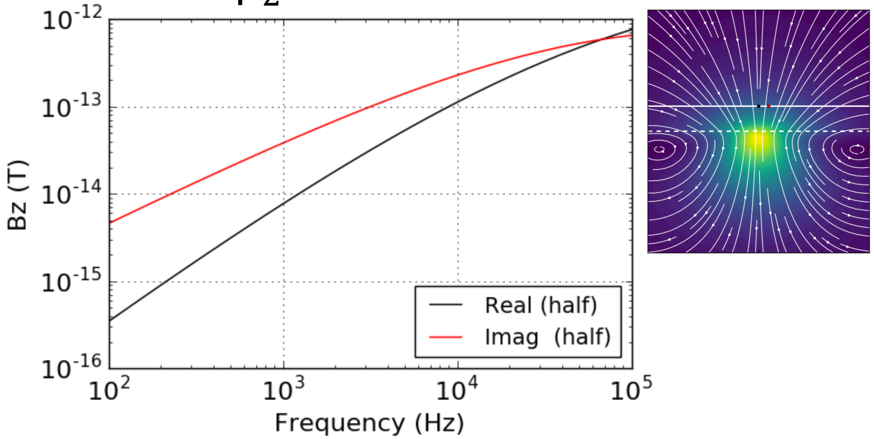


Magnetic flux density (**B** imag)

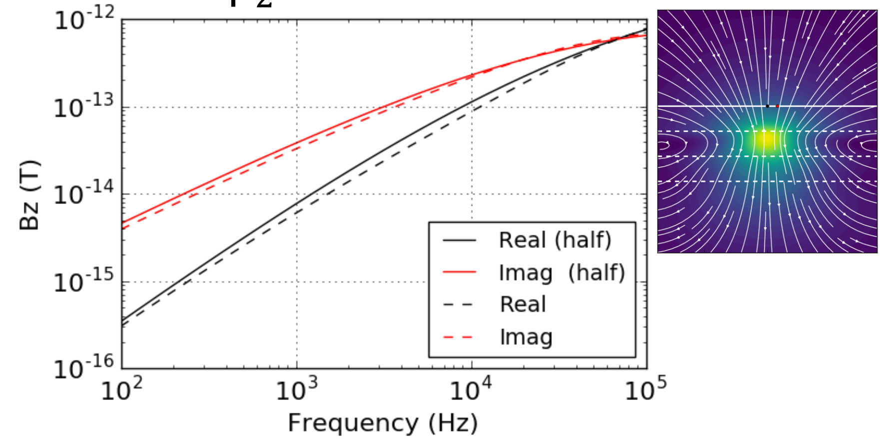


B_z sounding curves

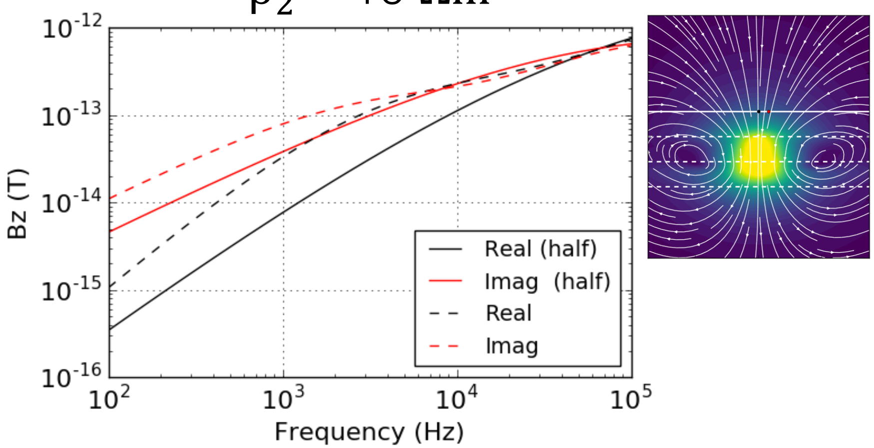
$\rho_2 = 100\ \Omega m$



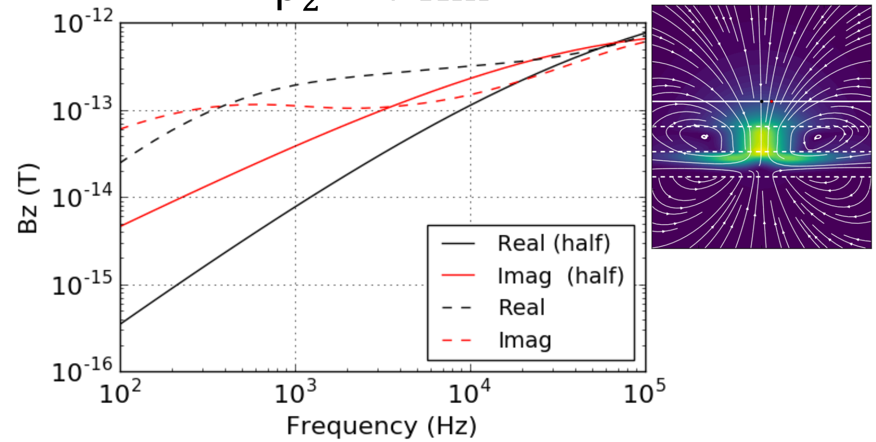
$\rho_2 = 1000\ \Omega m$



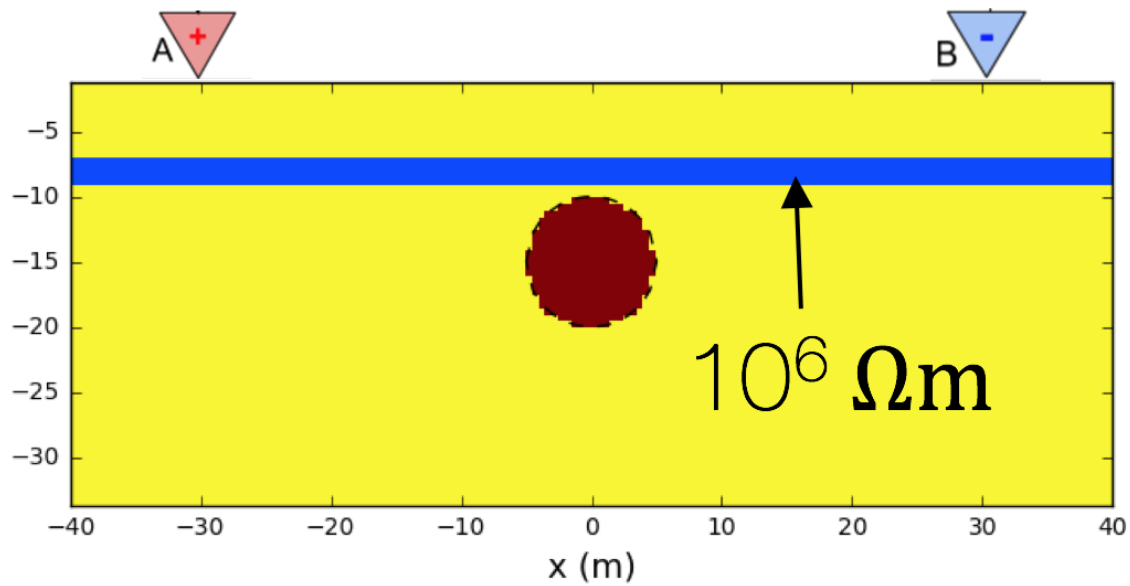
$\rho_2 = 10\ \Omega m$



$\rho_2 = 1\ \Omega m$

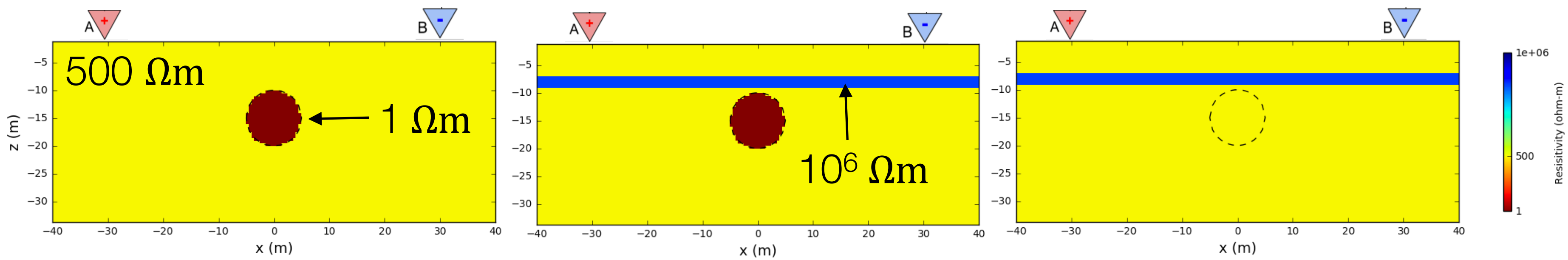


Back to the “shielding” problem

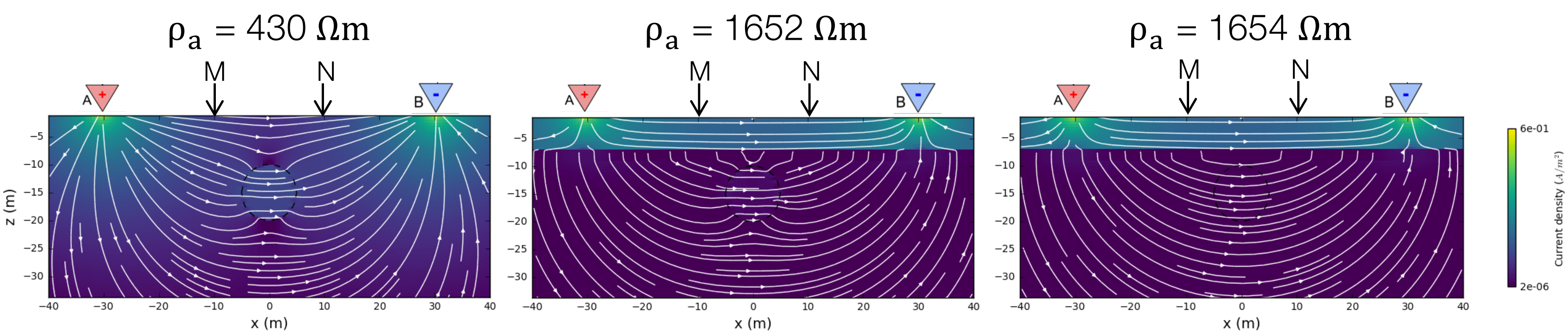


Shielding: DC with resistive layer

Resistivity models (thin **resistive** layer)

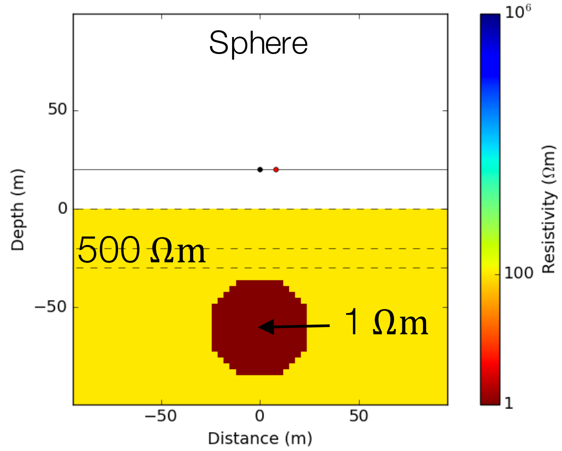


Currents and measured data at MN

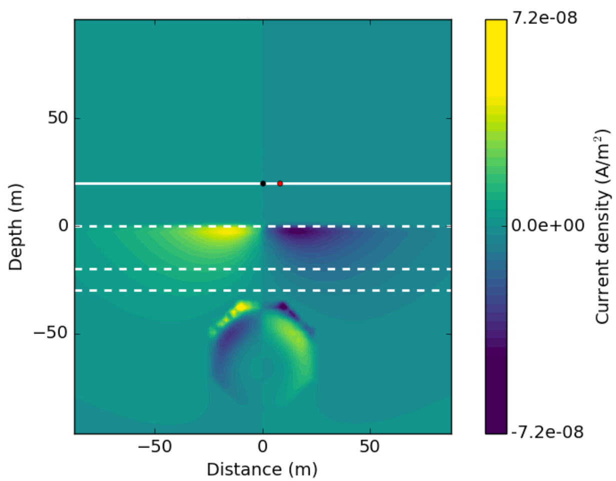


Shielding: EM with resistive layer

Resistivity models (thin **resistive** layer)

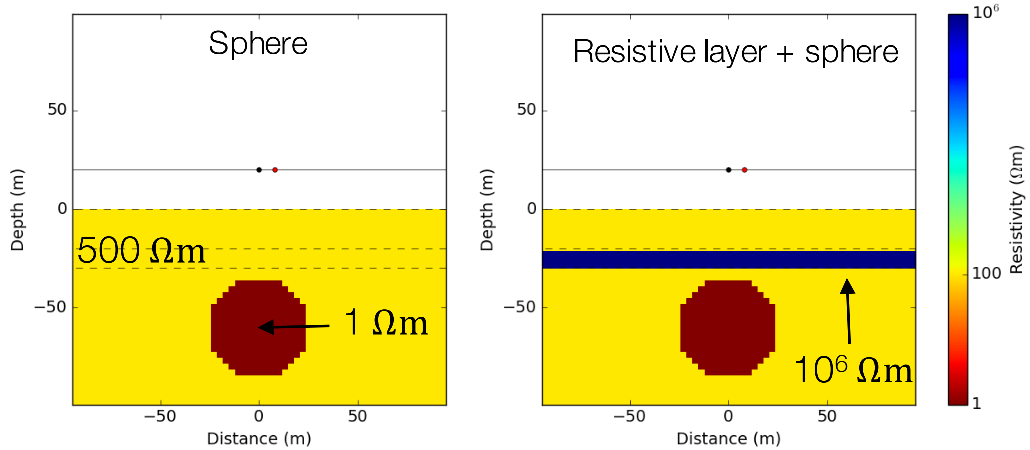


Currents (J_y imag)

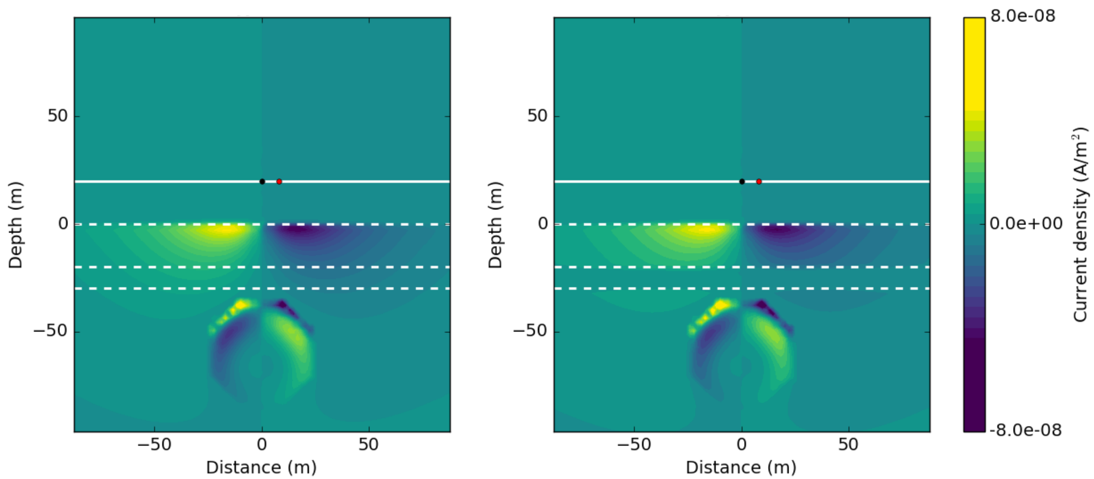


Shielding: EM with resistive layer

Resistivity models (thin **resistive** layer)

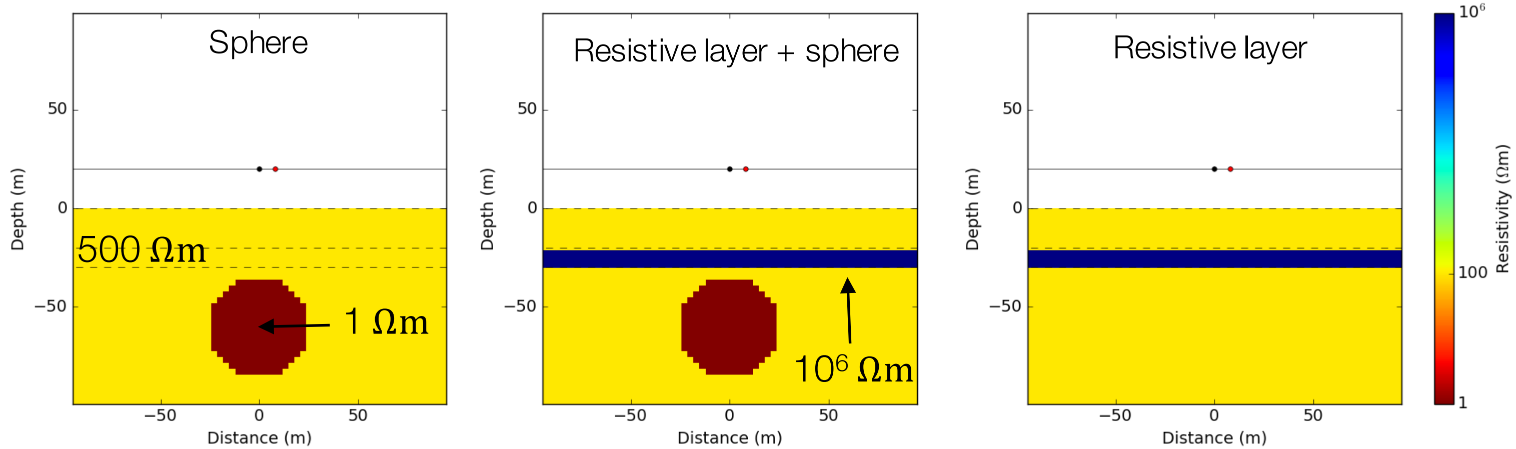


Currents (J_y imag)

