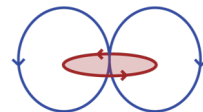
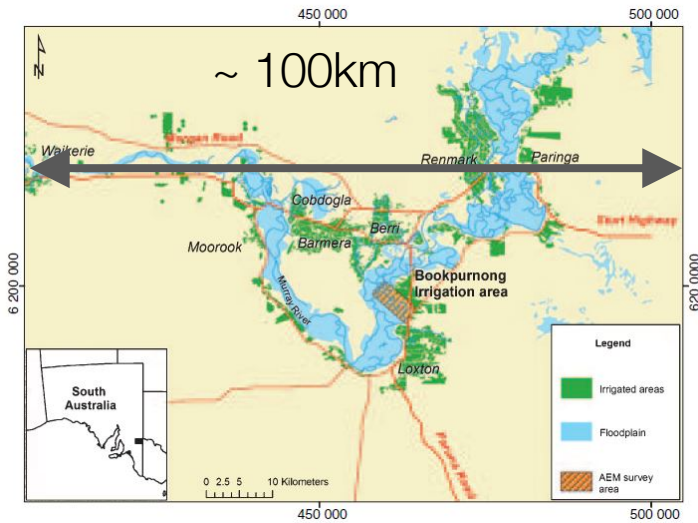


EM: Inductive Sources

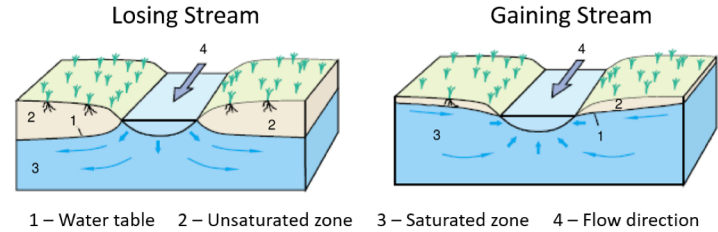


Motivation

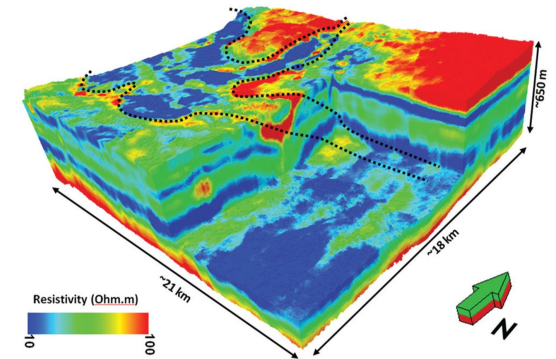
Large areas to be covered



Groundwater



High resolution near surface



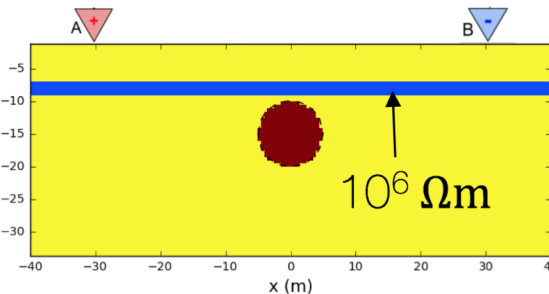
Rugged terrain



Minerals



Shielding problem



Outline

Setup

- Basic experiment
- Transmitters, Receivers

Time Domain EM

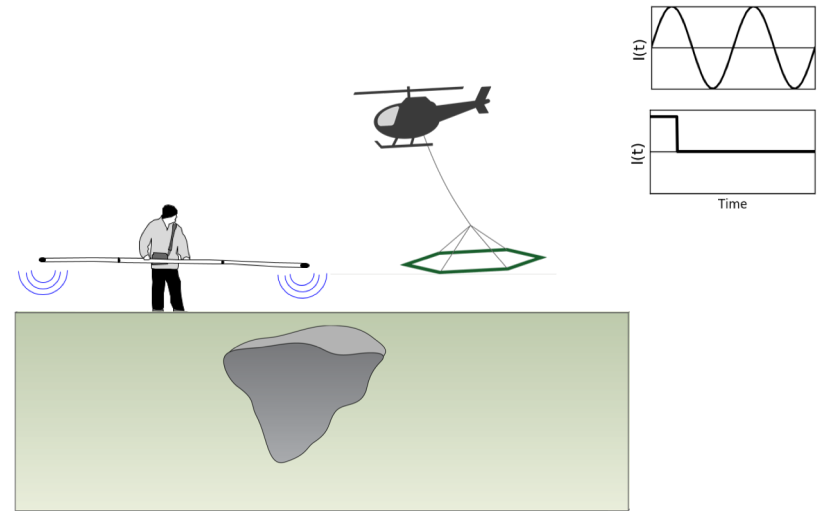
- Vertical Magnetic Dipole
- Propagation with Time
- Case History – Groundwater, Minerals, Hydrocarbons
- Horizontal magnetic dipole

Frequency Domain EM

- Vertical Magnetic Dipole
- Effects of Frequency
- Case History – Groundwater, Minerals

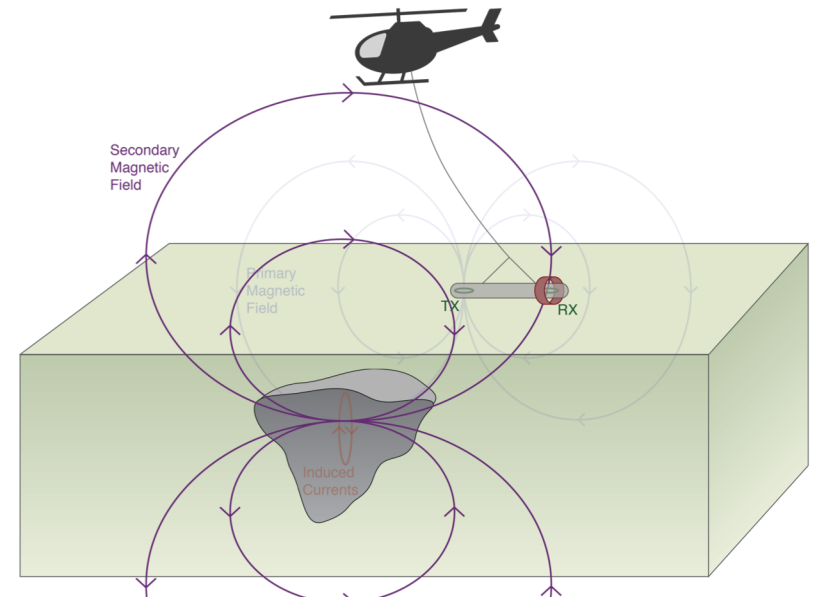
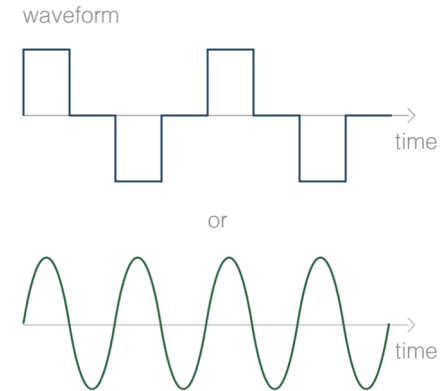
Important questions

- What is the target?
 - at the surface? At depth?. 1D, 2D, 3D?
- Transmitter
 - Location: surface? in the air?
 - Waveform: frequency or time?
 - “Size” and orientation?
- Exciting the target
 - Conductivity of the target and host
 - Geometry of the target (Coupling)
- Receiver and data
 - What fields to measure?
 - What instrument?
- Where to collect data? How many? How accurate?
- What is depth of investigation?
- What is the “footprint” of the transmitter”
 - These are questions of **SURVEY DESIGN**



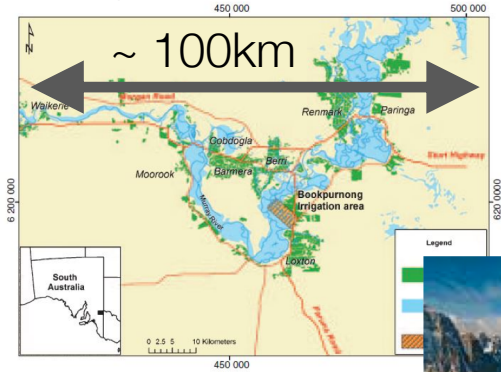
Basic Experiment

- **Transmitter:**
 - Produces a primary magnetic field
- **Exciting the target:**
 - Time varying magnetic fields generate electric fields everywhere
 - Producing currents in conductors
- **Receiver:**
 - Induced currents produce secondary magnetic fields



Transmitter

Large areas



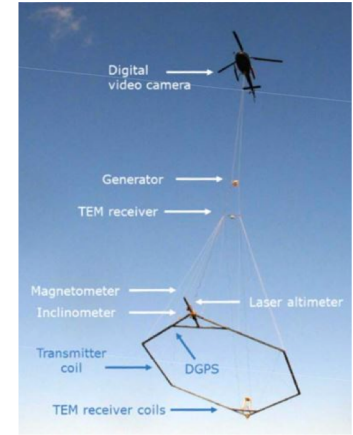
Rugged terrain



Airborne Survey

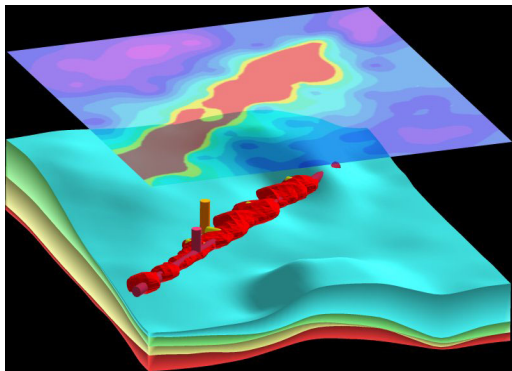


Resolve

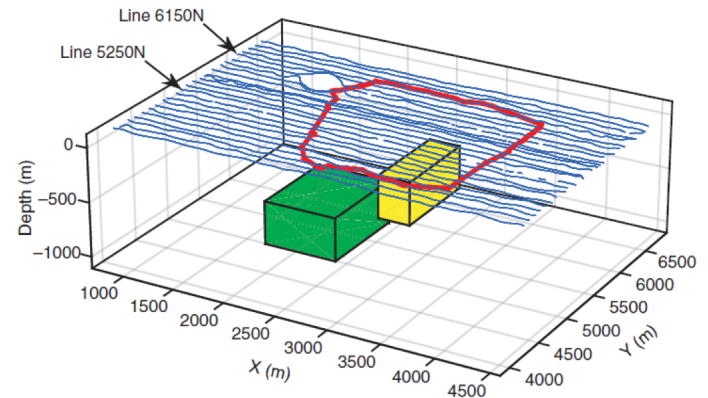


SkyTEM

Deep Targets

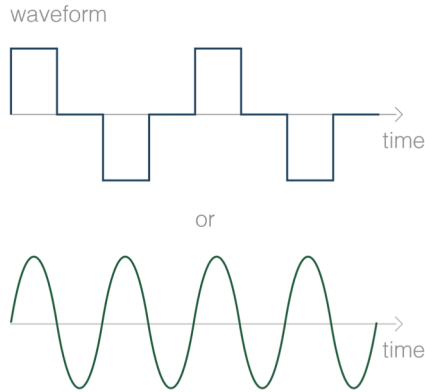


Large Loop



Transmitter

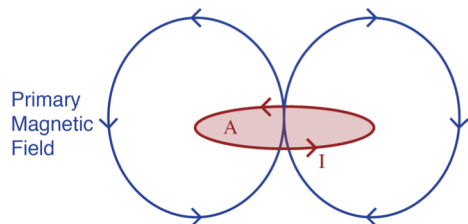
- Time or frequency?



- Key factor is moment

$$m = I \text{ (current)} A \text{ (area)} N \text{ (# of turns)}$$

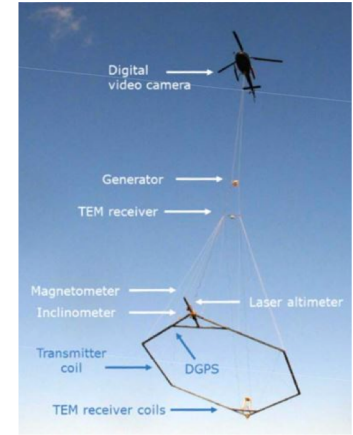
$$\mathbf{B}(\mathbf{r}) = \frac{\mu_0}{4\pi} \left(\frac{3\mathbf{r}(\mathbf{m} \cdot \mathbf{r})}{|\mathbf{r}|^5} - \frac{\mathbf{m}}{|\mathbf{r}|^3} \right)$$



Airborne Survey

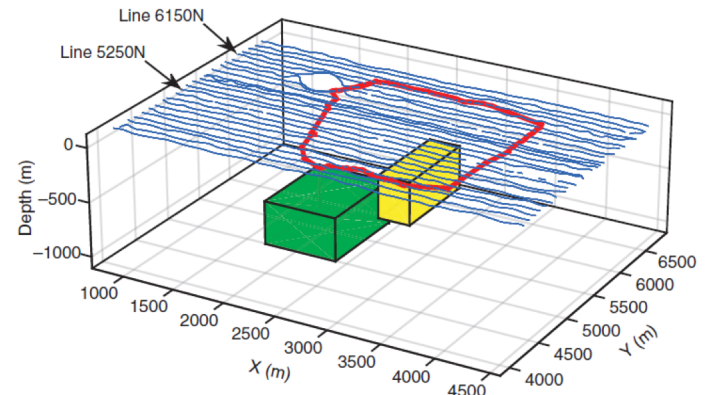


Resolve



SkyTEM

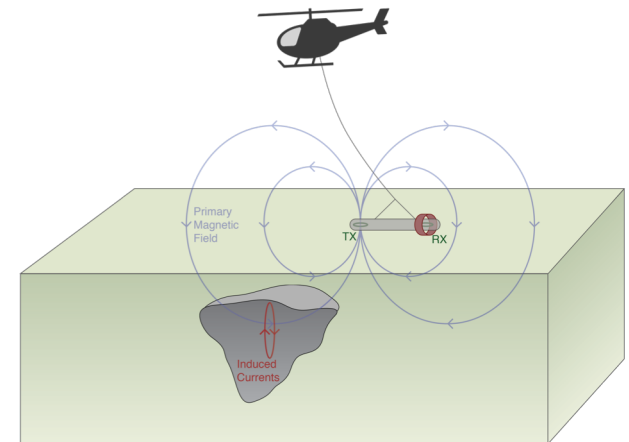
Large Loop



Exciting the target

- Primary field from a loop
- Fields fall off
 - $1/r^3$ geometric decay
 - Attenuation
- Want to be as close as possible to target
 - Ground based systems
 - Helicopter
 - Fixed wing aircraft
- Always concerned about coupling

$$\mathbf{B}(\mathbf{r}) = \frac{\mu_0}{4\pi} \left(\frac{3\mathbf{r}(\mathbf{m} \cdot \mathbf{r})}{|\mathbf{r}|^5} - \frac{\mathbf{m}}{|\mathbf{r}|^3} \right)$$

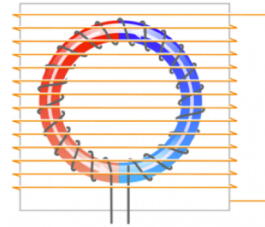


Receiver and Data

Magnetometer

- Measures:
 - Magnetic field
 - 3 components
- eg. 3-component fluxgate

$\mathbf{b}(t)$

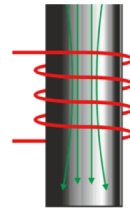


Fluxgate

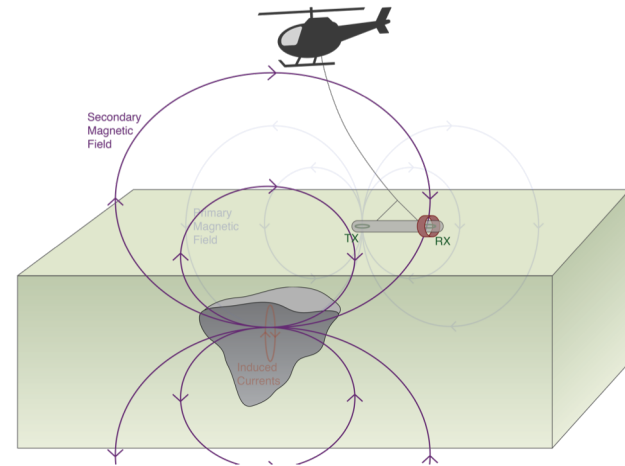
Coil

- Measures:
 - Voltage
 - Single component that depends on coil orientation
 - Coupling matters
- eg. airborne frequency domain.
 - ratio of H_s/H_p is the same as V_s/V_p

$$\frac{\partial b}{\partial t}$$

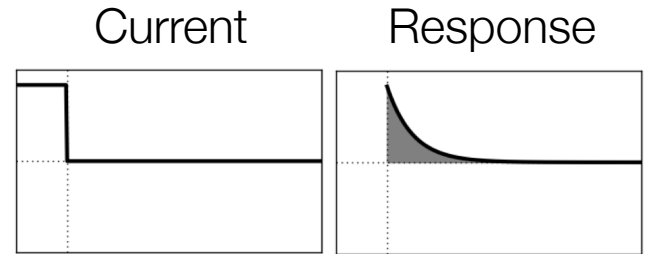


Coil

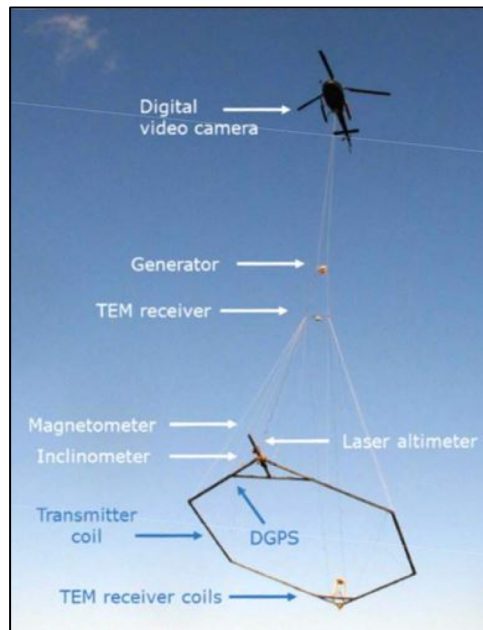


Receiver: Time Domain

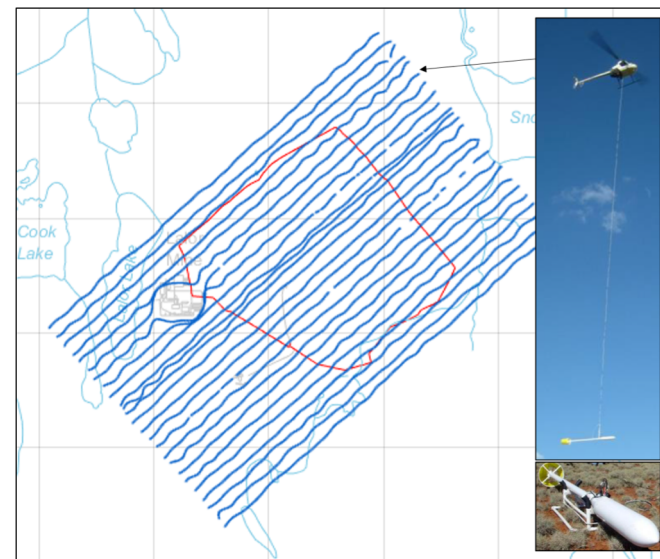
- Primary field has off-time
- Measure secondary fields
- Receivers can be mounted on transmitter loop or above it



SkyTEM

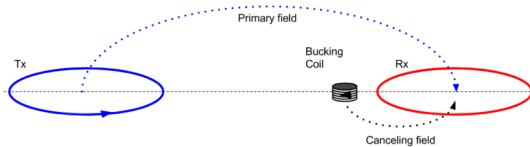


HeliSAM

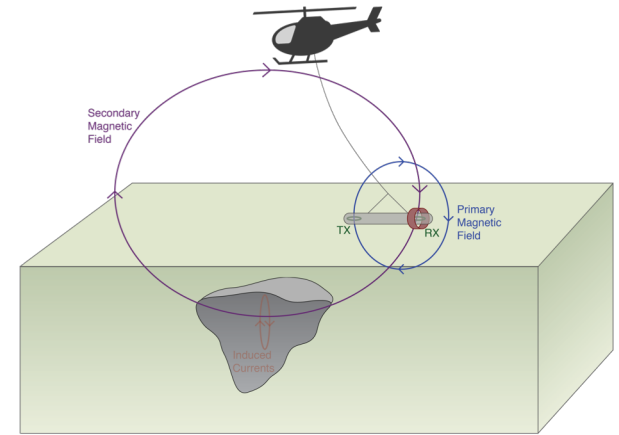


Receiver: Frequency Domain

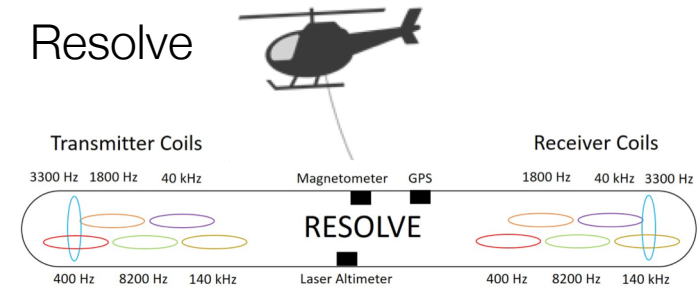
- Primary field
 - always “on”
 - large compared to secondary fields
- Primary removal
 - Compute and subtract
 - Bucking coil



- Main requirement:
 - Know positions of Tx and Rx
 - Keep them in one unit



Resolve

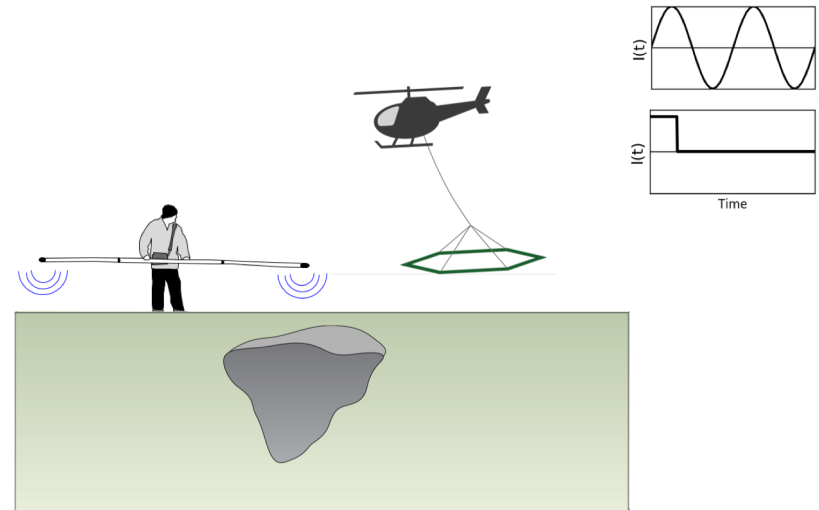


EM-31



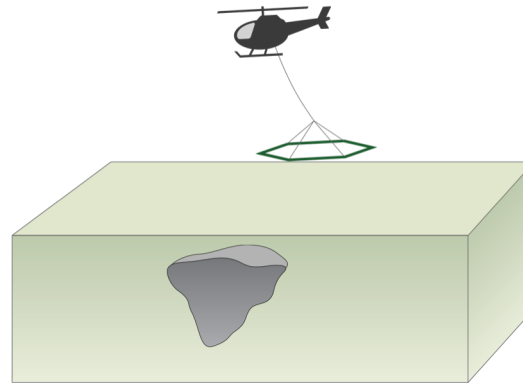
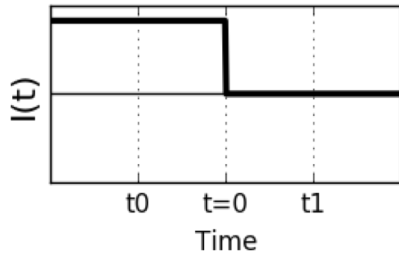
Important questions

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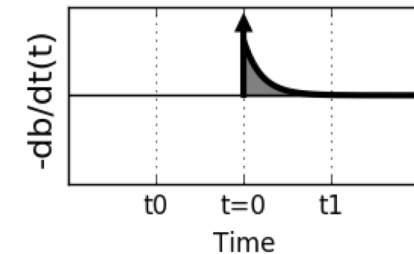
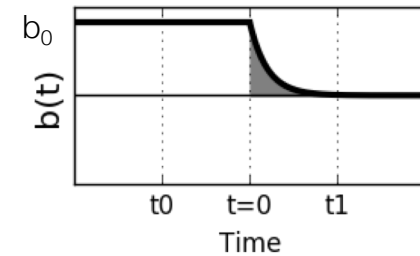


EM with Inductive Sources: Time Domain

Transmitter current



Receiver

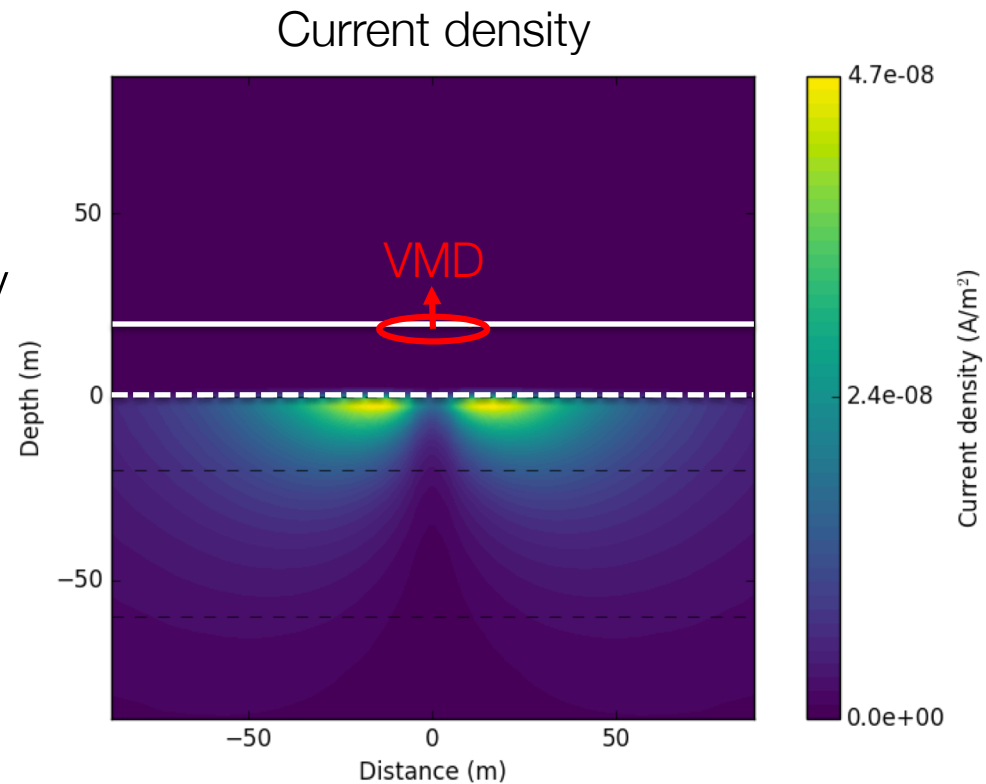


time	b	db/dt
$t < 0$	b_0	0
$t = 0$	b_0	$-b_0\delta(t)$
$t > 0$	secondary	secondary

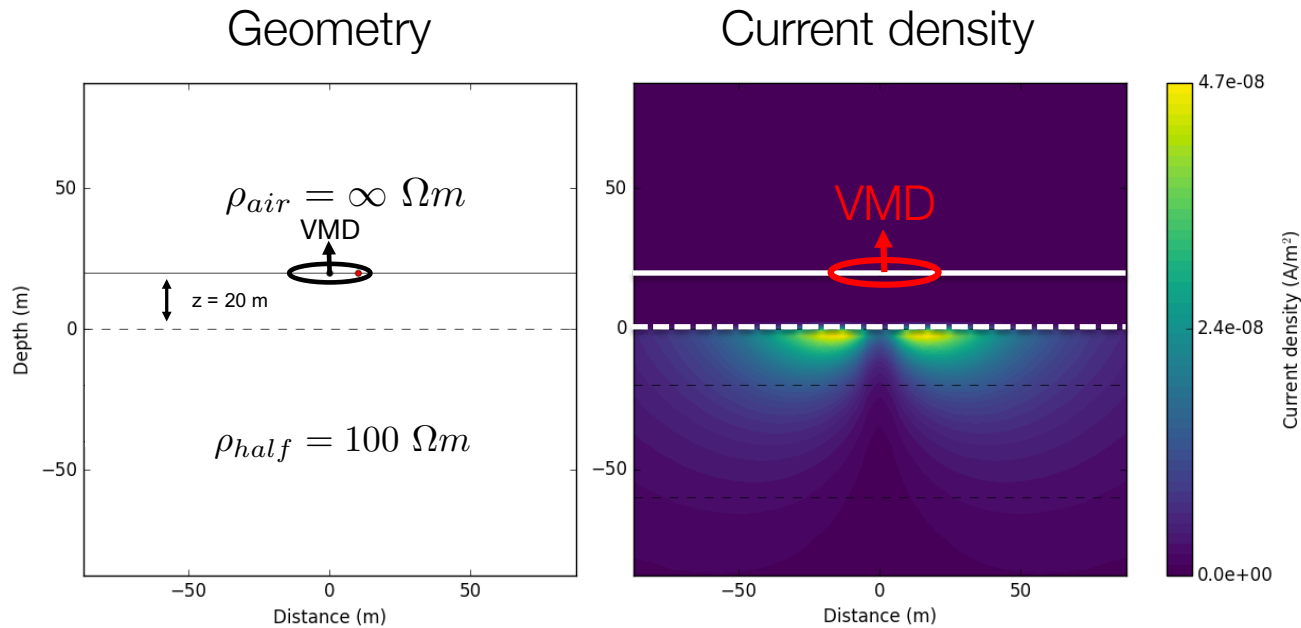
$\delta(t)$: Dirac-delta function

Footprint of Airborne EM system

- What volume of earth is “seen” by the airborne system?
 - Where are the currents?
- Currents depend on
 - Transmitter
 - Waveform: time or frequency
 - Background conductivity
- Simple case: loop source over homogeneous earth

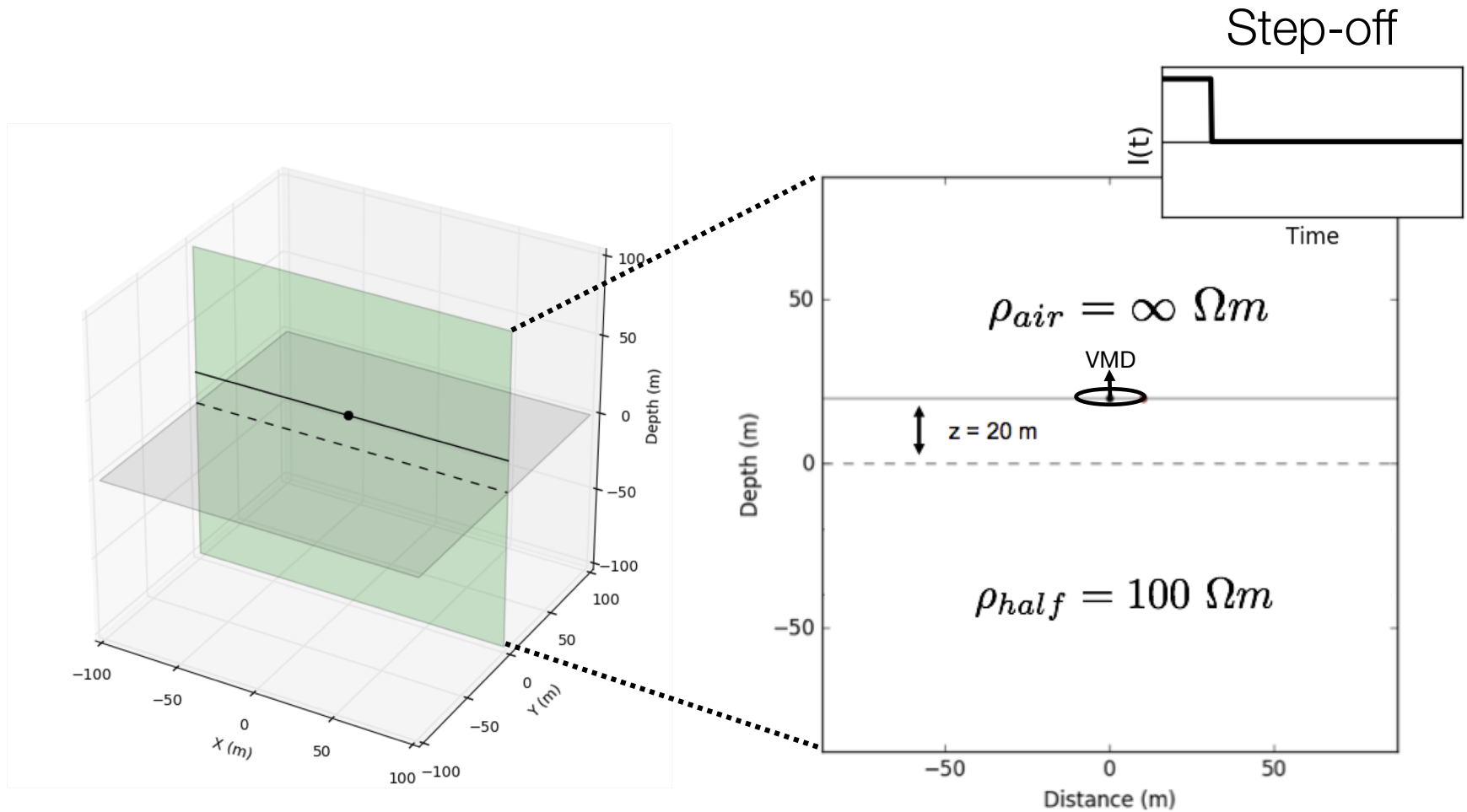


Vertical Magnetic Dipole (VMD)



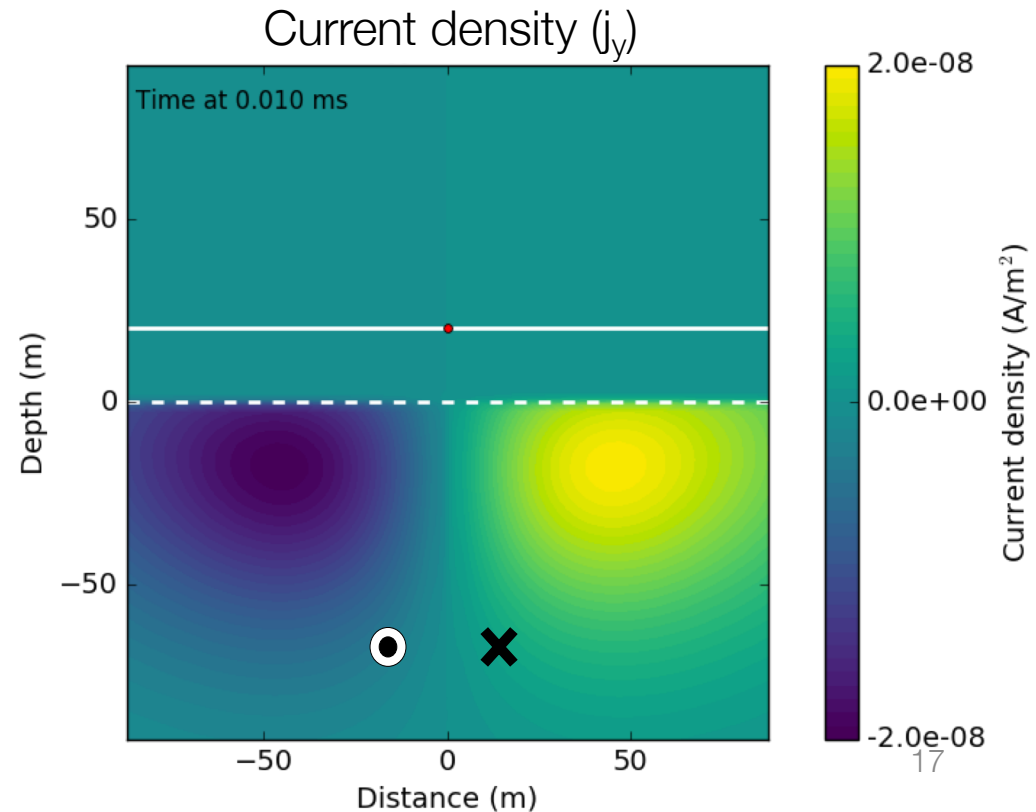
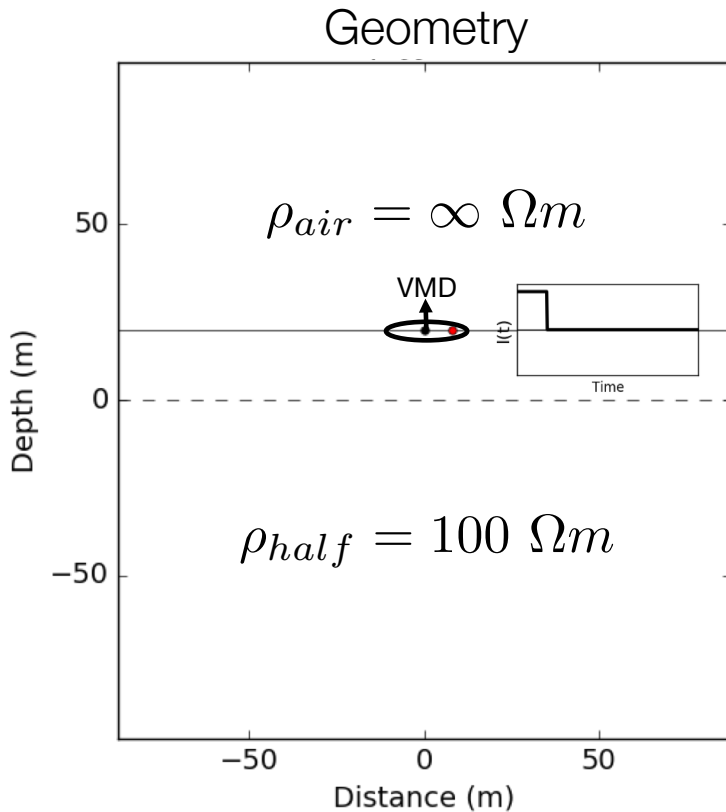
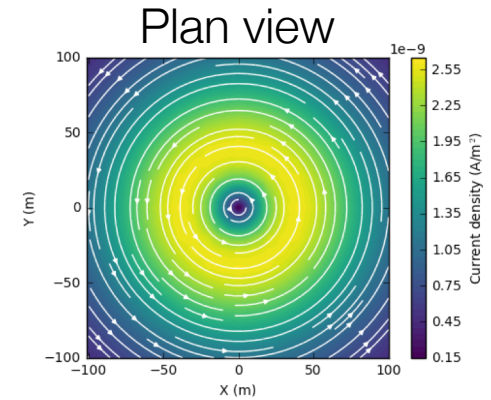
- Some questions
 - Where, and how strong, are the currents?
 - How do they depend upon the conductivity?
 - What do the resulting magnetic fields look like?