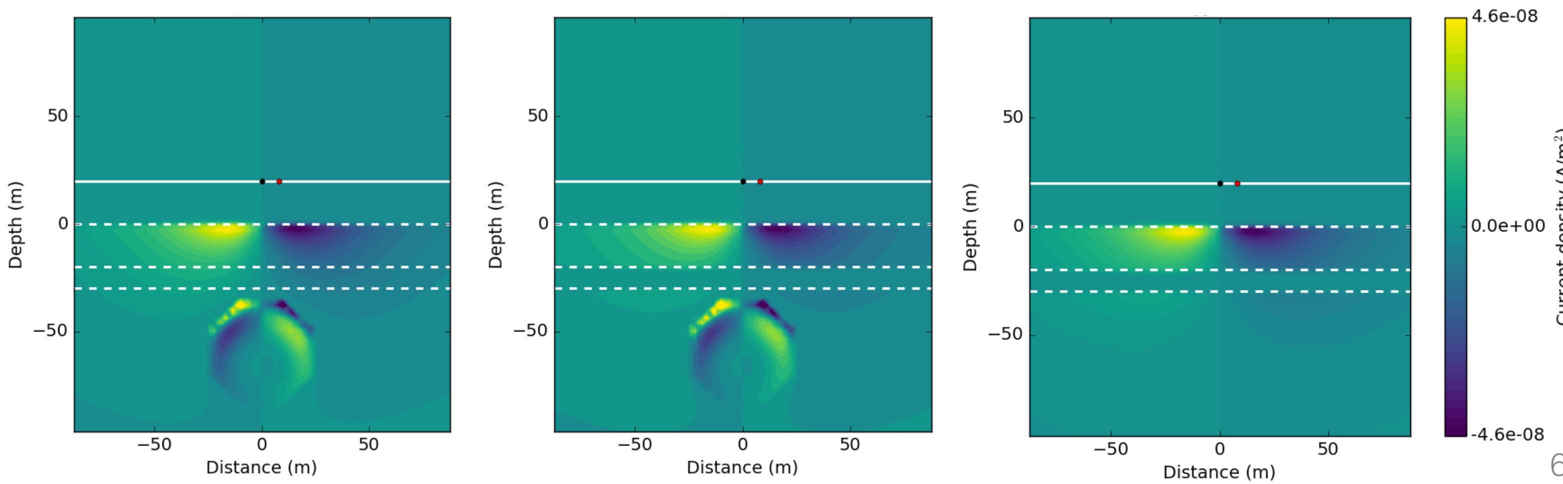


Shielding: EM with resistive layer

Resistivity models (thin **resistive** layer)

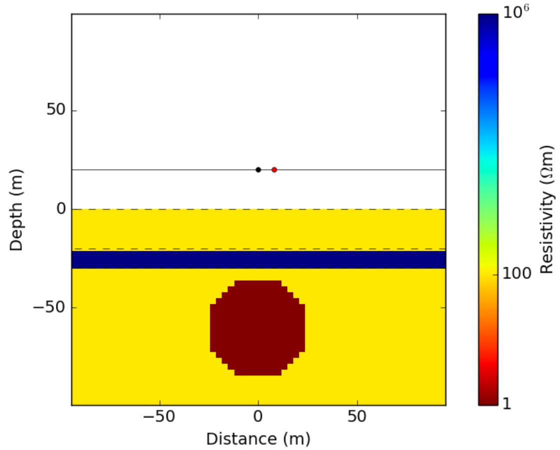


Currents (J_y imag)

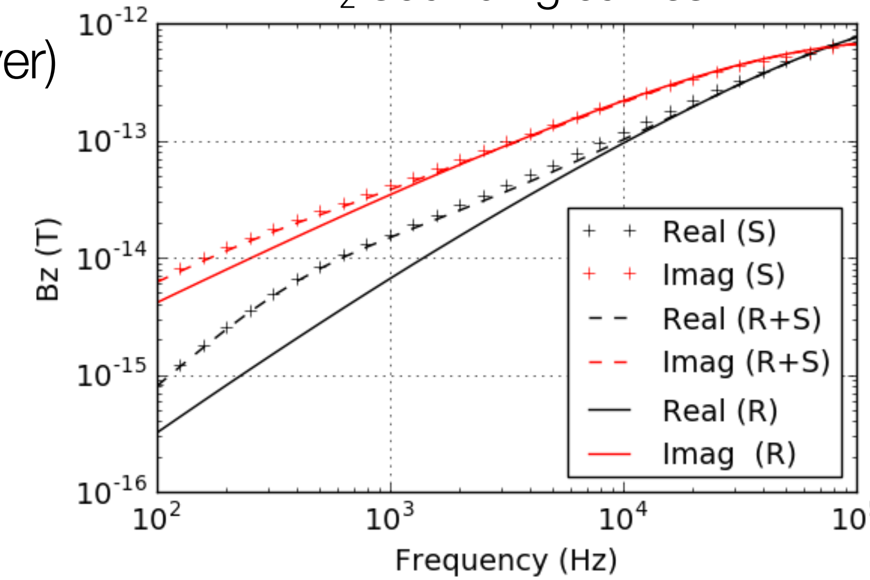


Shielding: EM with resistive layer

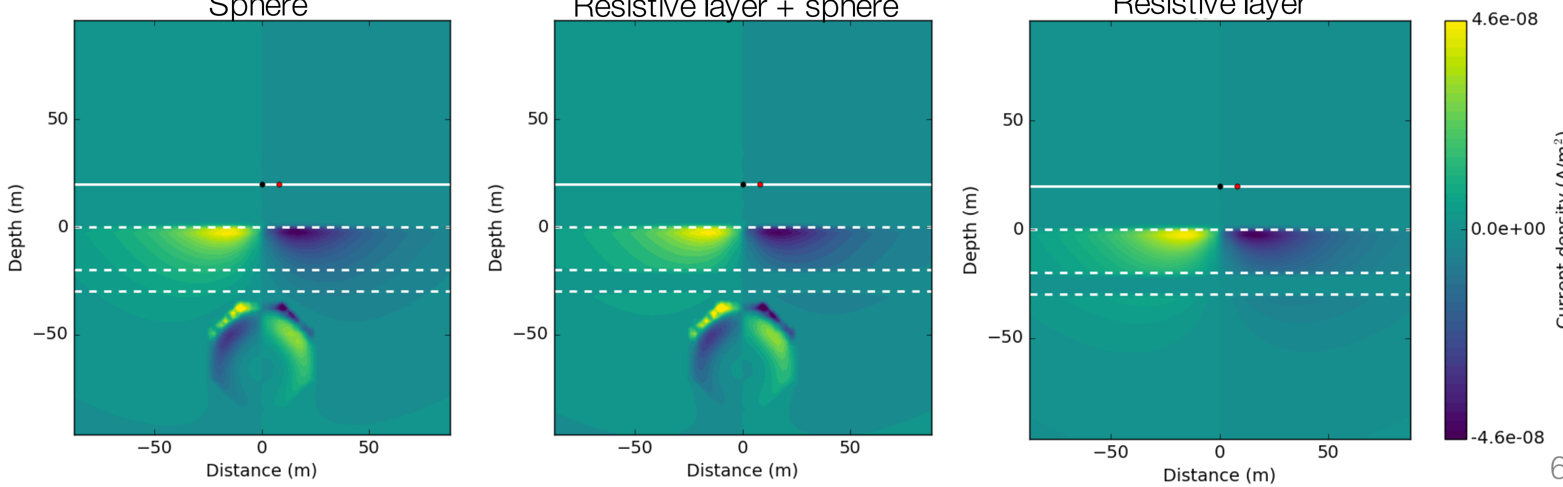
Resistivity models (thin **resistive** layer)



B_z sounding curves

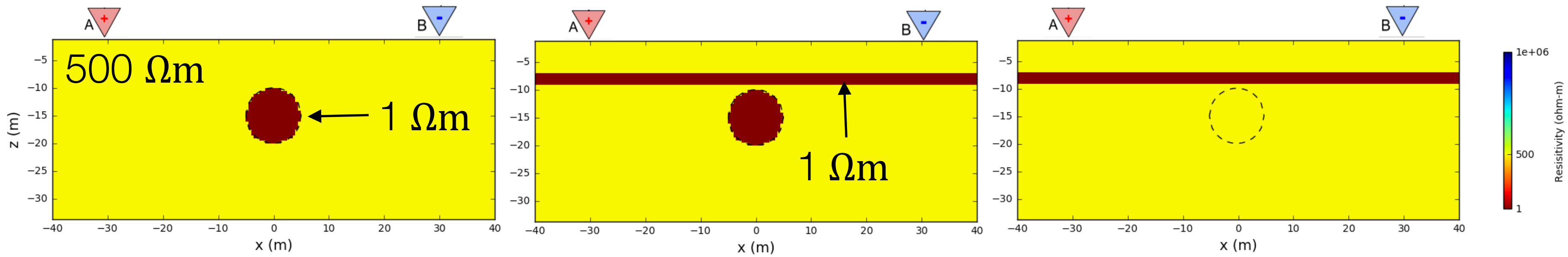


Currents (J_y imag)

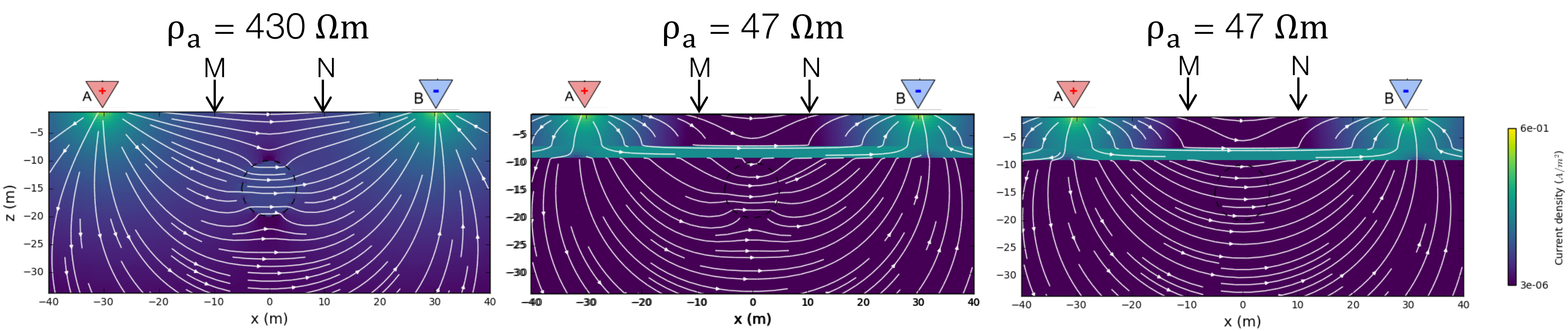


Shielding: DC with conductive layer

Resistivity models (thin **conductive** layer)

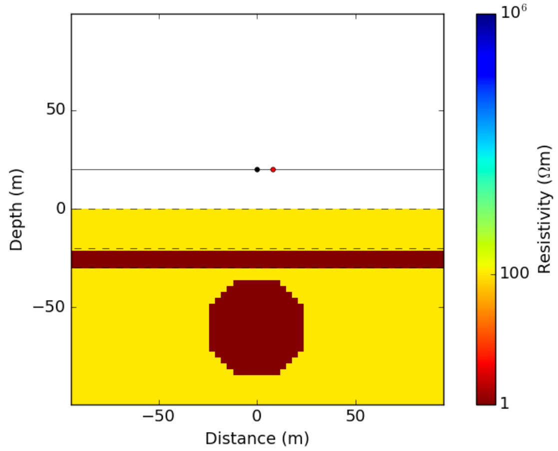


Currents and measured data at MN

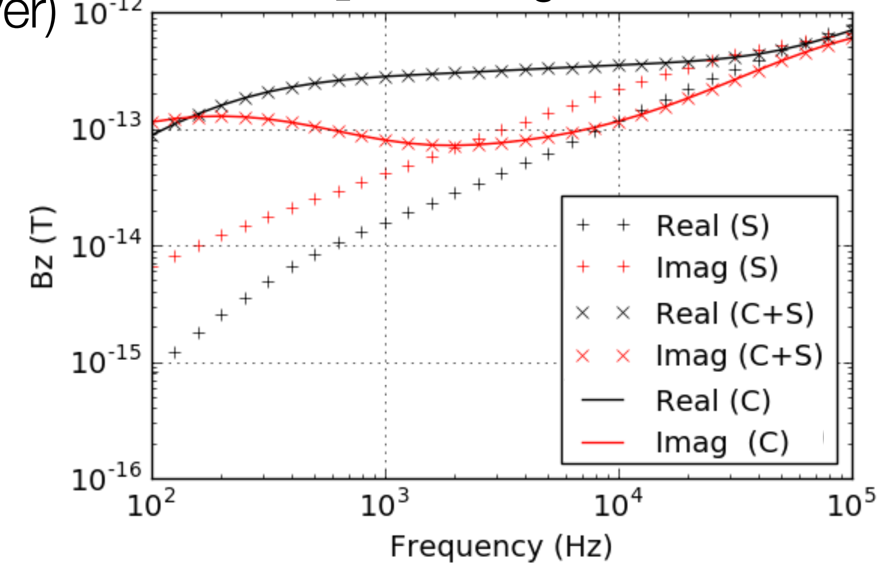


Shielding: EM with conductive layer

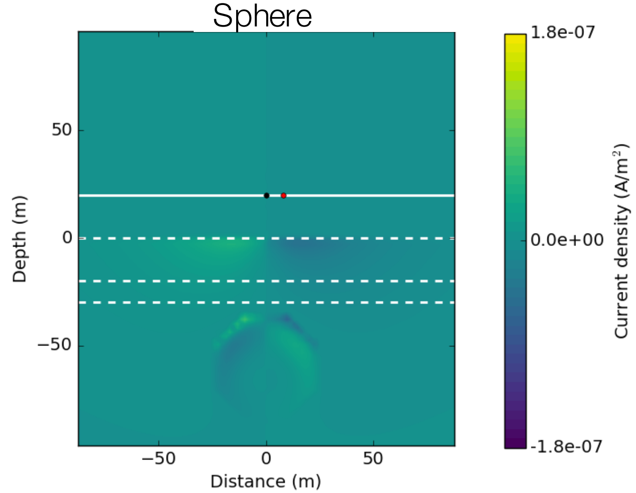
Conductivity models (thin **resistive** layer)



B_z sounding curves

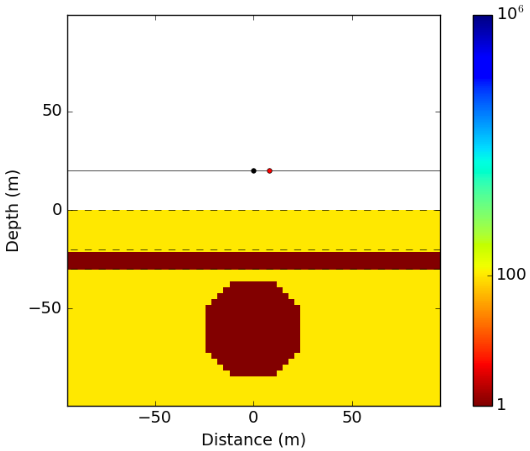


Currents (J_y imag)

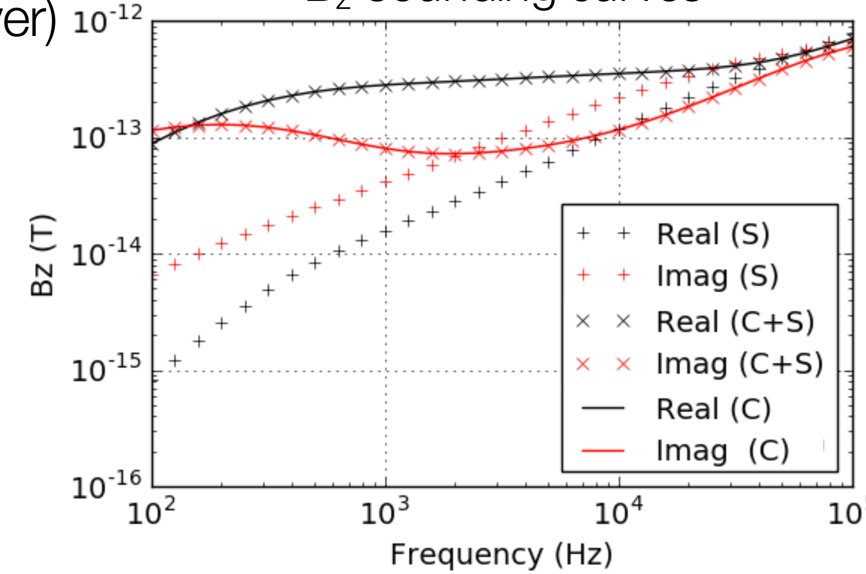


Shielding: EM with conductive layer

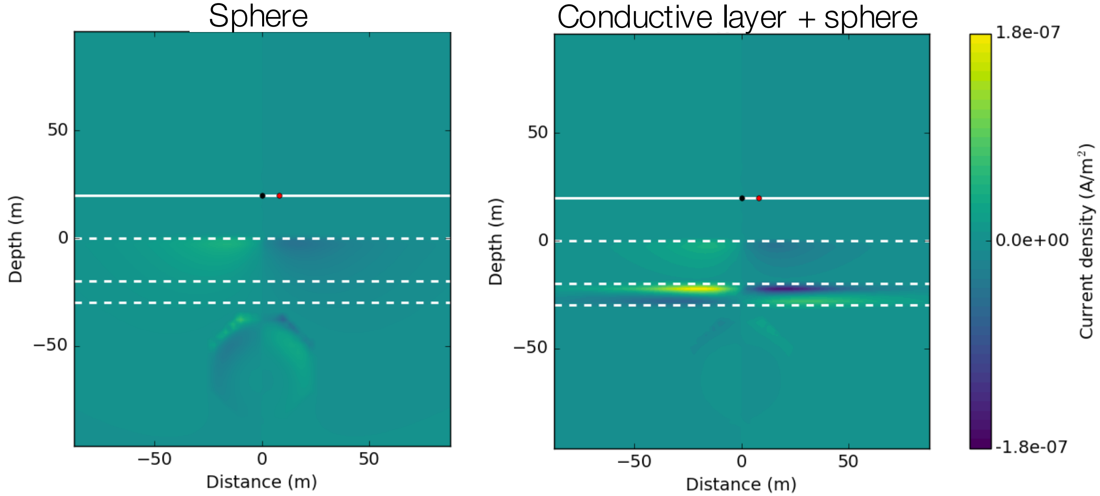
Conductivity models (thin **resistive** layer)



B_z sounding curves

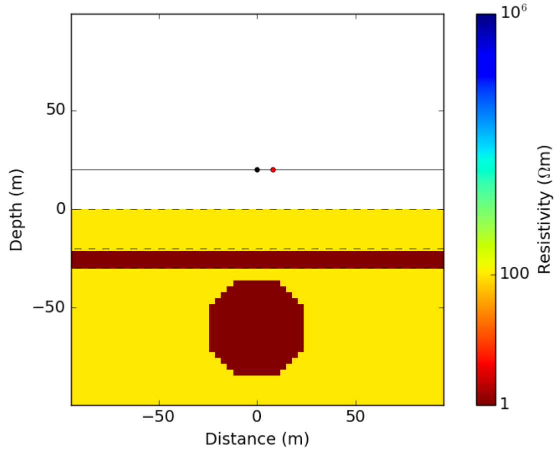


Currents (J_y imag)