Vertical Magnetic Dipole over a halfspace (TDEM)



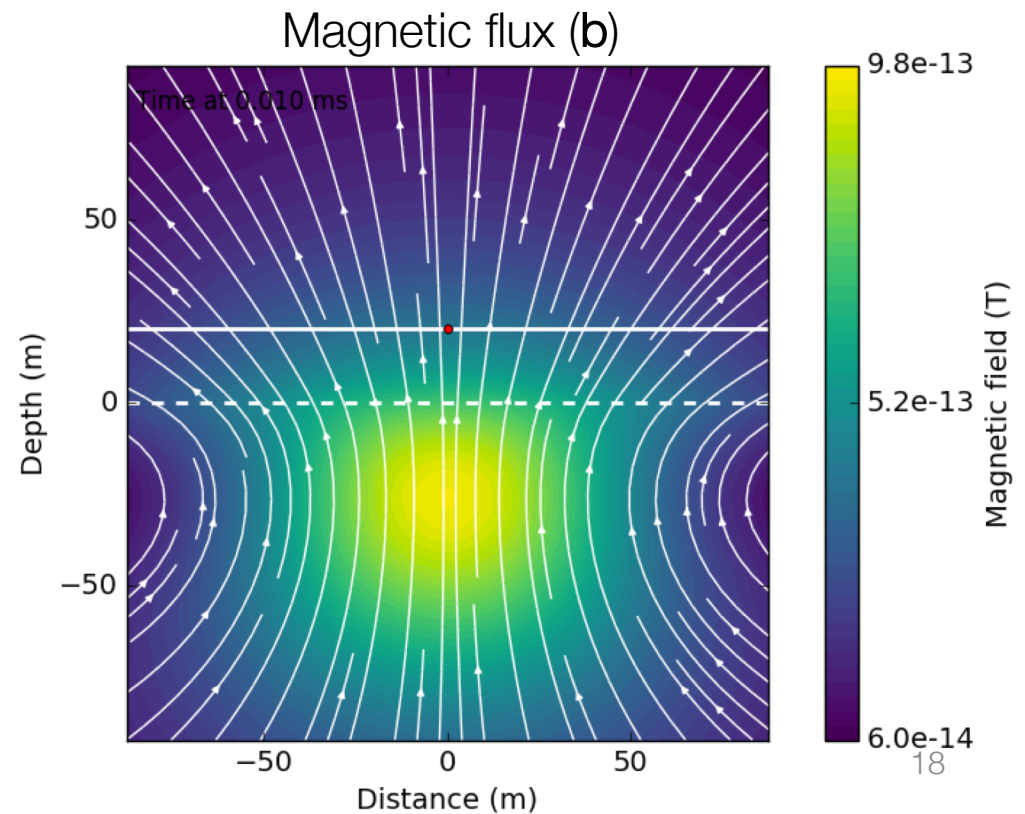
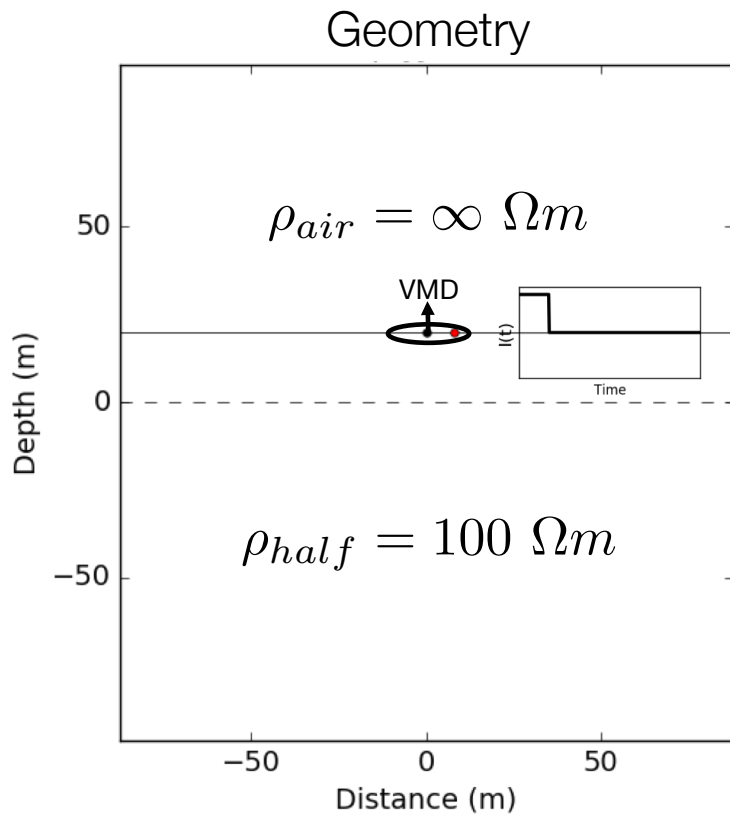
Current Density

- Time: 0.01ms



Magnetic flux density

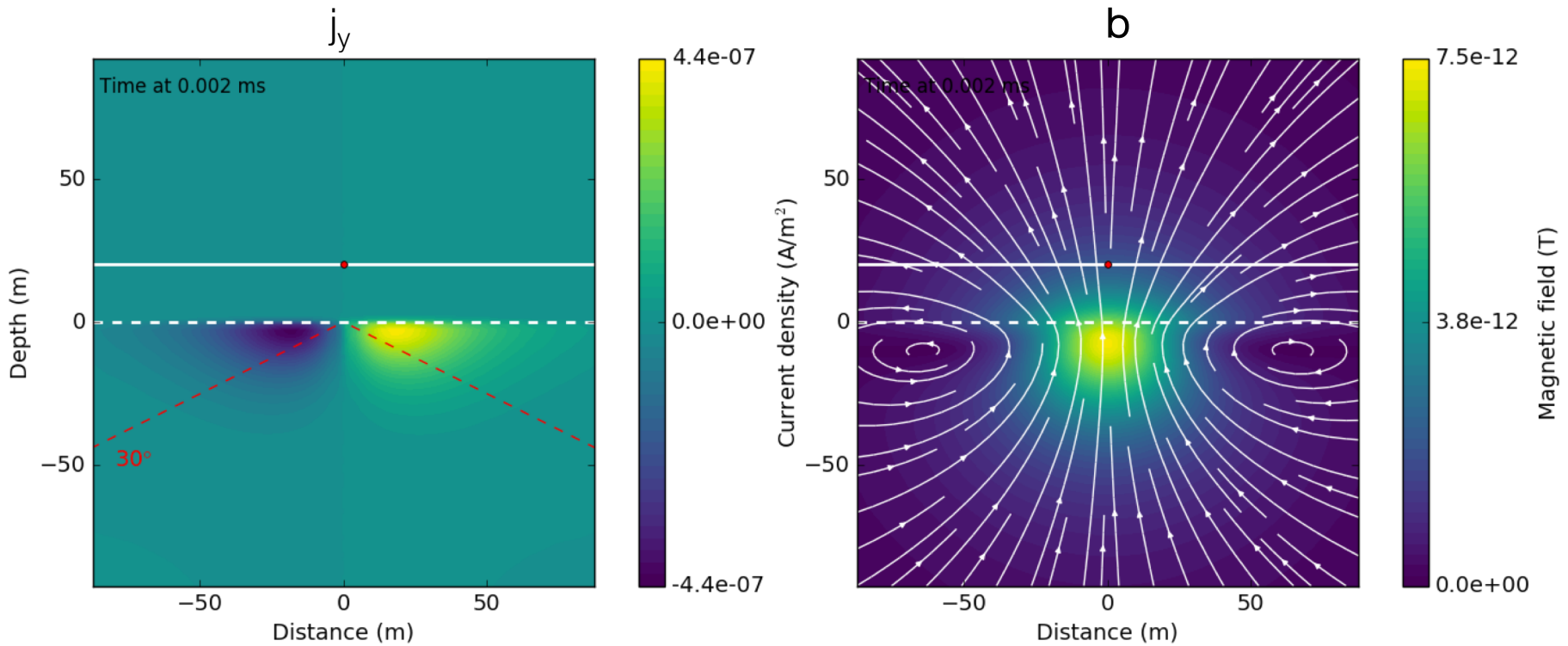
- Time: 0.01ms



Propagation through time

- Time: 0.002ms
- diffusion distance = 18 m

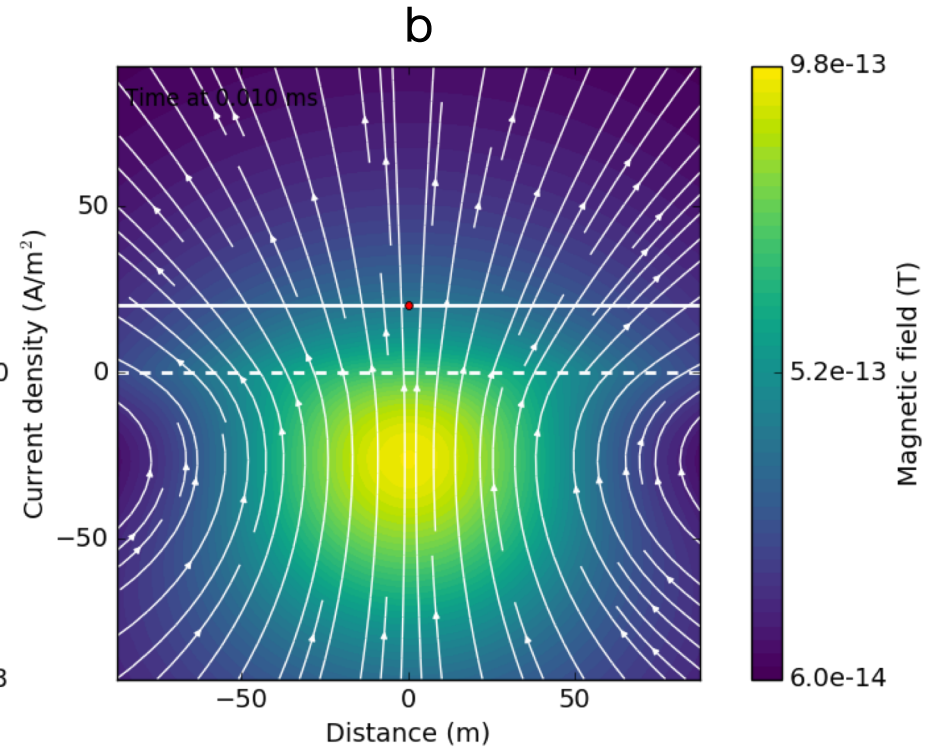
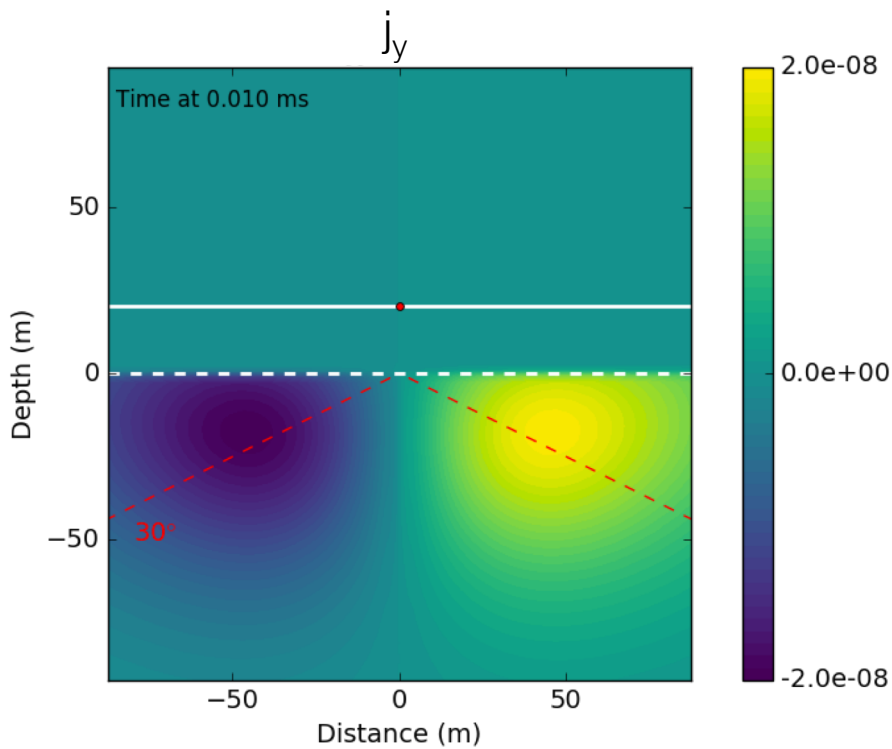
$$d = 1260\sqrt{t\rho}$$



Propagation through time

- Time: 0.01ms
- diffusion distance = 38 m

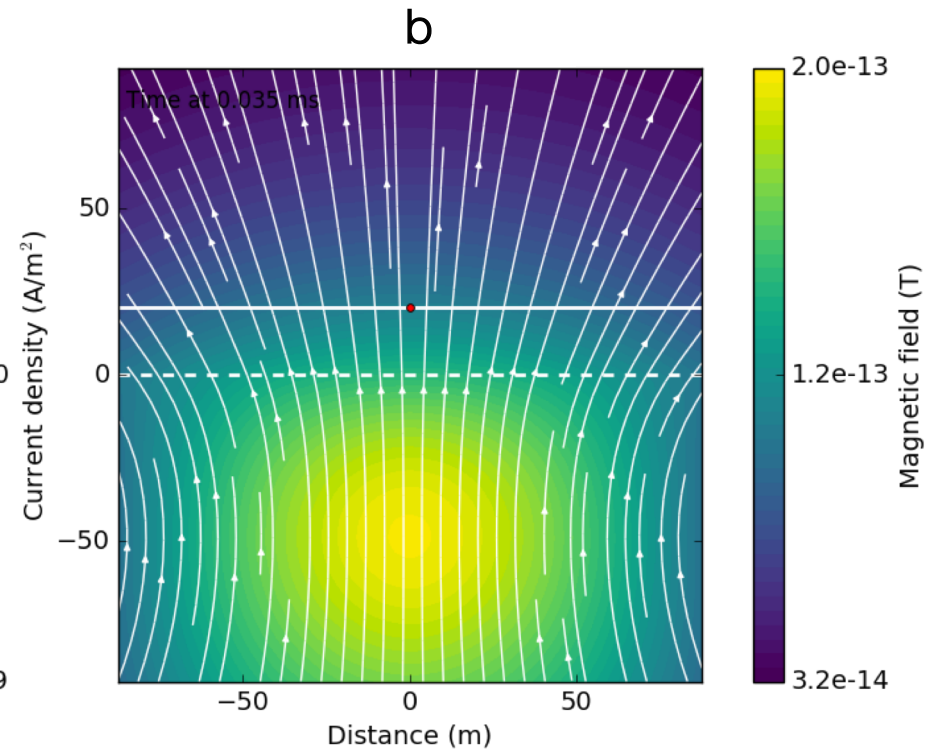
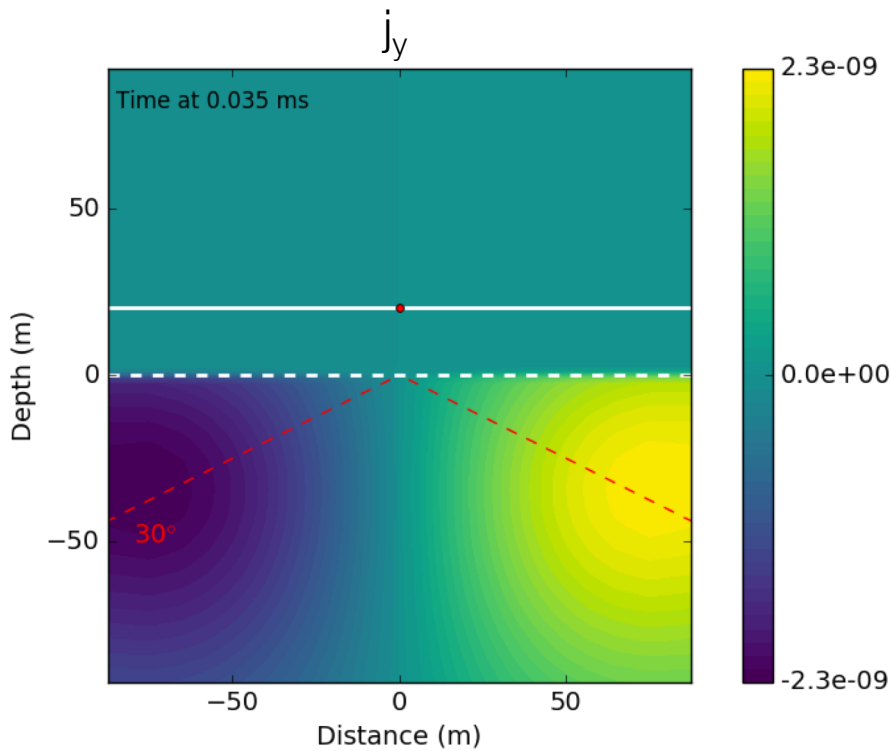
$$d = 1260\sqrt{t\rho}$$



Propagation through time

- Time: 0.035ms
- diffusion distance = 75 m

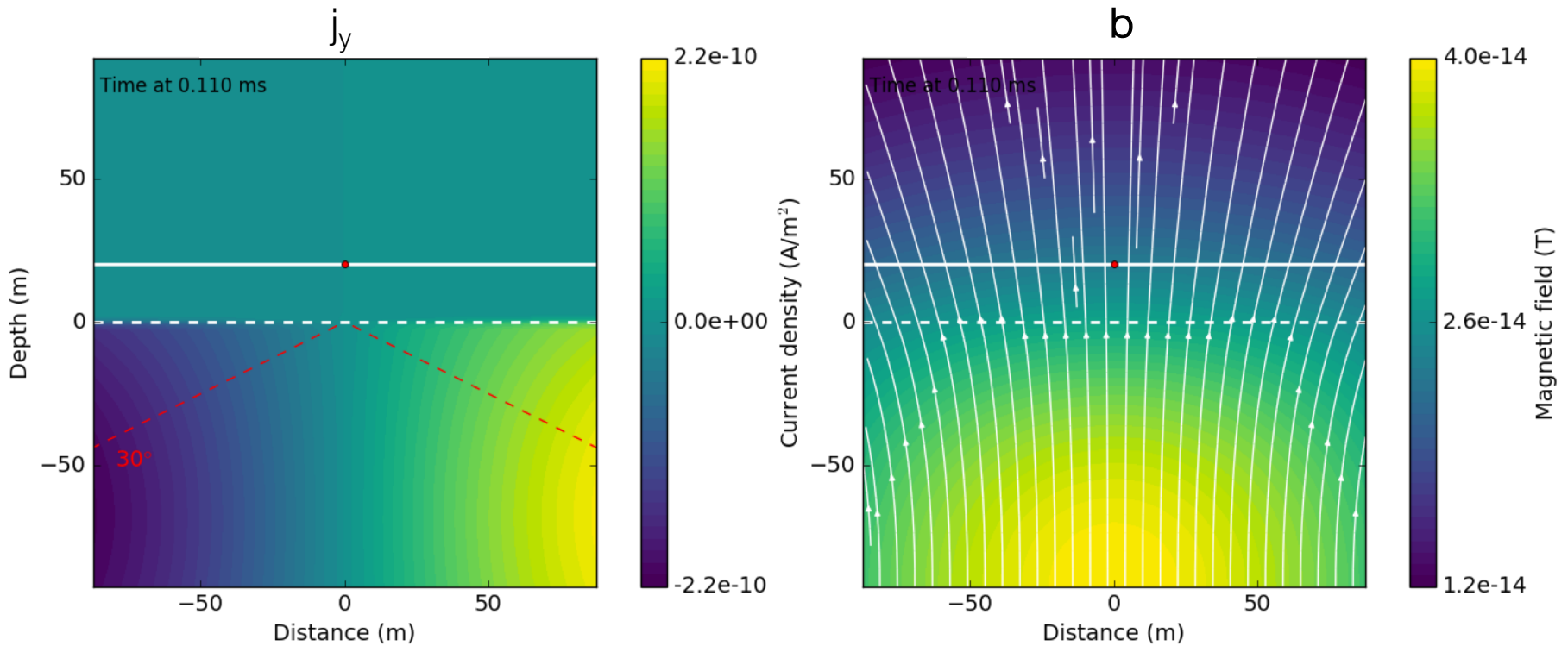
$$d = 1260\sqrt{t\rho}$$



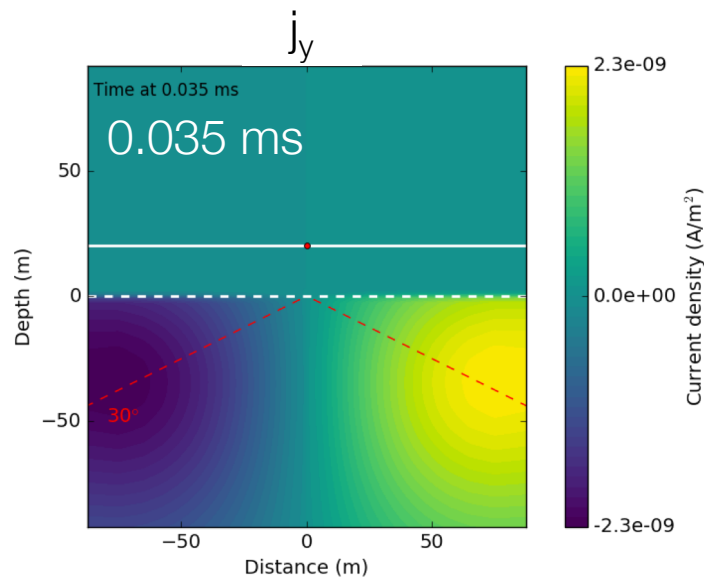
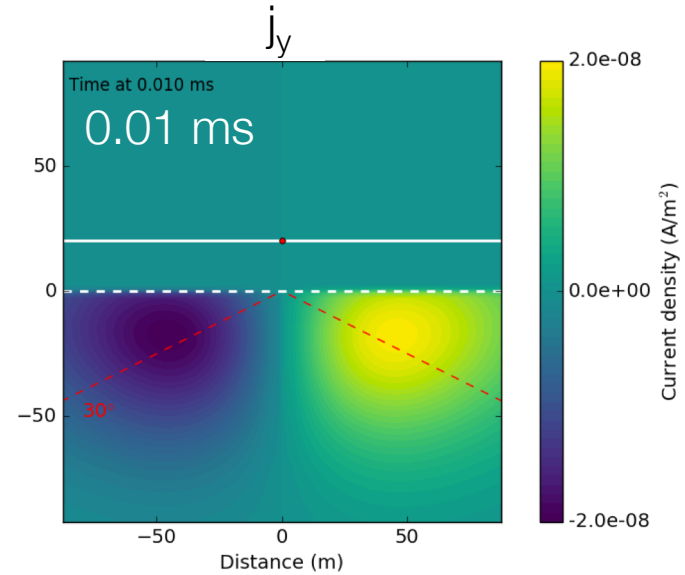
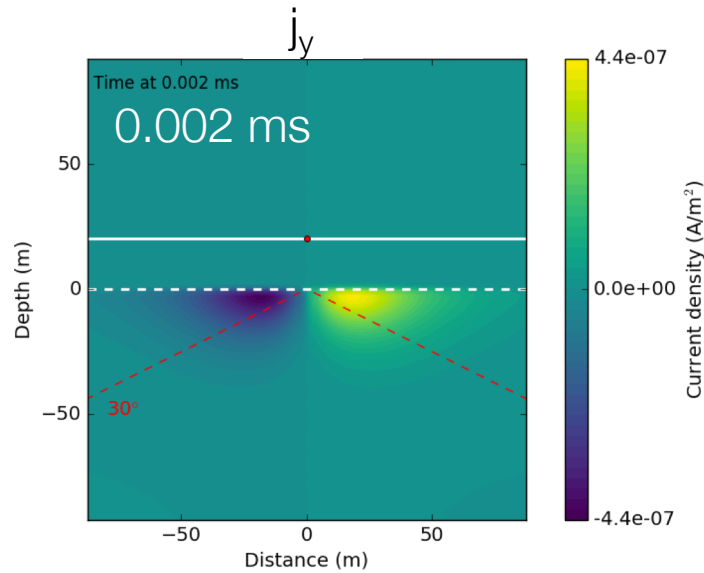
Propagation through time

- Time: 0.110ms
- diffusion distance = 132 m

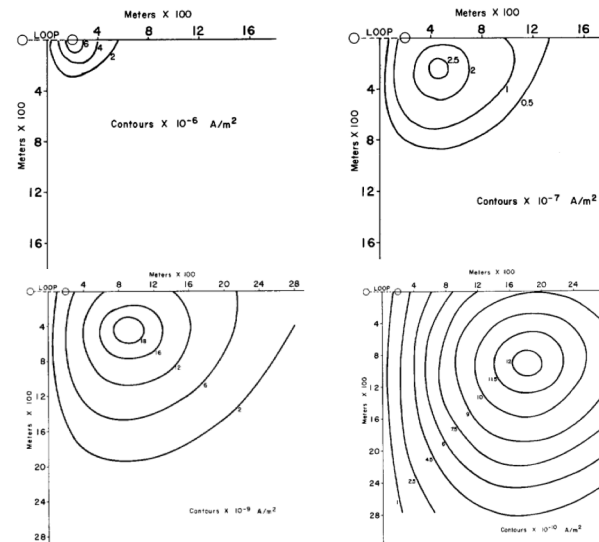
$$d = 1260\sqrt{t\rho}$$



Summary: propagation through time



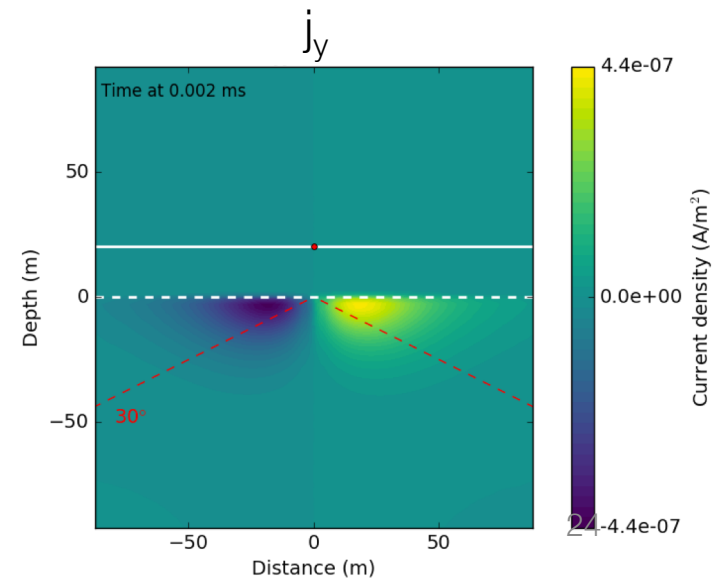
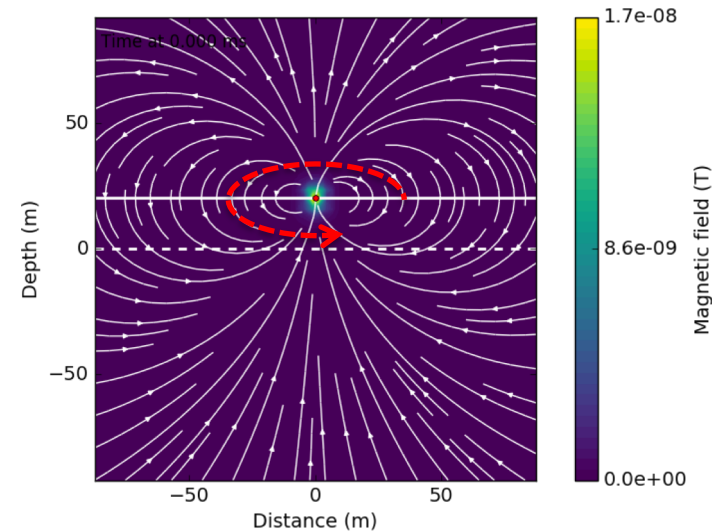
Nabighian (1979)



Important points

- Currents flow in same plane as transmitter currents
- Currents diffuse outward downward
- Each transmitter has a “footprint”
- Max resolution controlled by earliest time
- Depth of investigation controlled by latest time

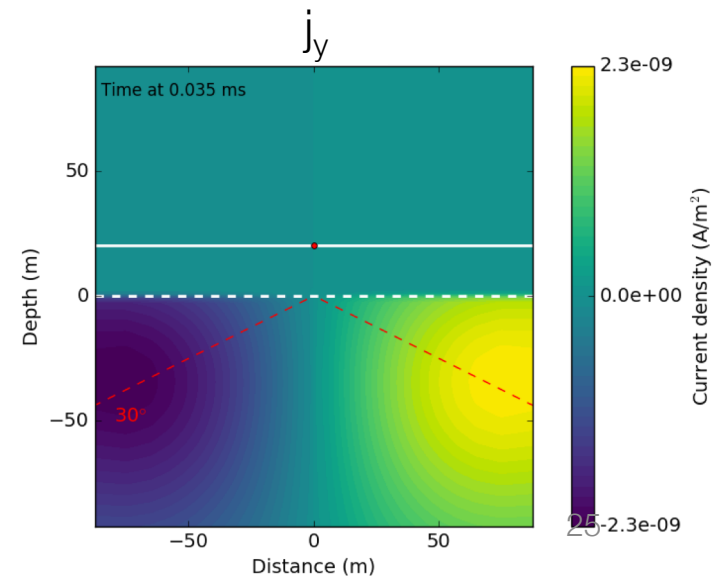
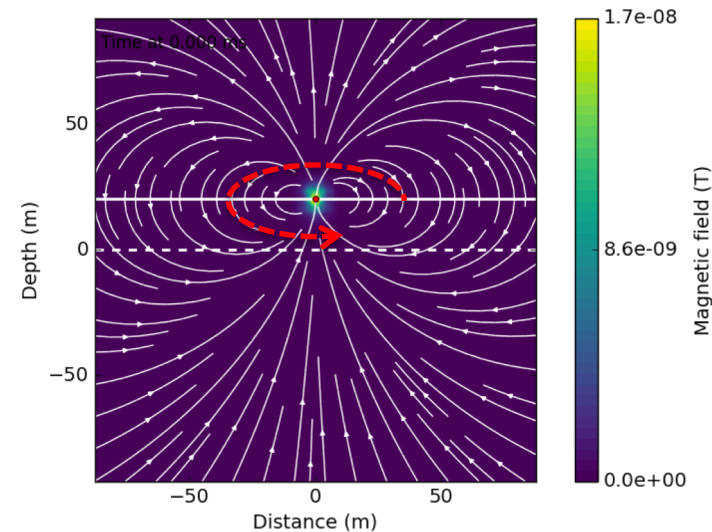
magnetic field (on-time)



Important points

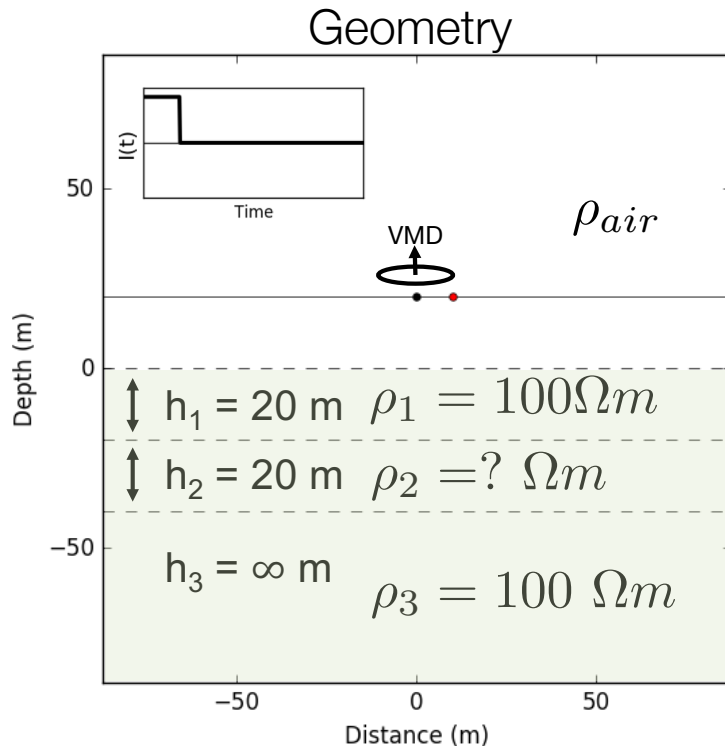
- Currents flow in same plane as transmitter currents
- Currents diffuse outward downward
- Each transmitter has a “footprint”
- Max resolution controlled by earliest time
- Depth of investigation controlled by latest time

magnetic field (on-time)