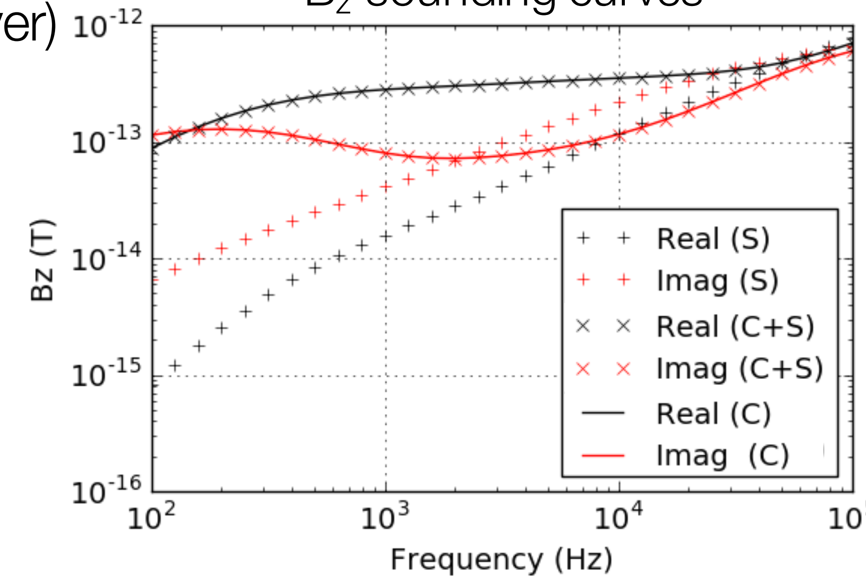


Shielding: EM with conductive layer

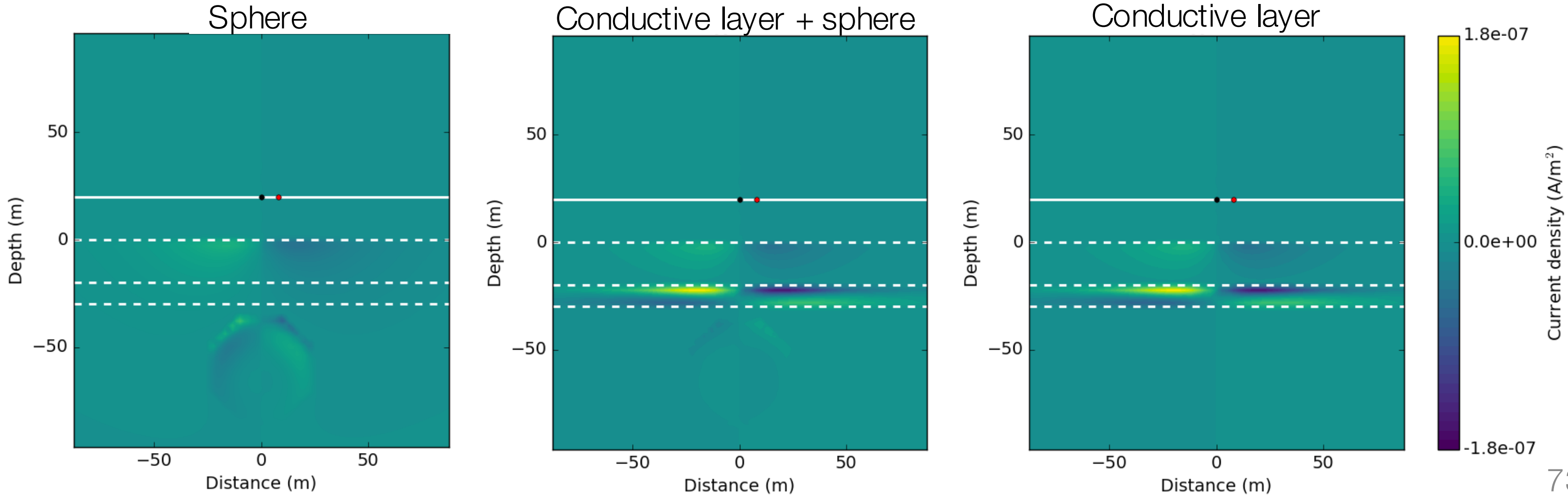
Conductivity models (thin **resistive** layer)



B_z sounding curves



Currents (J_y imag)



Outline

Setup

- Basic experiment
- Transmitters, Receivers

Time Domain EM

- Vertical Magnetic Dipole
- Propagation with Time
- Case History

Frequency Domain EM

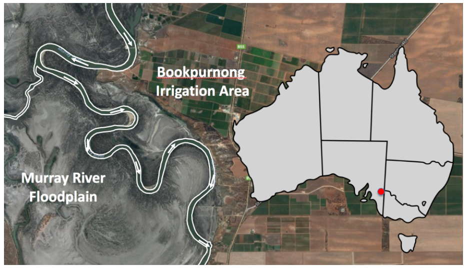
- Vertical Magnetic Dipole
- Effects of Frequency
- Case History – Groundwater

Questions

Case History: Bookpurnong

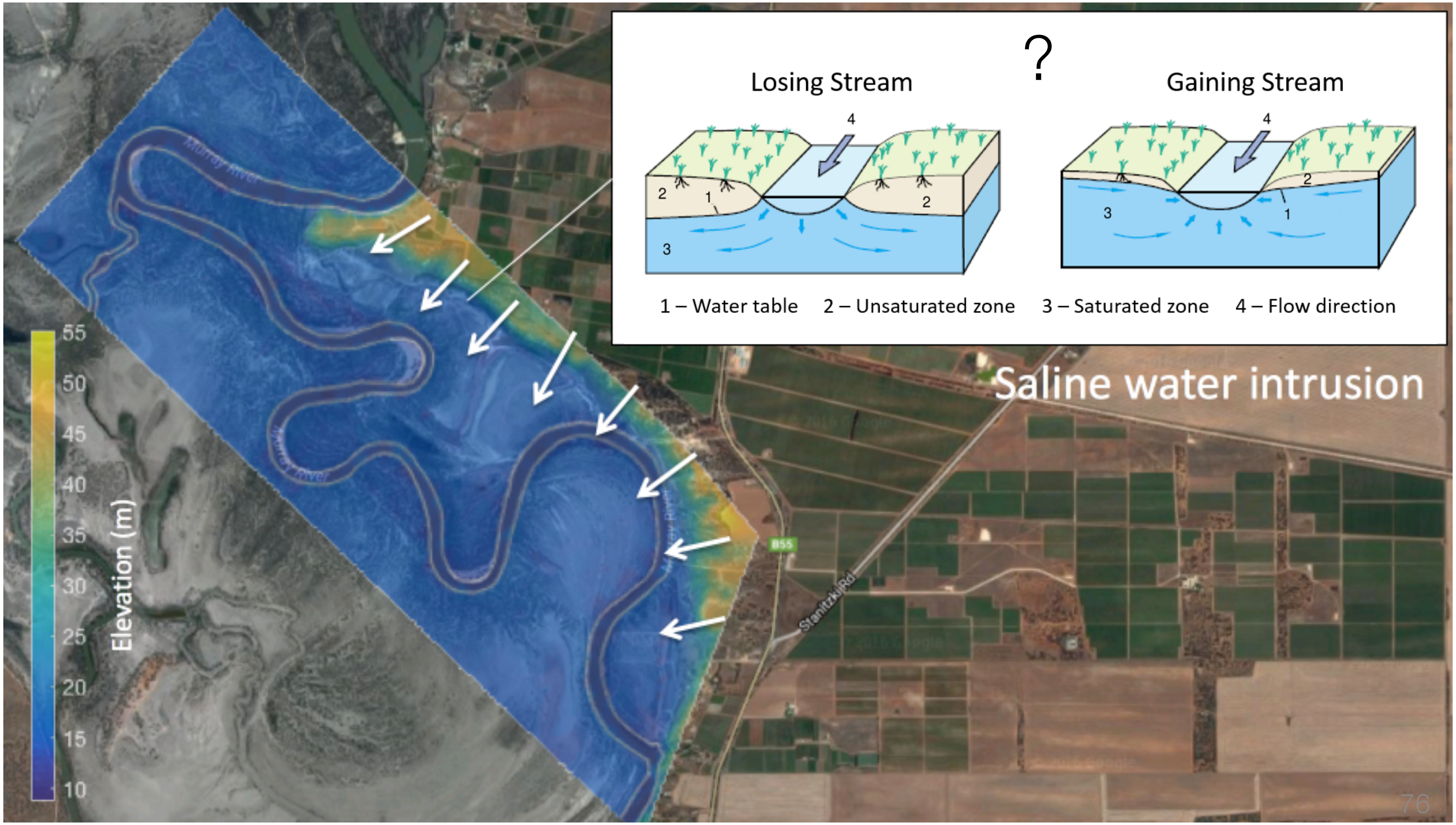
Viezzoli et al., 2009

Setup



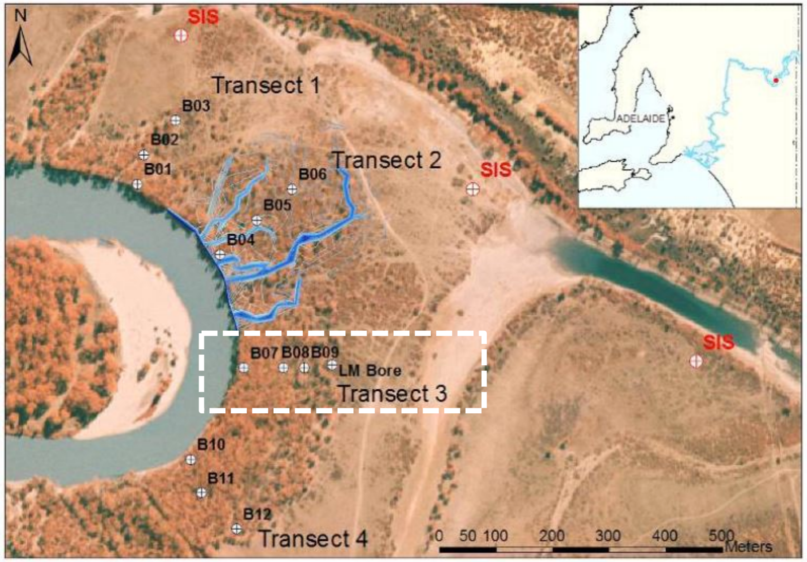
Geoscience Australia project

- Characterizing river salination



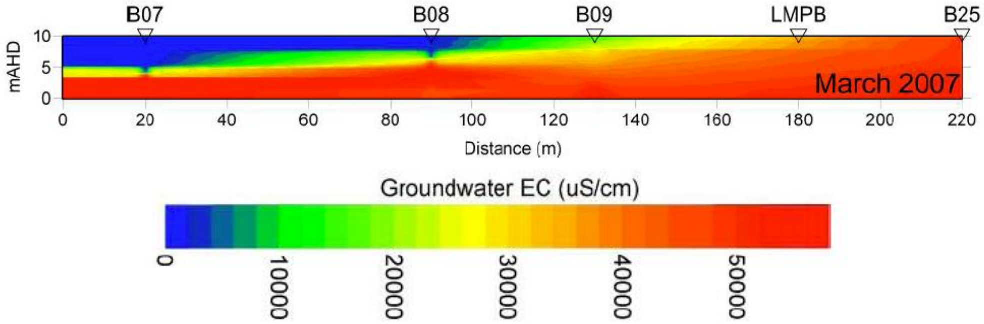
Properties

Location map for salinity measurements



Unit	Conductivity
Saline water	High, 3 - 5 S/m
Fresh water	Low, 0.01 S/m

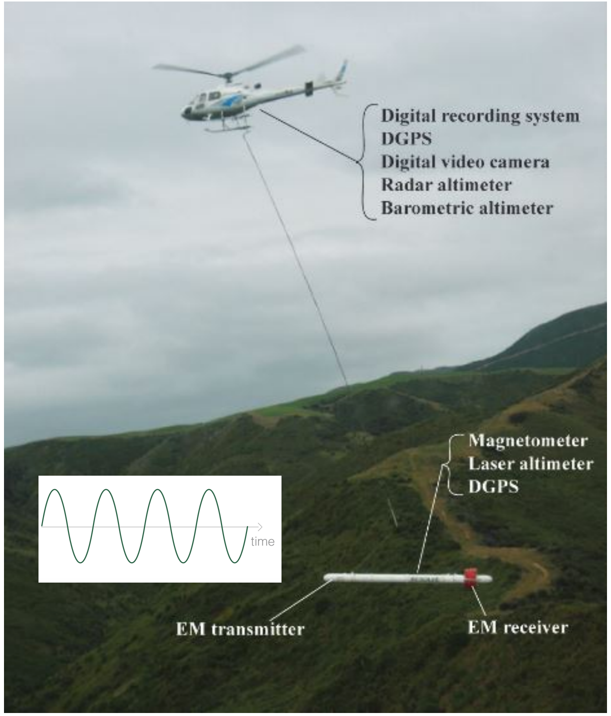
Conductivity from salinity measurements



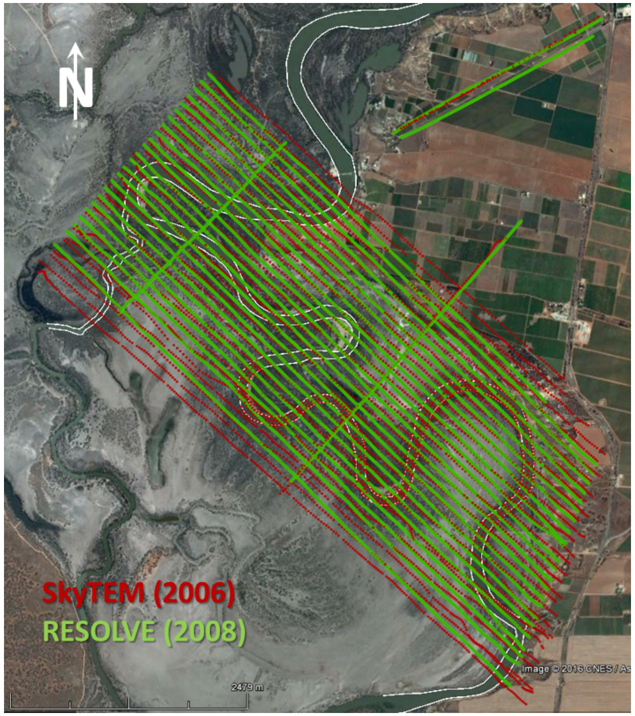
Holland et. al., 2008.

Survey

Resolve system (2008)

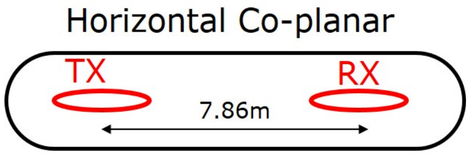


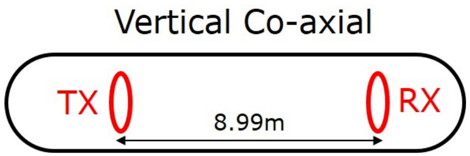
Flight lines



Horizontal Co-planar (HCP) frequencies:
- 382, 1822, 7970, 35920 and 130100 Hz

Vertical Co-axial (VCA) frequencies:
- 3258 Hz



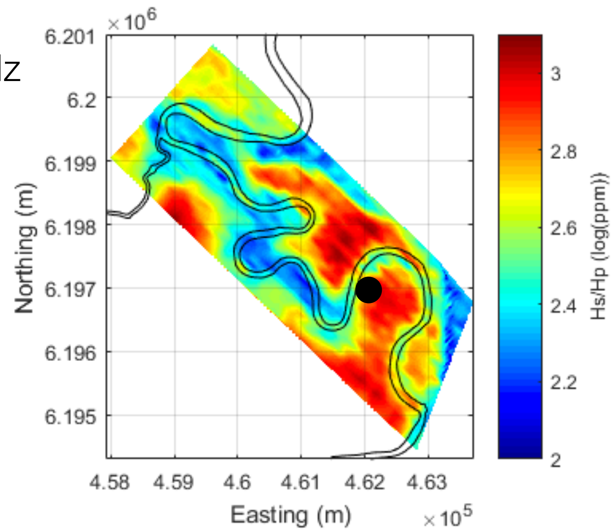


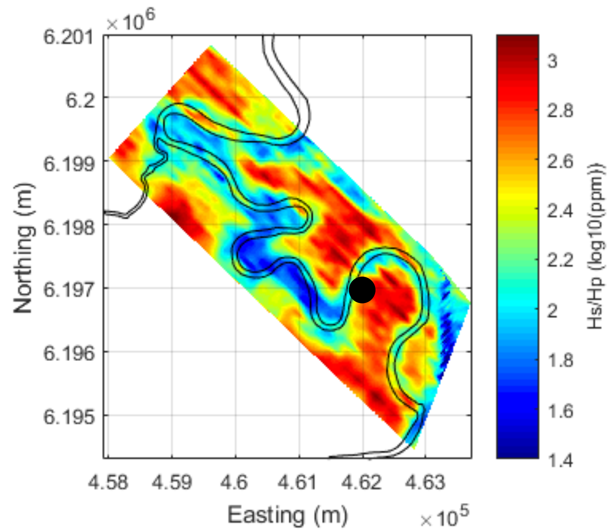
Horizontal Co-planar (HCP) data

In-Phase (Real)

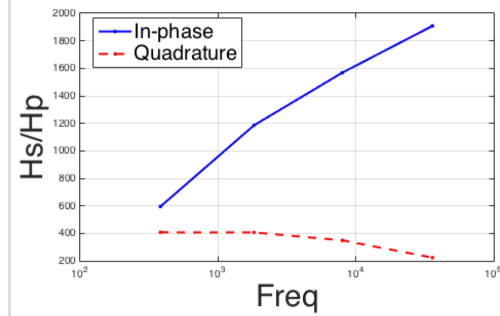
Quadrature (Imaginary)