Layered earth

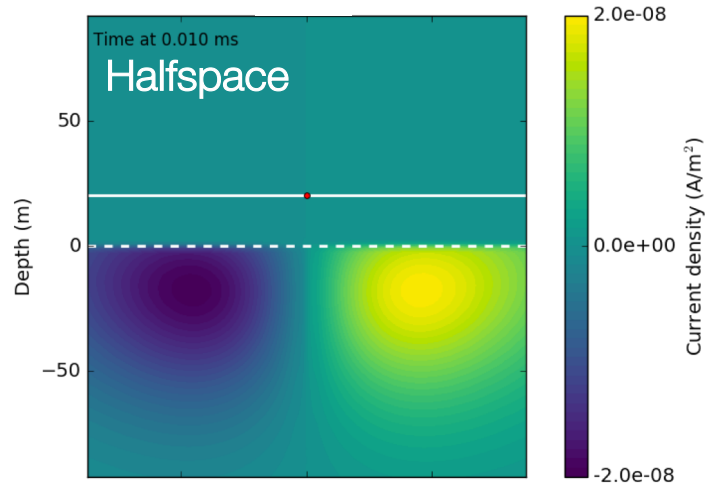
- 3 layers + air,
- ρ_2 varies



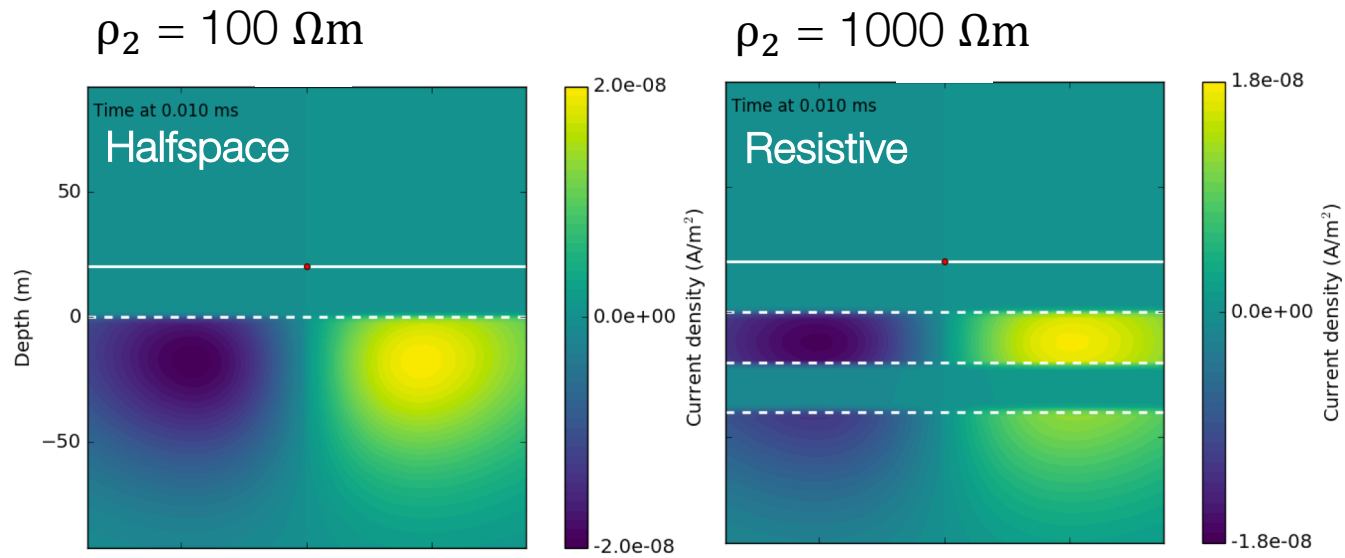
- Four different cases:
 - Halfspace
 - $\rho_2 = 100 \Omega m$
 - Resistive
 - $\rho_2 = 1000 \Omega m$
 - Conductive
 - $\rho_2 = 10 \Omega m$
 - Very conductive
 - $\rho_2 = 1 \Omega m$
- Fields
 - j_y off-time
 - b off-time

Layered earth currents (j_y)

$\rho_2 = 100 \Omega\text{m}$

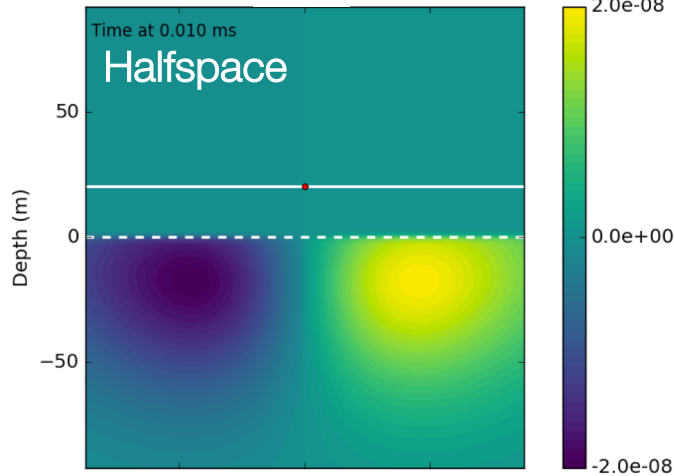


Layered earth currents (j_y)

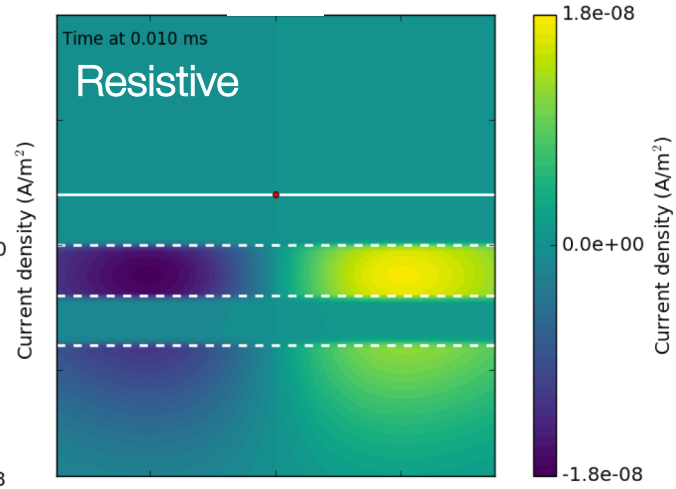


Layered earth currents (j_y)

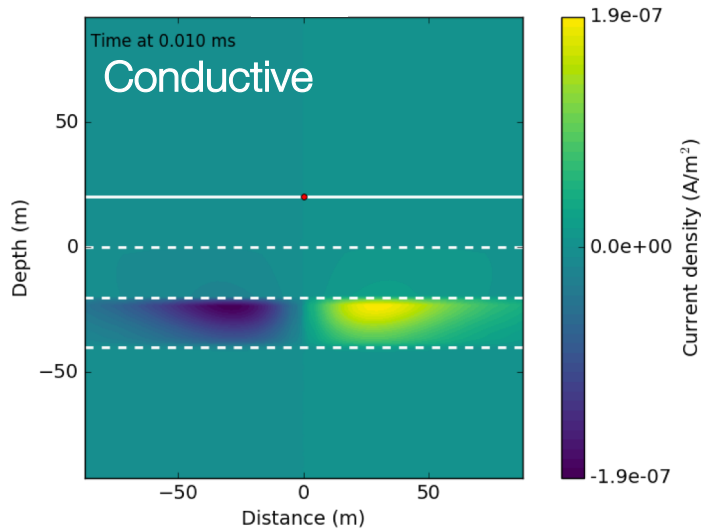
$\rho_2 = 100 \Omega\text{m}$



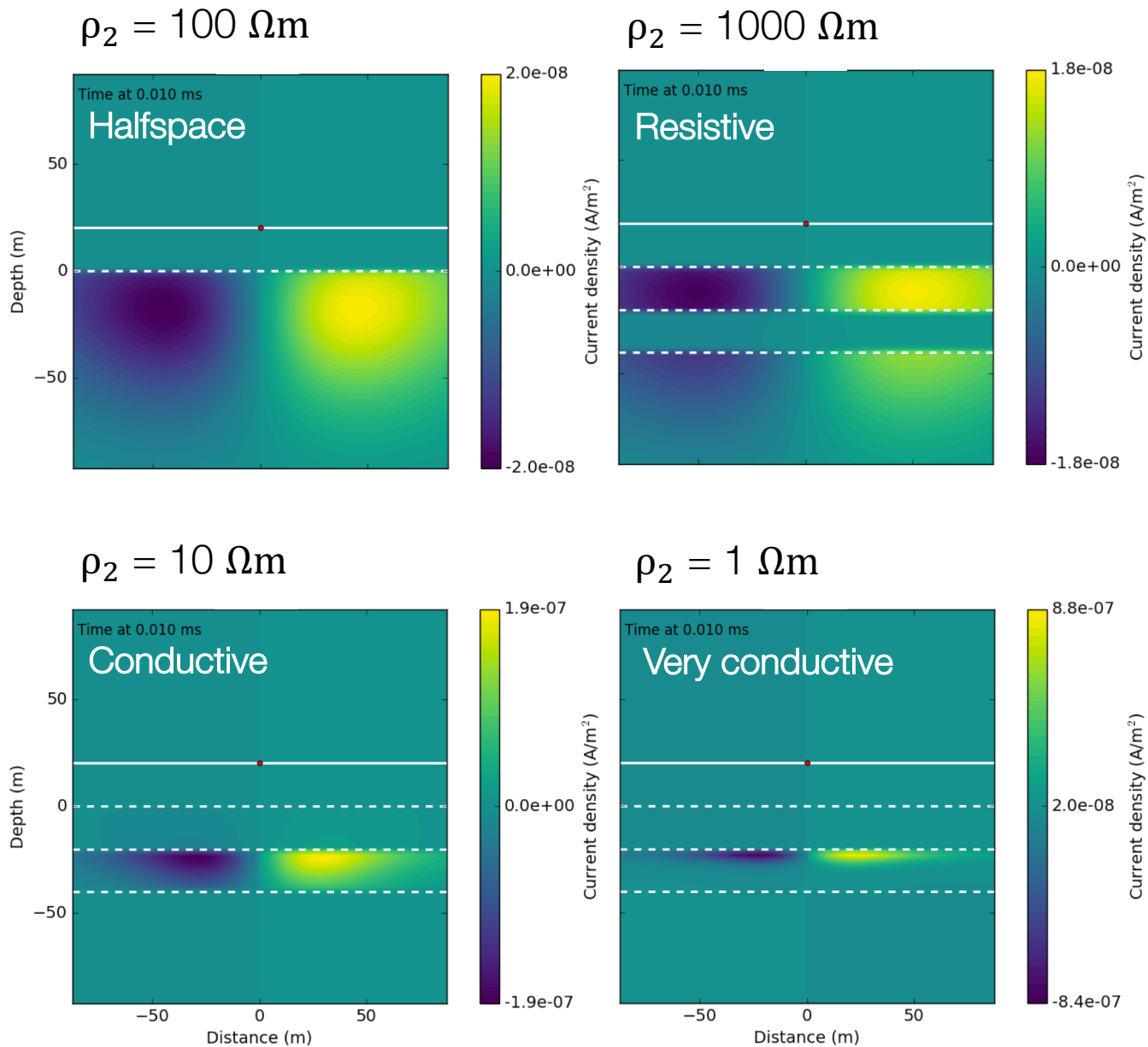
$\rho_2 = 1000 \Omega\text{m}$



$\rho_2 = 10 \Omega\text{m}$

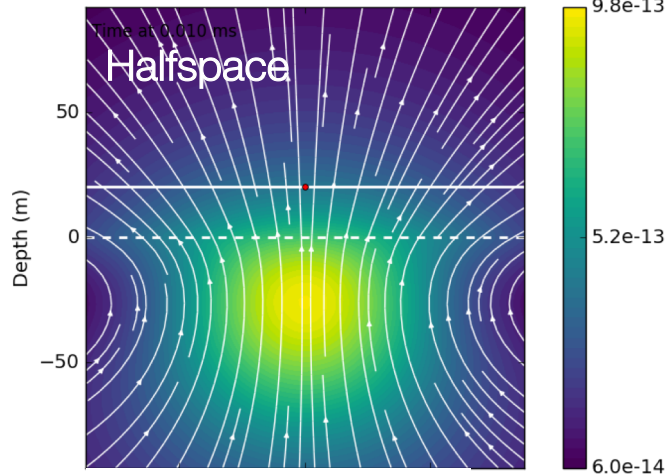


Layered earth currents (j_y)

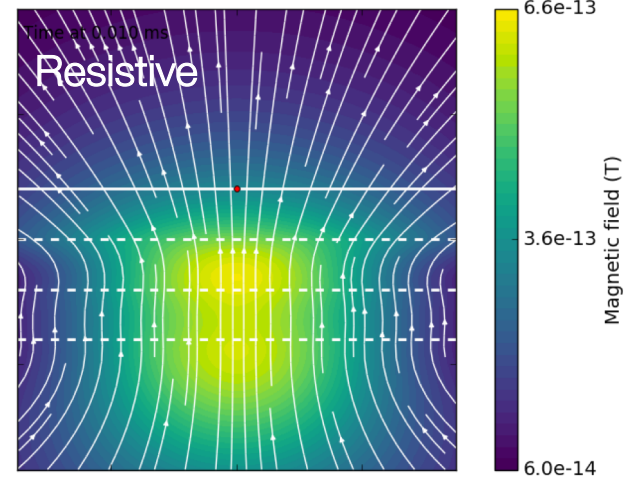


Layered earth mag. fields (b)

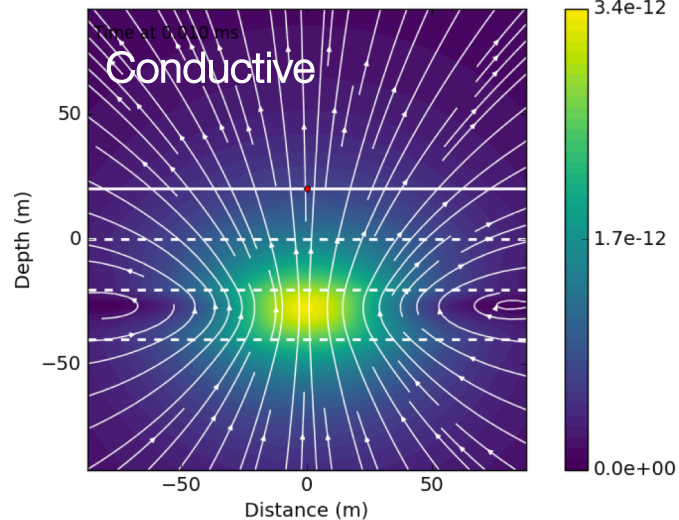
$\rho_2 = 100 \Omega\text{m}$



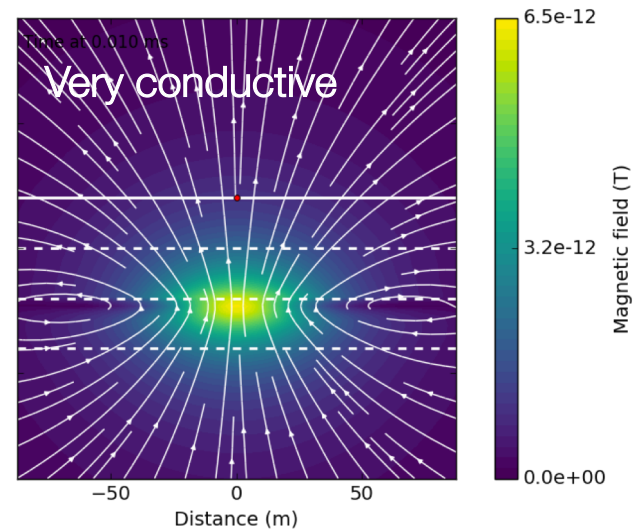
$\rho_2 = 1000 \Omega\text{m}$



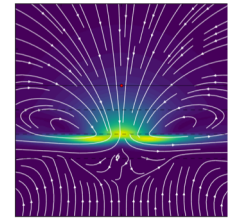
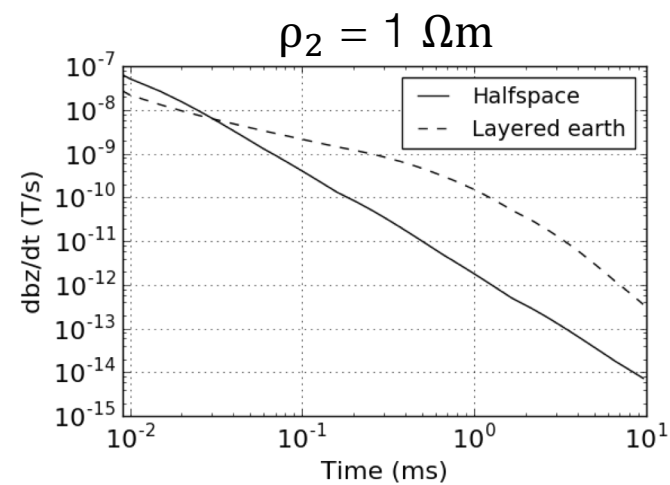
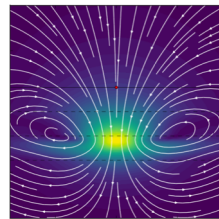
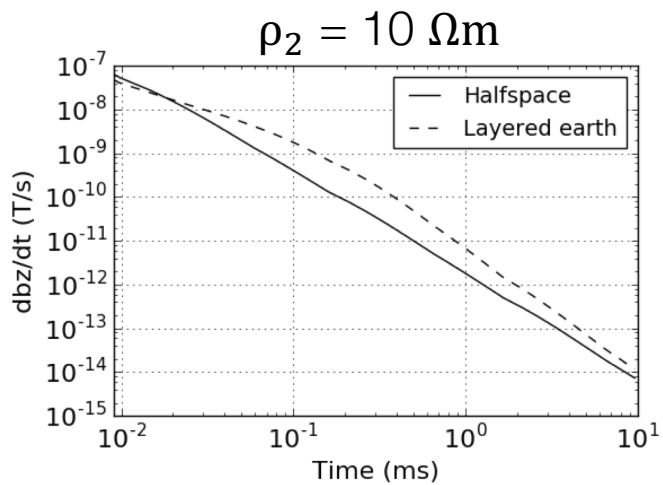
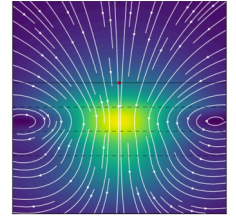
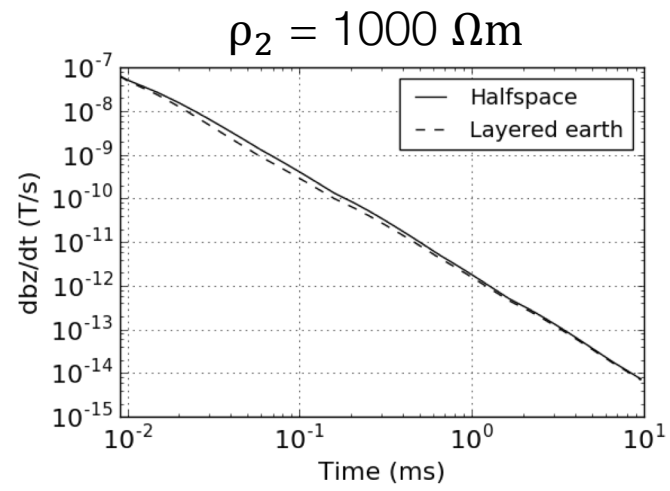
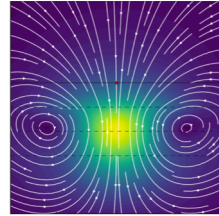
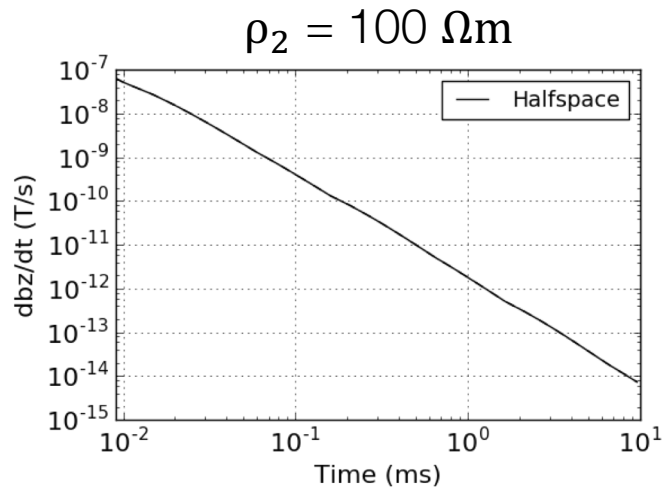
$\rho_2 = 10 \Omega\text{m}$



$\rho_2 = 1 \Omega\text{m}$

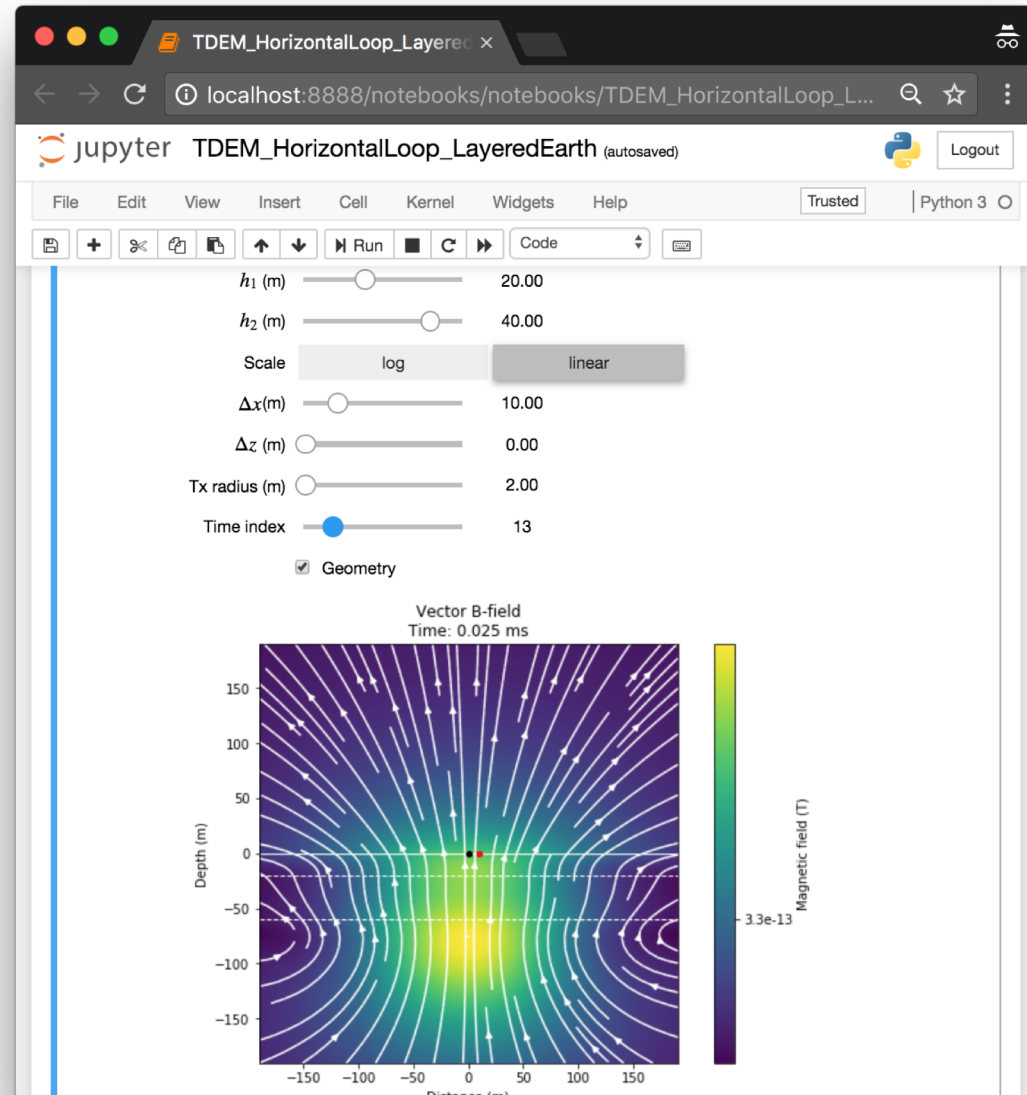


db_z/dt sounding curves



App: VMD over a layered earth (demo)

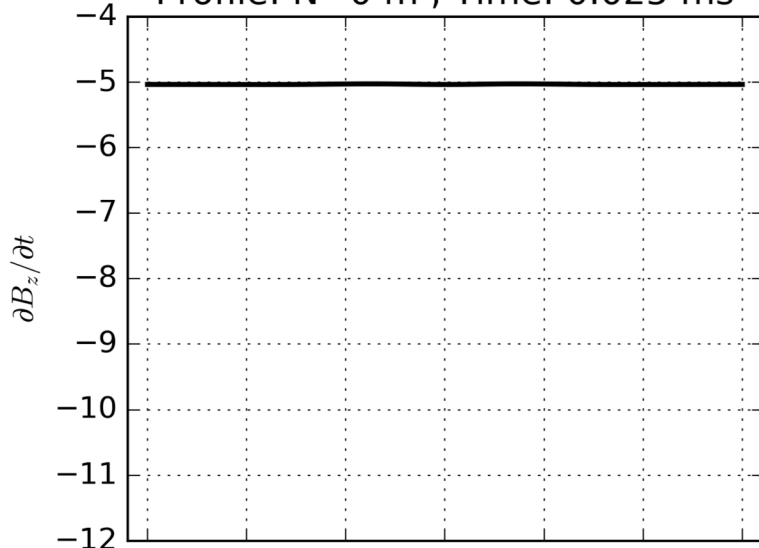
- TDEM_HorizontalLoop_Layered Earth
- Parameters:
 - Layer resistivities
 - Layer thicknesses
 - Source height, source receiver separation
- View:
 - Model
 - Electric field, magnetic field, current density through time
 - data



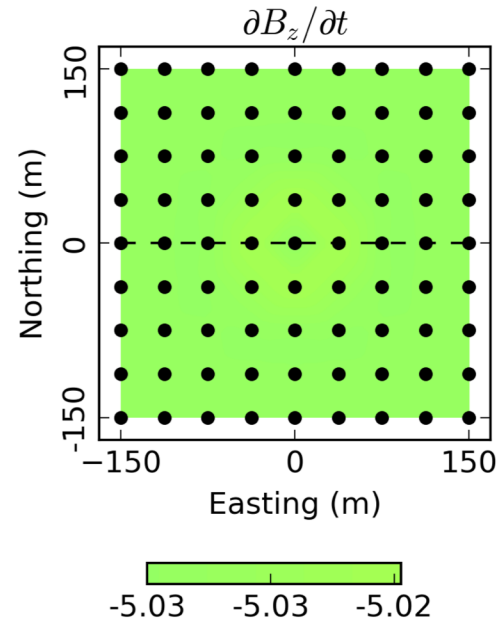
Airborne example: conductive sphere

Data profile

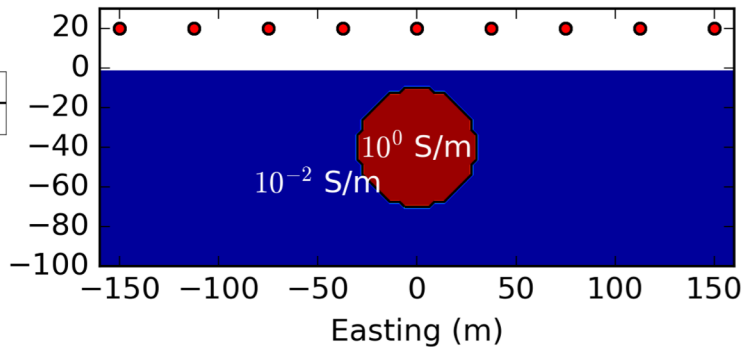
Profile: N=0 m , Time: 0.025 ms



Data map



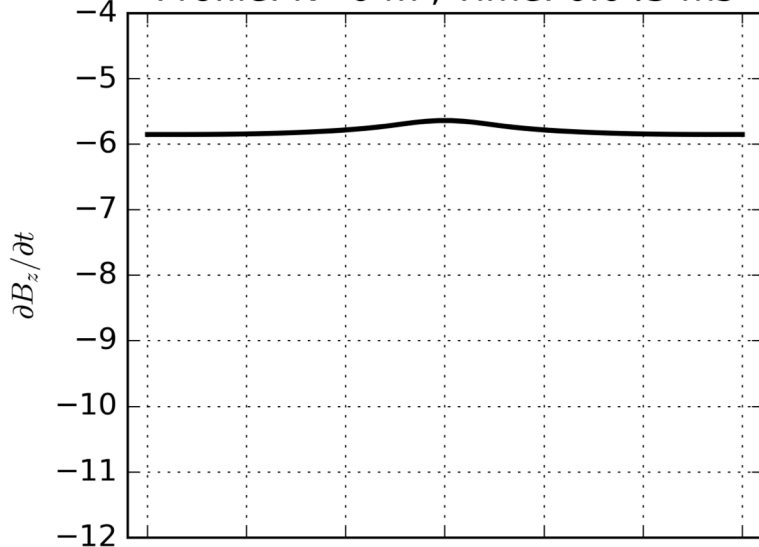
Conductivity



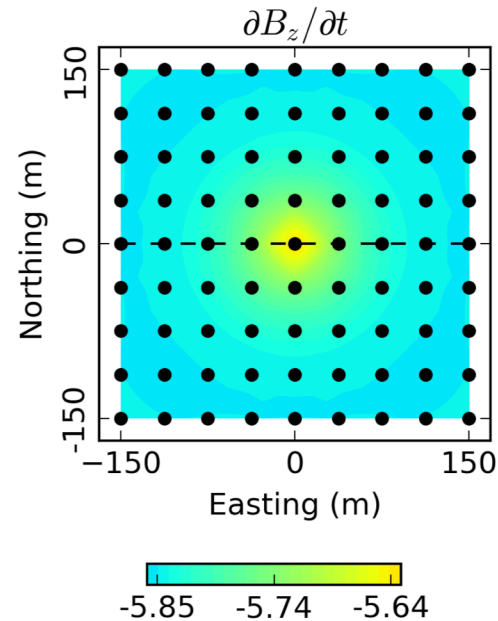
Airborne example: conductive sphere

Data profile

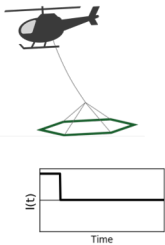
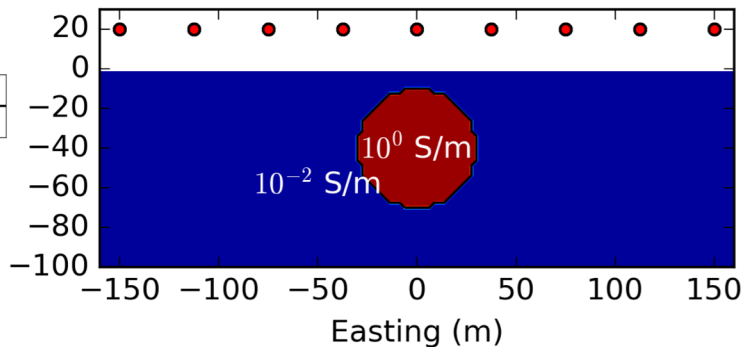
Profile: N=0 m , Time: 0.045 ms



Data map



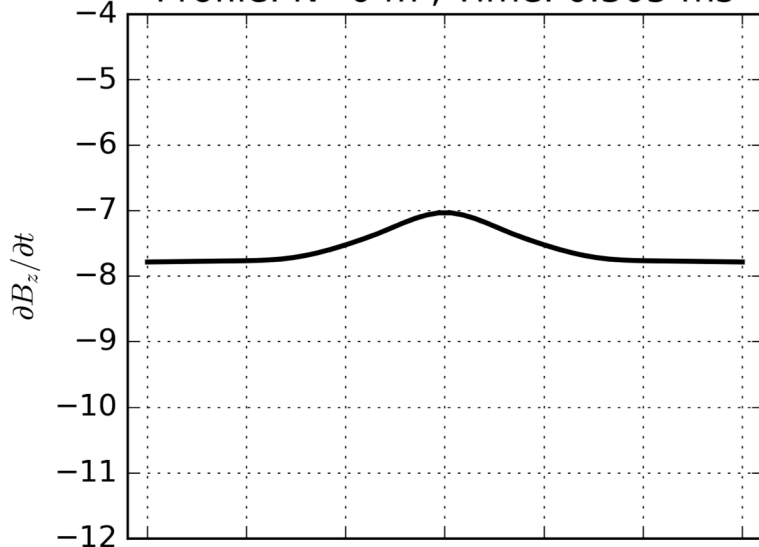
Conductivity



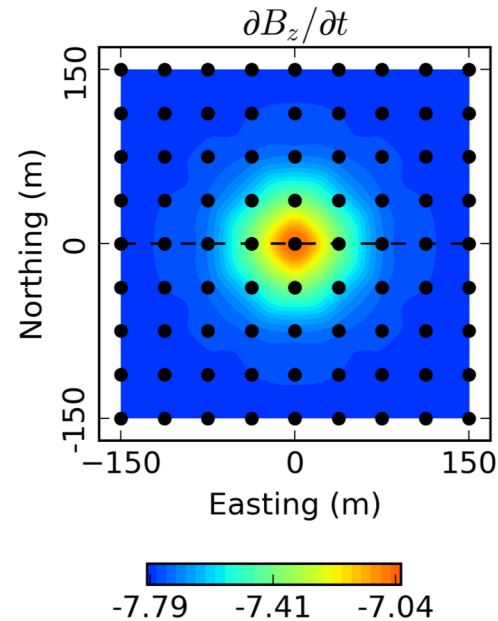
Airborne example: conductive sphere

Data profile

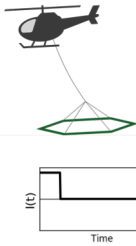
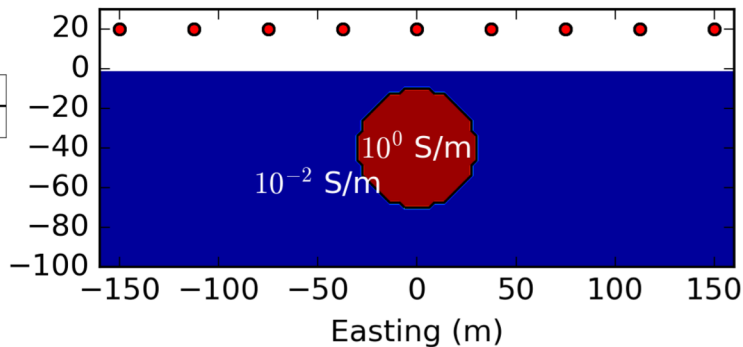
Profile: N=0 m , Time: 0.305 ms