382 Hz

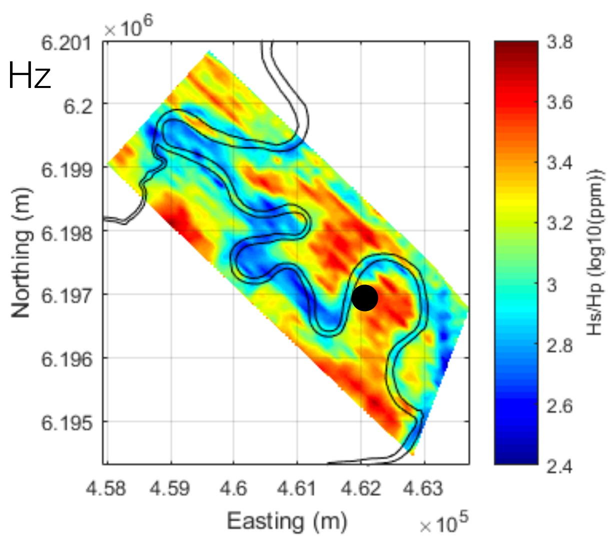


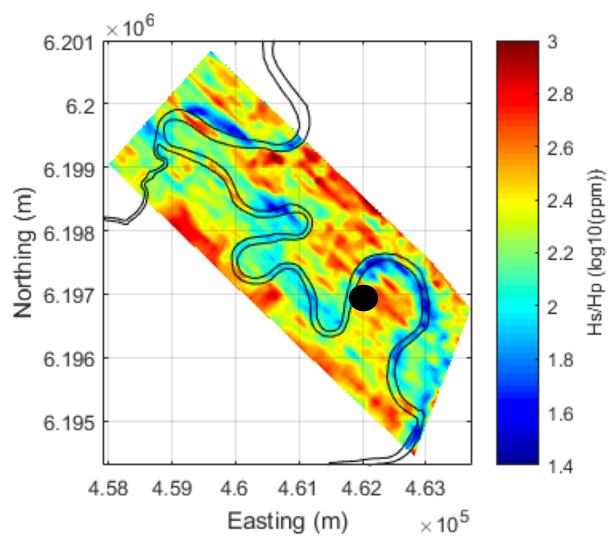


Sounding curve

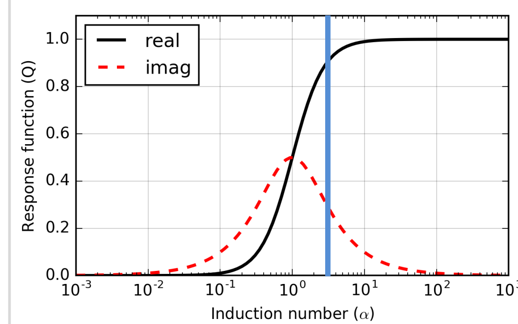


35920 Hz



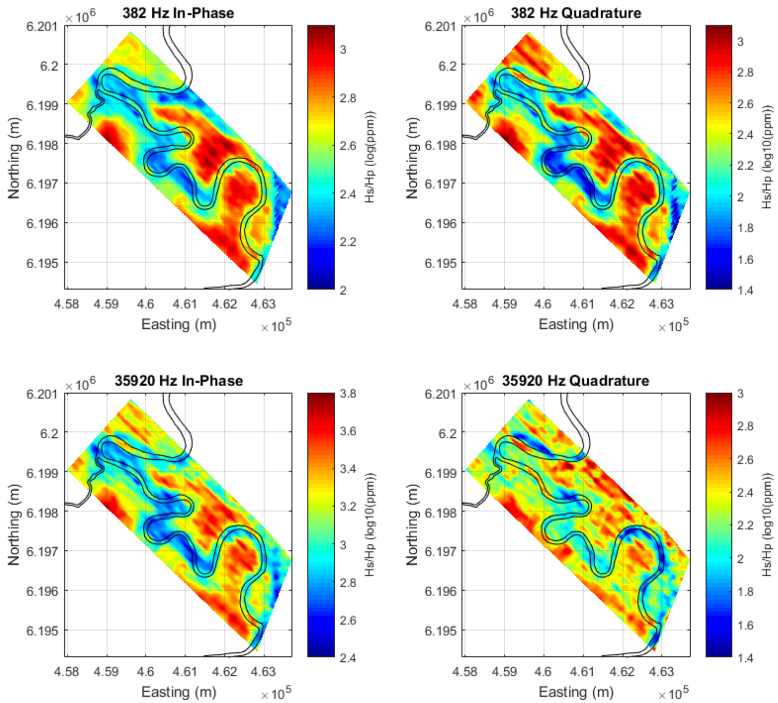


Response curve

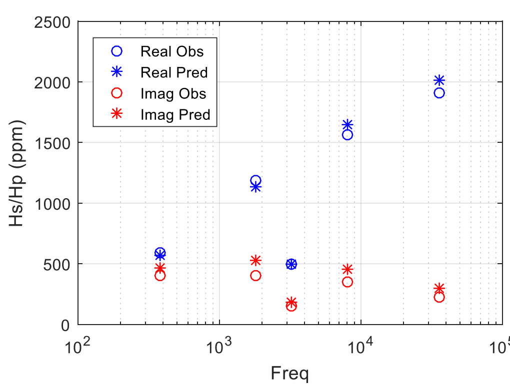


Processing: 1D inversion

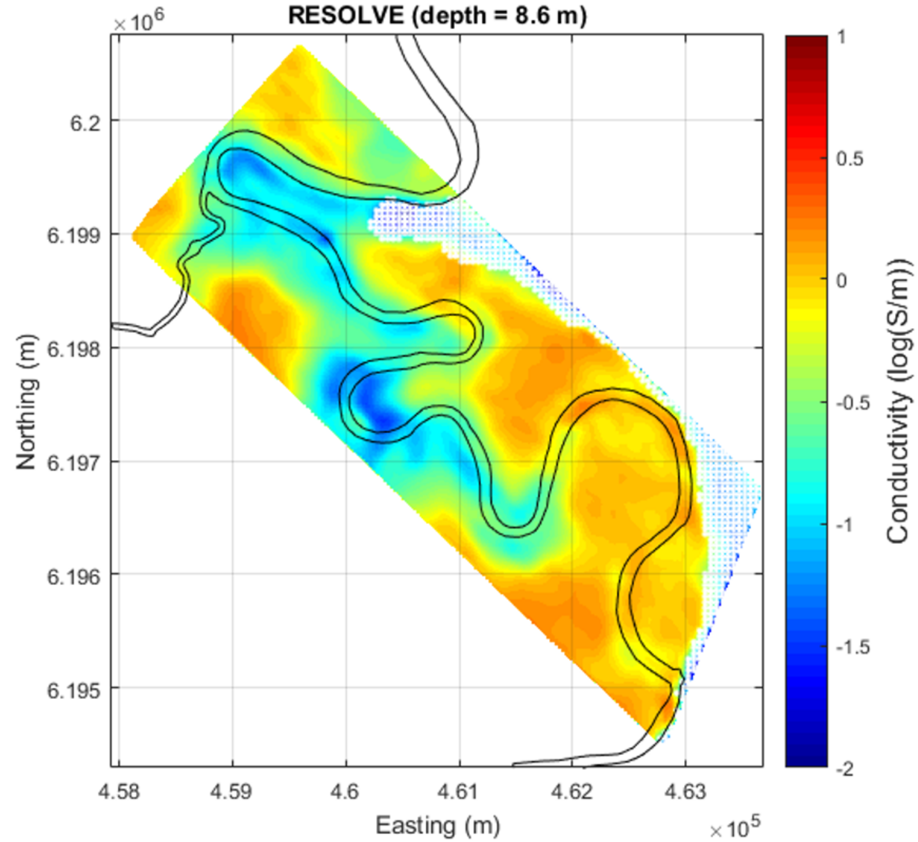
Data



Data fit

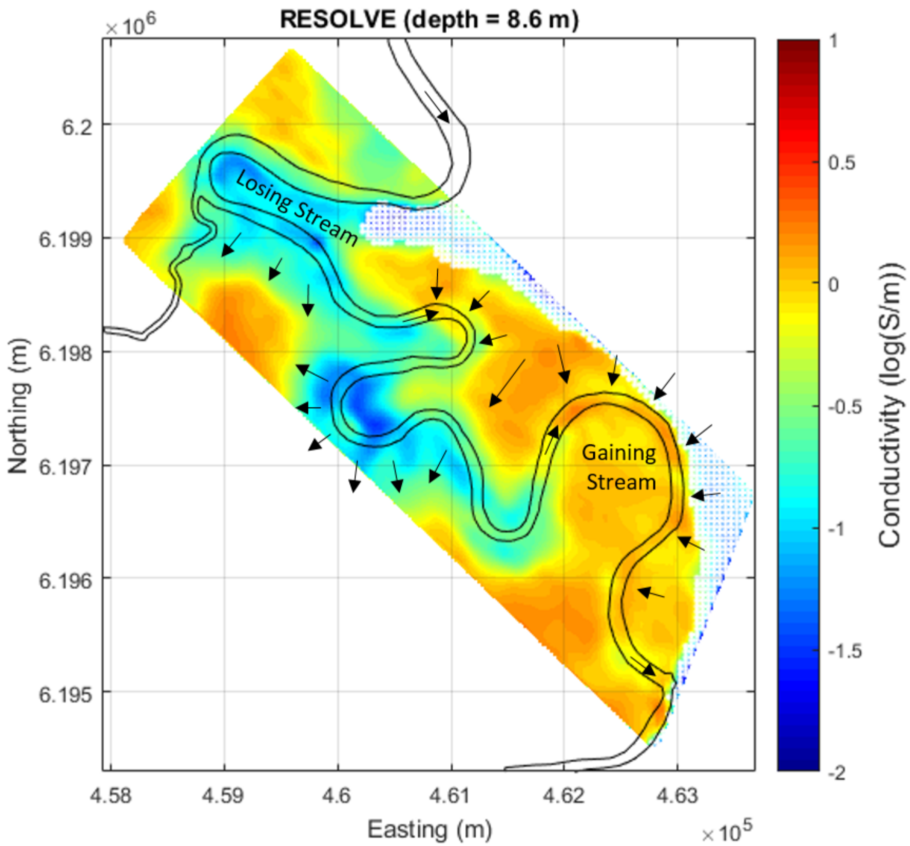


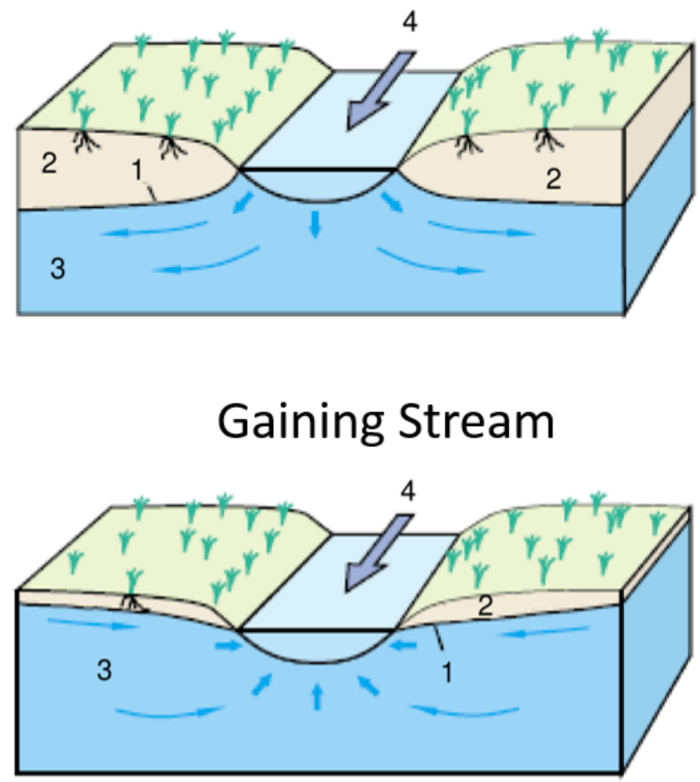
Conductivity model (stitched)



Interpretation

Conductivity model (stitched)



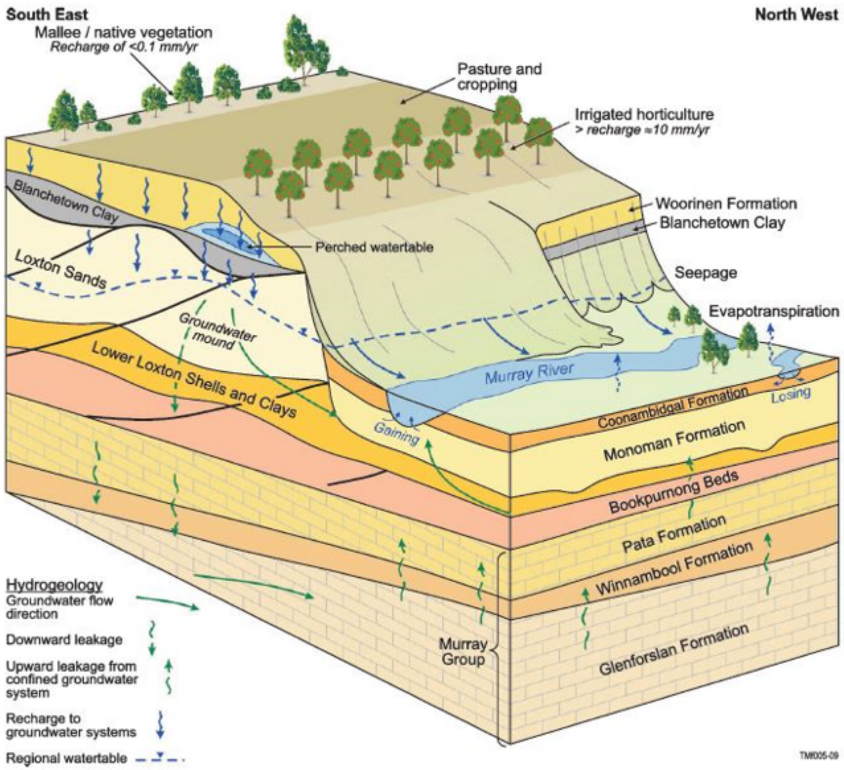


1 – Water table 2 – Unsaturated zone

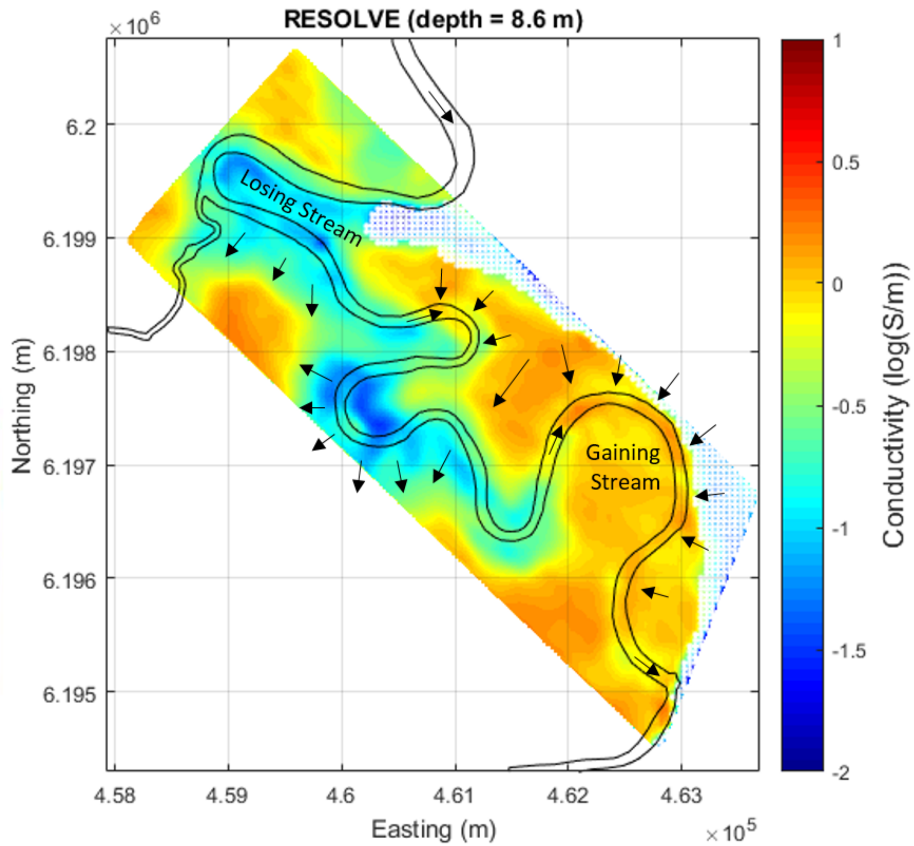
3 – Saturated zone 4 – Flow direction

Synthesis

Hydrological model

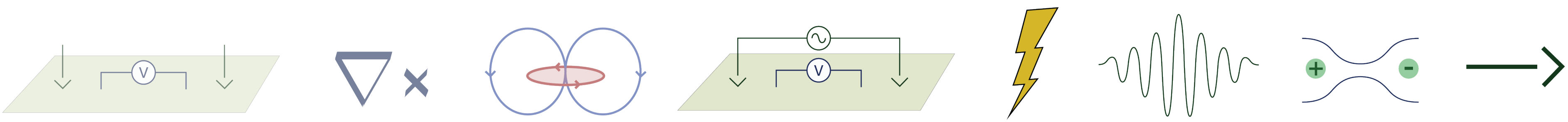


Conductivity model (stitched)