Data map



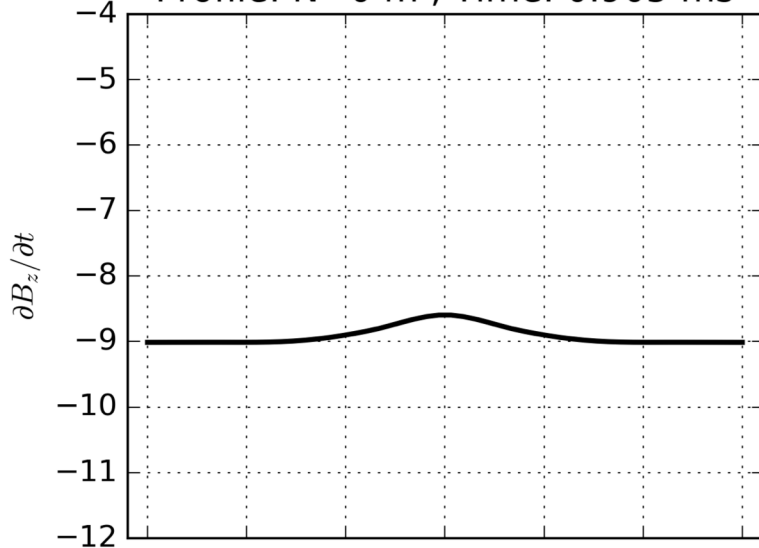
Conductivity



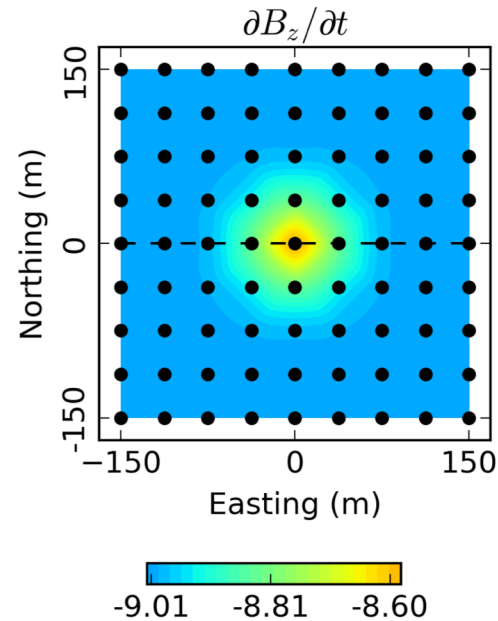
Airborne example: conductive sphere

Data profile

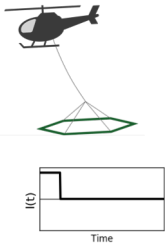
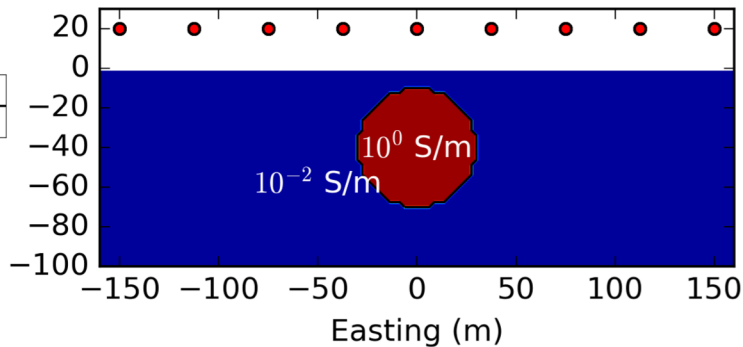
Profile: N=0 m , Time: 0.905 ms



Data map



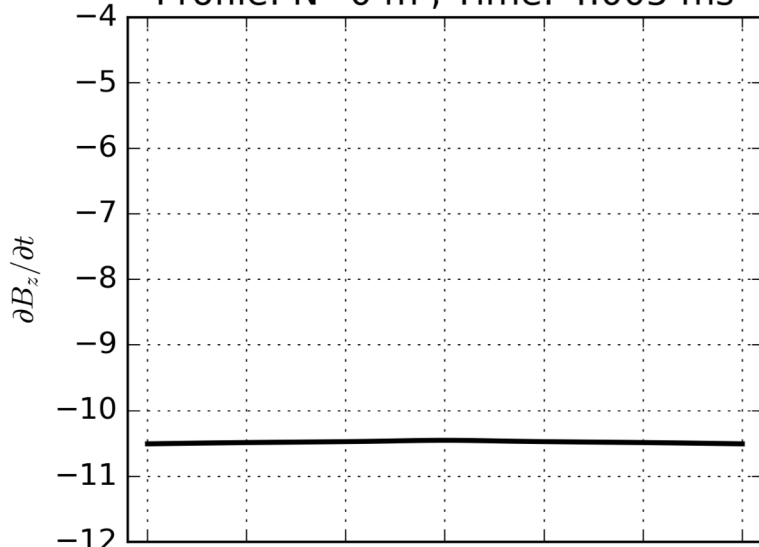
Conductivity



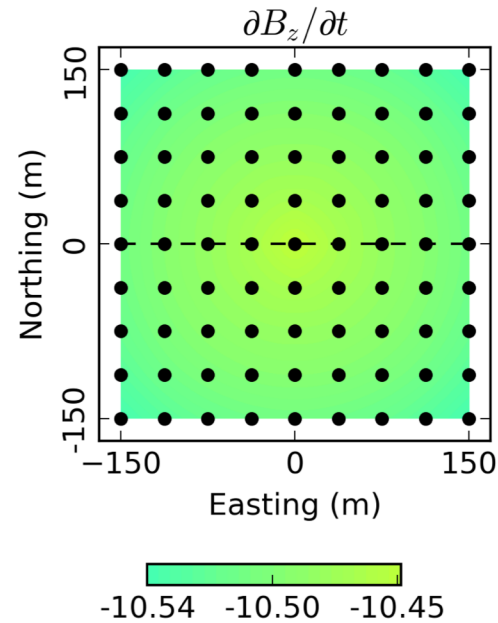
Airborne example: conductive sphere

Data profile

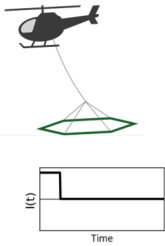
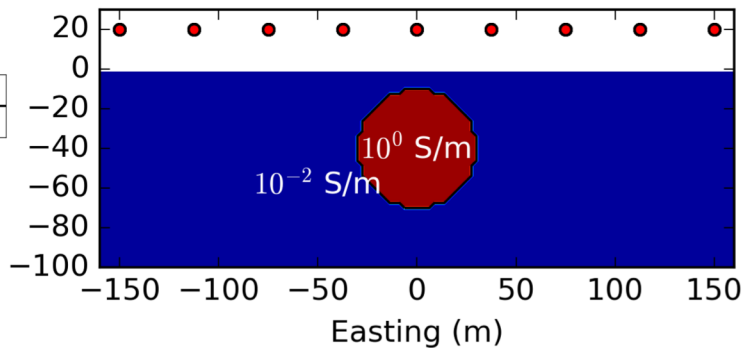
Profile: N=0 m , Time: 4.005 ms



Data map

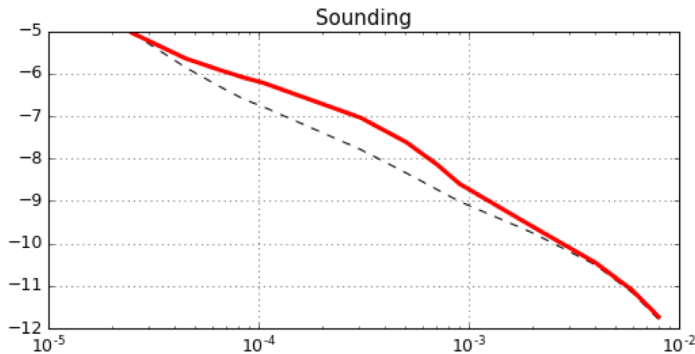
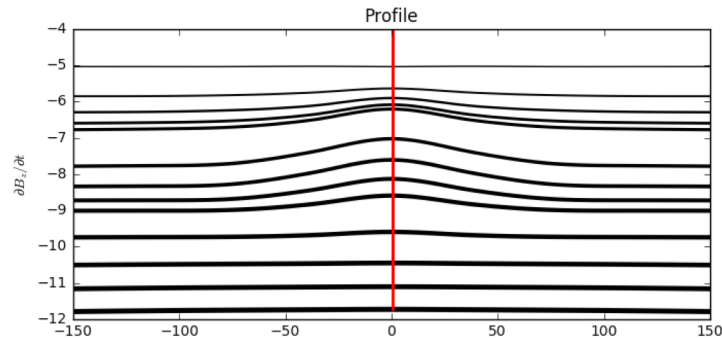


Conductivity

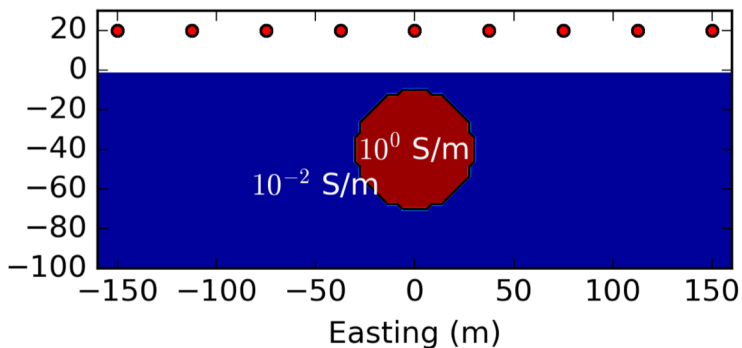


Summary: airborne example

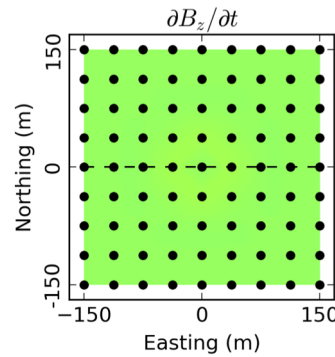
Data profile



Conductivity

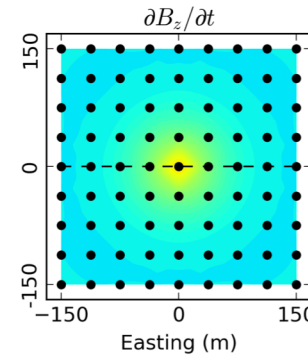


0.025 ms



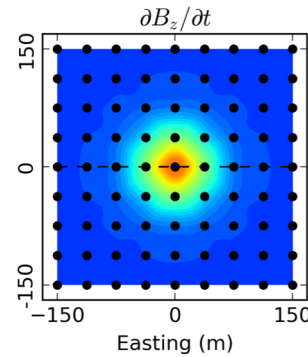
-5.03 -5.03 -5.02

0.045 ms



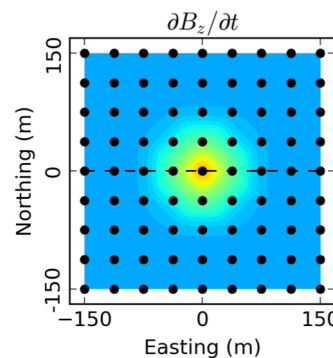
-5.85 -5.74 -5.64

0.305 ms



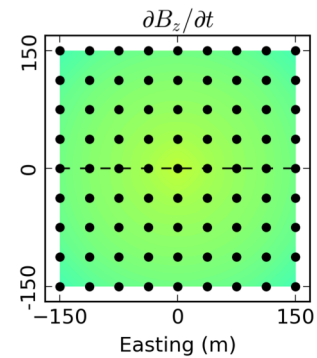
-7.79 -7.41 -7.04

0.905 ms



-9.01 -8.81 -8.60

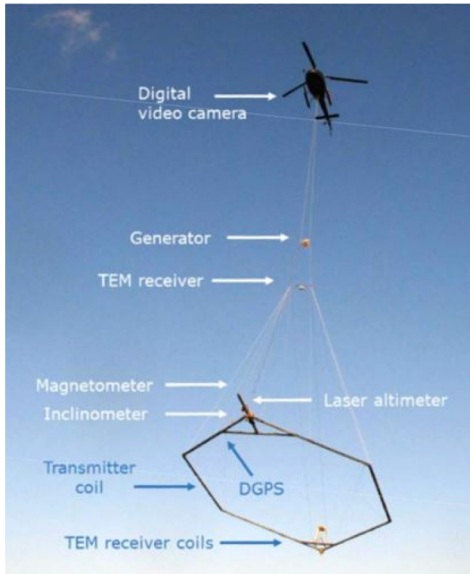
4.005 ms



-10.54 -10.50 -10.45

Some Airborne TDEM Systems

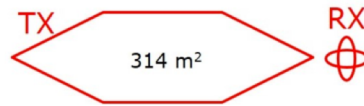
SkyTEM (2006)



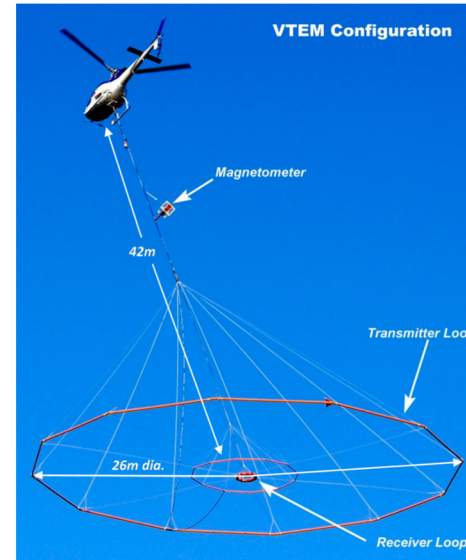
Area = 314 m²

Peak dipole moment:

- HM: 113040 NIA
- LM: 12560 NIA



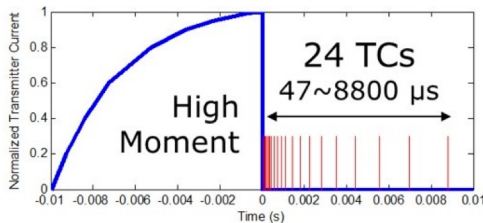
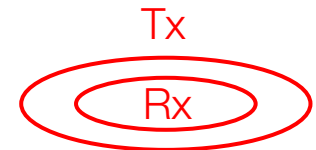
VTEM (2007)



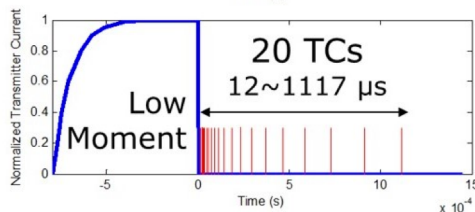
Area = 535 m²

Peak dipole moment:

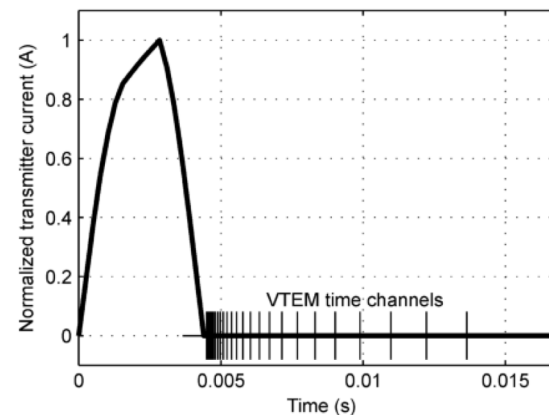
- 503,100 NIA



Peak current: 90 A
Turns: 4
On-time: 10 ms
Off-time: 10 ms

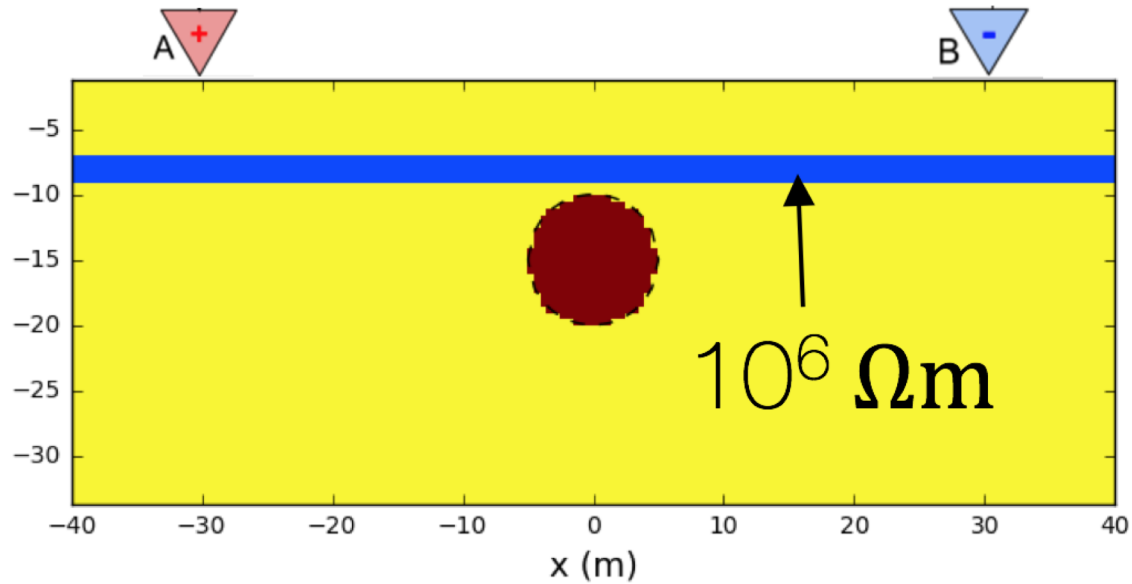


Peak current: 40 A
Turns: 1
On-time: 0.8 ms
Off-time: 1.45 ms



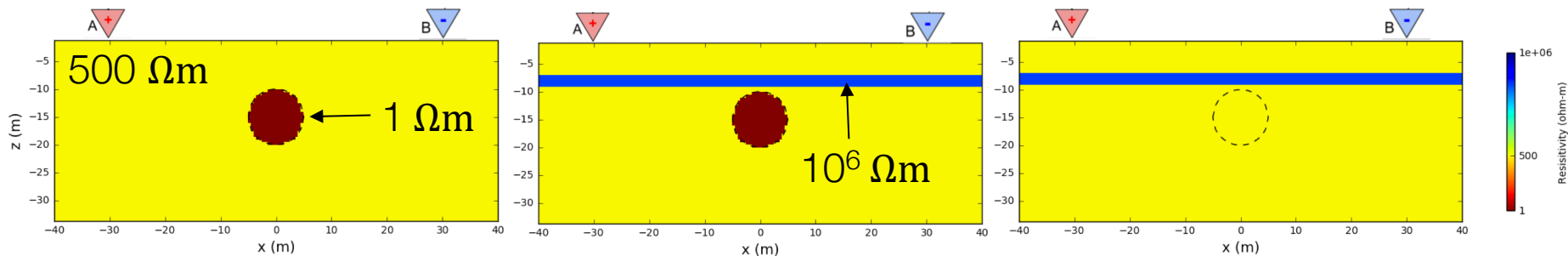
Peak current: 235 A
Turns: 4
On-time: 4.5 ms
Off-time: 9.1 ms

Back to the “shielding” problem

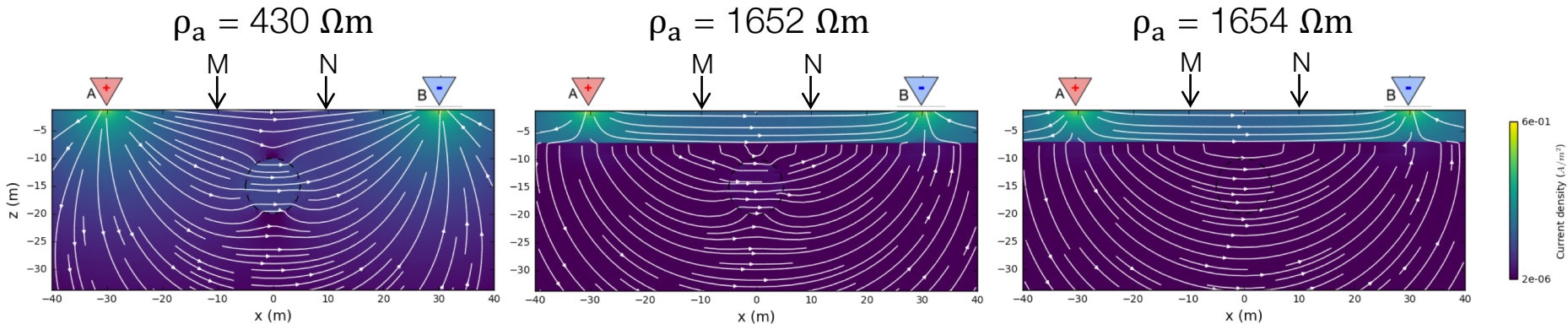


Shielding: DC with resistive layer

Resistivity models (thin resistive layer)

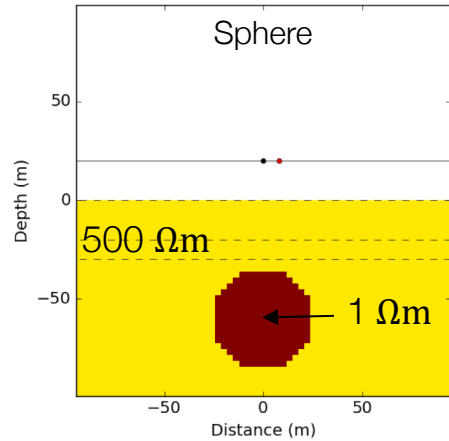


Currents and measured data at MN

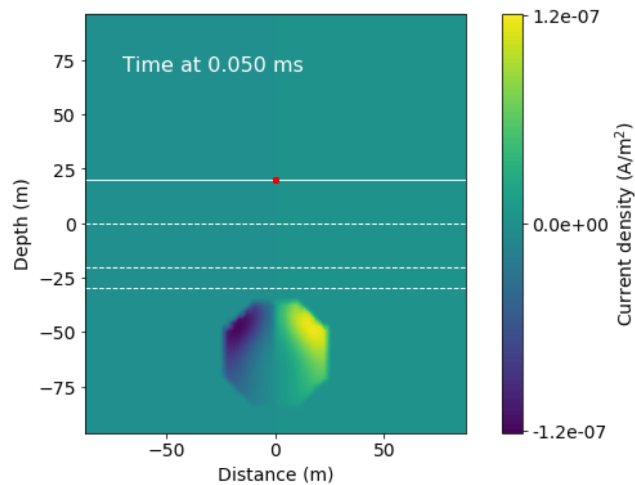


Shielding: EM with resistive layer

Resistivity models (thin resistive layer)

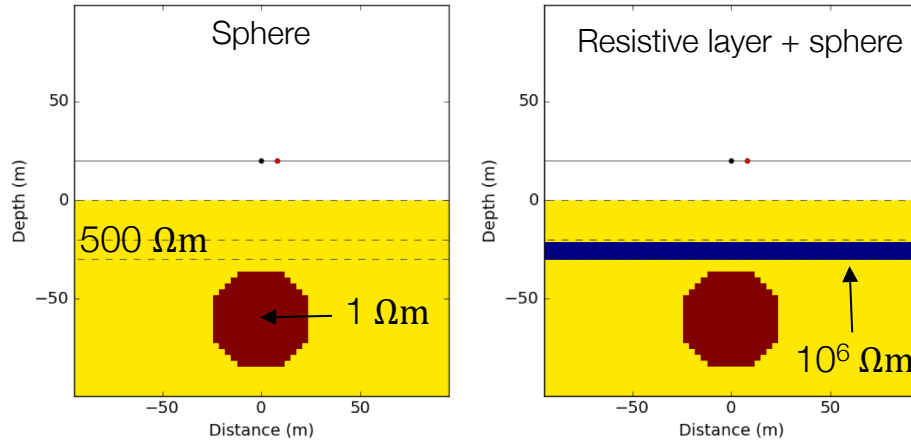


Currents (J_y)

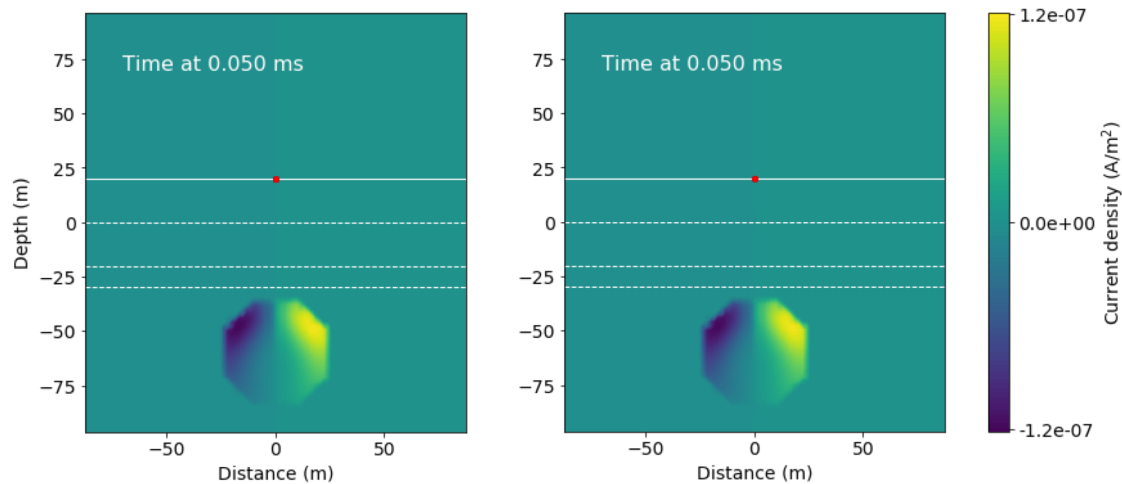


Shielding: EM with resistive layer

Resistivity models (thin resistive layer)

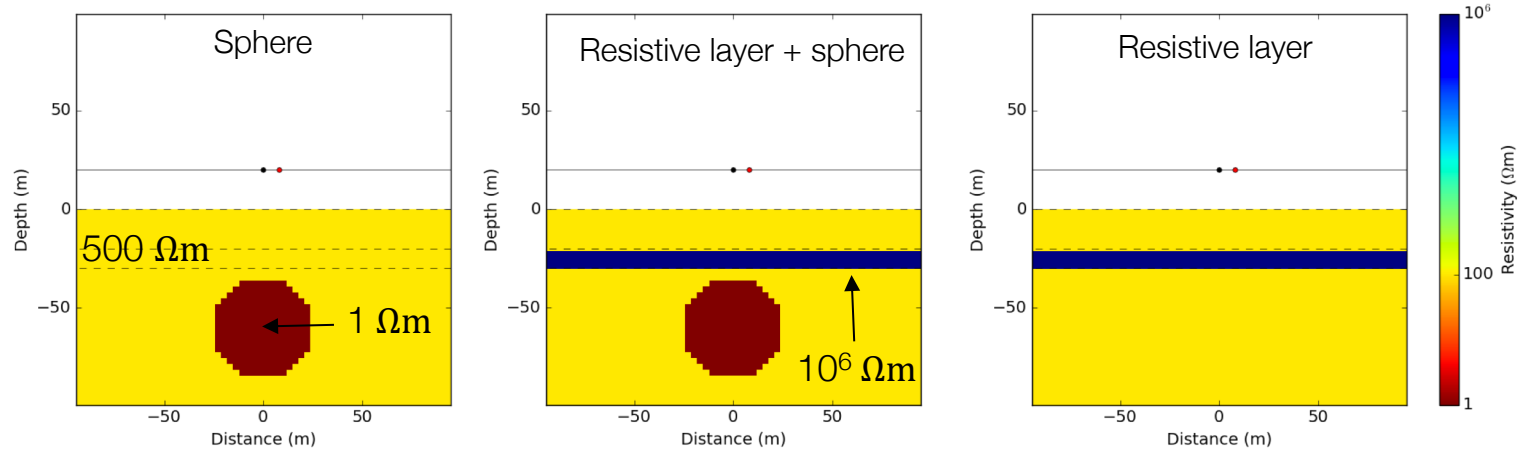


Currents (J_y)

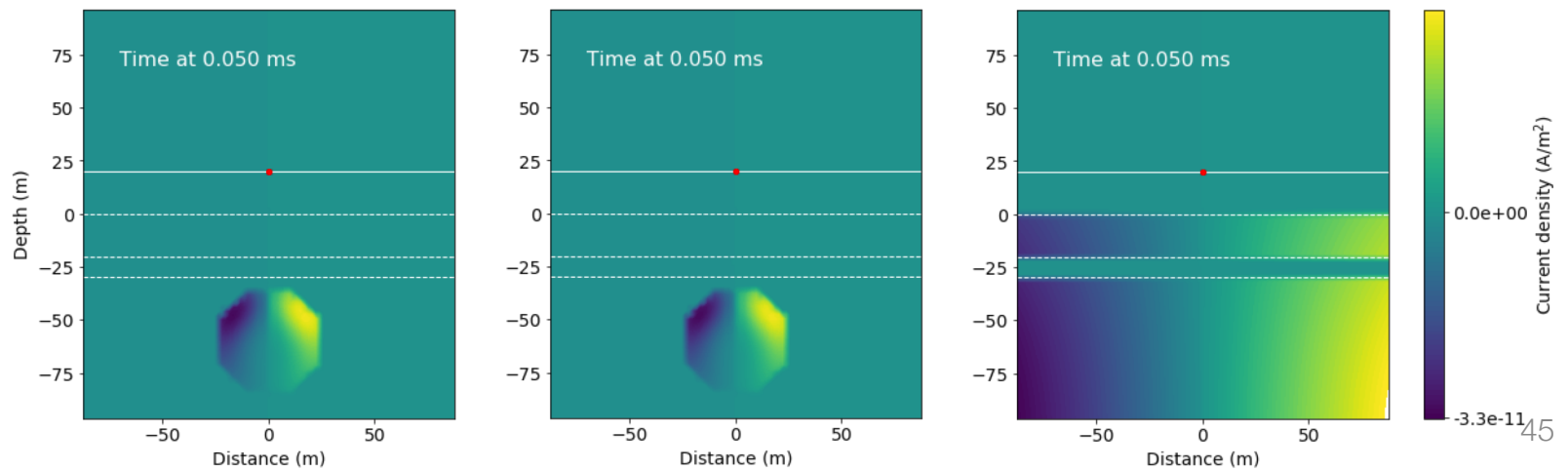


Shielding: EM with resistive layer

Resistivity models (thin resistive layer)

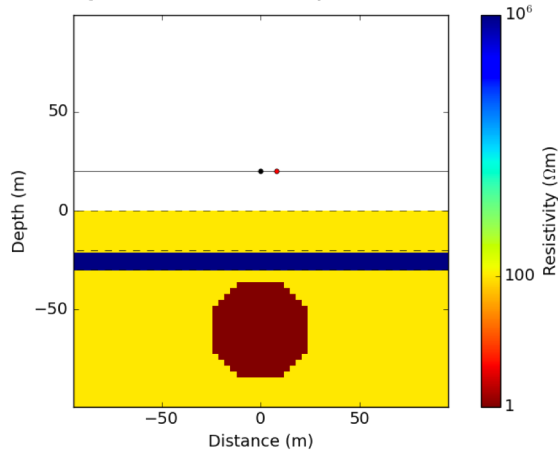


Currents (J_V)

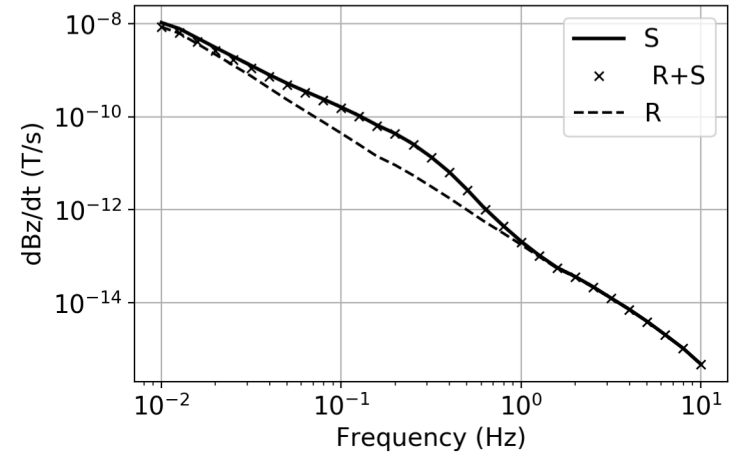


Shielding: EM with resistive layer

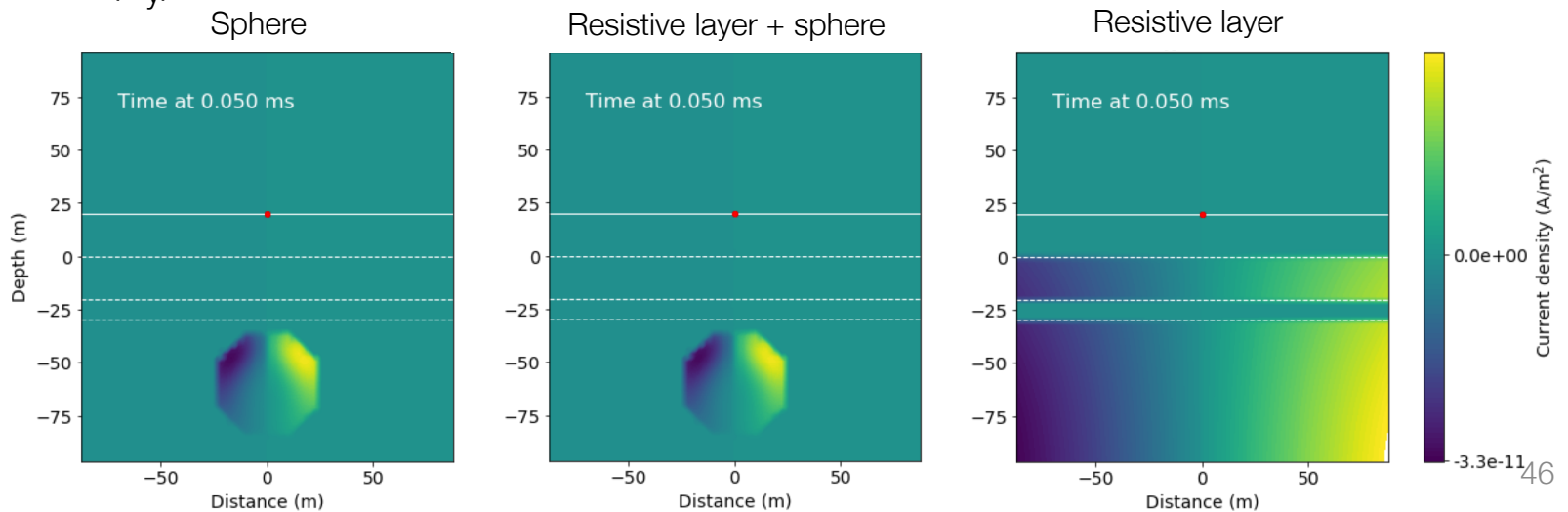
Resistivity models (thin resistive layer)



dB_z/dt sounding curves

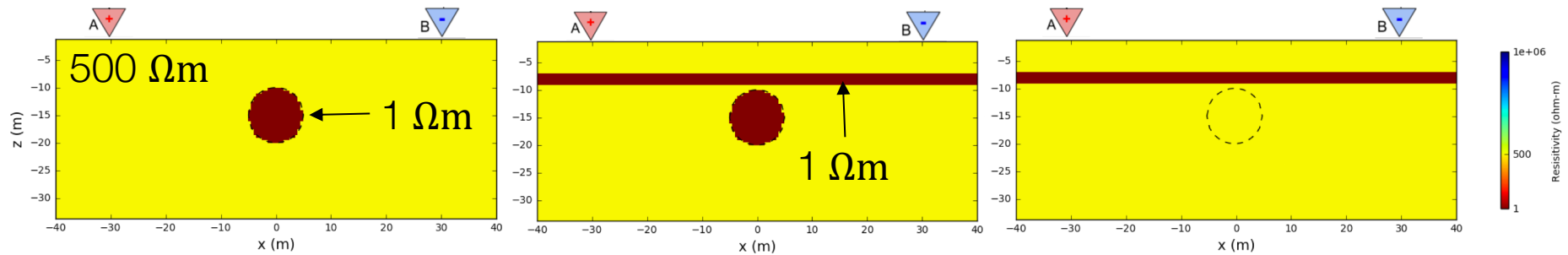


Currents (J_y)

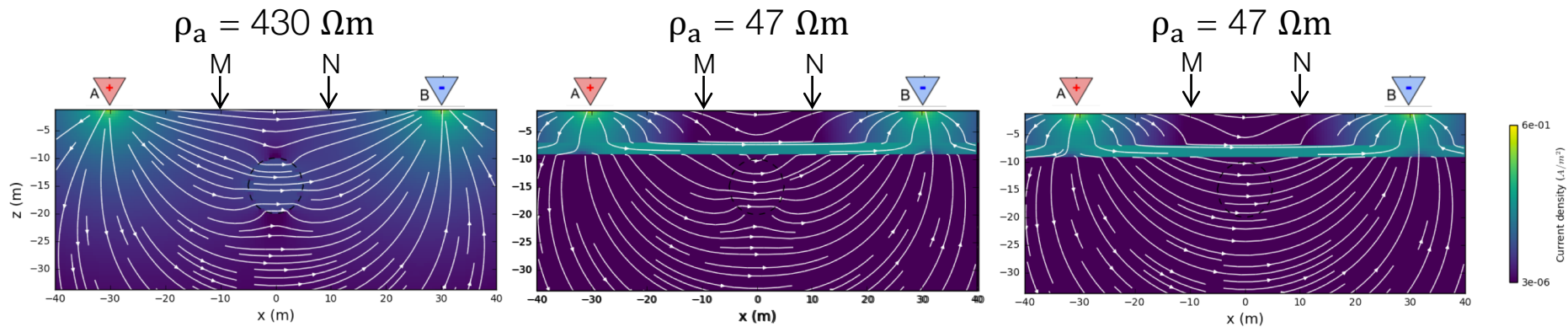


Shielding: DC with conductive layer

Resistivity models (thin conductive layer)

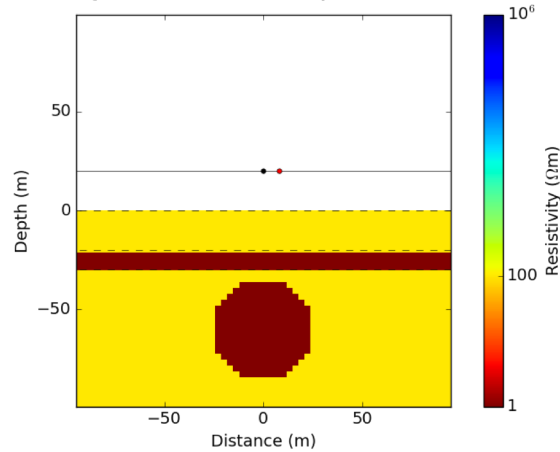


Currents and measured data at MN

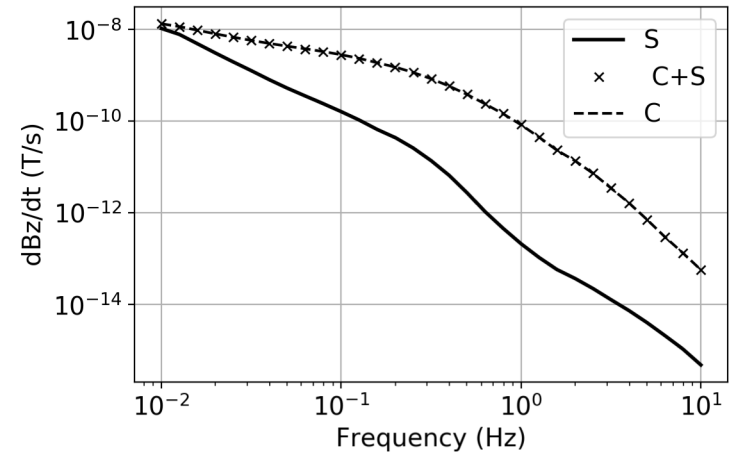


Shielding: EM with conductive layer

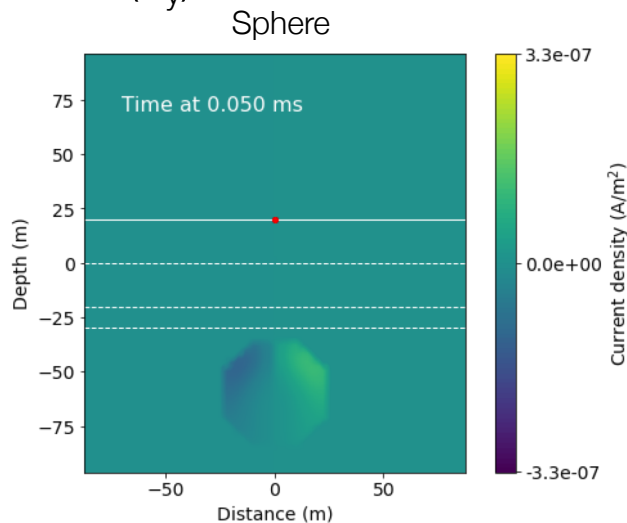
Resistivity models (thin conductive layer)



dB_z/dt sounding curves

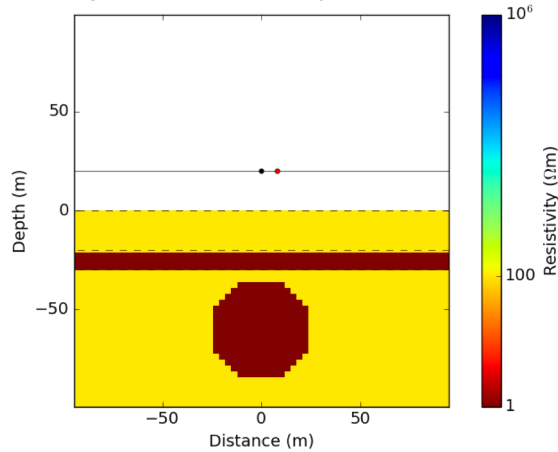


Currents (J_y)

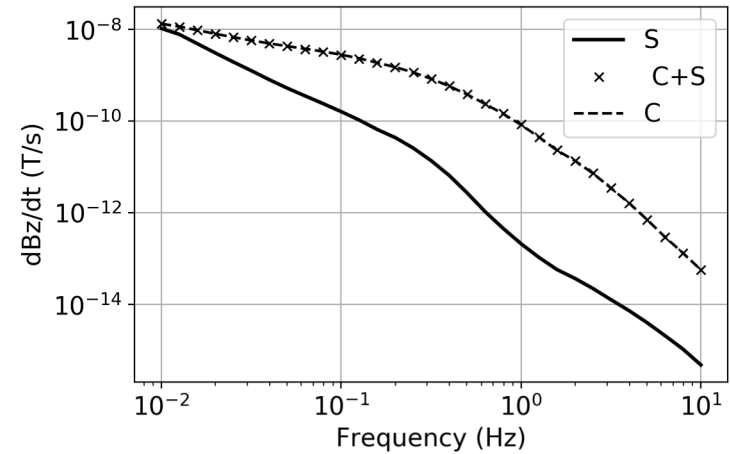


Shielding: EM with conductive layer

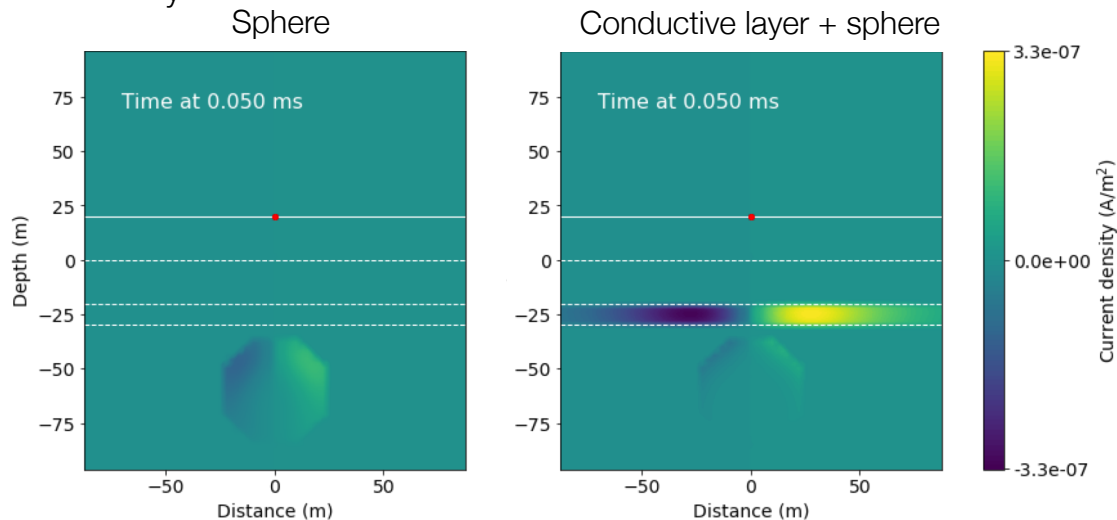
Resistivity models (thin conductive layer)



dB_z/dt sounding curves

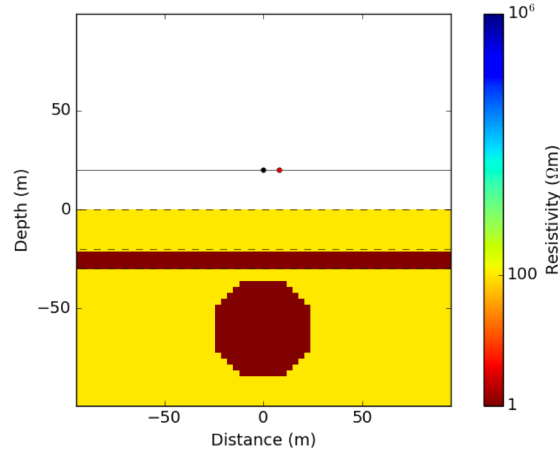


Currents (J_y)

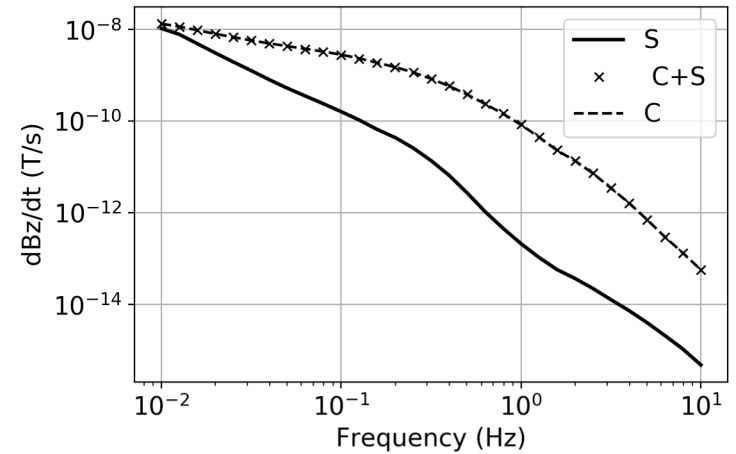


Shielding: EM with conductive layer

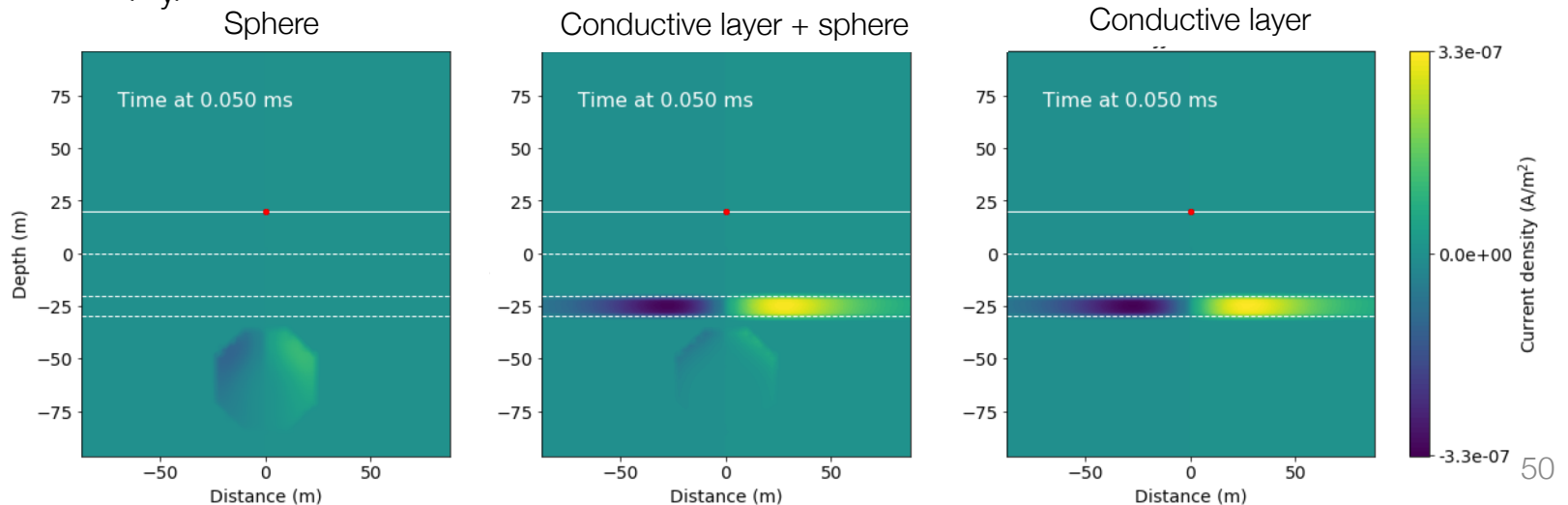
Resistivity models (thin conductive layer)



dB_z/dt sounding curves



Currents (J_y)



App: VMD over sphere with overburden

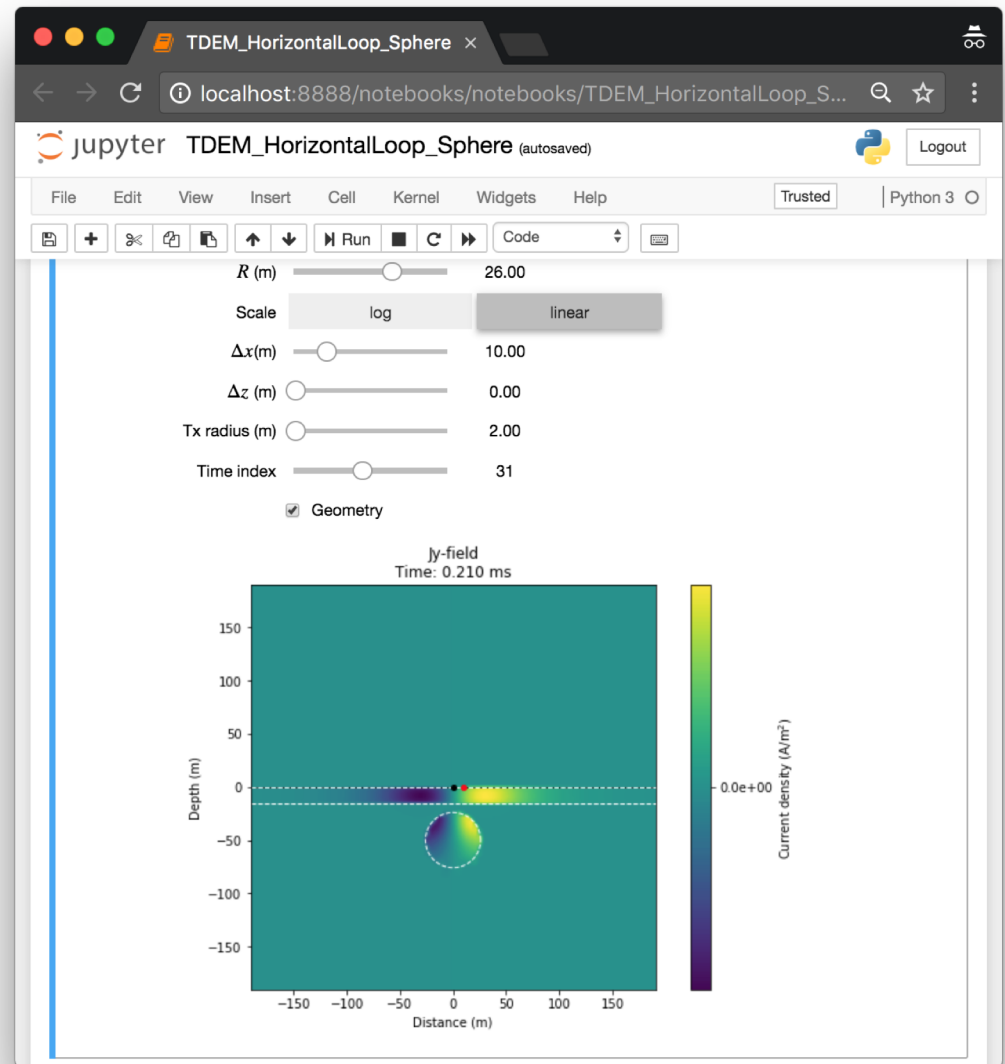
- TDEM_HorizontalLoop_Sphere

- Parameters:

- Resistivities of layer, background sphere
- Geometry of sphere, layer
- Source height, source receiver separation

- View:

- Model
- Electric field, magnetic field, current density through time
- data



Outline

Setup

Time Domain EM

- Vertical Magnetic Dipole
- Propagation with Time
- Effects of Background Conductivity
- Transmitters and receivers
- Decay Curves
- Case History: Groundwater, Minerals, Hydrocarbons
- Horizontal magnetic Dipole

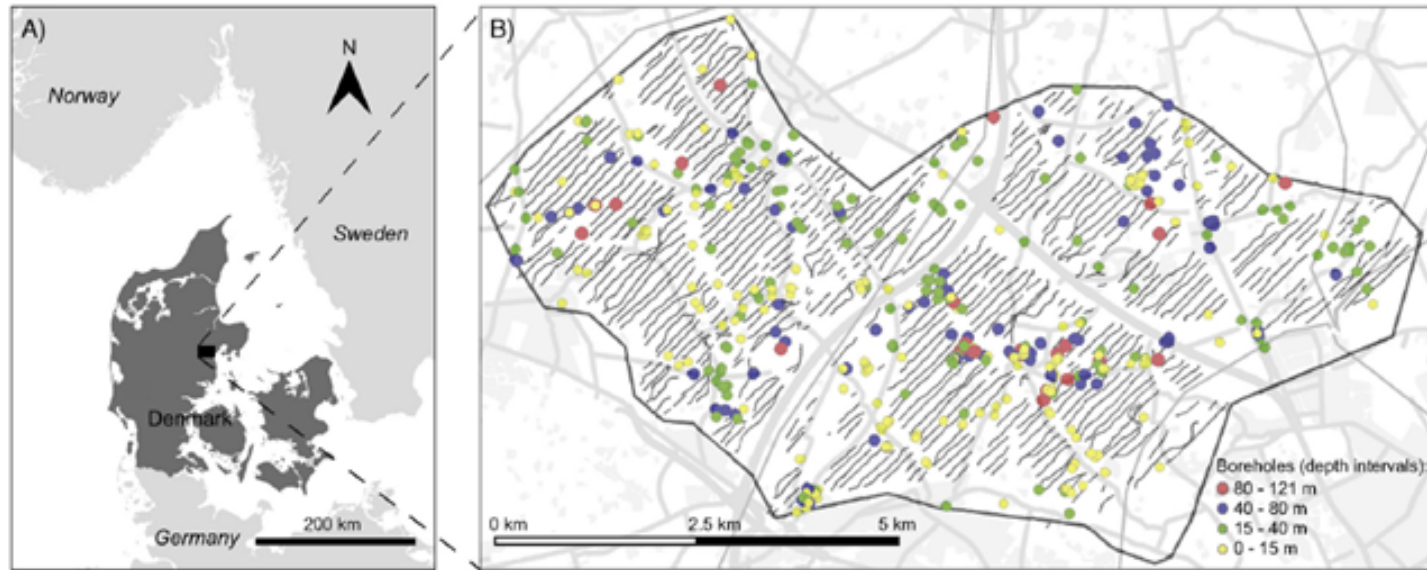
Case History: Kasted

Vilhelmsen et al. (2016)

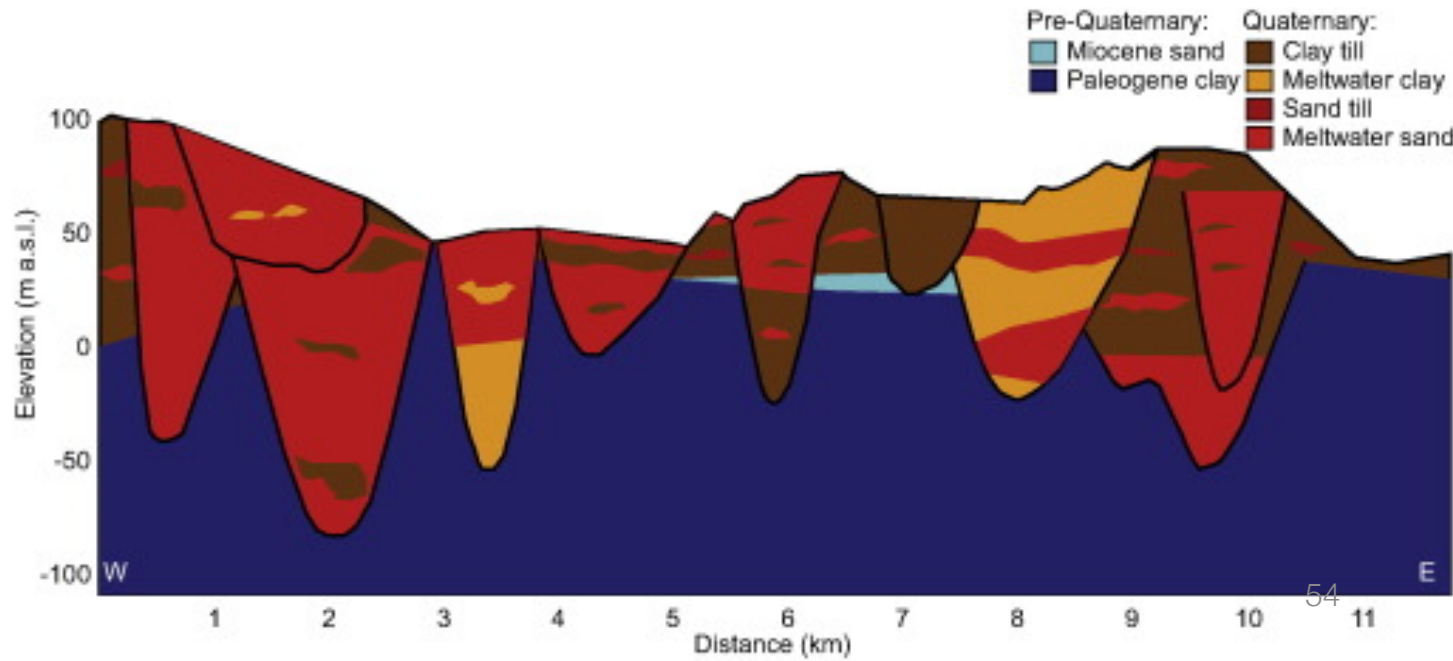
Setup

A) Survey Area:
Kasted, Denmark

B) Borehole
locations

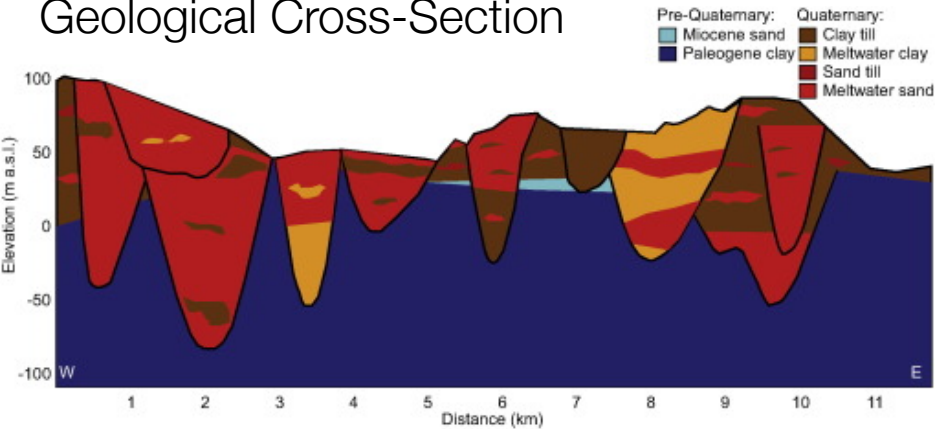


Local Geology:
W-E cross-section



Properties

Geological Cross-Section

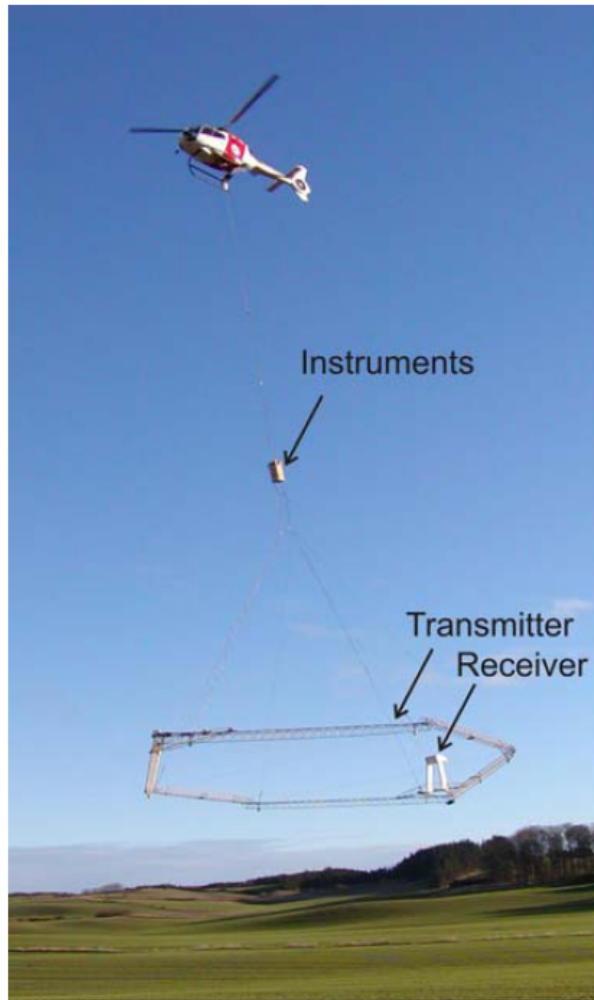


- Buried valleys with clays beneath
- Infill (water-bearing): coarse sand and gravel
- Clays are conductive (1-40 Ωm)
- Water-bearing sands and gravels are more resistive (>40 Ωm)

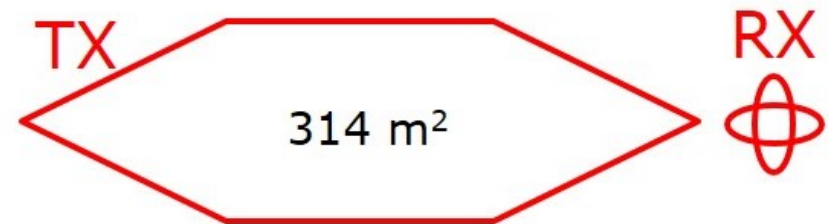
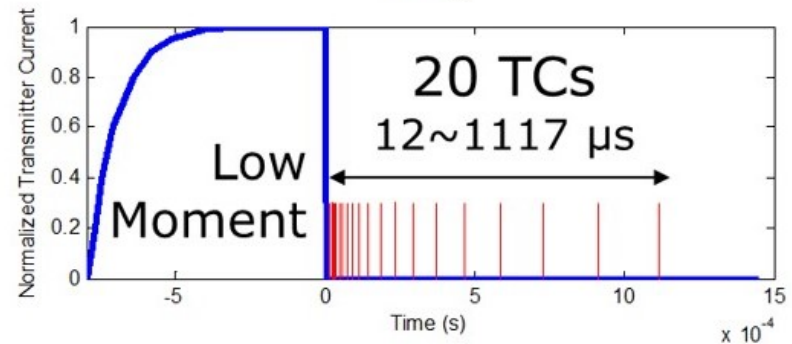
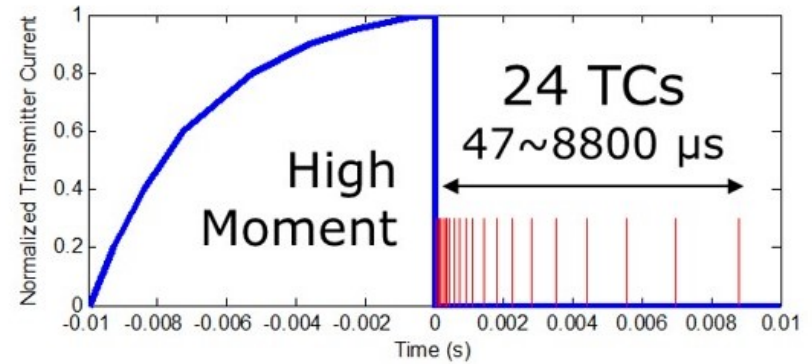
Geological Units	Resistivity (Ωm)
Palaeogene Clay	1-10
Clay Till	25-60
Sand Till	>50
Meltwater Sand and Gravel	>60
Glaciolacustrine Clay	10-40
Miocene Silt and Sand	>40
Miocene Clay	10-40
Sand	>40
Clay	1-60

Survey

SkyTEM System



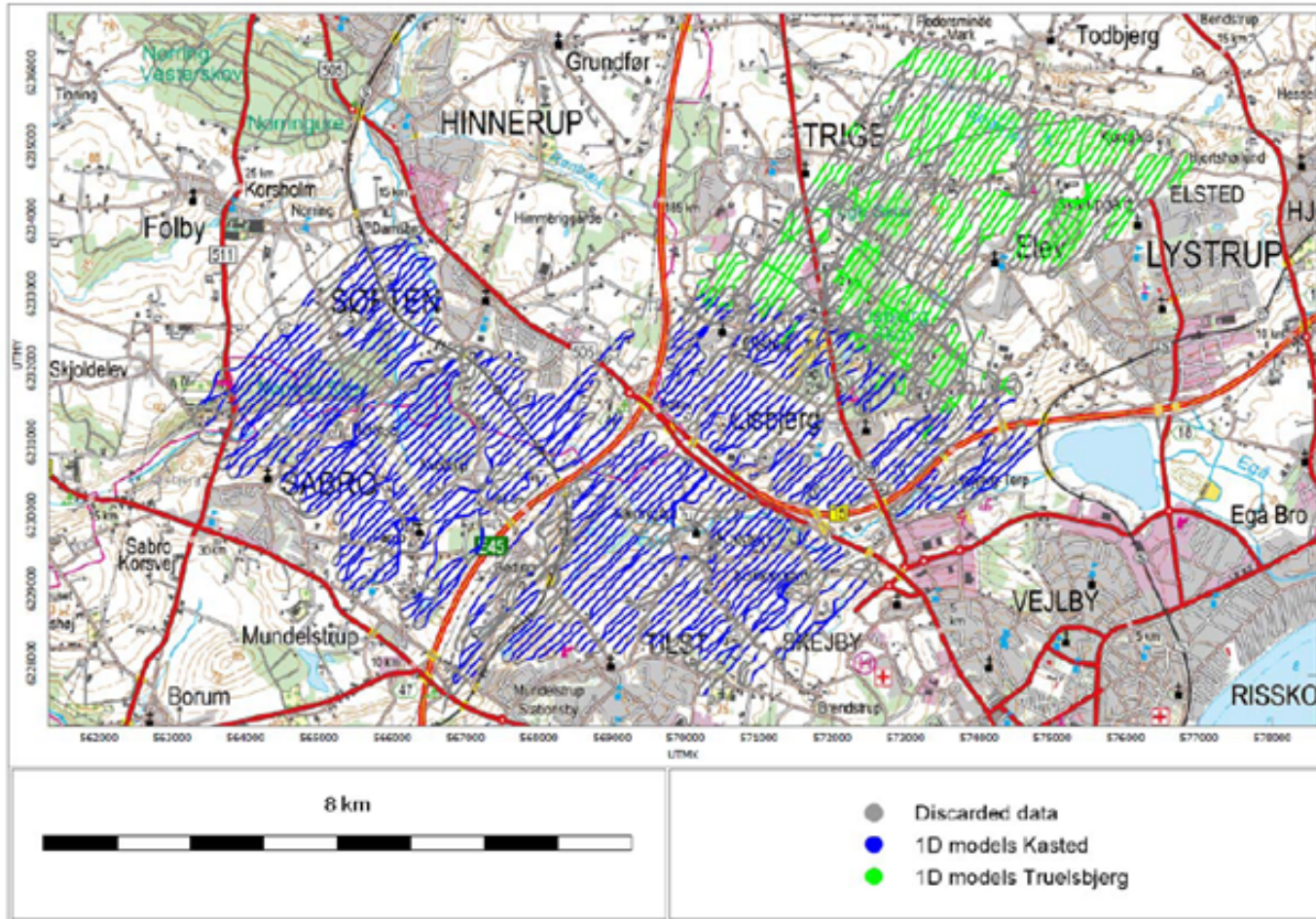
System Configuration



- Low moment (LM) used to image near surface structures
- High moment (HM) used to image deeper structures

Data

Blue: data used for Kasted study

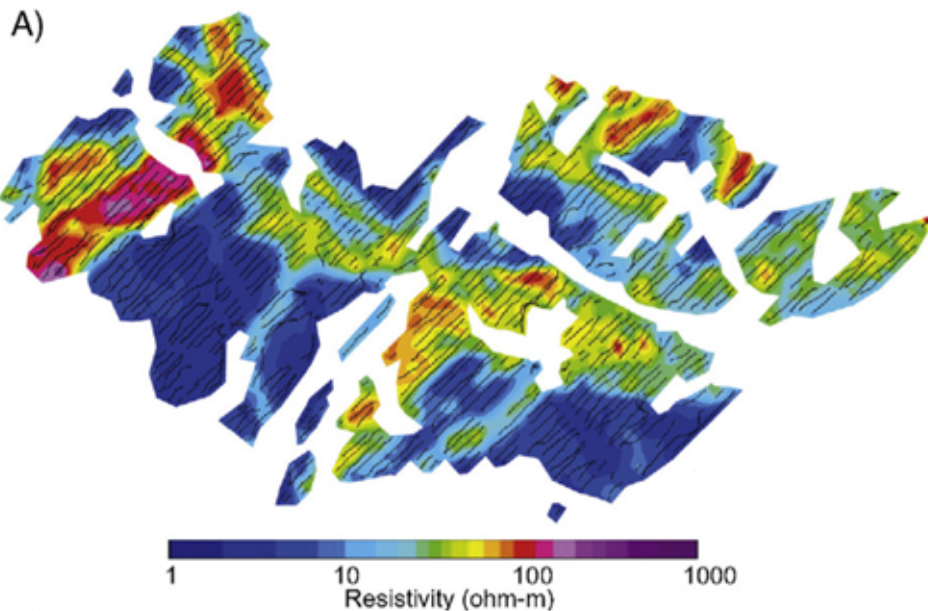


- 333 line km of data, 100 m line-spacing
- Data points with strong coupling to cultural noise were removed (~30%)