End of Inductive Sources

Next up



DC Resistivity

EM Fundamentals

Inductive Sources

Grounded Sources

Natural Sources

GPR

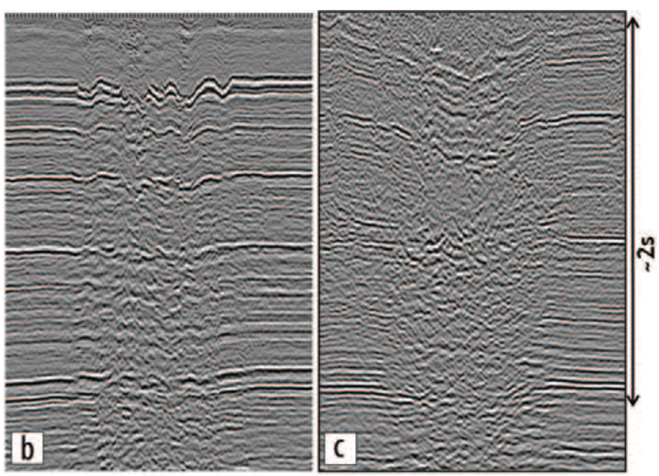
Induced Polarization

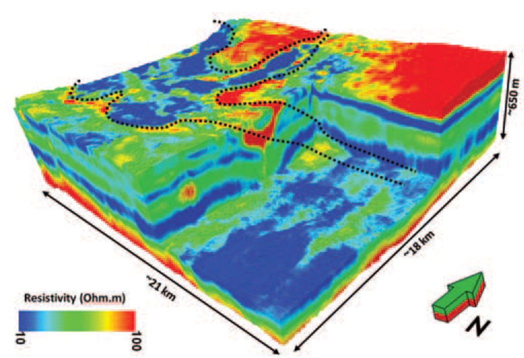
The Future

Lunch: Play with apps

Case History: Wadi Sahba

Colombo et al. 2016

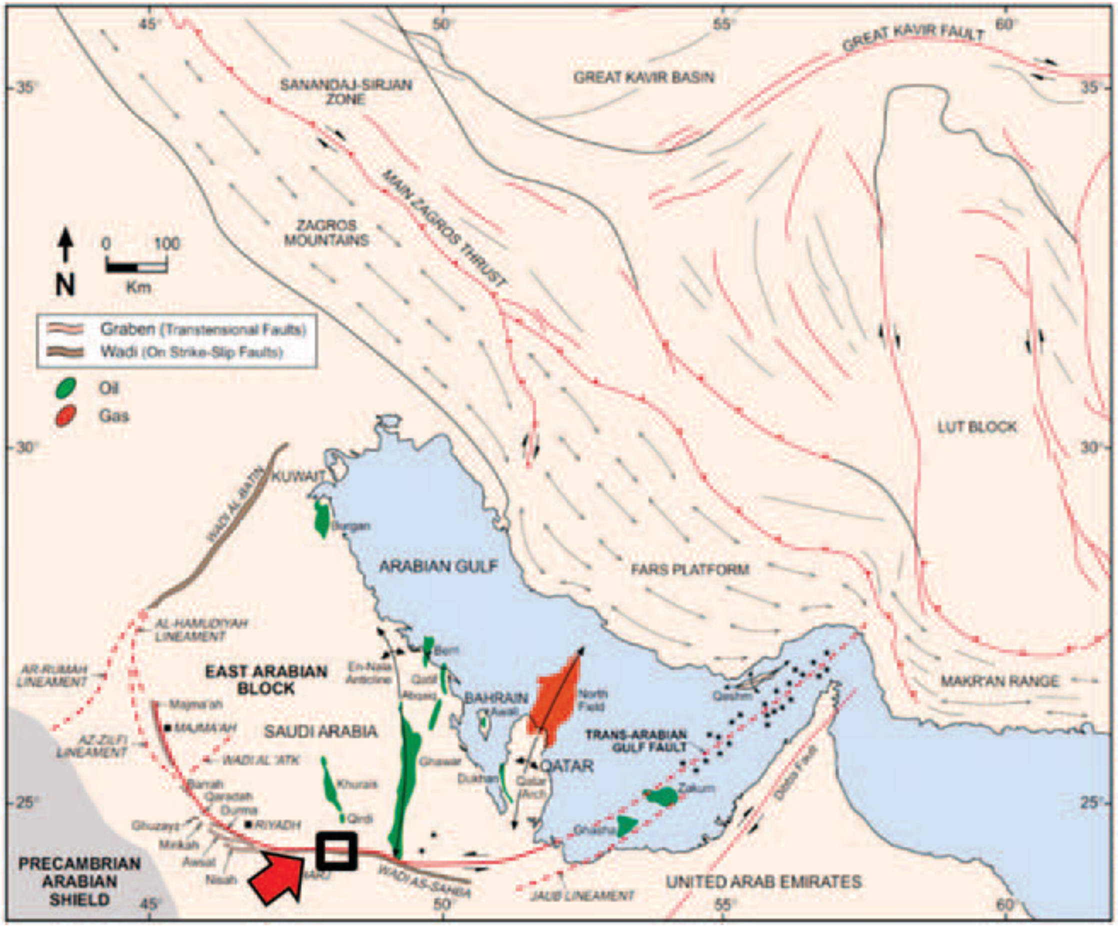




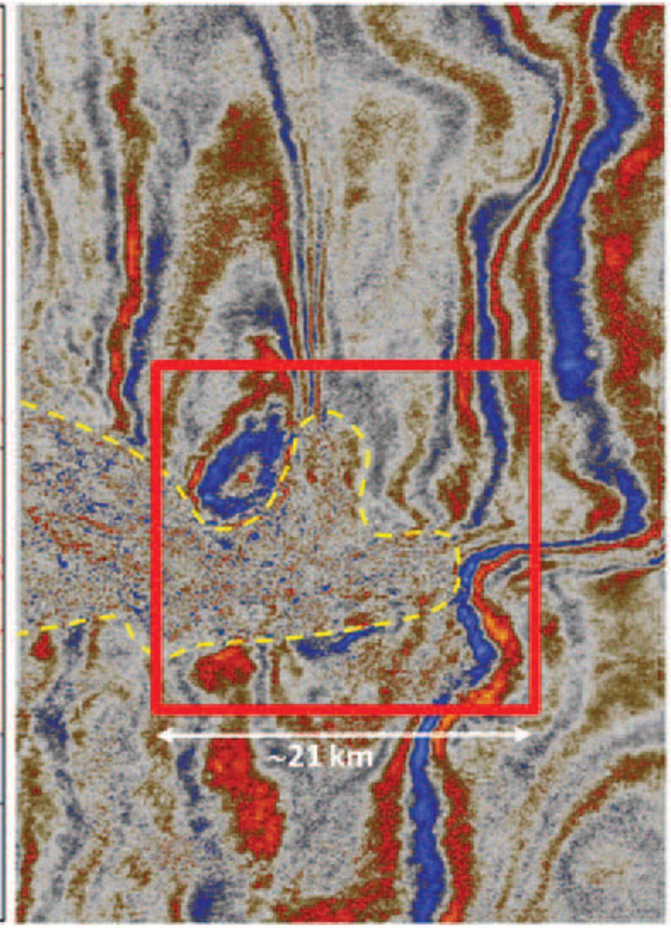
?

Setup

Location of Wadi area, Saudi Arabia



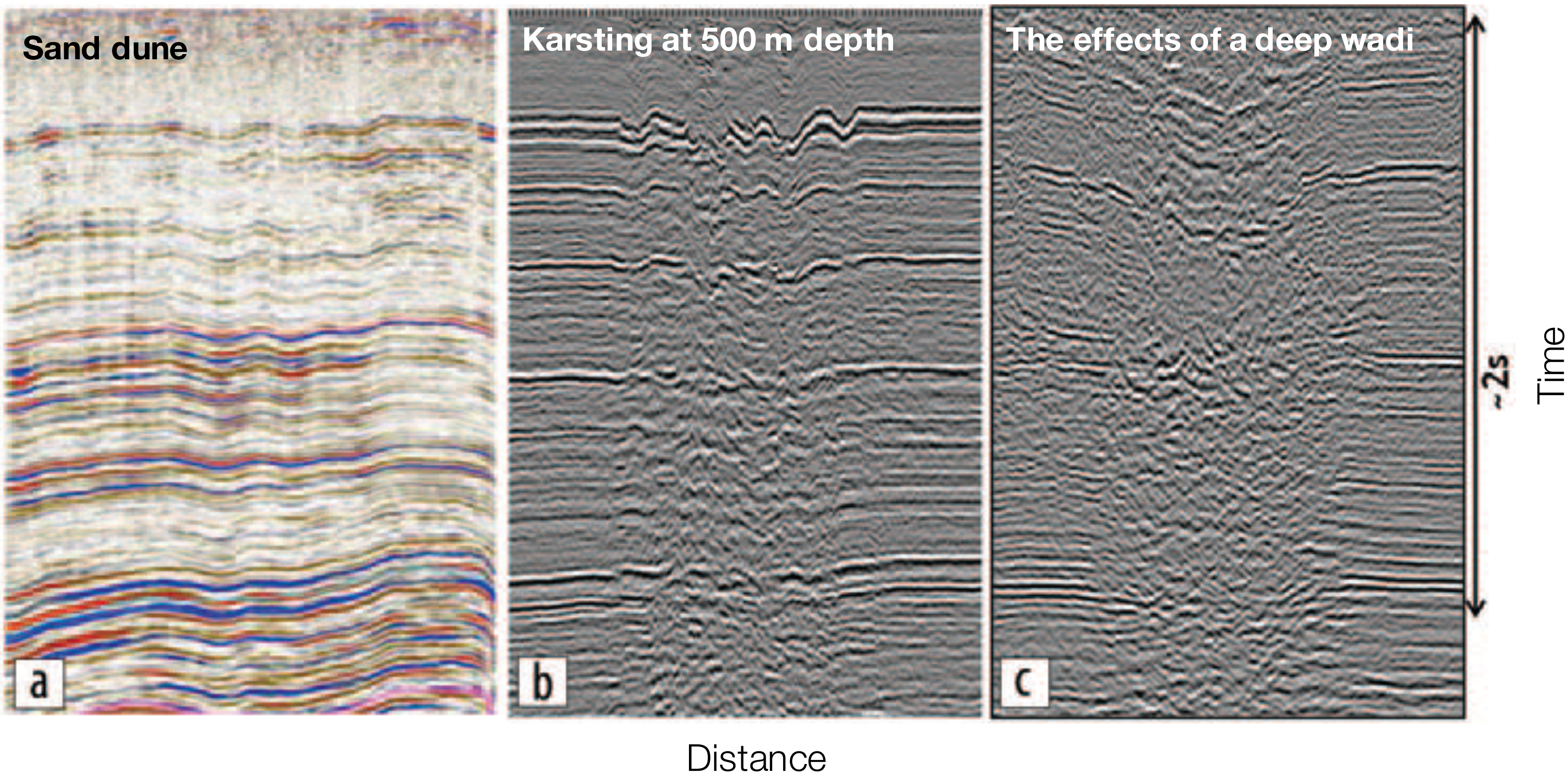
Seismic slice



- Oil and gas exploration in the Middle East: Major structures to stratigraphic traps and low relief structures

Challenges for processing seismic data

Example seismic sections



- Strong effects from near surface anomalies even after static corrections

Properties

- P-velocity and conductivity:

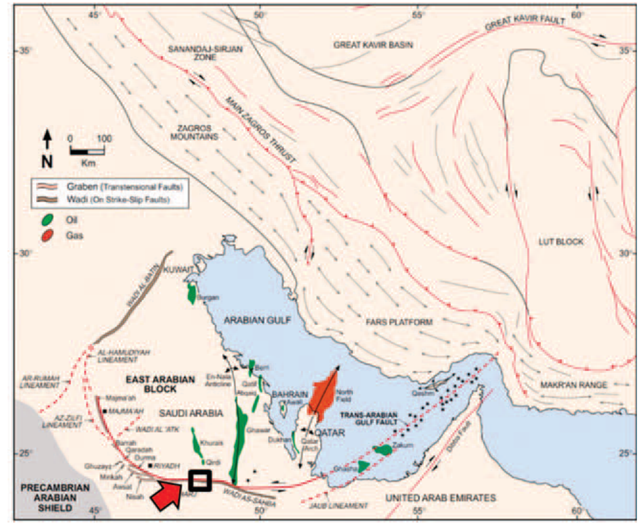
$$v_p = g(\phi)$$
$$\sigma = f(\phi)$$

v_p: P-velocity
ϕ: porosity

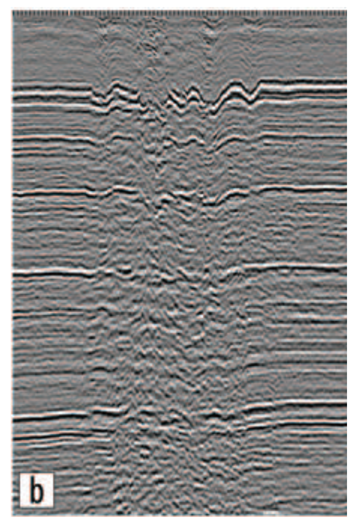
- Poor seismic data:
  - strong scattering effects probably caused by flower faults
  - velocity inversions (high to low v_p)

- From previous multi-physics analyses:
  - strong structural similarity between the inverted resistivity, and the existing seismic results

Geologic map



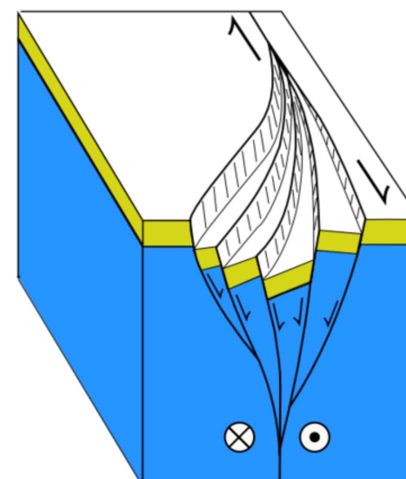
Seismic section



Time

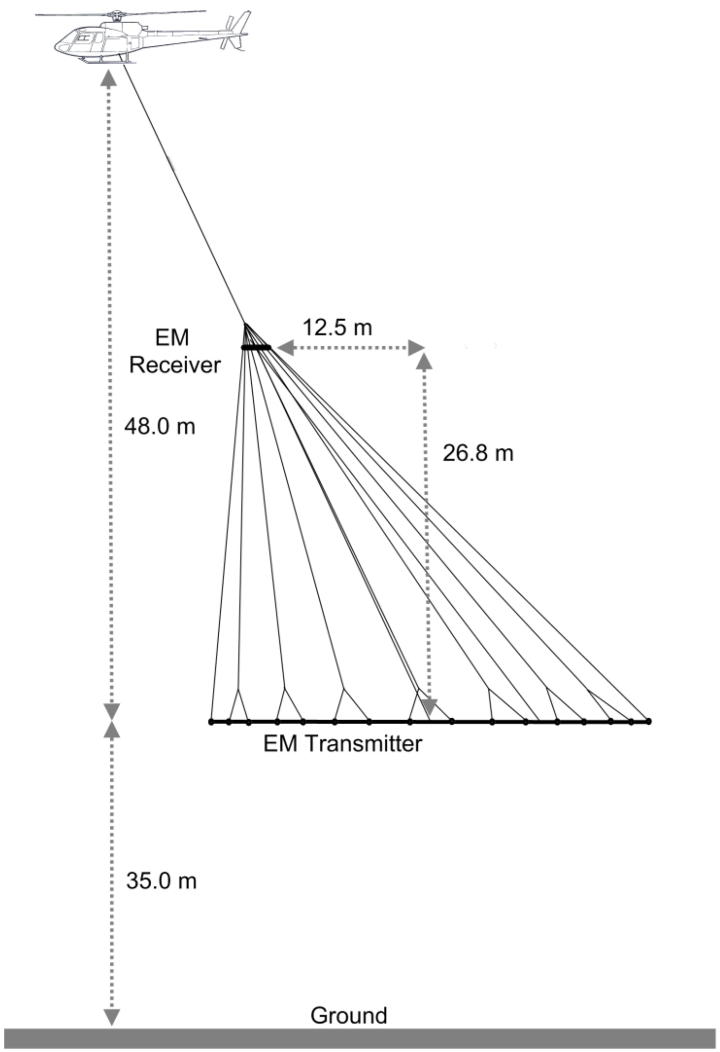
Distance

Flower faults

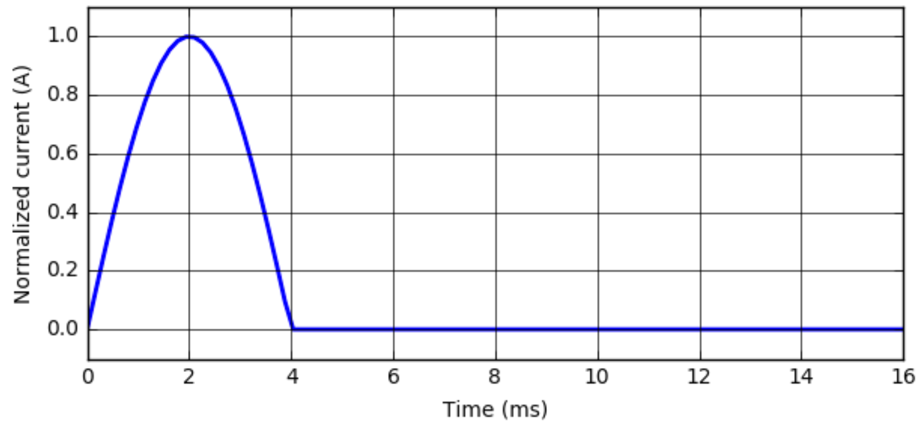


Survey

HELITEM

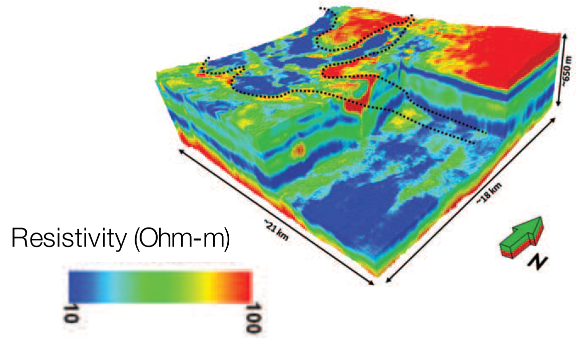


System Configuration

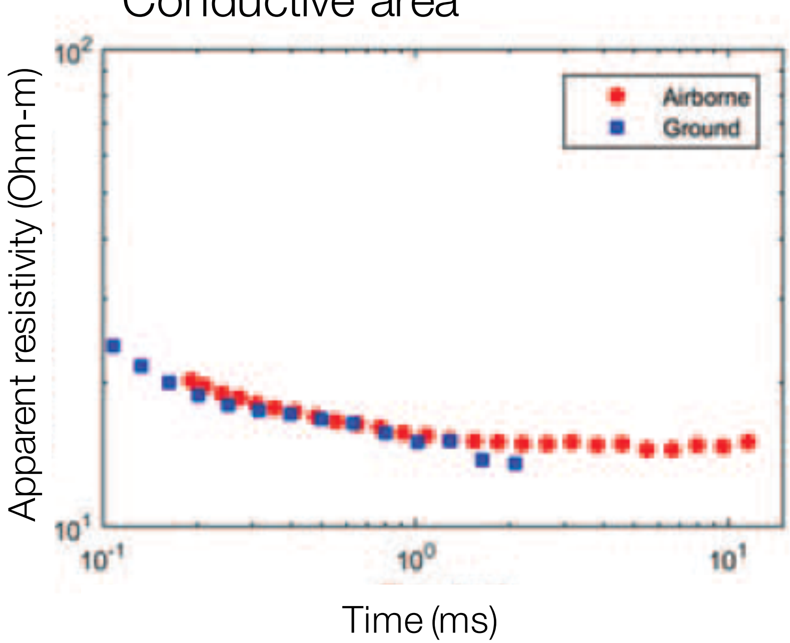


- Peak Tx current: 1200 A
- Dipole moment: 1.7×10^6 A-m^2
- Stacked TEM curve spacing: ~2.7 m
- Total soundings: ~1.6 million