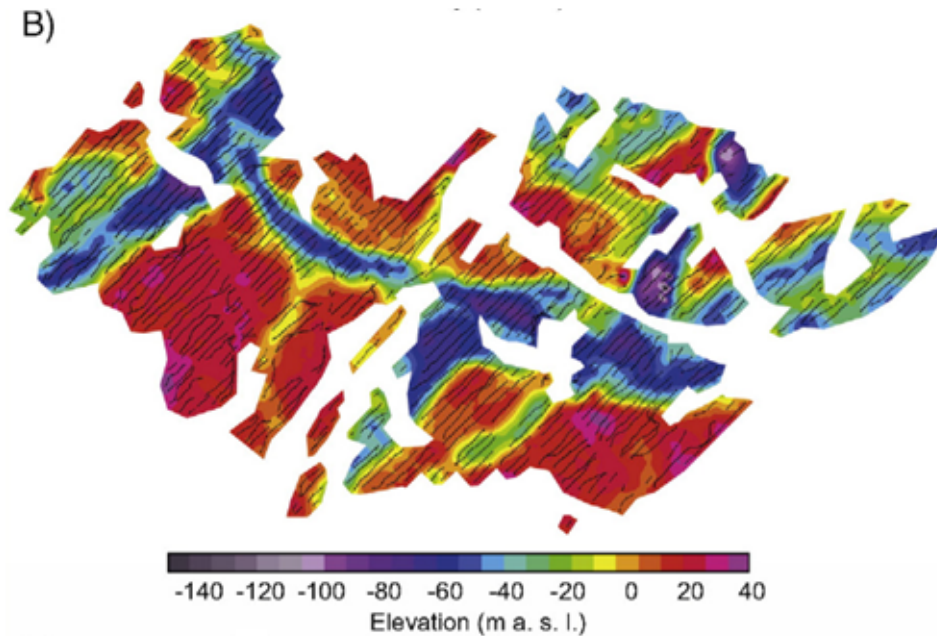
Processing (inversion)

- Spatially constrained 1D inversion → quasi-3D approach
- 9,500 soundings were inverted using 25 layers

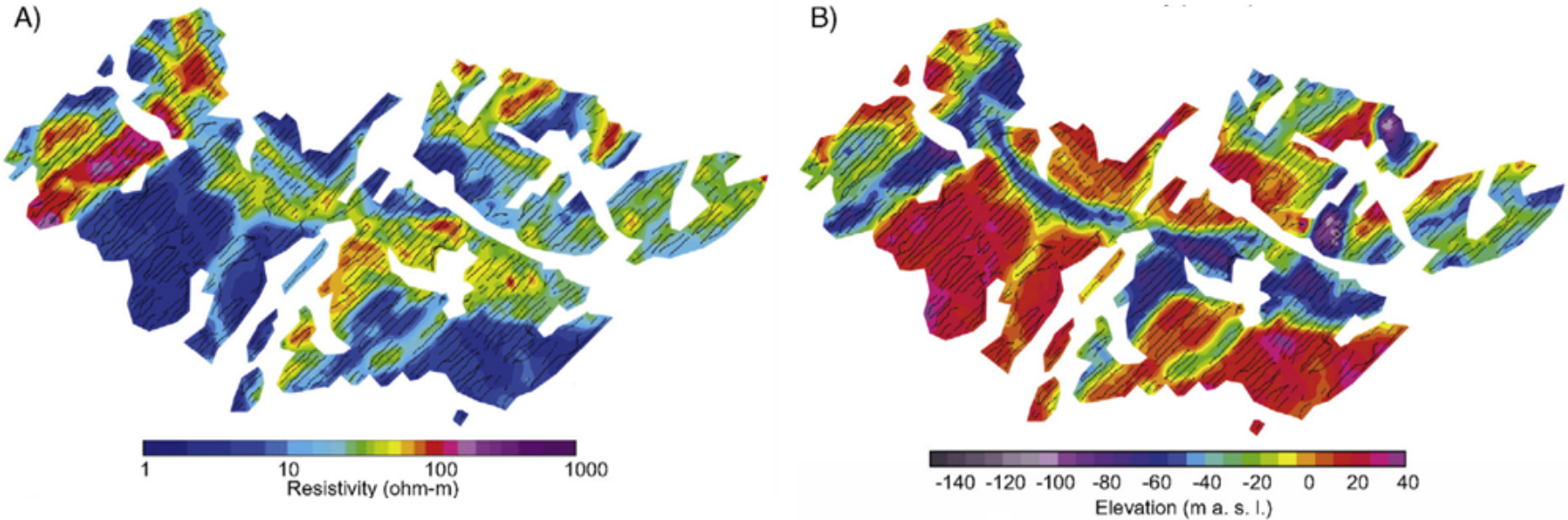
Depth slice 5 m above sea-level



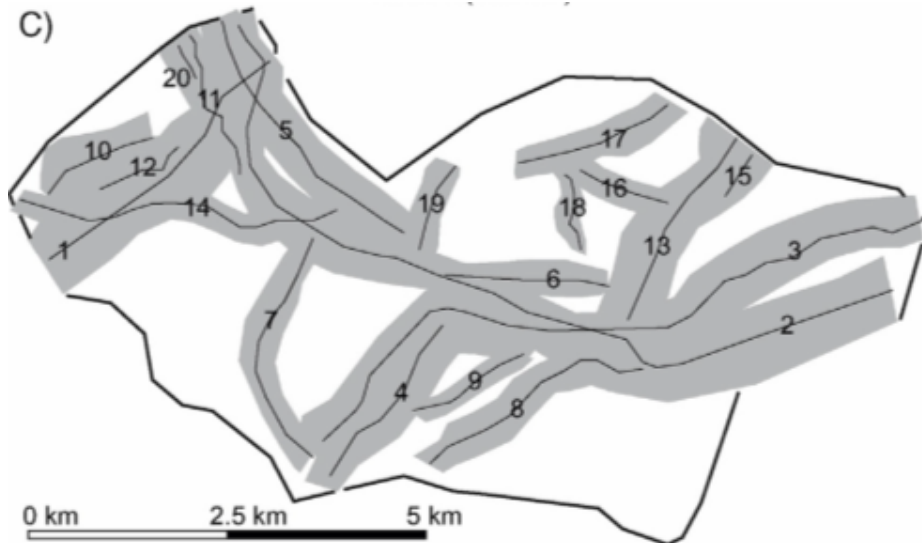
Approximate depth to the top of Paleogene clay layer



Interpretation

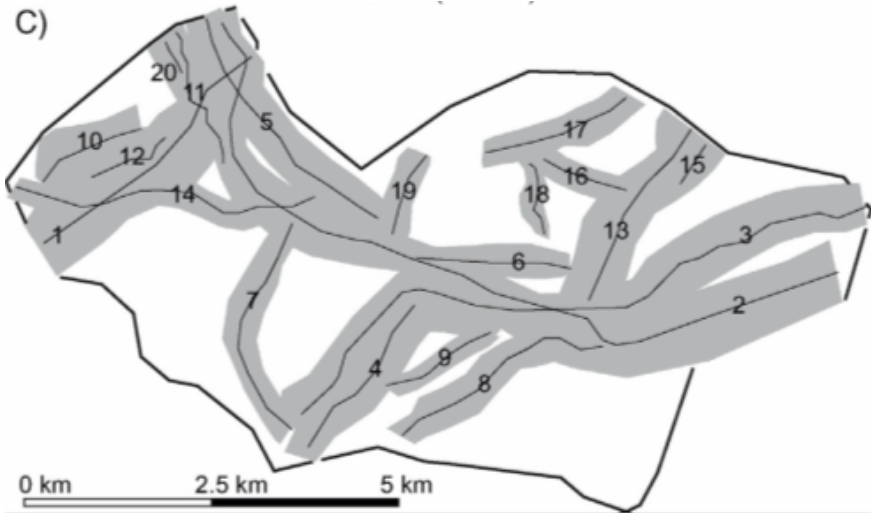


Delineation of valley structures



- Inversion results used to construct geological model.
- Delineated 20 buried and cross-cutting valley structures.

Synthesis

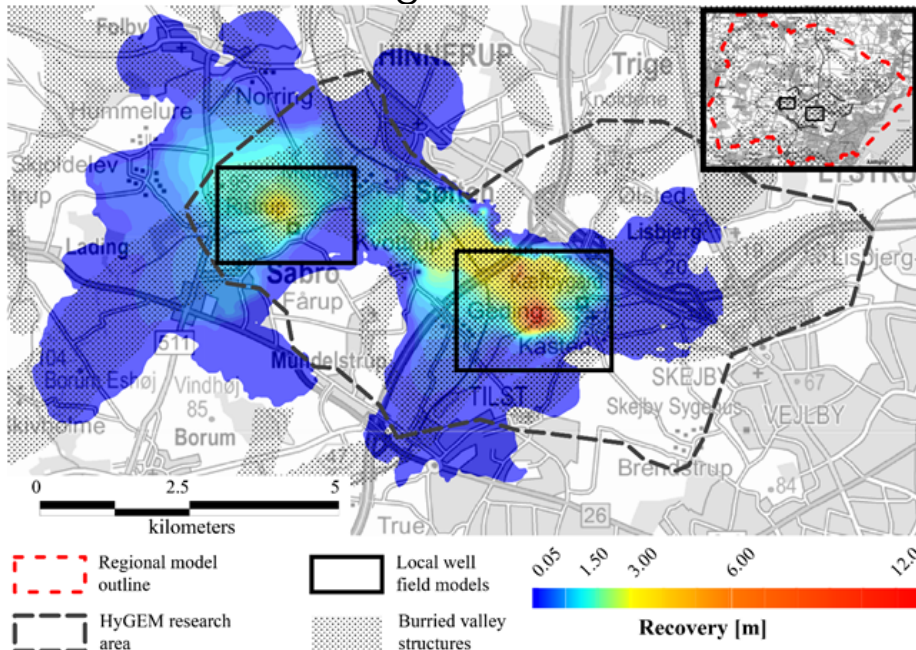


- 3D geologic model incorporated into MODFLOW-USG groundwater modeling tool

- Extracted water from 2 wells.

MODFLOW-USG groundwater model

- Dewatering between the two wells correlated with the resistive valley structures



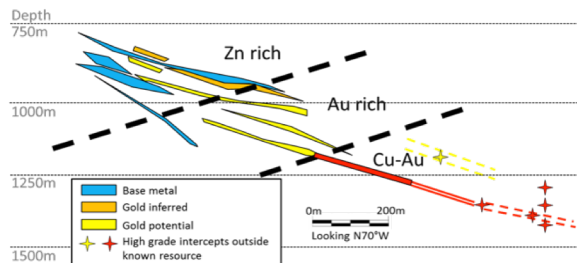
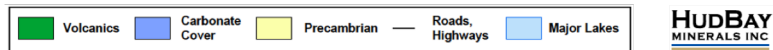
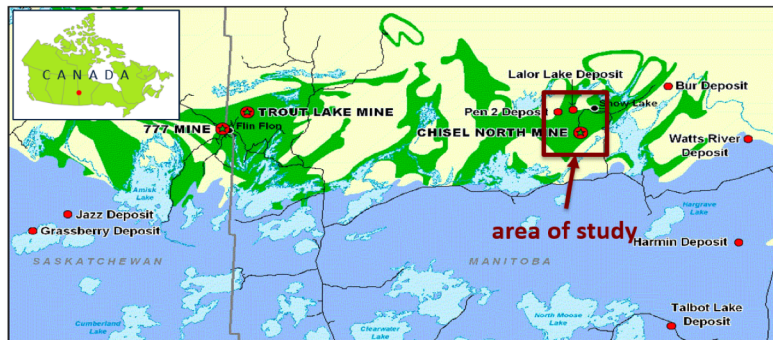
Case History: HeliSAM at Lalore

Yang & Oldenburg, 2016

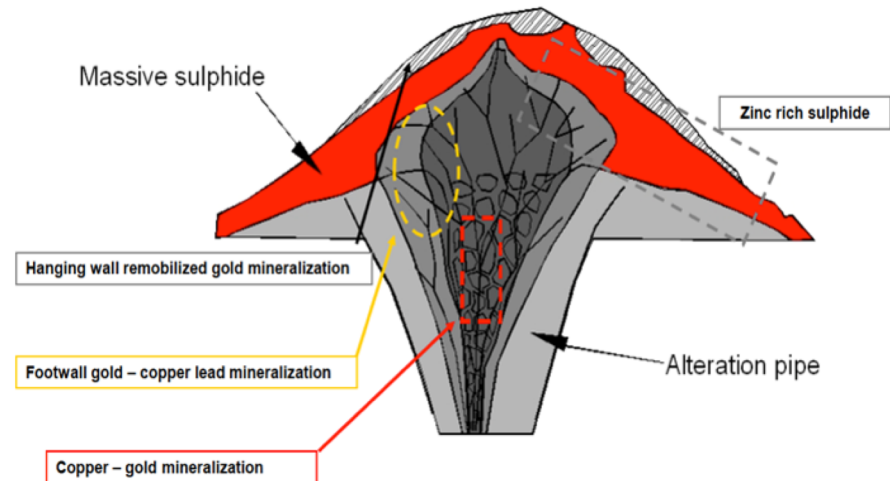
Setup

Geological framework

- Zinc-rich massive sulfides (Cap)
- Cu-Au sulfides: (stringers) within pipe
- Disseminated sulfides around deposit



Typical cross-section

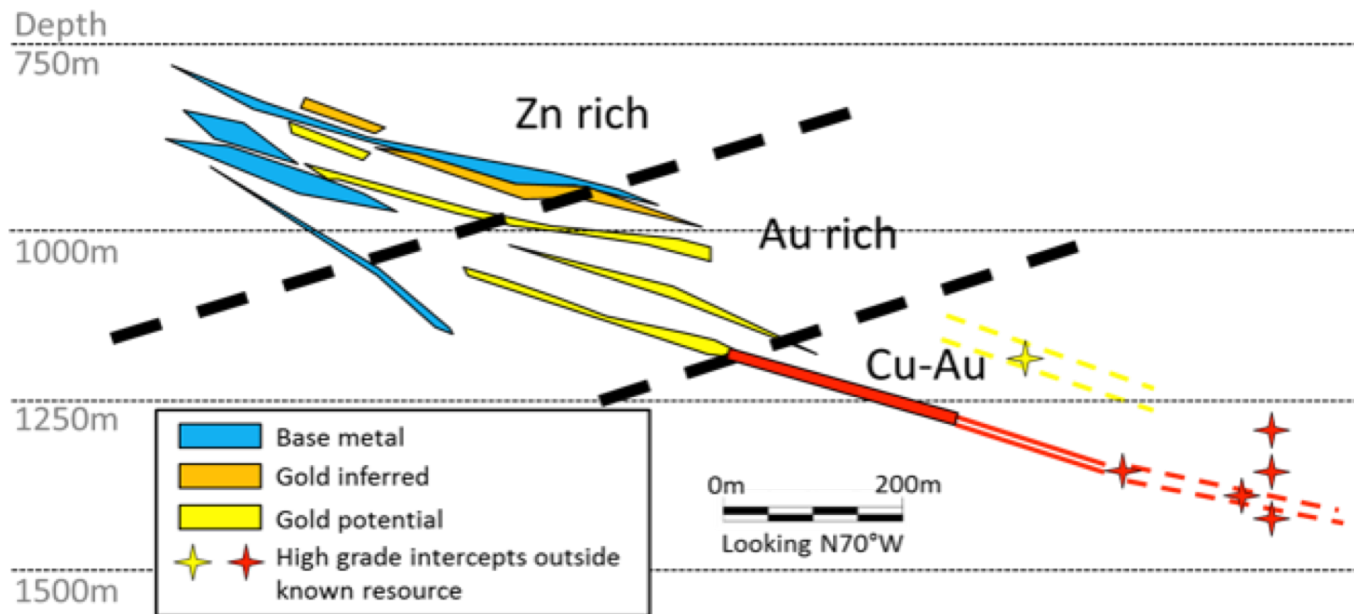


Goal:

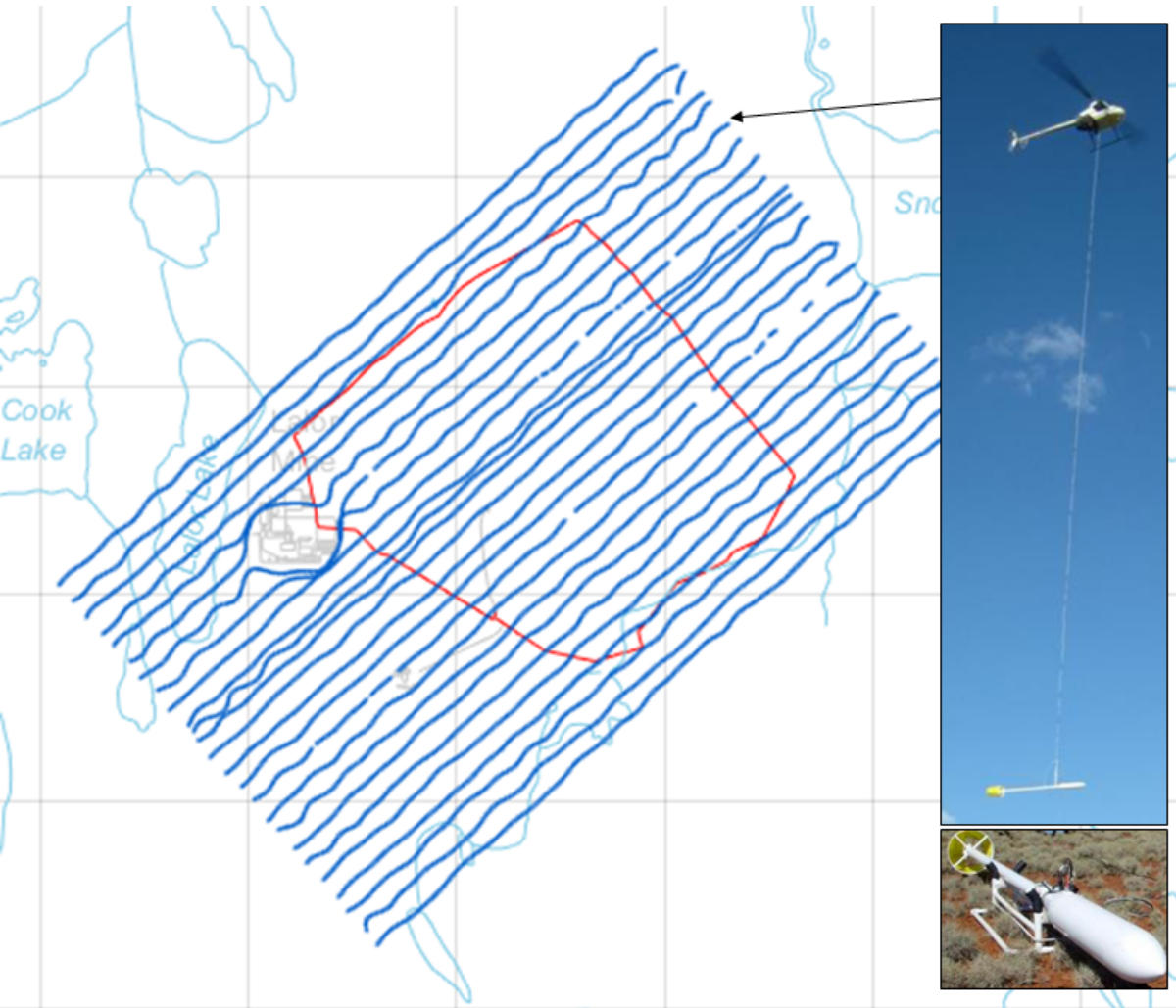
- Find deposits
- TDEM to find deeper off-hole targets

Properties

Rocks/minerals	Resistivity
volcanics	$\sim 1000 \Omega\text{m}$
sulfides	$\sim 1 \Omega\text{m}$



Survey: HeliSAM



Transmitter: (Red)

- Ground loop (~2km)
- Waveform: 7.5 Hz, 50%
- Ramp turn-off 0.4ms

Receiver:

- Cesium Vapor Mag
- 16 Time Ch: 0.42-27 ms

Flight lines: (Blue)

- 100 m spacing,
- Data every 5 m

Data

- Measure total field

$$\mathbf{B} = \mathbf{B}_0 + \mathbf{B}_a + \mathbf{B}_{em}$$

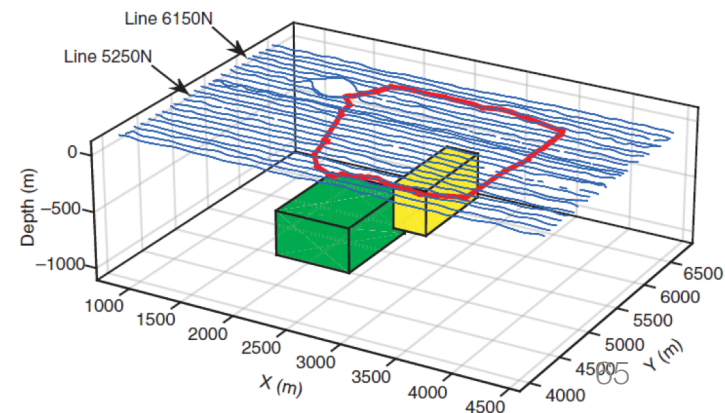
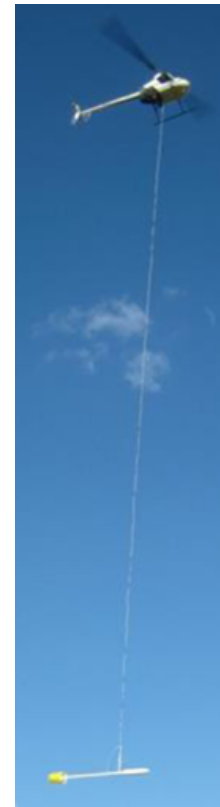
earth's magnetic field anomalous earth's field induced EM field

- Project secondary fields onto $\hat{\mathbf{B}}_0$

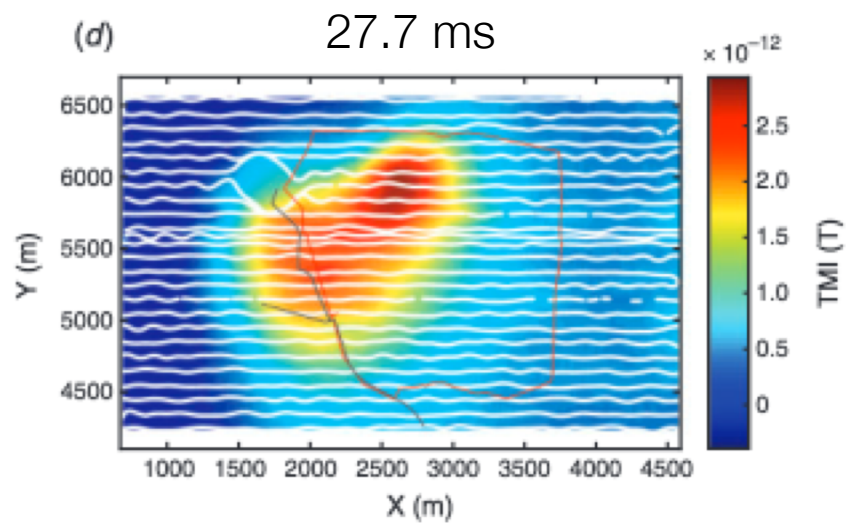
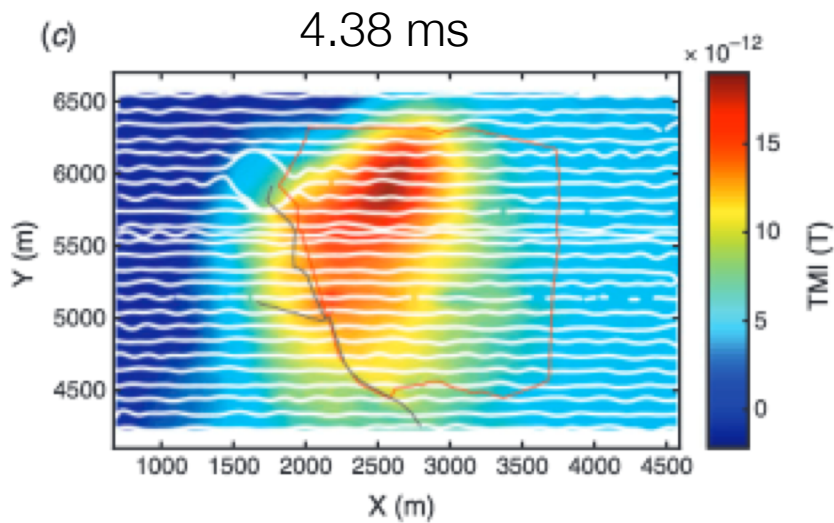
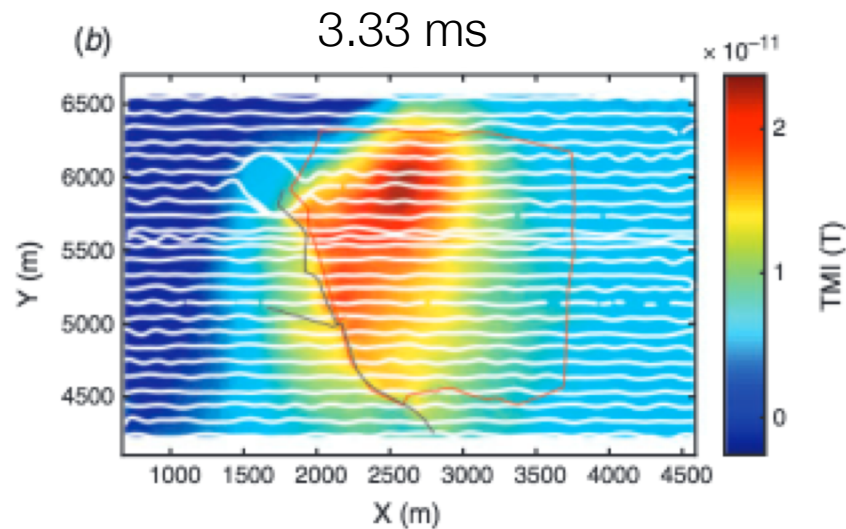
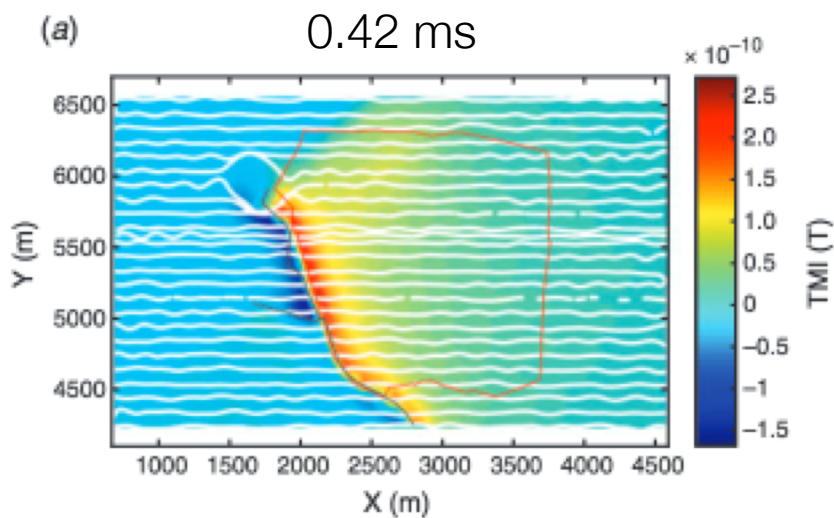
$$\begin{aligned}\Delta|\mathbf{B}| &= |\mathbf{B}_0 + \mathbf{B}_a + \mathbf{B}_{em}| - |\mathbf{B}_0| \\ &\approx (\mathbf{B}_a + \mathbf{B}_{em}) \cdot \hat{\mathbf{B}}_0\end{aligned}$$

- Change polarity on TX
- Subtract to obtain HeliSAM data

$$\Delta|\mathbf{B}| \approx \mathbf{B}_{em} \cdot \hat{\mathbf{B}}_0$$

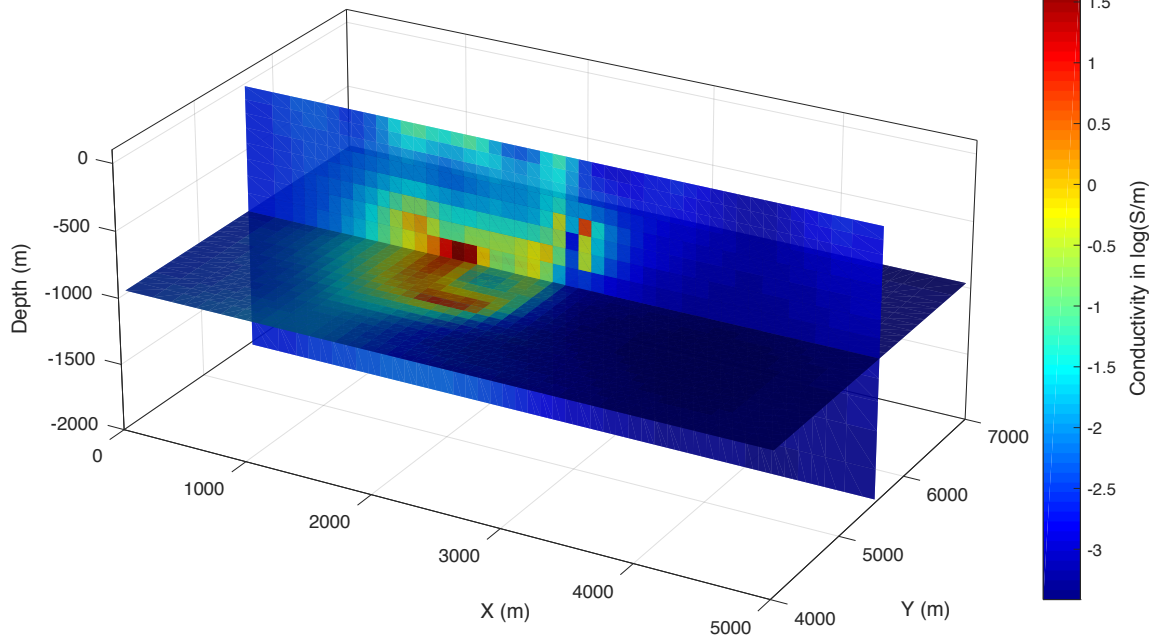
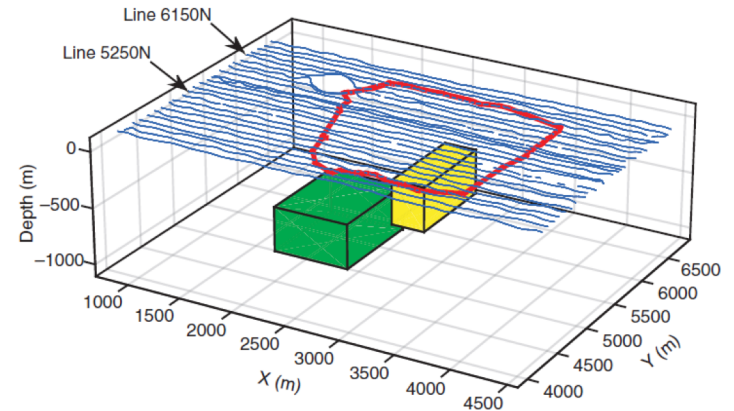


Data



Processing: Inversion of Late Time Data

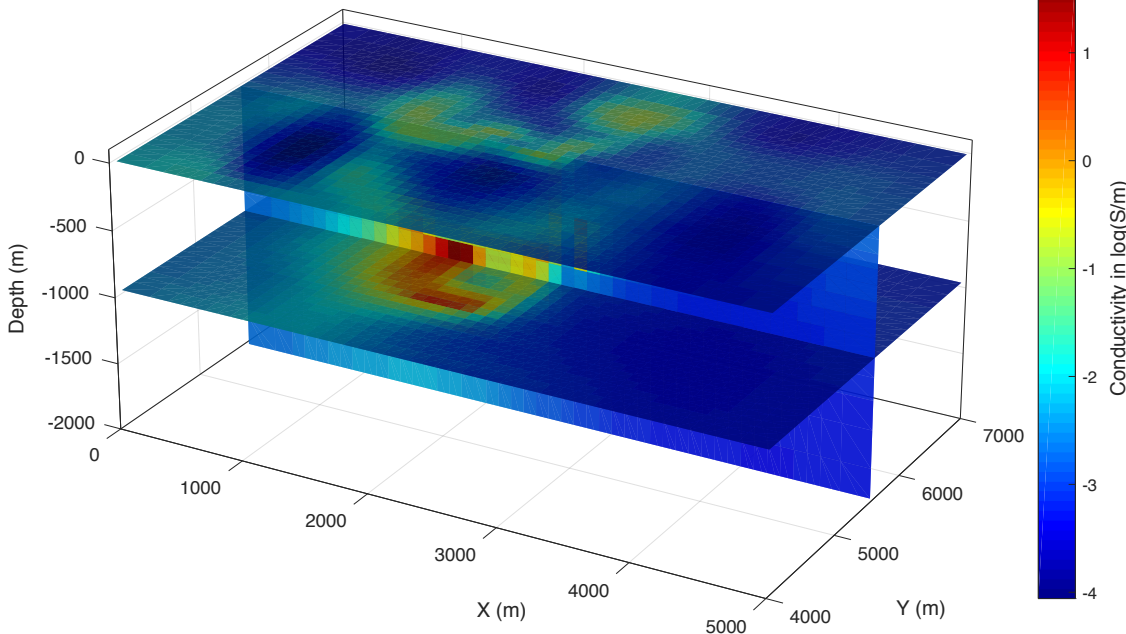
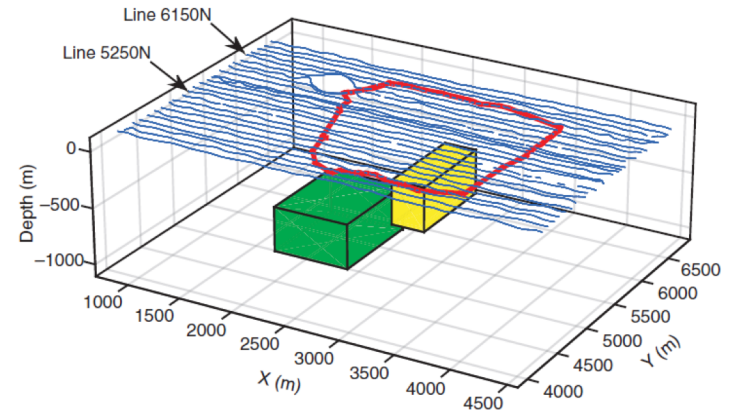
- Discard early time data
 - Contaminated by infrastructure
- Invert Time Ch 8-16 (4.44-28 ms)
- Inversion needs a “warm start”
 - Maxwell used to generate 2 prisms



- Image deep structure

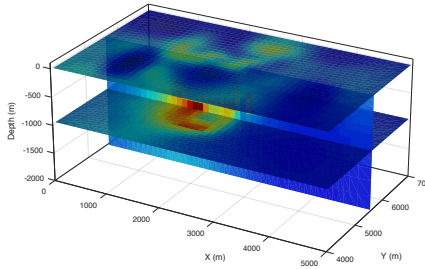
Processing: Inversion of Late Time Data

- Discard early time data
 - Contaminated by infrastructure
- Invert Time Ch 8-16 (4.44-28 ms)
- Inversion needs a “warm start”
 - Maxwell used to generate 2 prisms



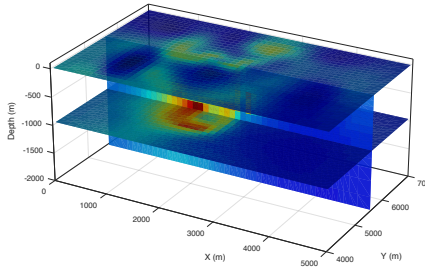
- Image deep structure
- See near surface conductive features

Processing: Inversion of Early Time Data



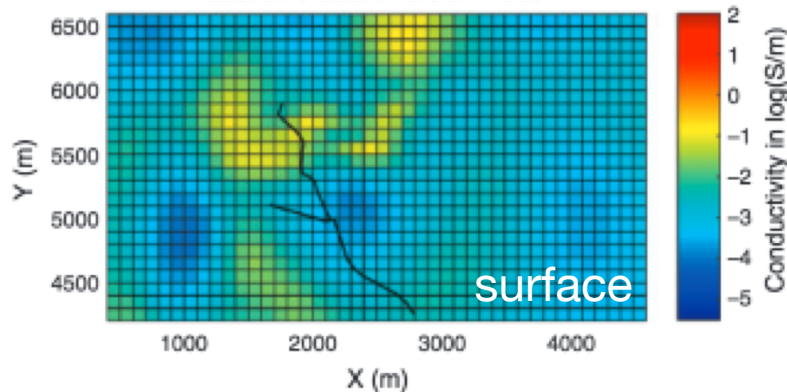
- Late-time inversion sees deep structure
- Some conductive features near surface
- **What is the effect of throwing away the early time data?**

Processing: Inversion of Early Time Data



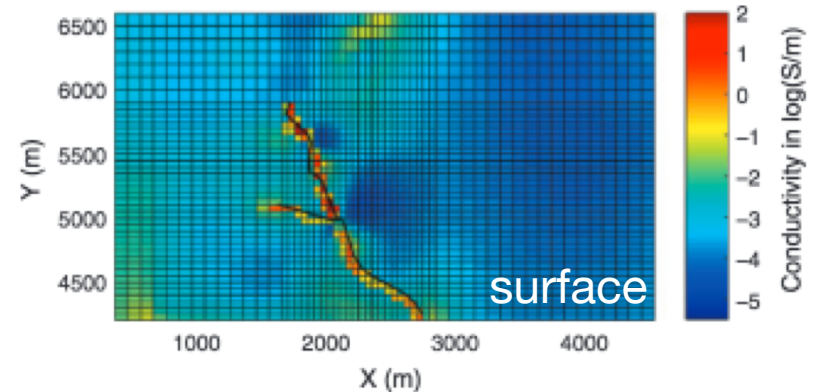
- Late-time inversion sees deep structure
- Some conductive features near surface
- **What is the effect of throwing away the early time data?**

Inverting late time data
TC 8 – 15 [4.4 – 28 ms]



- erroneous near surface structure

Inverting early time data
TC 1-7 [0.4 – 3.3 ms]

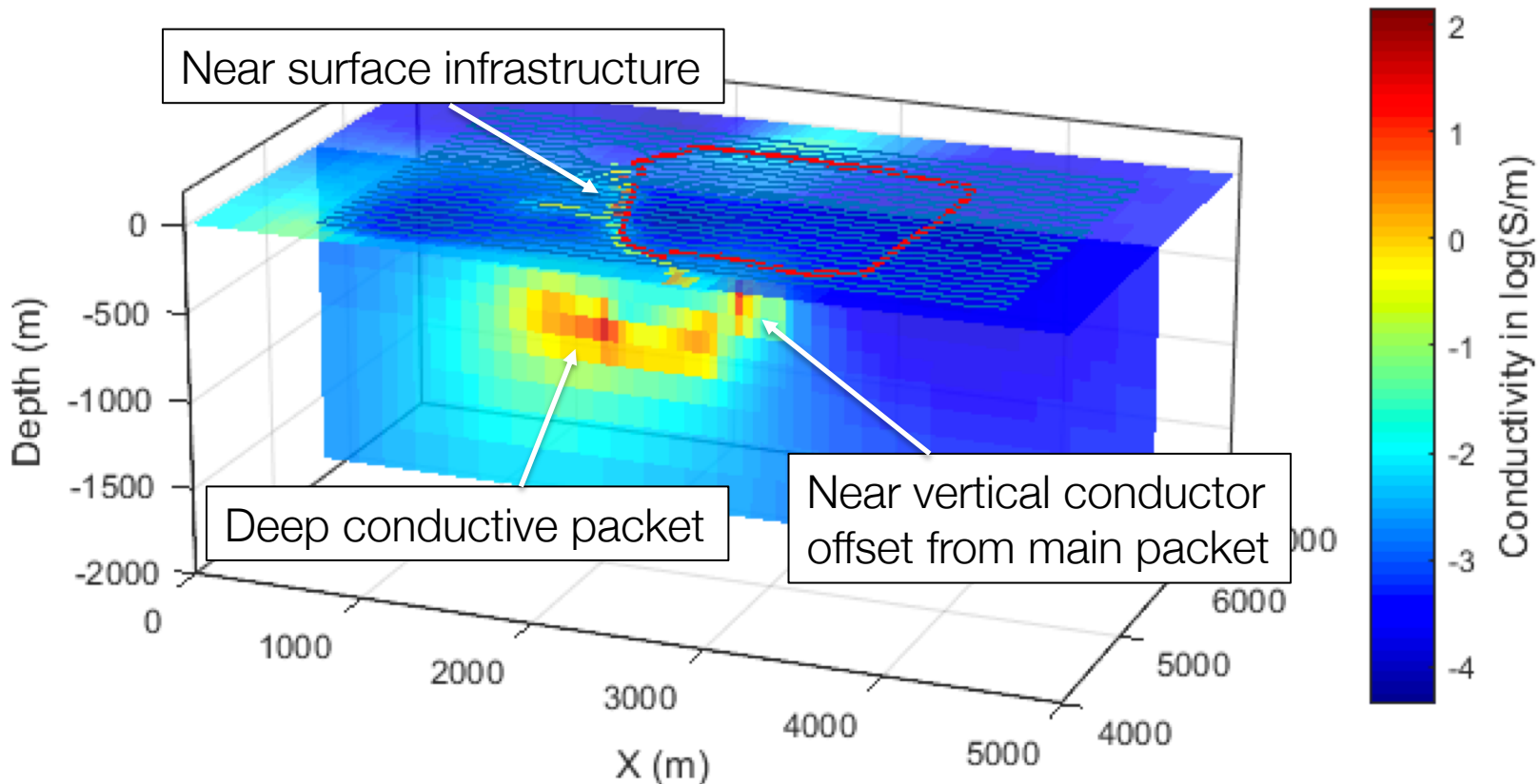


- information about infrastructure and near-surface conductivity

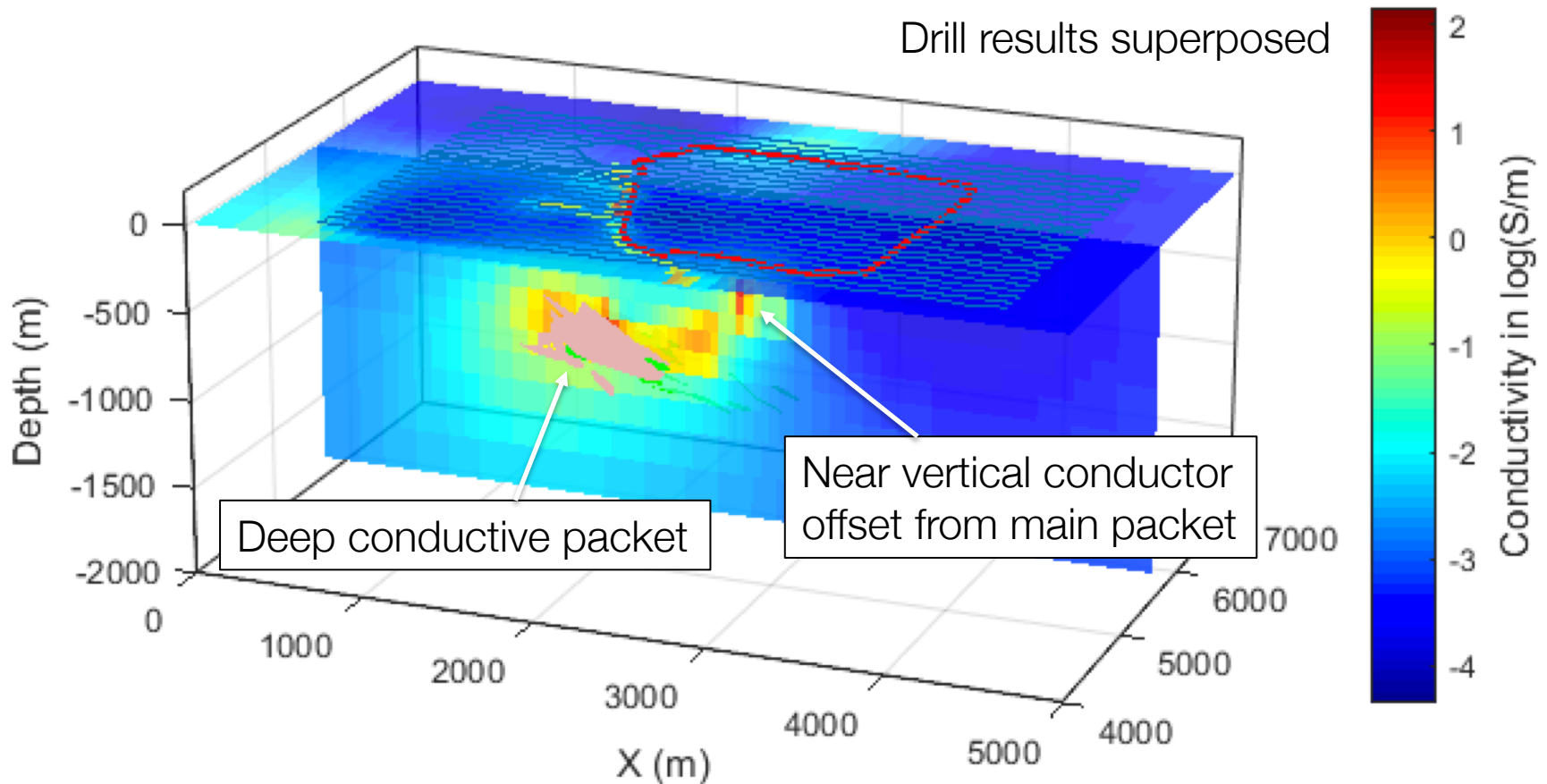
Processing: Inversion of all time channels

Starting and reference model:

- High conductivity from early time inversion
- Two conductive blocks

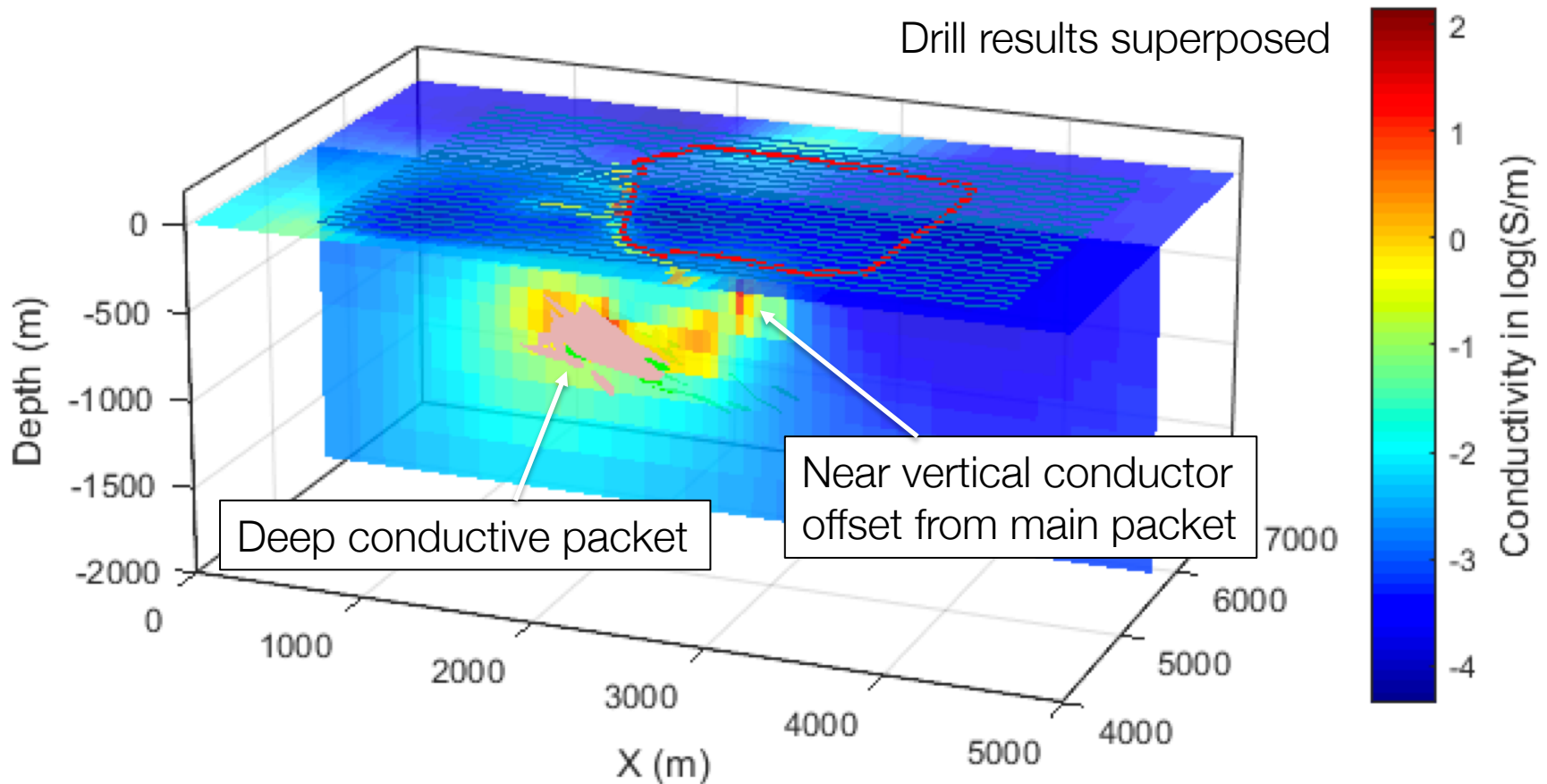


Synthesis



- Imaged main known conductive bodies
- Second conductor: recently drilled and contained sulfides (argillite)

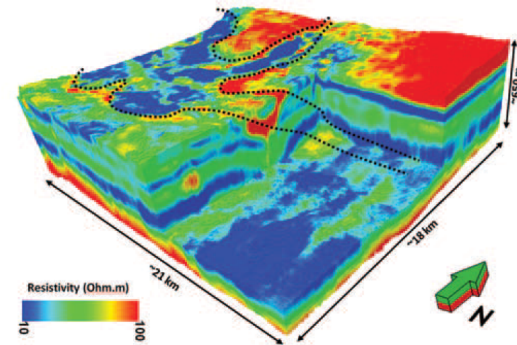
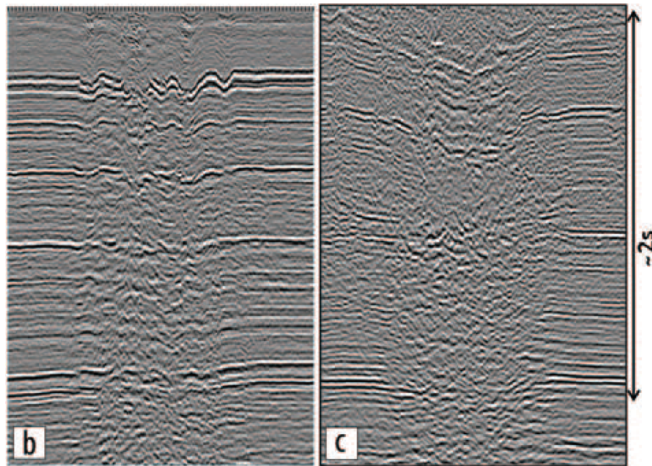
Takeaways



- Early time data:
 - constrain near surface structure infrastructure
 - Improved inversion for late time
- Warm start of inversion was necessary for deep conductors

Case History: Wadi Sahba

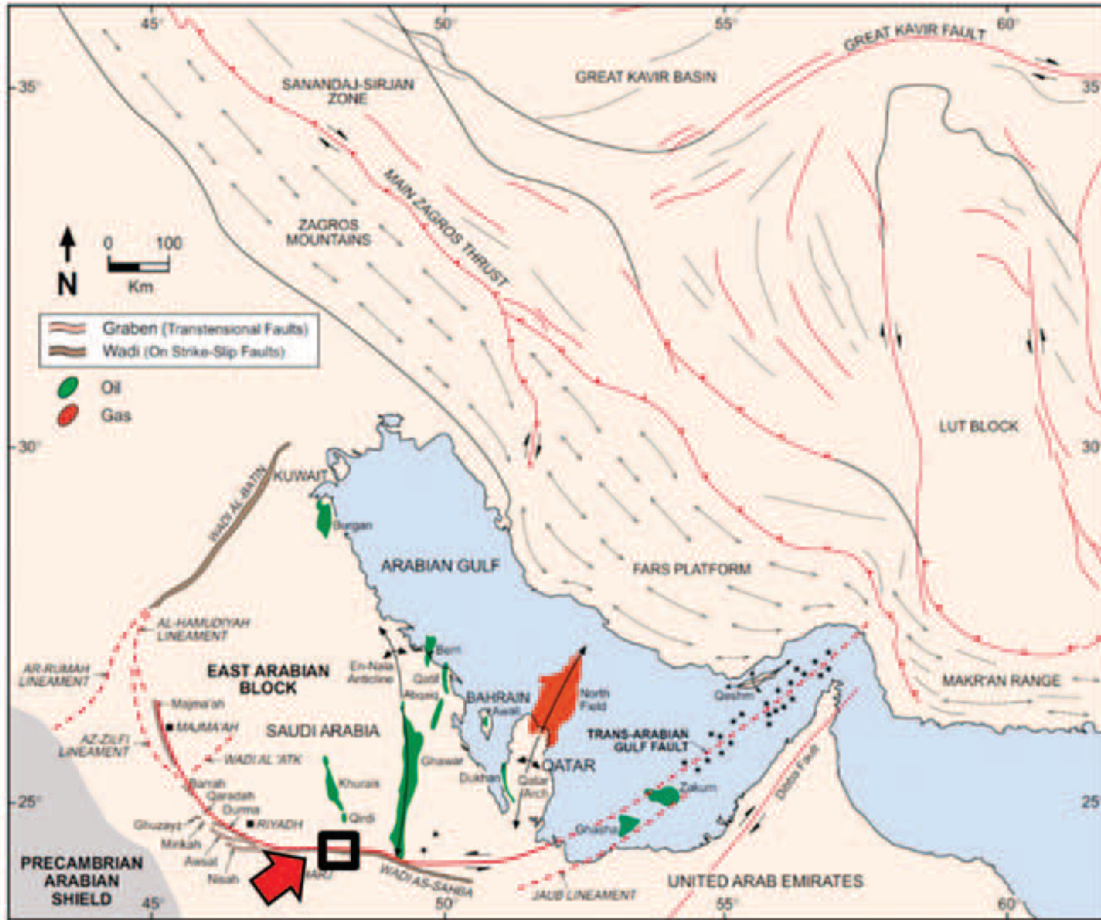
Colombo et al. 2016



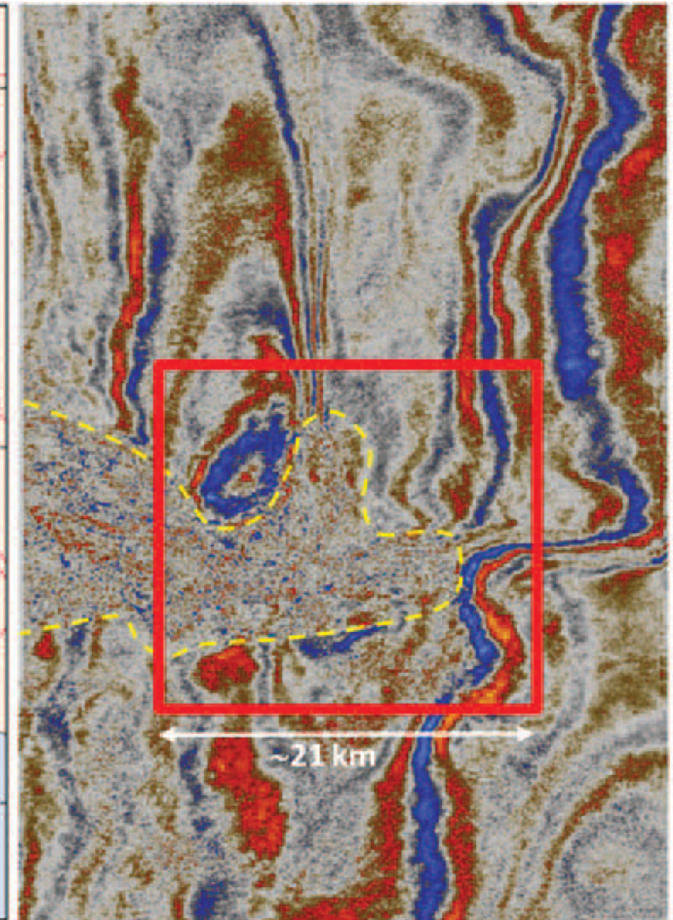
?

Setup

Location of Wadi area, Saudi Arabia



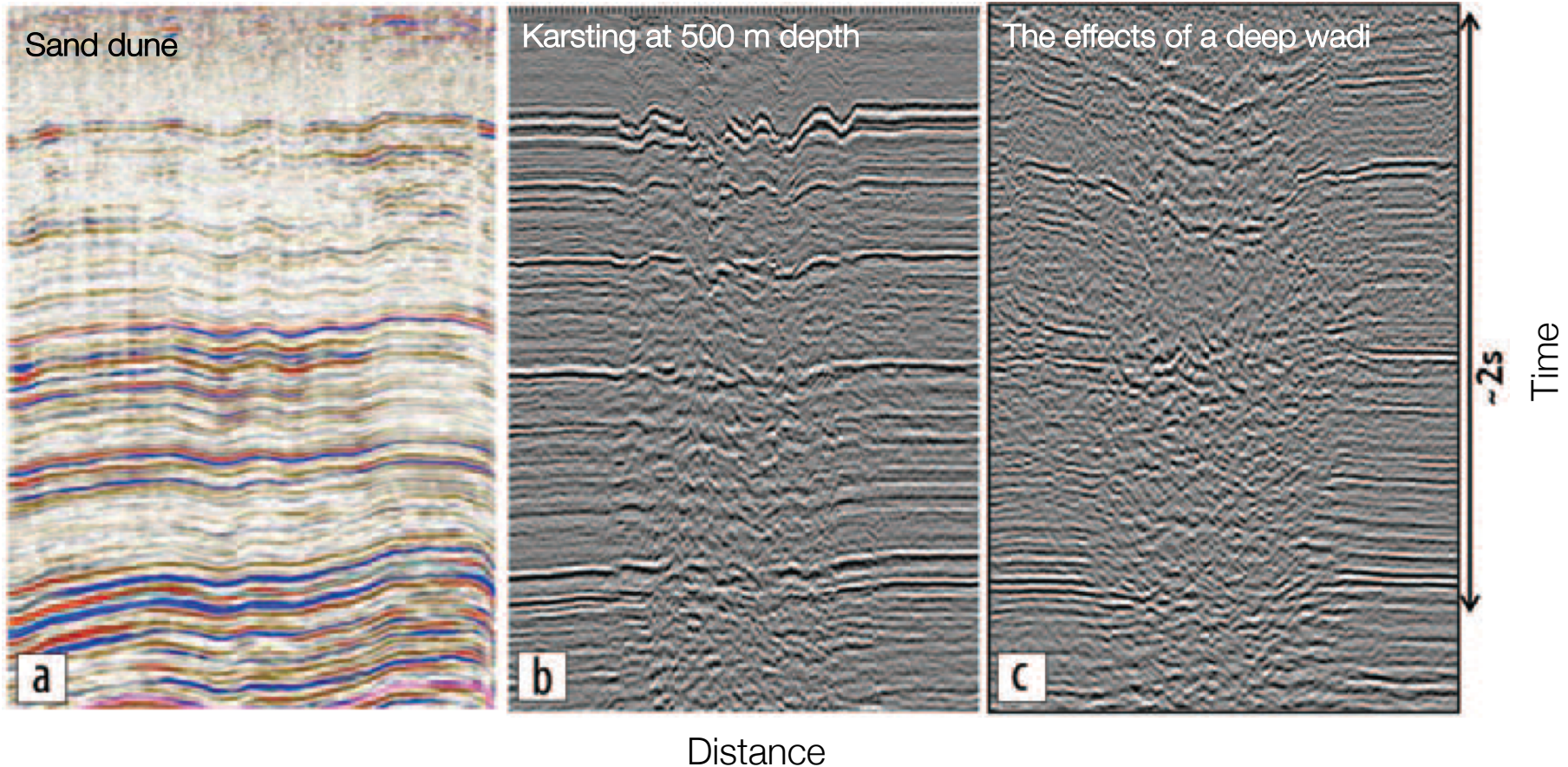
Seismic slice



- Oil and gas exploration in the Middle East: Focus is now stratigraphic traps and low relief structures

Challenges for processing seismic data

Example seismic sections



- Strong effects from near surface anomalies even after static corrections

Properties

- P-velocity and conductivity:

$$v_p = g(\phi)$$

v_p : P-velocity

$$\sigma = f(\phi)$$

ϕ : porosity

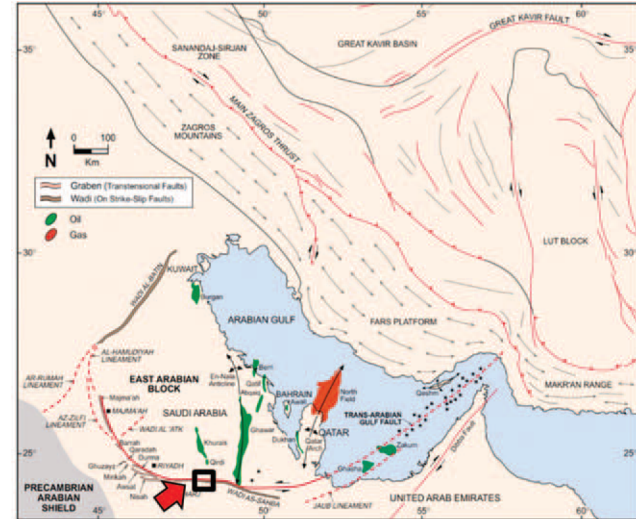
- Poor seismic data:

- strong scattering effects probably caused by flower faults
- velocity inversions (high to low v_p)

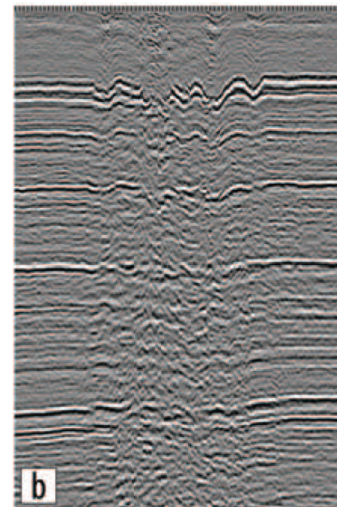
- From previous multi-physics analyses:

- strong structural similarity between the inverted resistivity, and the existing seismic results

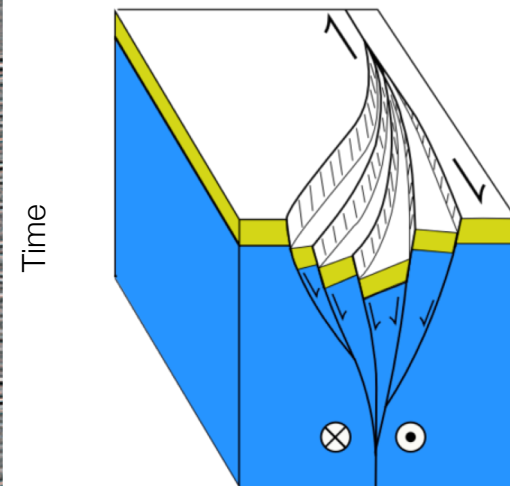
Geologic map



Seismic section



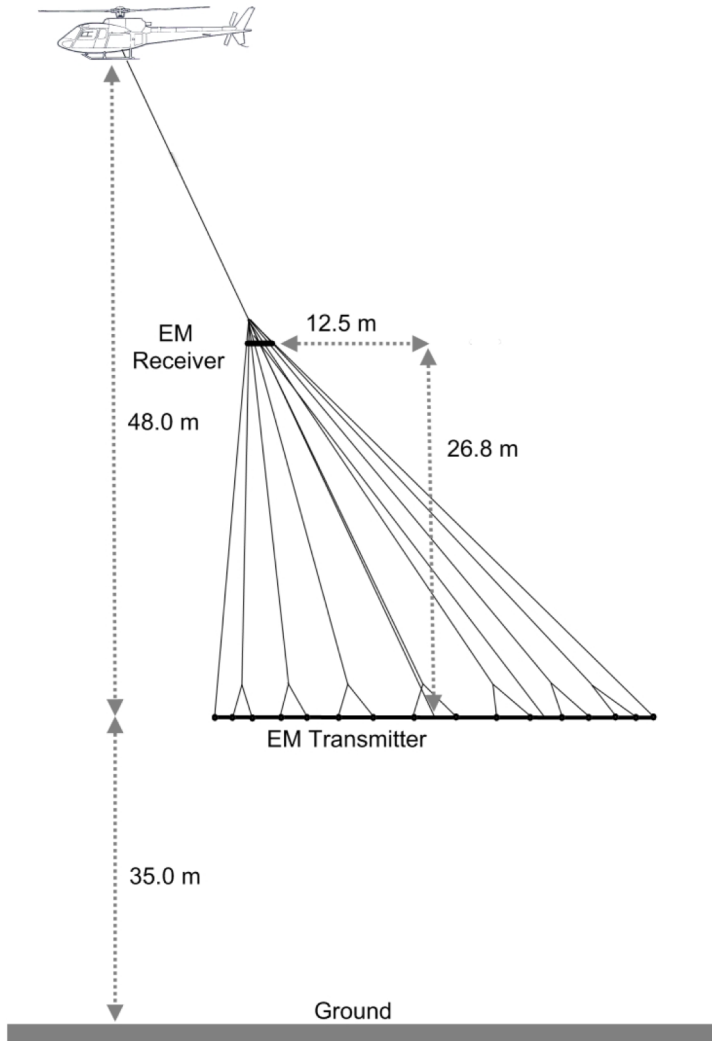
Flower faults



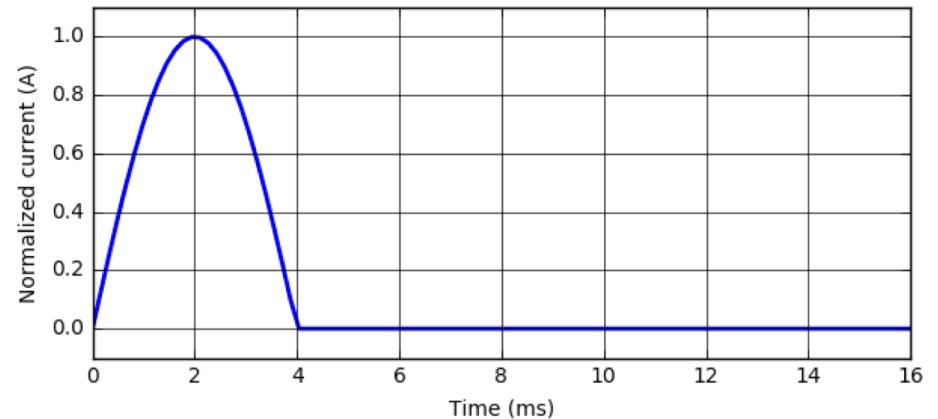
Distance

Survey

HELITEM

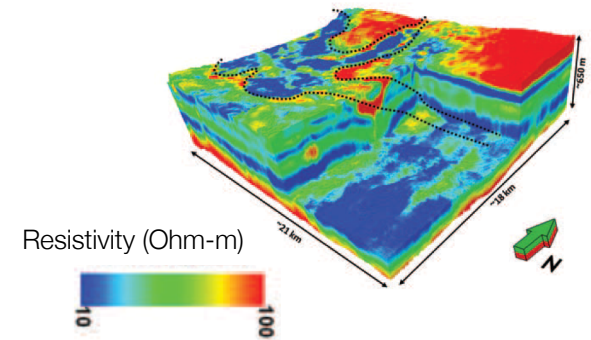


System Configuration

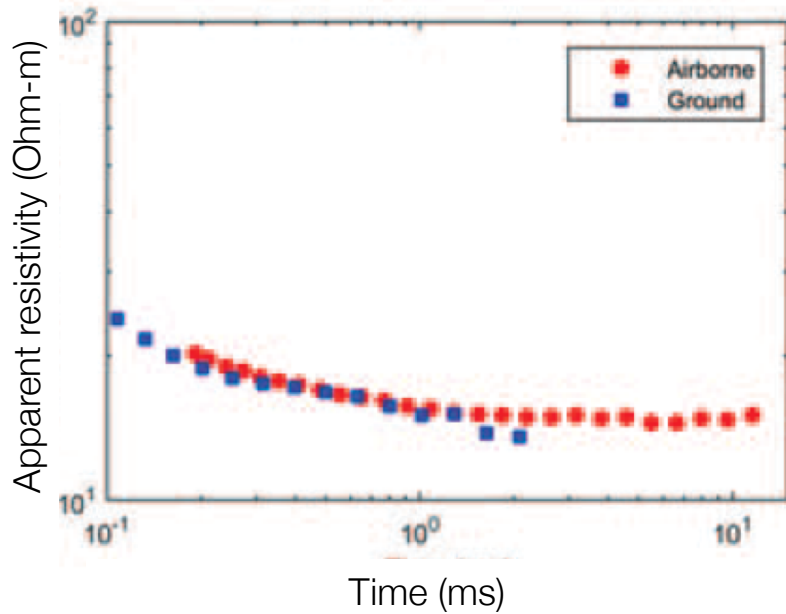


- Peak Tx current: 1200 A
- Dipole moment: 1.7×10^6 A-m²
- Stacked TEM curve spacing: ~2.7 m
- Total soundings: ~1.6 million