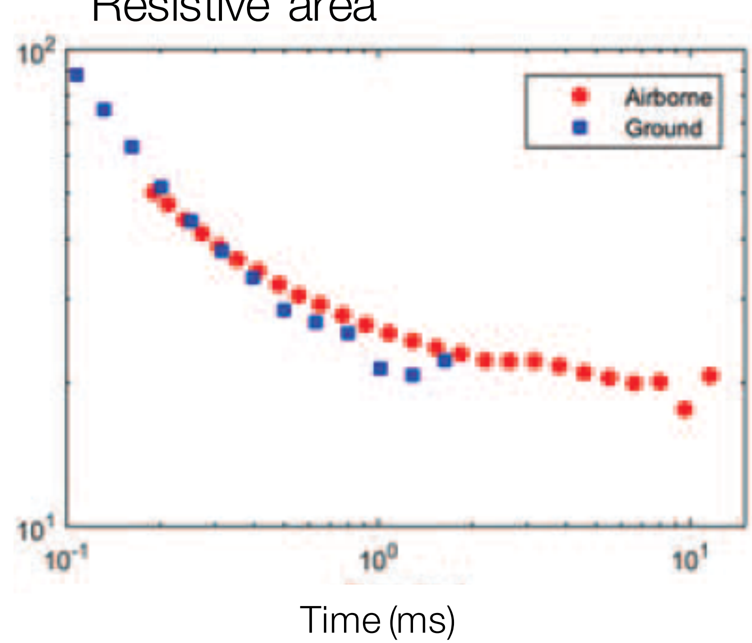
Comparisons: airborne and ground EM



Conductive area

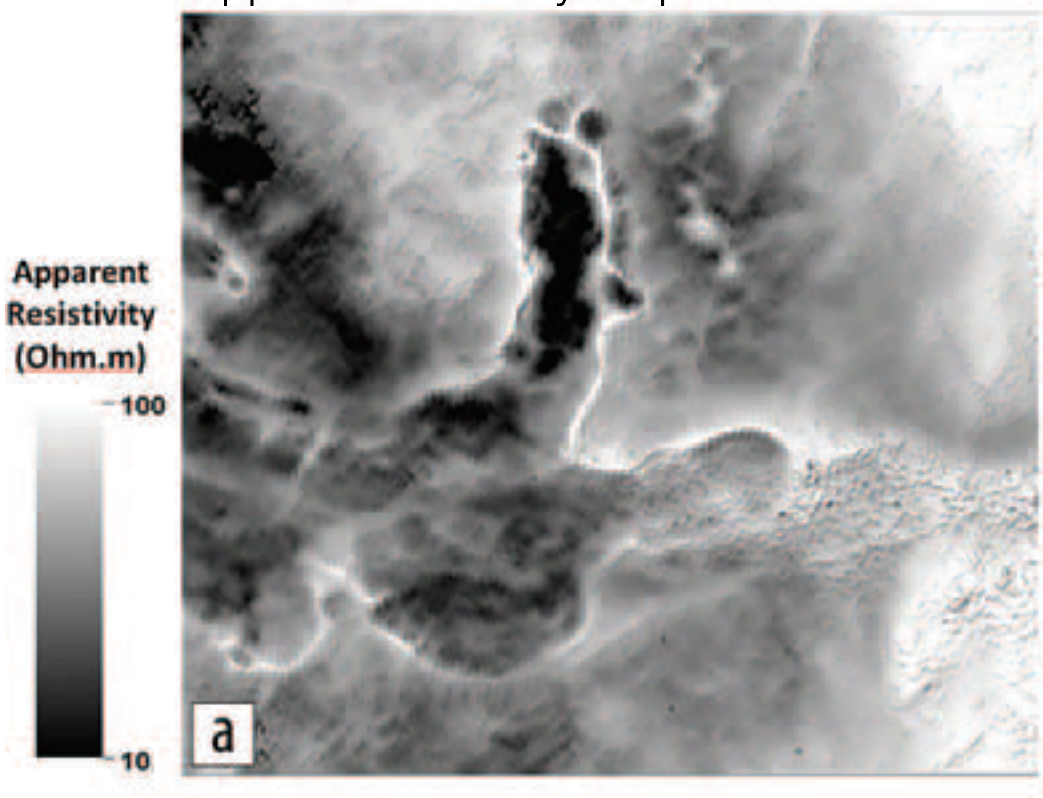


Resistive area

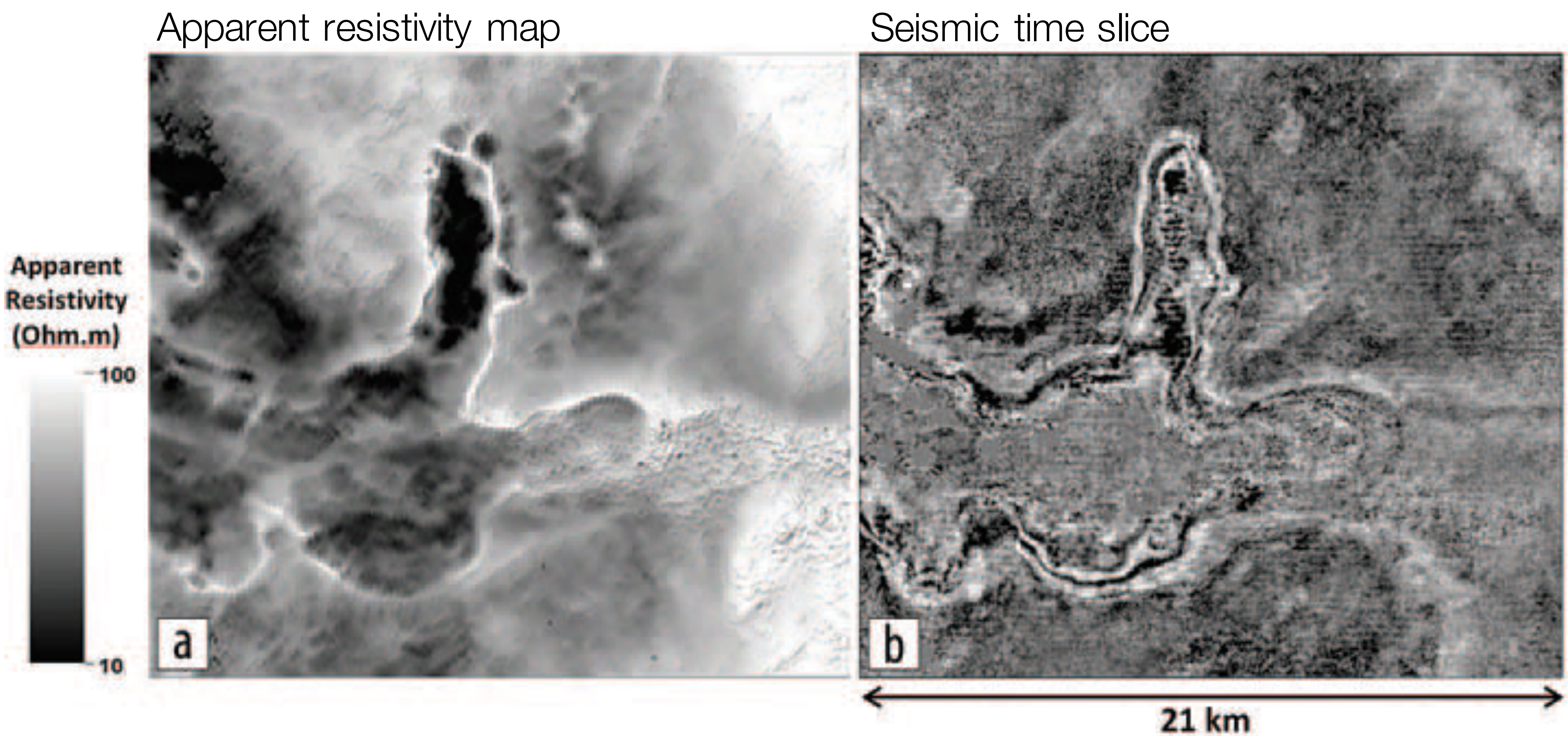


EM data

Apparent resistivity map

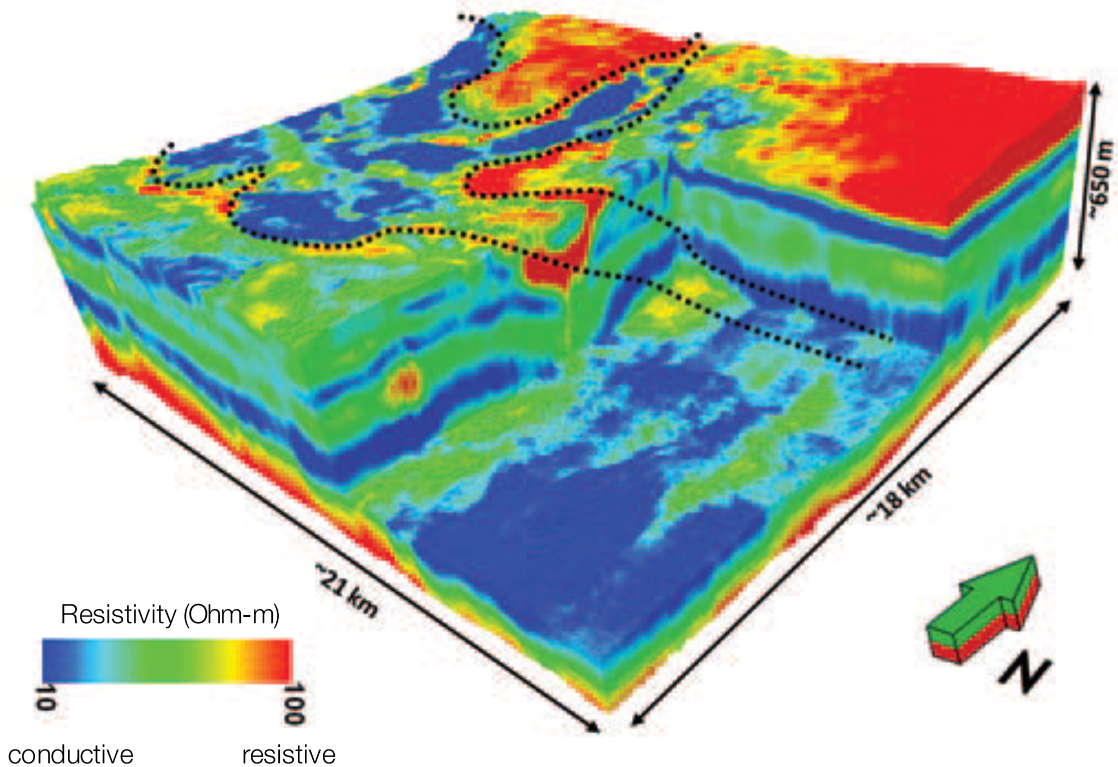


Comparison: EM and Seismic data



Processing: EM inversion

Conductivity model



- 1D inversion for each sounding location
- Lateral constraint is used

Cooperative inversion: Seismic + EM

- How **EM** can help seismic tomography inversion?

Velocity (v_p): high to low (significant challenge)

Conductivity ($\boldsymbol{\sigma}$): high to low

$$v_p = g(\phi)$$
$$\sigma = f(\phi)$$

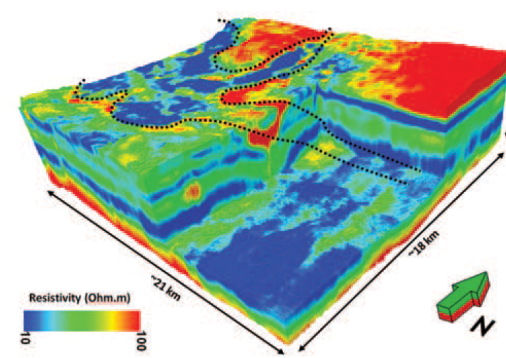
ϕ: porosity

$\mathbf{m}_s$: Slowness
$\mathbf{m}_\sigma$: Conductivity

$$\psi(\mathbf{m_s}, \mathbf{m}_\sigma) = \psi_m(\mathbf{m}_s) + \frac{1}{\lambda_1}\psi_d(\mathbf{m}_s) + \underline{\frac{1}{\lambda_2}\psi_x(\mathbf{m}_s, \mathbf{m}_\sigma)} + \frac{1}{\lambda_3}\psi_{rp}(\mathbf{m}_s, \mathbf{m}_\sigma)$$

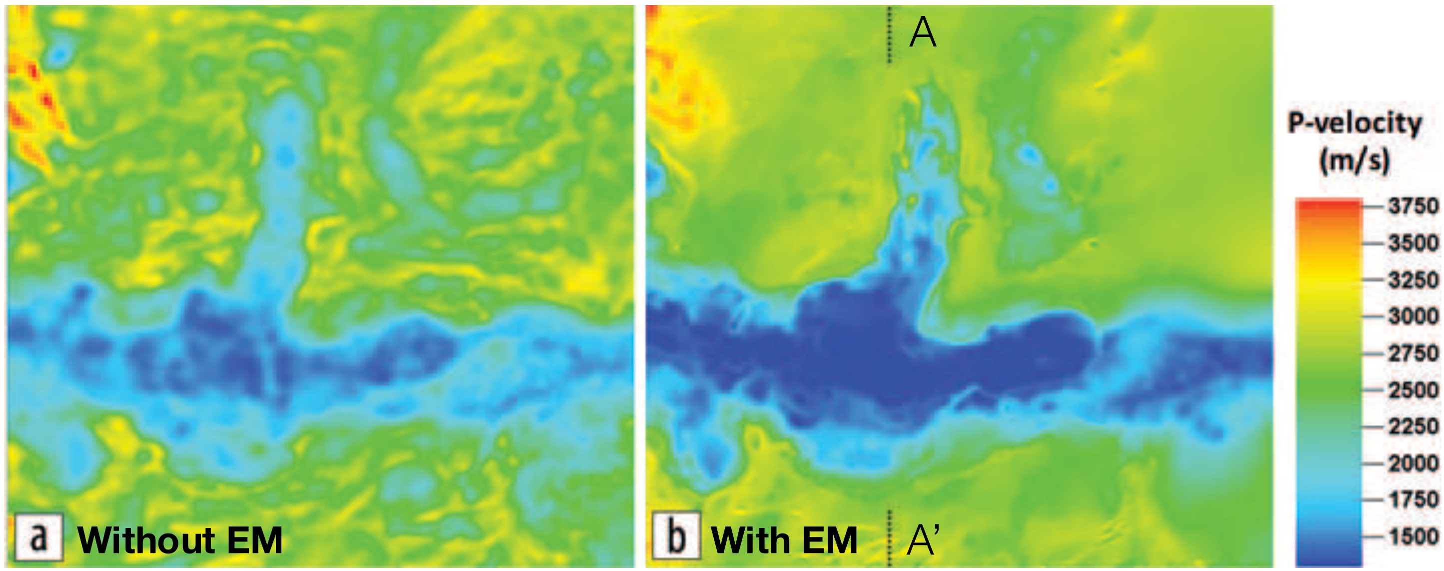
$$||\nabla \mathbf{m_s} \times \nabla \mathbf{m}_\sigma||_2^2$$

Gallardo and Meju, 2004



Cooperative inversion: Seismic + EM

V_p depth slices at 340 m below sea level

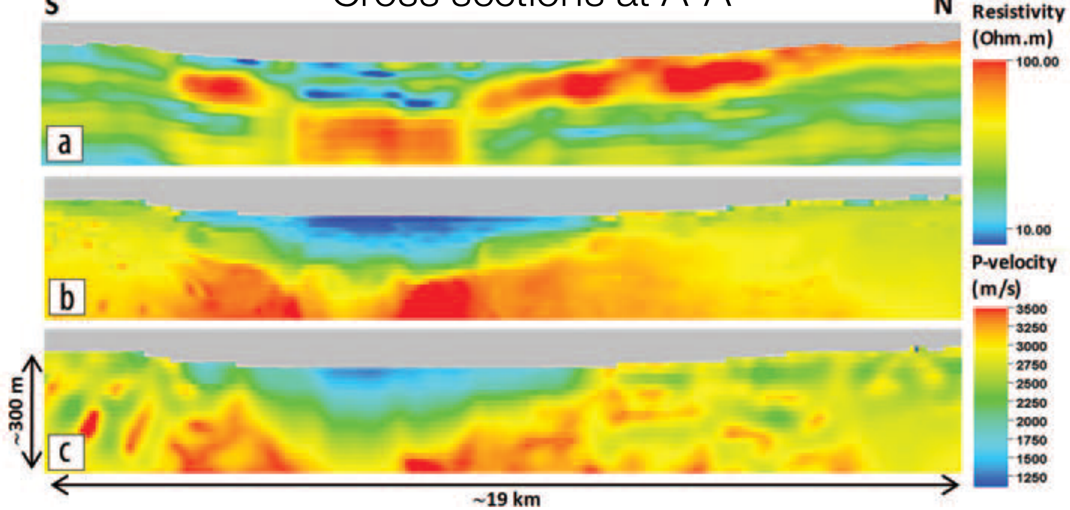


Cross sections at A-A'

Conductivity

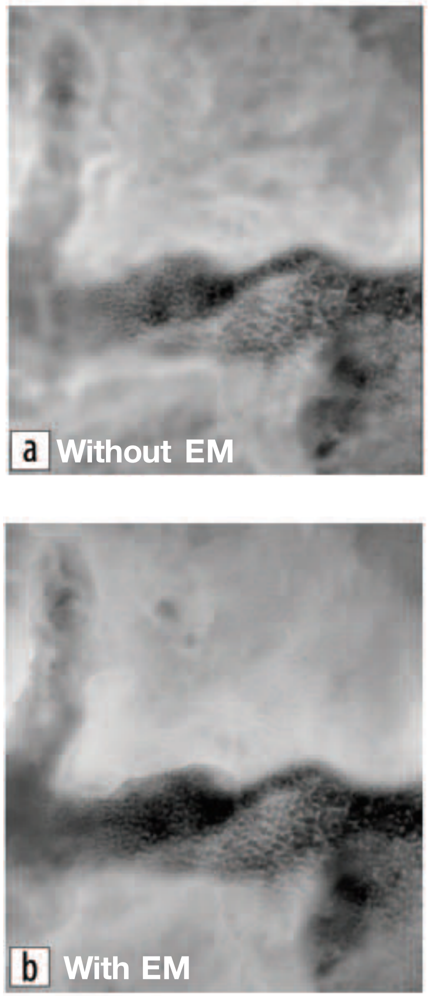
v_p **with EM**

v_p **without EM**

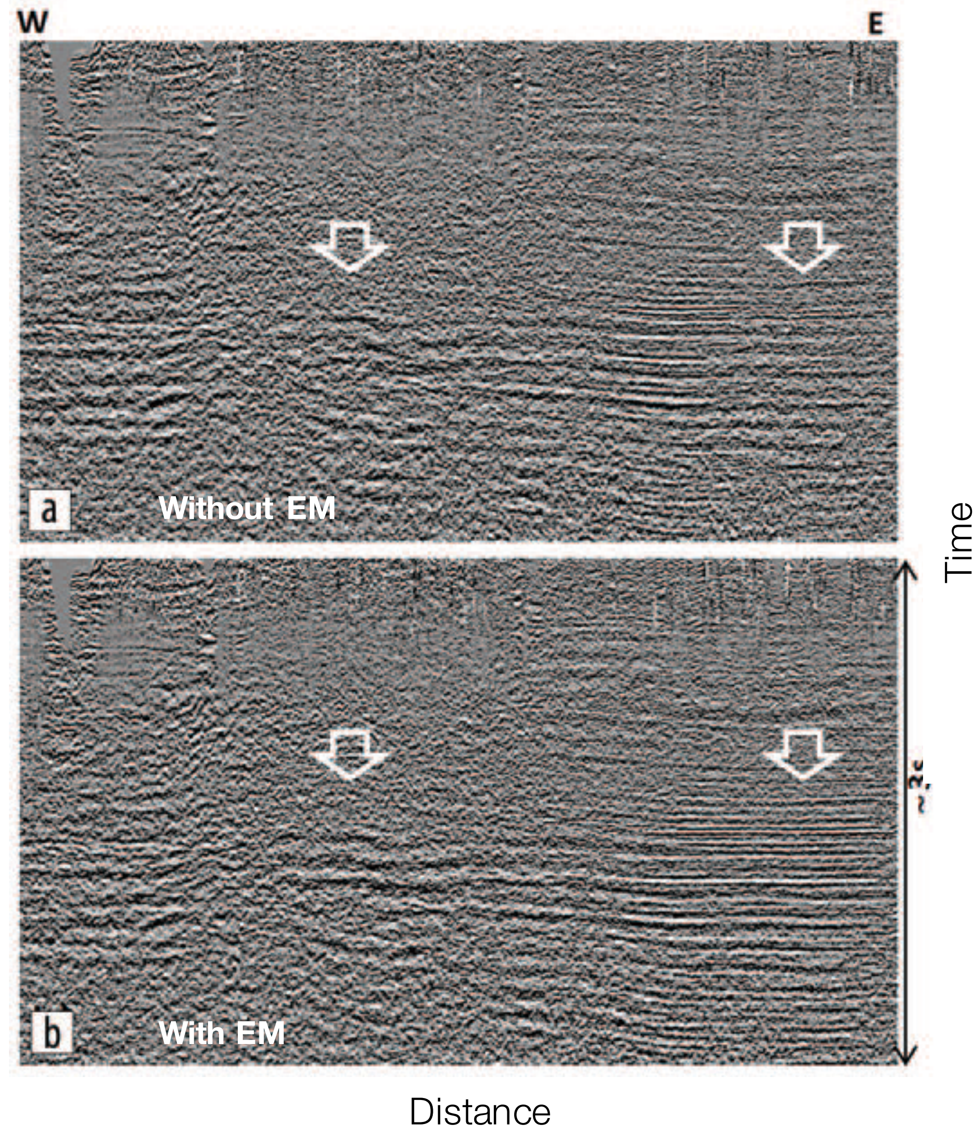


Static correction

Estimated statics on plan map



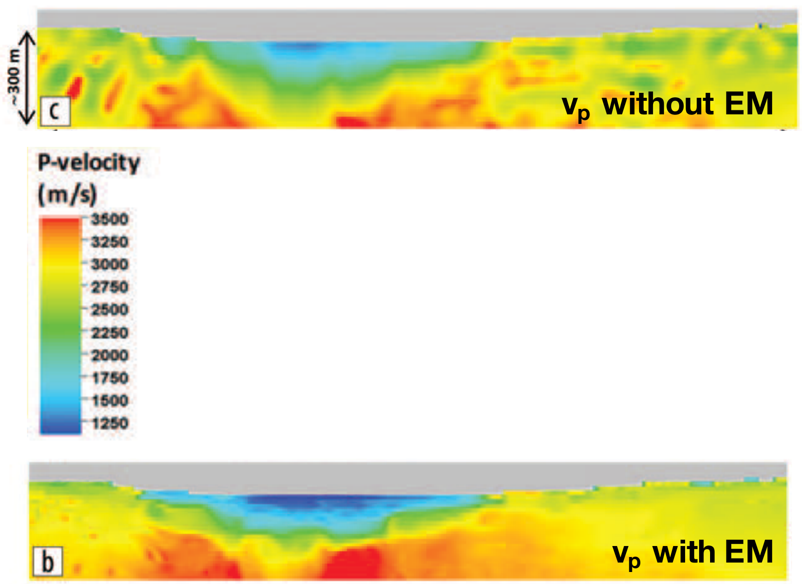
Static corrected sections



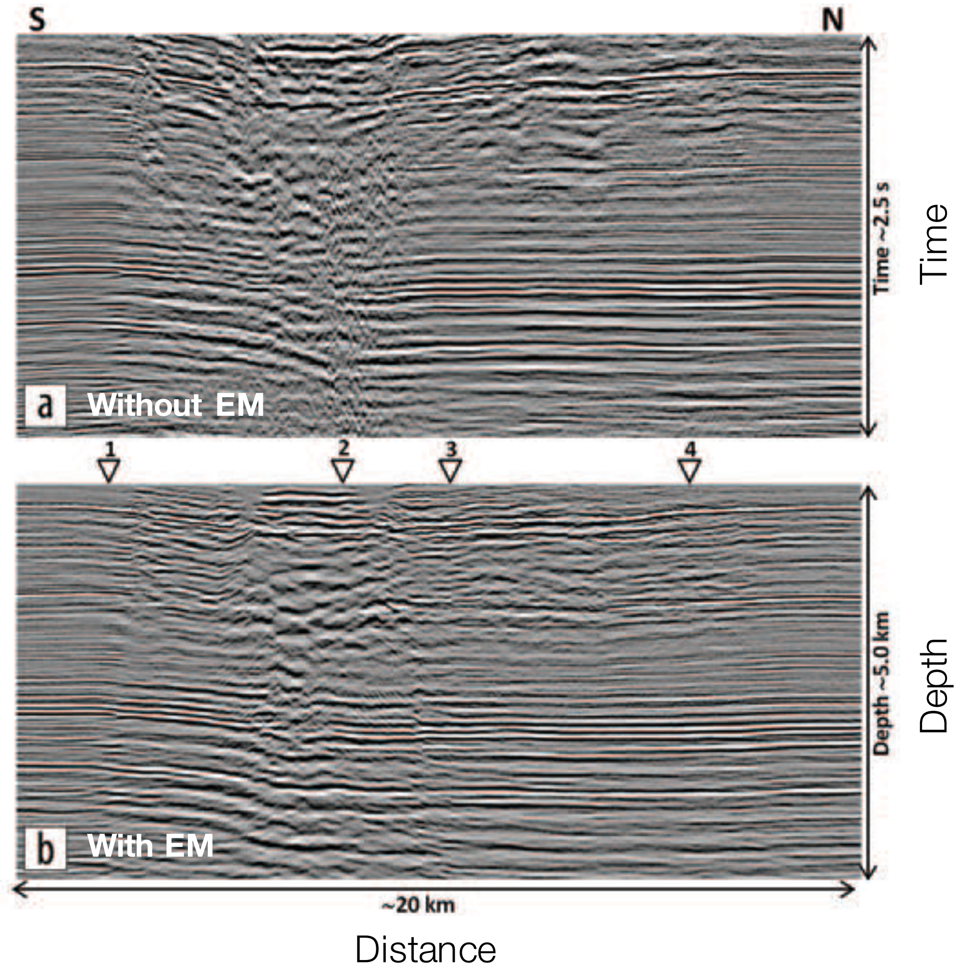
Pre-stack depth migration

- Impact of the improved v_p model to a pre-stack depth migration:

v_p cross sections at A-A'



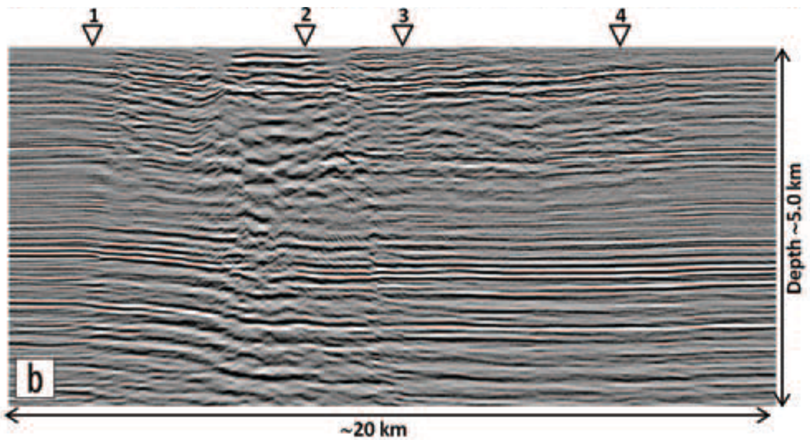
Cross sections at A-A'



Interpretation and Synthesis

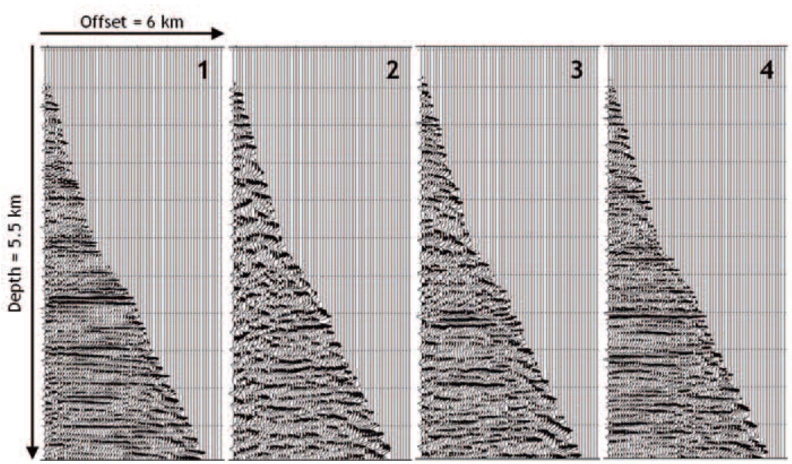
Depth section at A-A'

Depth



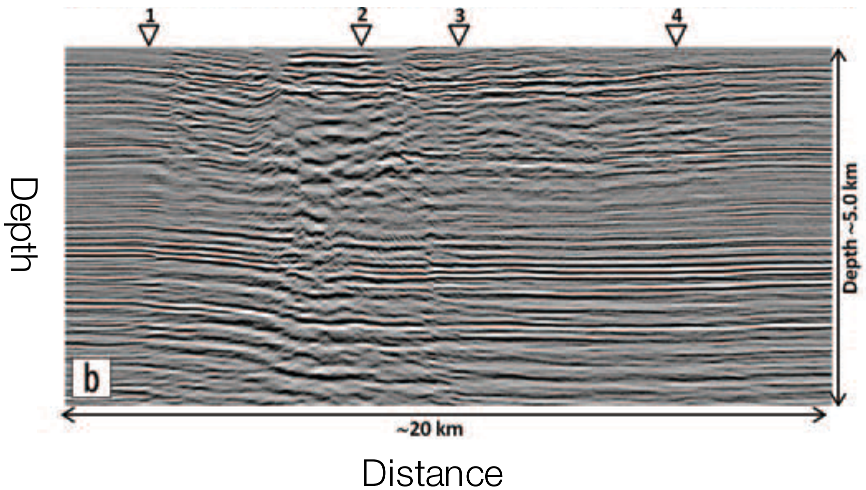
Distance

Common image gathers

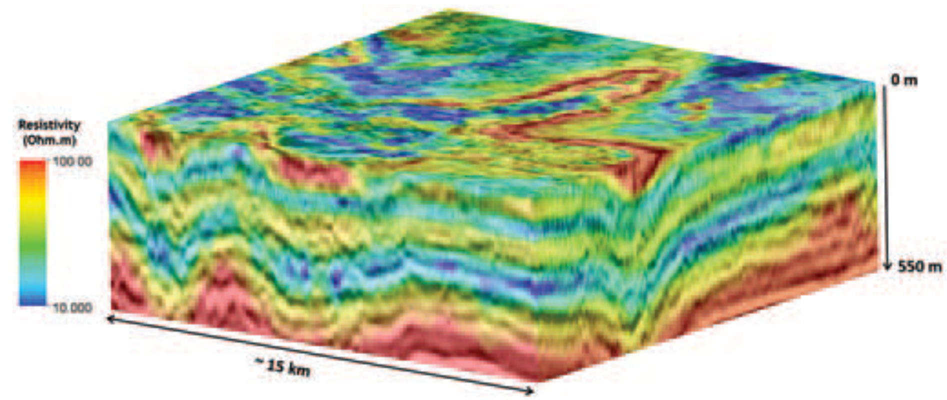


Interpretation and Synthesis

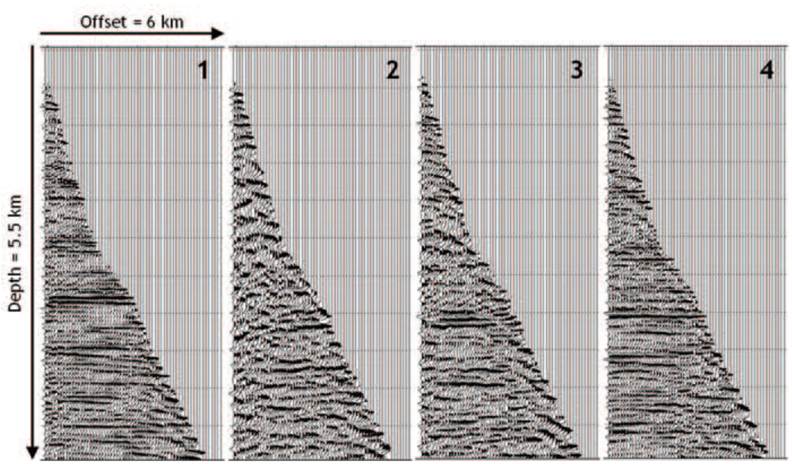
Depth section at A-A'



3D prestack depth migration co-rendered with EM



Common image gathers



- High resolution near surface conductivity from EM improves velocity model
- Helps seismic imaging:
  - Static correction
  - Pre-stack depth migration

Unexploded Ordnance
(UXO)

Unexploded Ordnance (UXO)

Definition: a munition that was armed, fired and remains unexploded

Sources:
- Regions of military conflict
- Munitions/bombing ranges
- Avalanche control

Countries Significantly Impacted by UXOs

http://www.nohowinc.com/

http://www.centennialofflight.gov

http://www.dma.state.mn.us/

Various Types of UXO

- Landmines
- Bombs
- Bombies (from cluster bombs)
- Rocket-propelled grenades (RPG)
- Hand-held grenades
- Mortars

How do we find UXO?

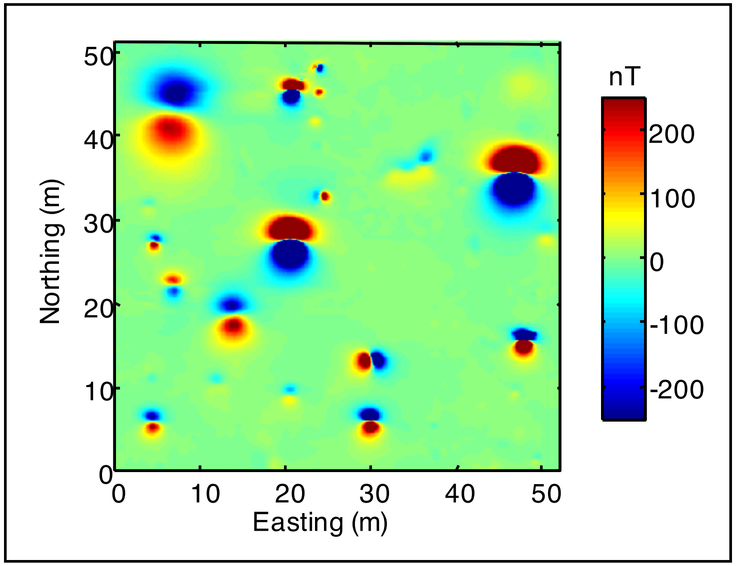
Magnetic Surveys: Locate Anomalies

- Analogue data
- Flag anomaly locations

Ferrex

- Digital data
- Look for magnetic dipoles

TM4



Magnetic Survey: Dig Anomalies

Digital UXO Location and Classification

Problem

- Most anomalies are not UXO
- Digging every anomaly is expensive

Goal

- Classify anomalies
- Dig only UXOs

Strategy

- Need more information than provided by magnetics
- UXO: composed of steel
  - conductive and magnetic

➡ Use electromagnetics

Fundamental Physics: EM Survey

- Controlled source generates primary magnetic field
- Primary field induces eddy currents within UXO
- Eddy currents decay over time
- Eddy current produce a secondary field which decays over time

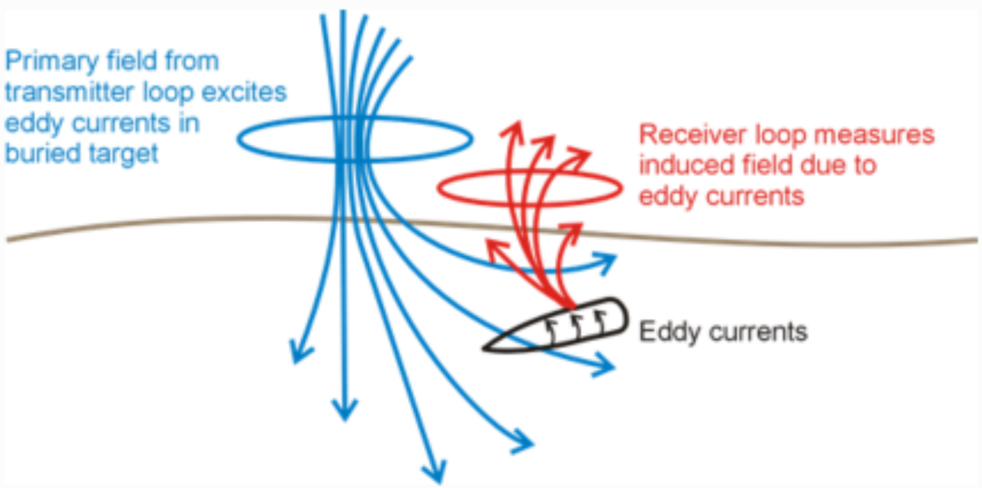
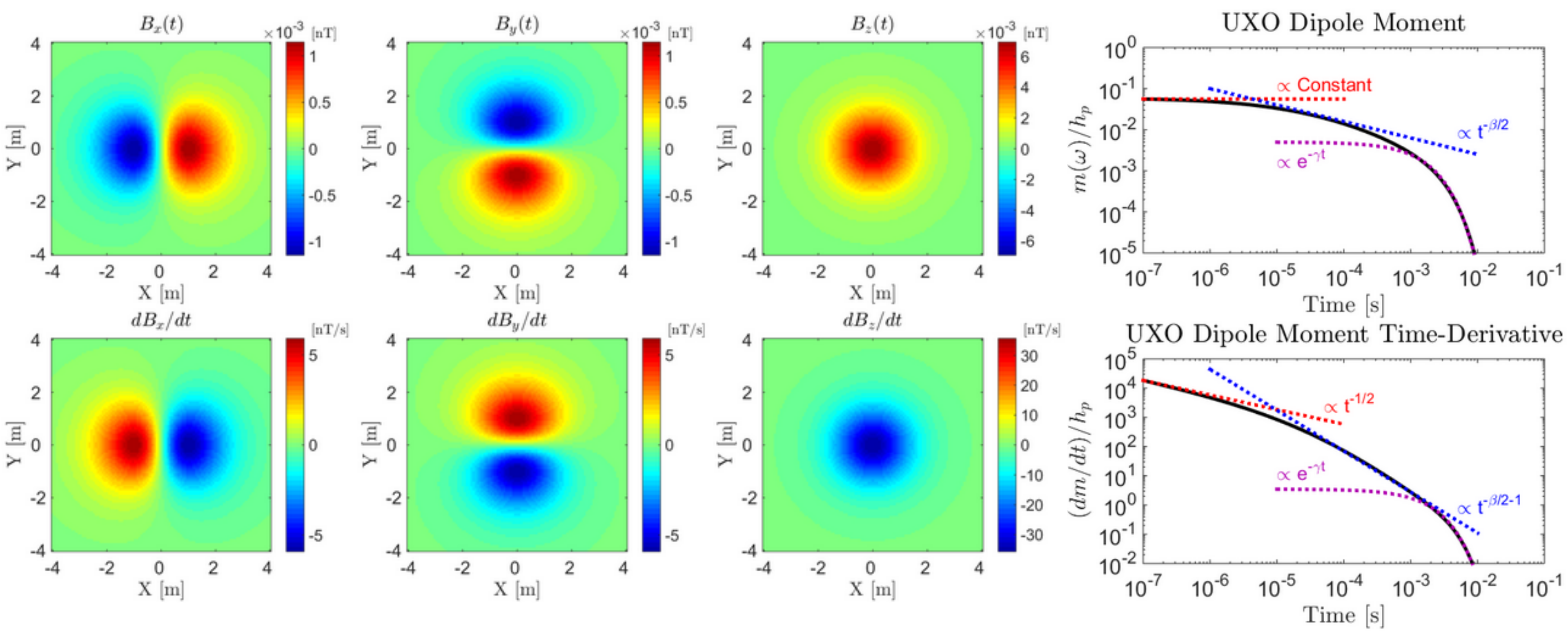


Fig. 260 Electromagnetic induction (EMI) survey for UXO location.

Fundamental Physics: EM Survey

- UXO responses modeled as magnetic dipoles
- Dipoles decay with time
- Rate of decay is indicative of the type of object
- UXOs have characteristic early, mid and late-time decay behaviours



Dipole Model and Polarization Tensor

- UXO response modeled as dipole:

$$\mathbf{b_s}(t) = \frac{\mu_0}{4\pi}\left[\frac{3\mathbf{r}\left[\mathbf{r}\cdot\mathbf{m}(t)\right]}{r^5} - \frac{\mathbf{m}(t)}{r^3}\right]$$

- m(t) is dipole moment (decays with time)

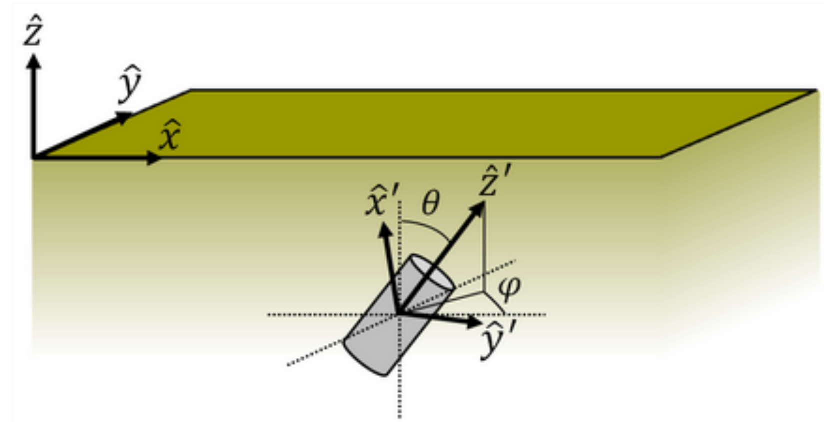
- m(t) depends on:
  1. Orientation of the inducing field
  2. The polarization tensor

$$\mathbf{m}(t) = \mathbf{A^T L A h_p}$$

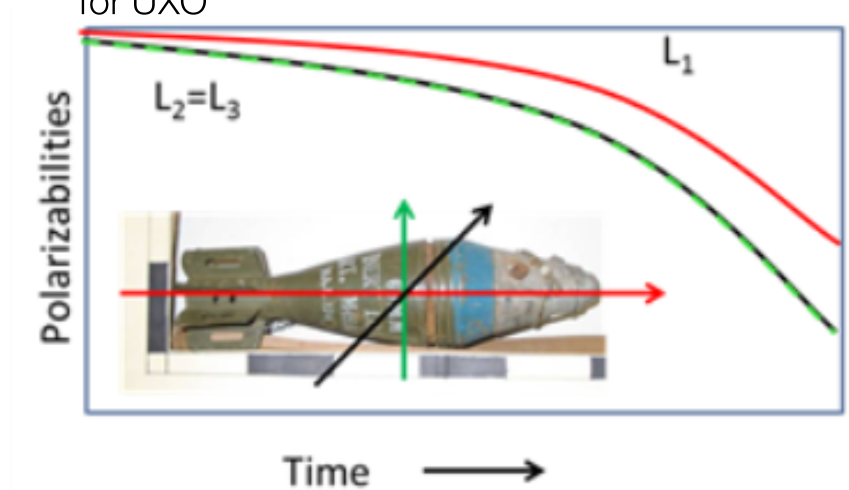
- The polarization tensor L:

$$\mathbf{L}(t) = \begin{bmatrix} L_1(t) & 0 & 0 \\ 0 & L_2(t) & 0 \\ 0 & 0 & L_3(t) \end{bmatrix}$$

Field and UXO coordinate systems



Primary (L1) and secondary (L2,L3) polarizations for UXO



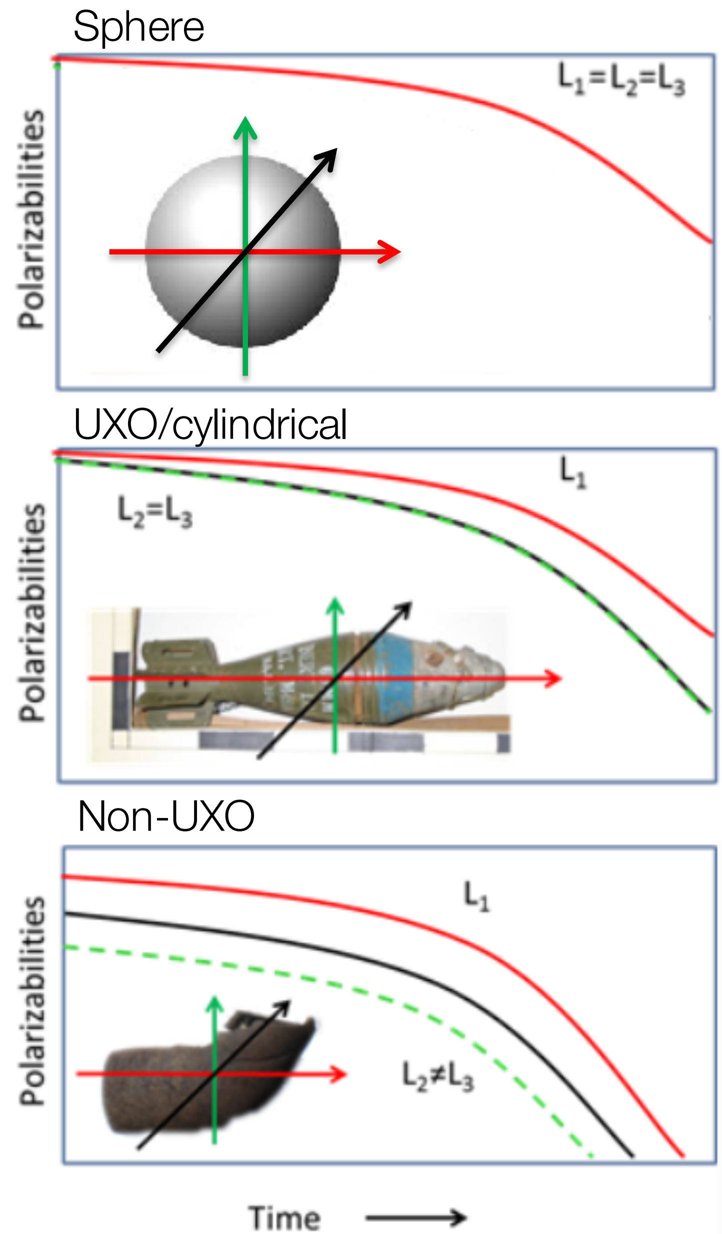
Objects and Polarization Tensors

- Polarization tensor characterizes decay and provides information about dimensionality

- Sphere:
  - Polarization strength independent of primary field direction
  - $\mathrm{L1} = \mathrm{L2} = \mathrm{L3}$

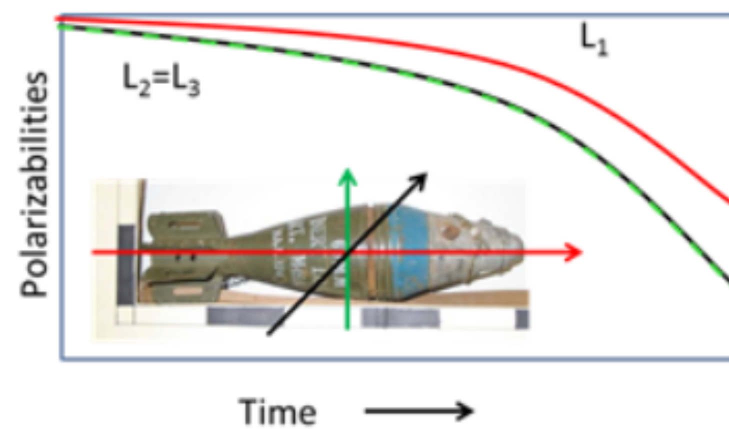
- UXO:
  - Cylindrical in shape
  - Stronger polarization along primary axis
  - $\mathrm{L1} > \mathrm{L2} = \mathrm{L3}$

- Non-UXO:
  - Arbitrary shape
  - Polarization different along different orientations
  - $\mathrm{L1} \neq \mathrm{L2} \neq \mathrm{L3}$



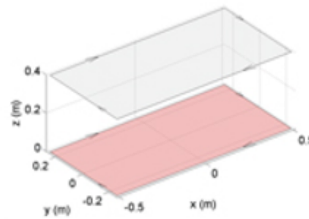
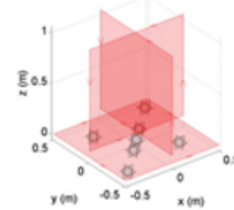
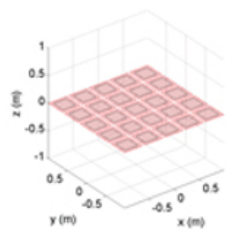
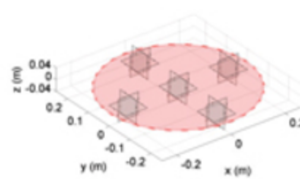
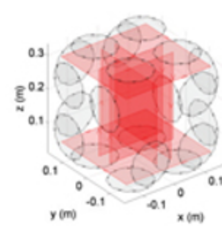
UXO Classification in Practice

- Survey area and pick targets
- Collect high-resolution data over a target
- Recover the elements of the polarization tensor
- Use the polarization tensor to infer information about the object's shape
- Match the recovered polarization tensor to those of object stored in a library to classify



To carry out inversion for polarization tensor need data:
- multiple transmitters (orientations)
- multiple components of data

Common Systems

	Sensor	Geometry	Time channels
EM-61	EM-61 		$t_{min} = 0.2\ ms$ $t_{max} = 1.5\ ms$ $N = 4$
MetalMapper	MetalMapper 		$t_{min} = 0.1\ ms$ $t_{max} = 10\ ms$ $N = 42$
TEMTADS	TEMTADS 		$t_{min} = 0.1\ ms$ $t_{max} = 20\ ms$ $N = 115$
MPV	MPV 		$t_{min} = 0.1\ ms$ $t_{max} = 20\ ms$ $N = 32$
BUD	BUD 		$t_{min} = 0.1\ \text{ms}$ $t_{max} = 1.5\ ms$ $N = 45$

Survey Design

Line and Station Spacing:

- Depends on dimensions and depth of targets and system being used.
- Insufficient sampling makes locating and classifying targets more challenging.

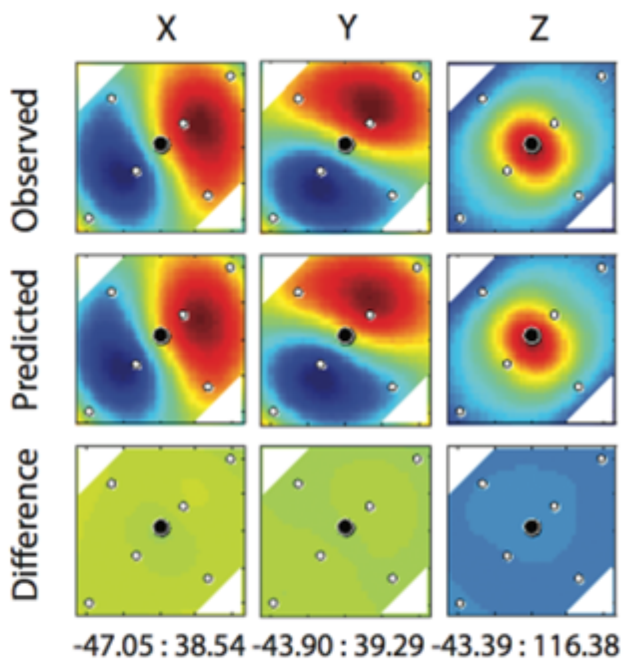
Excitation Orientation

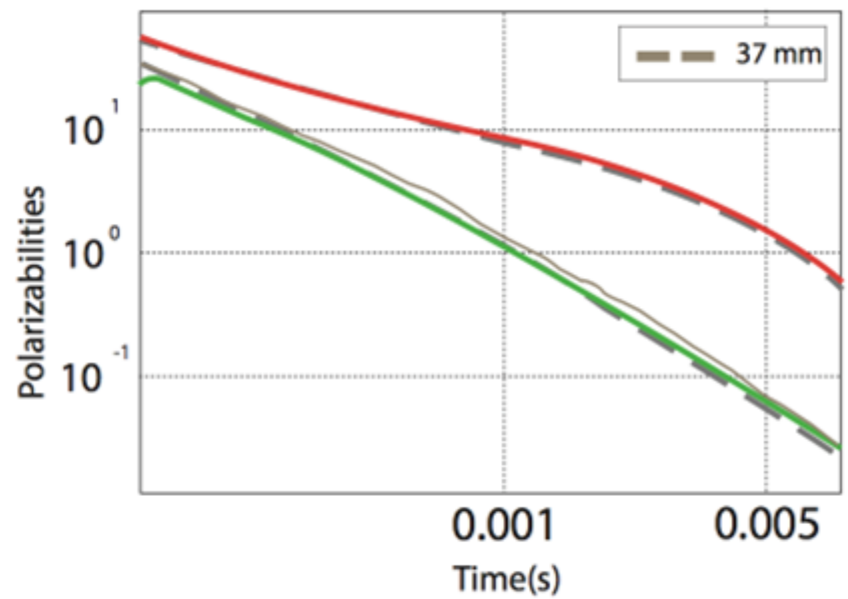
- To recover polarization tensor, target must be polarized from as many angles as possible.
- May require multiple passes with single transmitter or use of multi-transmitter system.

Time Channels

- Sufficient time-channels required to characterize decay behaviour.

Example: Metal Mapper Data

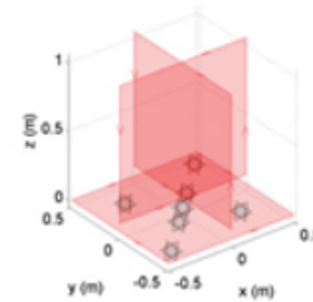


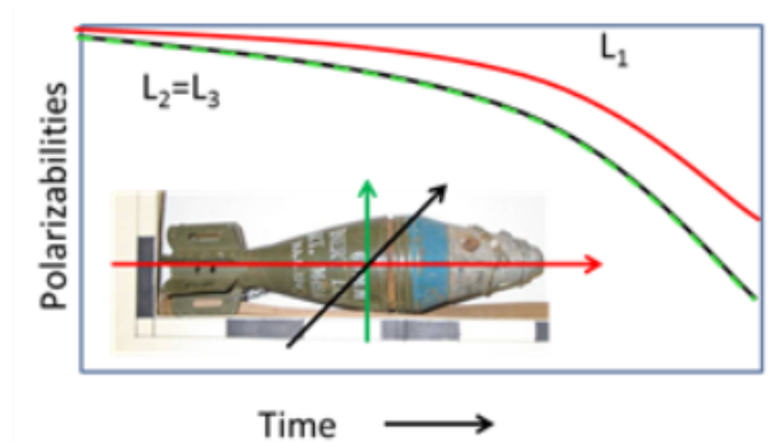


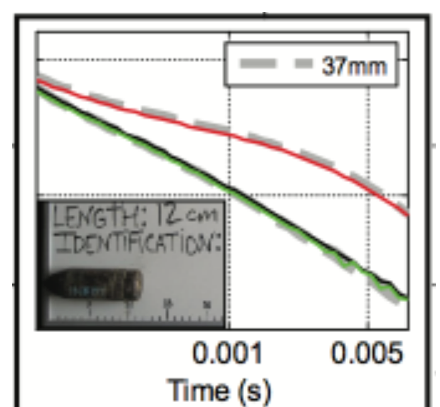
- Polarizations indicate a cylindrical object
- Predicted data using recovered polarization tensor fits the observed data
- Recovered polarizations match those of a 37 mm projectile

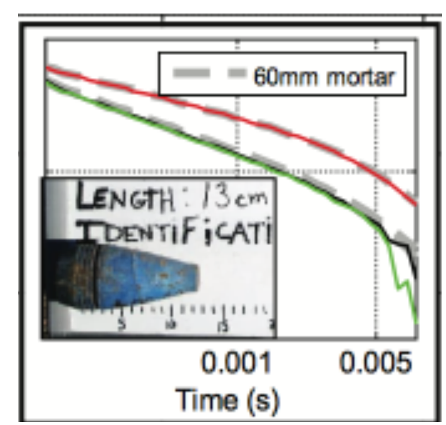
Summary

- UXO are compact conductive permeable objects
- EM is ideal survey
- Requires multiple transmitters and receivers
- Processing yields polarization curves
- Discrimination









Field Example: Pole Mountain

History

- Periods of military use 1897-1961
- Many types of munitions (explosive projectiles, mortars, small arms)
- Land reclamation currently not possible

Goals:

- Test classification algorithm on different objects
- Determine dig/no dig list for targets

Location of Pole Mt., Wyoming, US

Field Example: Pole Mountain

EM61-MK2:

- Efficient over rugged terrain
- Single Tx and Rx loops
- Located 2,368 anomalies

Metal Mapper:

- Multiple Tx and Rx loops
- Cued interrogation data over anomalies
- Data used for classification and prioritize dig list

EM61-MK2 (locate anomalies)

Metal Mapper (cued interrogation)

Field example: Pole Mountain

- All 2,368 TEM anomalies were dug to verify
- 1,829 correctly identified as clutter or assigned to no dig through classification
- Only 453 non-munition items dug before all 160 munition items dug.
- 99% of munition items located within first ~300 digs
- Correctly identified all types of munititons.