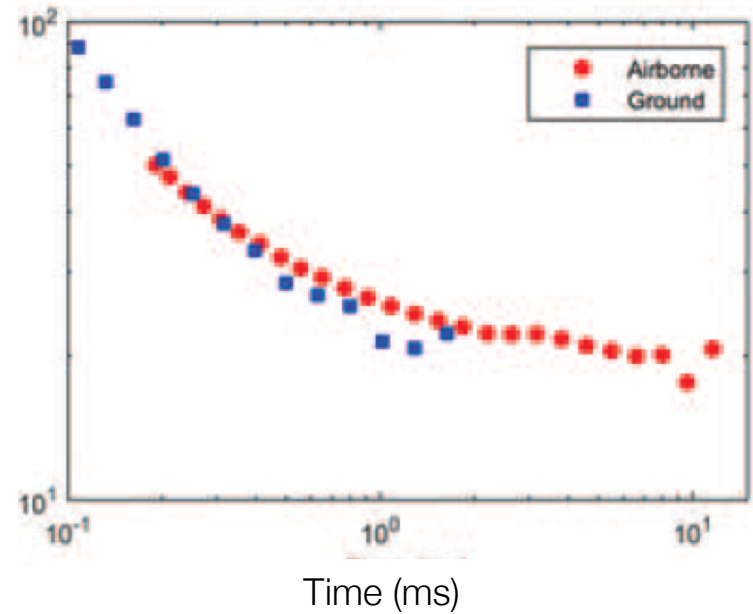
Comparisons: airborne and ground EM



Conductive area

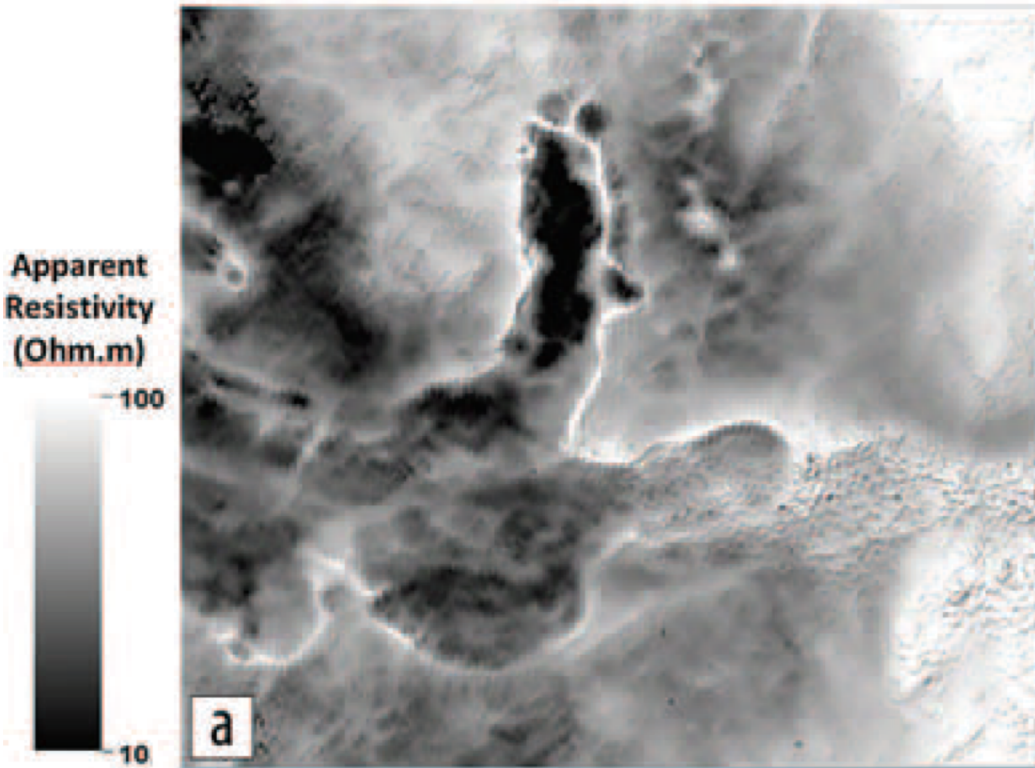


Resistive area



EM data

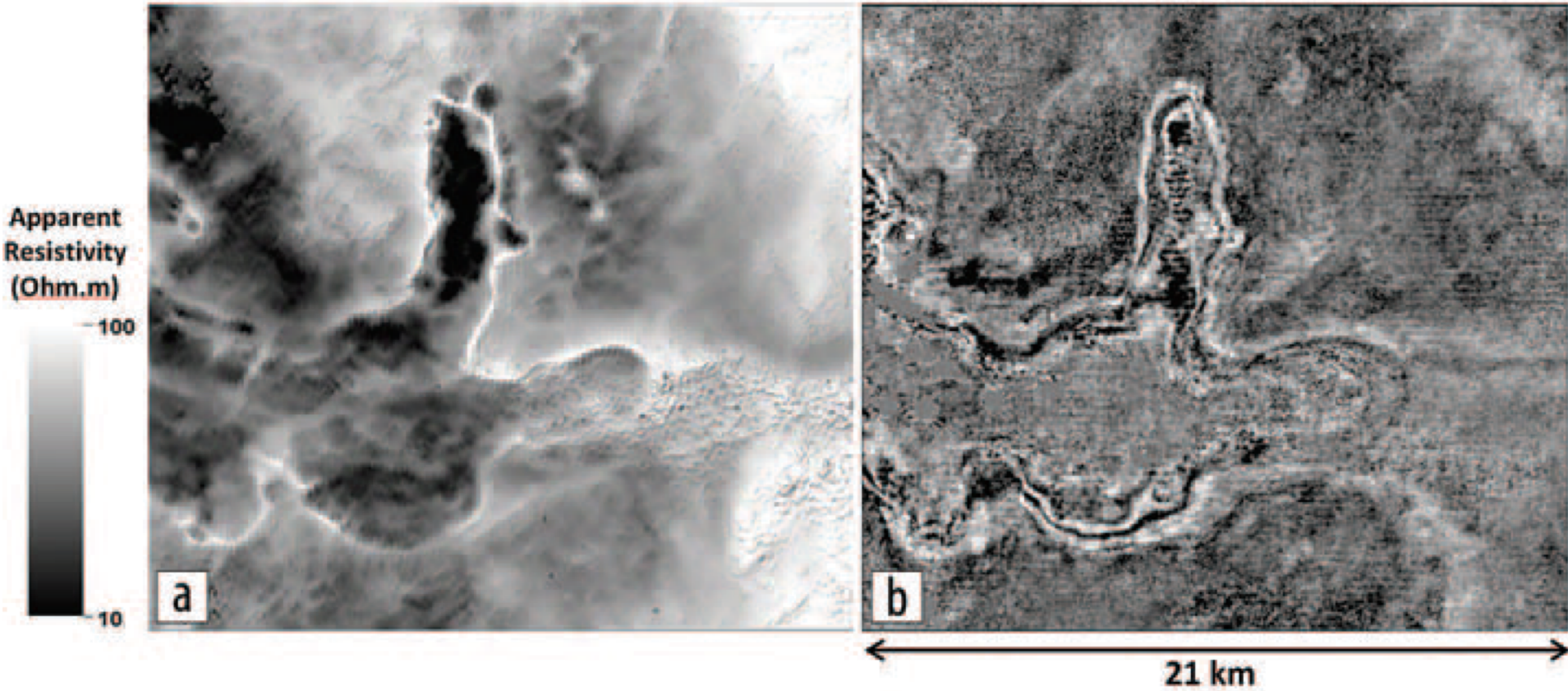
Apparent resistivity map



Comparison: EM and Seismic data

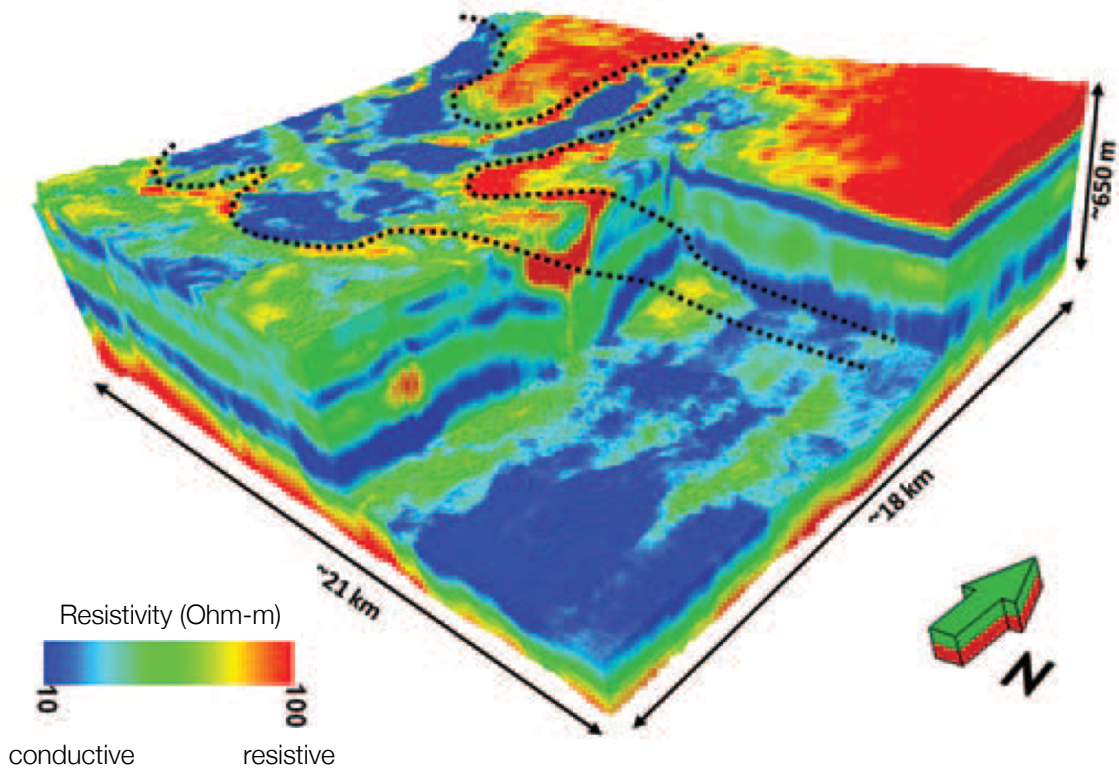
Apparent resistivity map

Seismic time slice



Processing: EM inversion

Conductivity model



- 1D inversion for each sounding location
- Lateral constraint is used

Cooperative inversion: Seismic + EM

- How EM can help seismic tomography inversion?

Velocity (v_p): high to low (significant challenge)

Conductivity (σ): high to low

$$v_p = g(\phi)$$
$$\sigma = f(\phi)$$

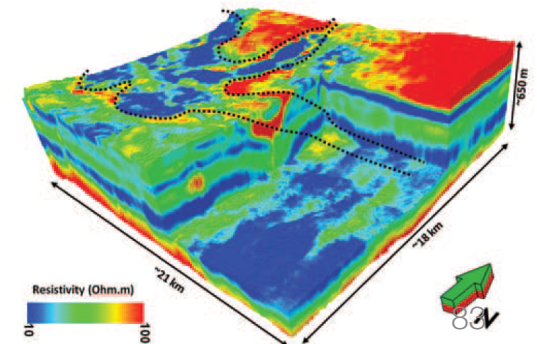
ϕ : porosity

\mathbf{m}_s : Slowness
 \mathbf{m}_σ : Conductivity

$$\psi(\mathbf{m}_s, \mathbf{m}_\sigma) = \psi_m(\mathbf{m}_s) + \frac{1}{\lambda_1} \psi_d(\mathbf{m}_s) + \frac{1}{\lambda_2} \psi_x(\mathbf{m}_s, \mathbf{m}_\sigma) + \frac{1}{\lambda_3} \psi_{rp}(\mathbf{m}_s, \mathbf{m}_\sigma)$$

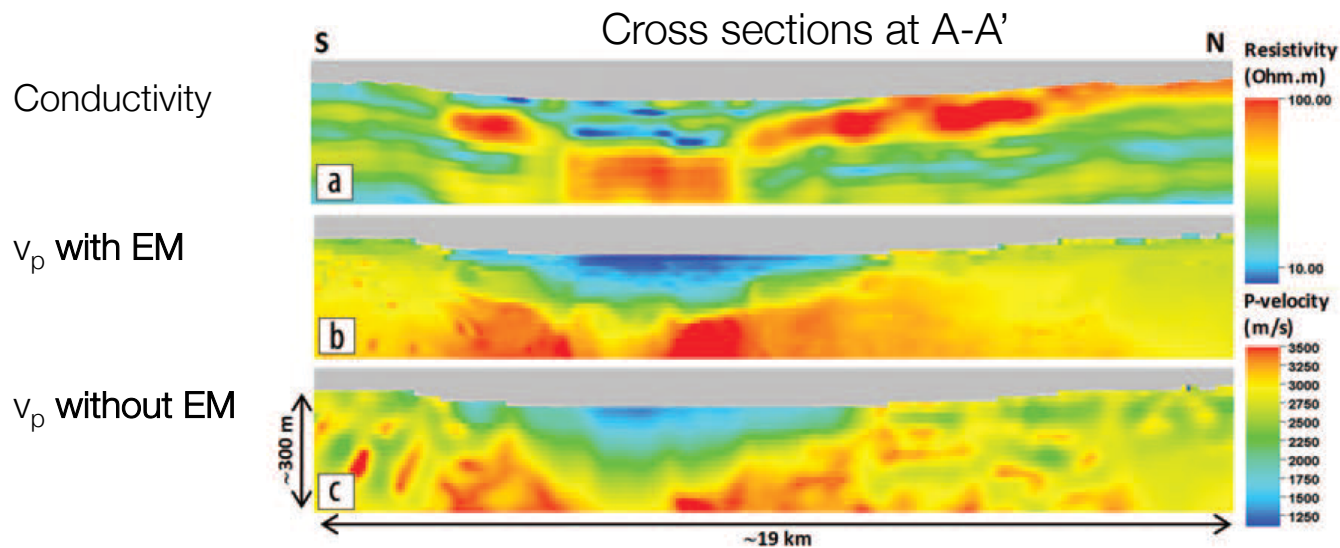
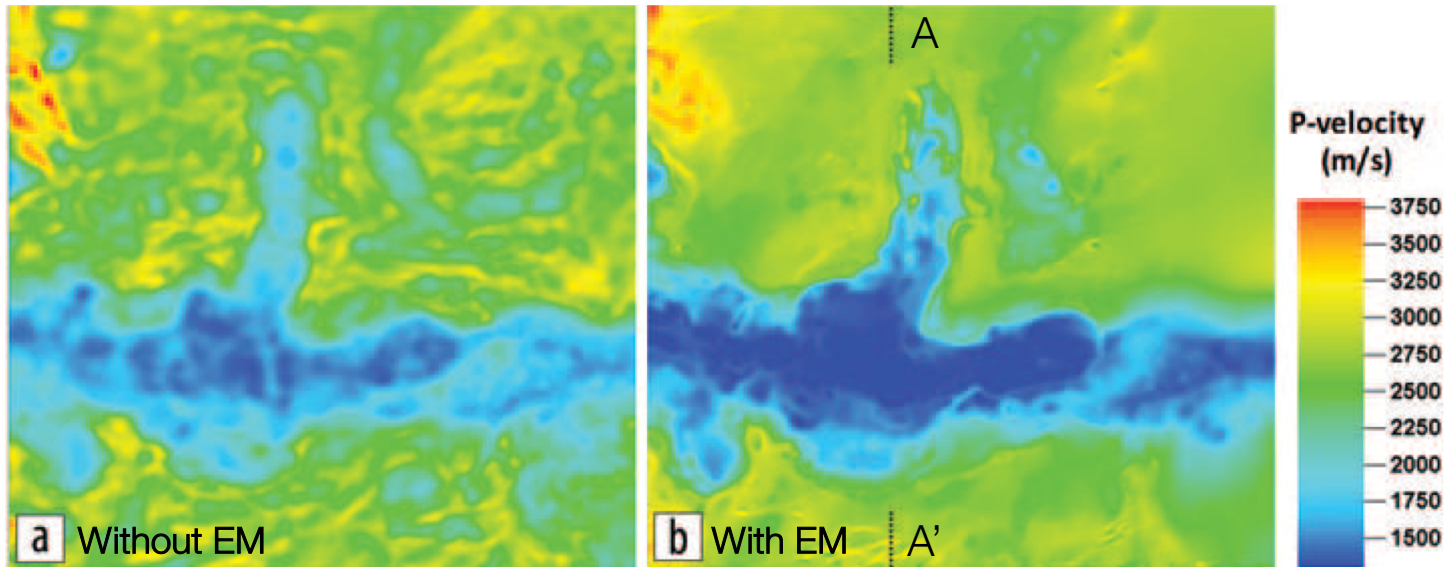
$$\|\nabla \mathbf{m}_s \times \nabla \mathbf{m}_\sigma\|_2^2$$

Gallardo and Meju, 2004



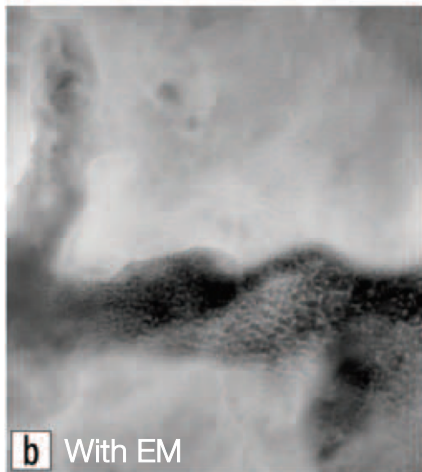
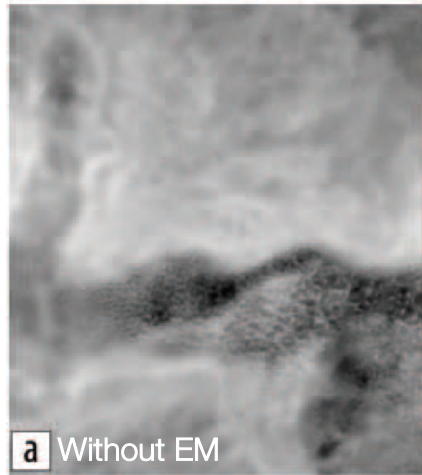
Cooperative inversion: Seismic + EM

V_p depth slices at 340 m below sea level

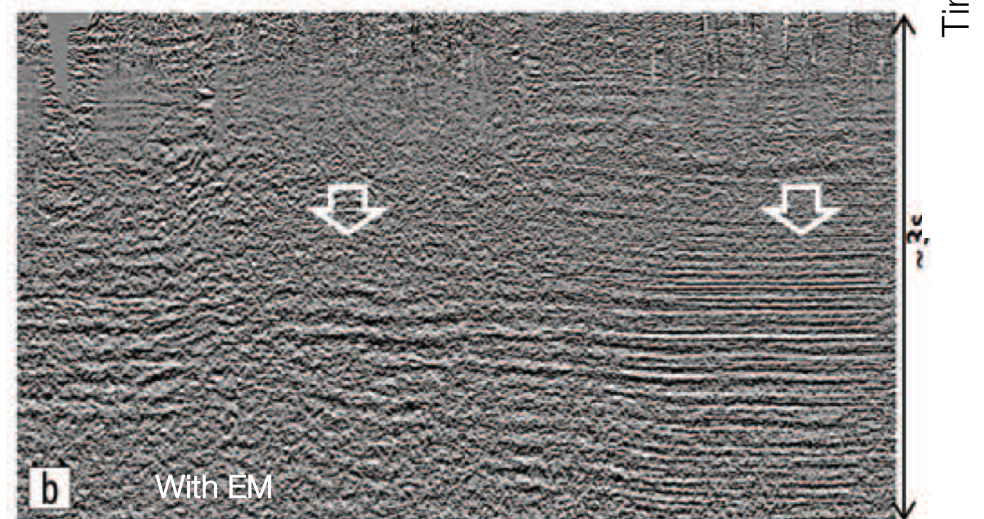
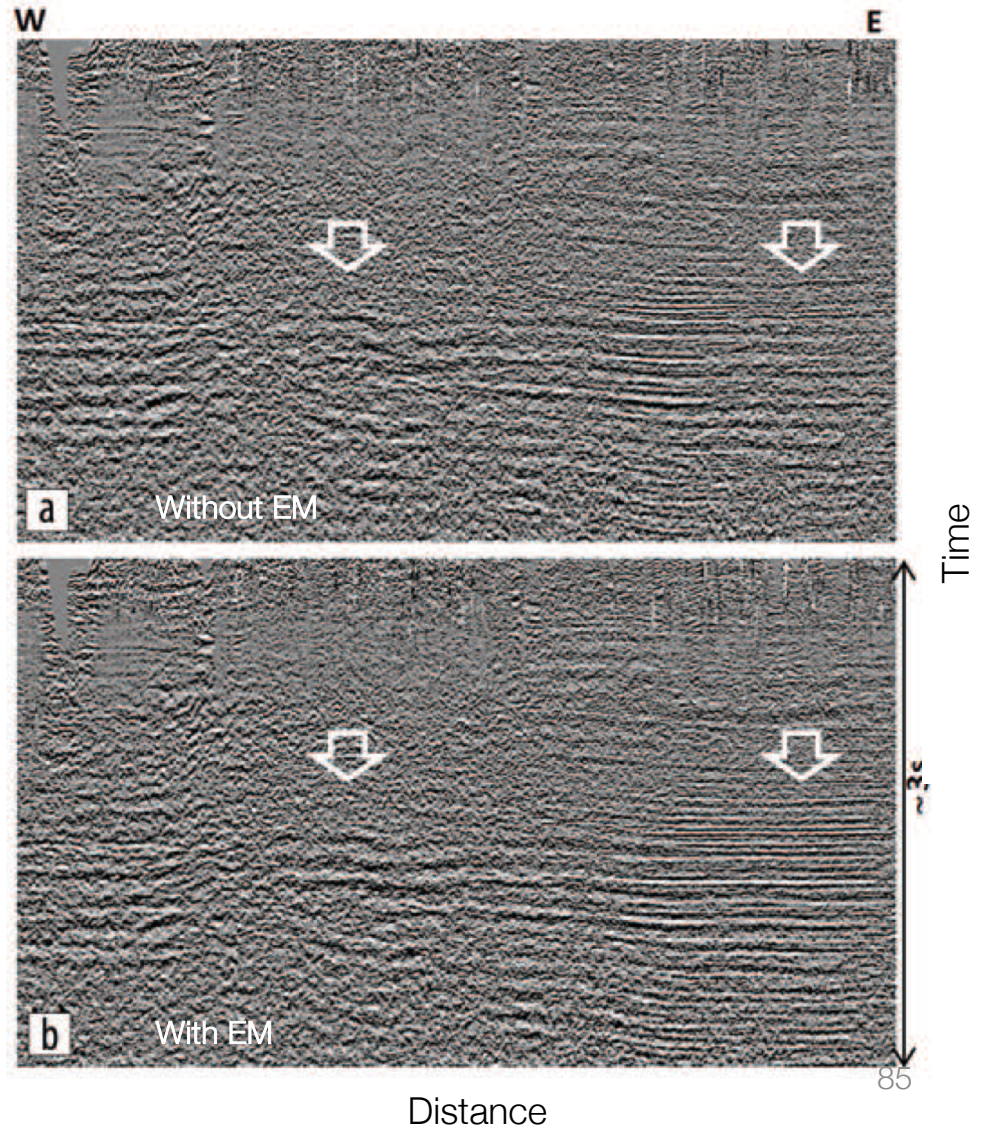


Static correction

Estimated statics on plan map



Static corrected sections

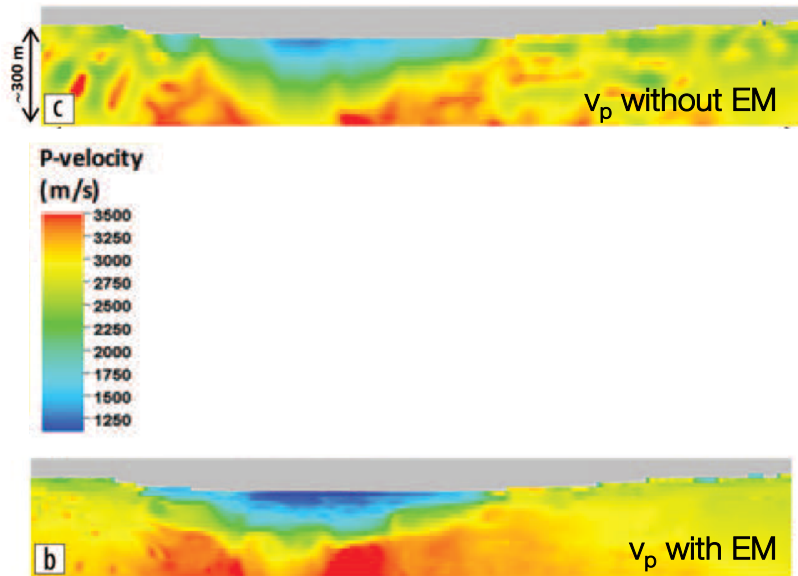


Distance

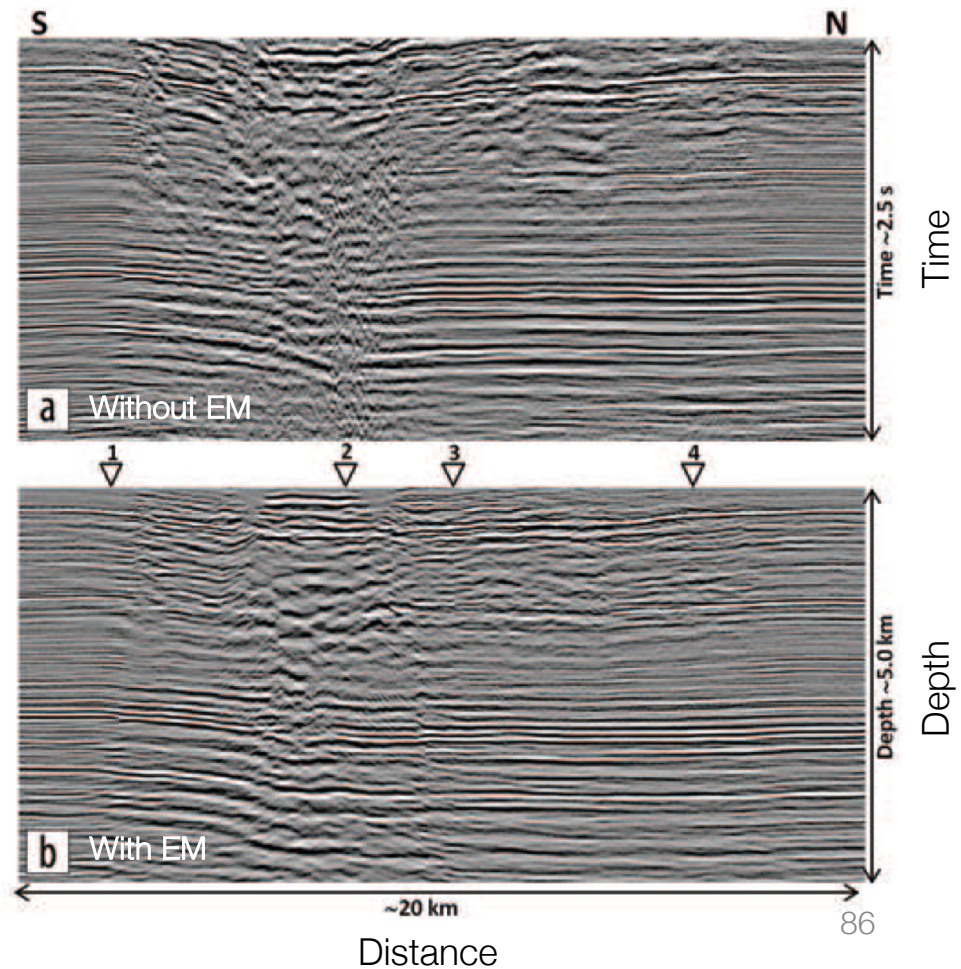
Pre-stack depth migration

- Impact of the improved v_p model to a pre-stack depth migration:

v_p cross sections at A-A'

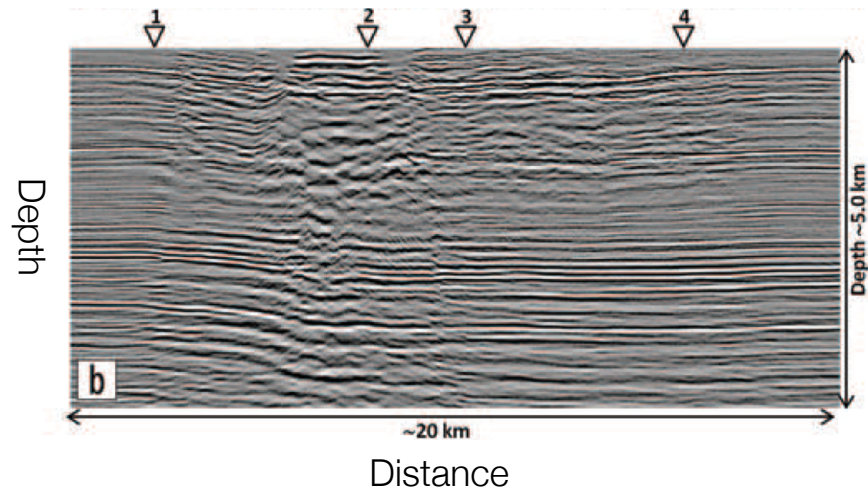


Cross sections at A-A'

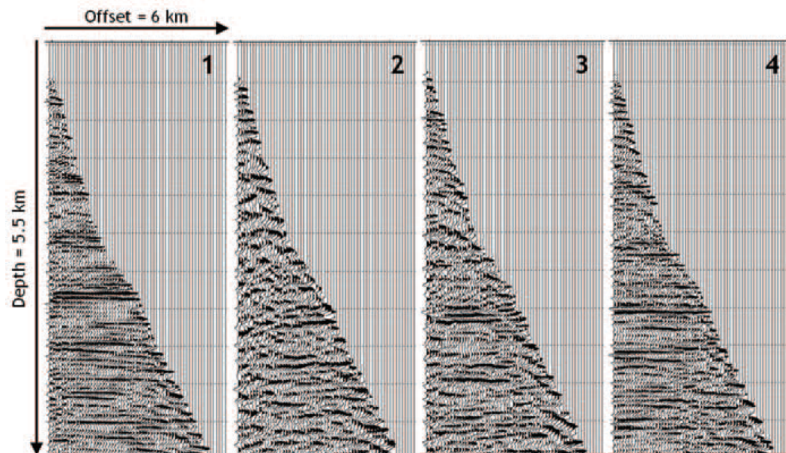


Interpretation and Synthesis

Depth section at A-A'

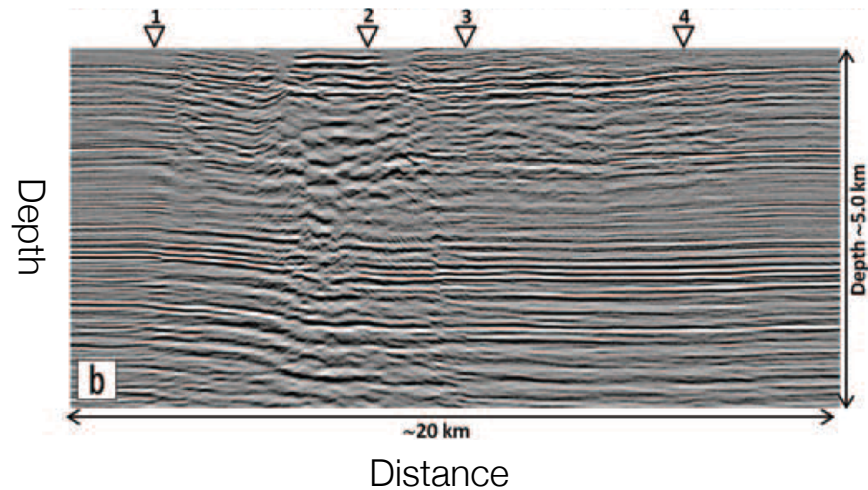


Common image gathers

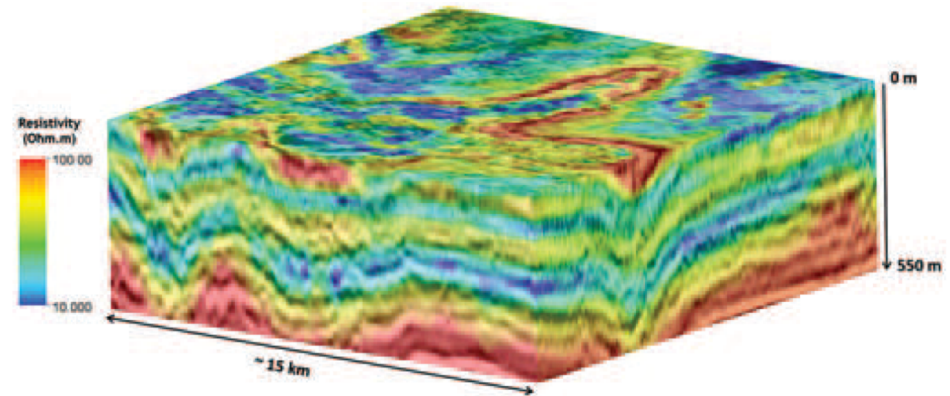


Interpretation and Synthesis

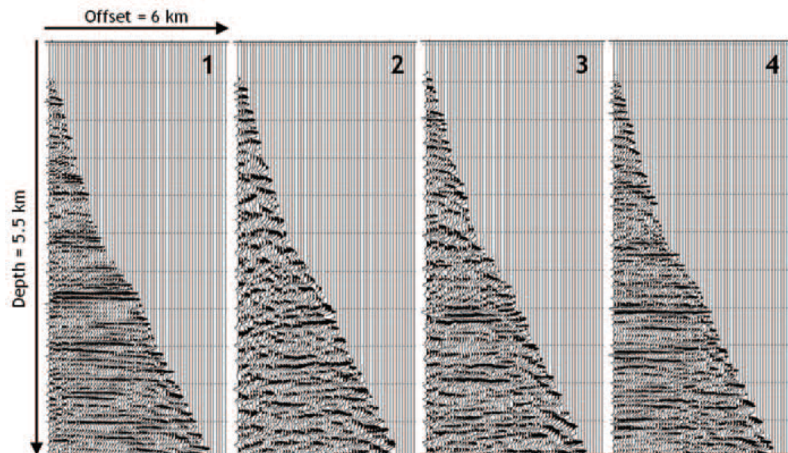
Depth section at A-A'



3D prestack depth migration co-rendered with EM

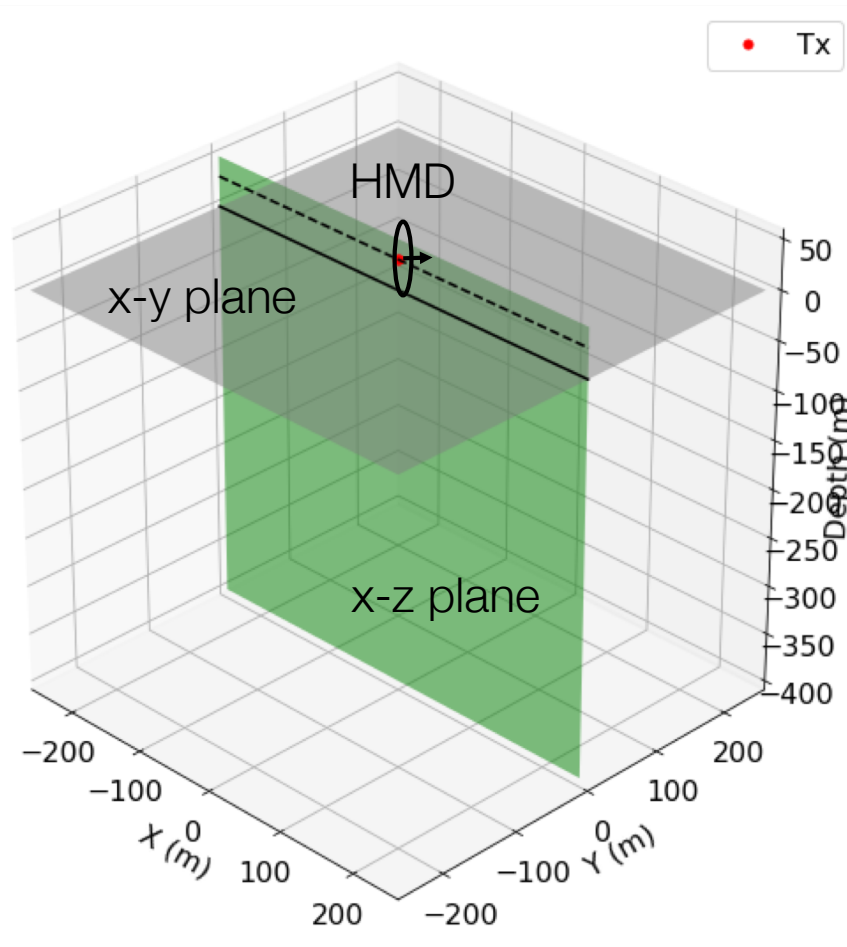


Common image gathers



- High resolution near surface conductivity from EM improves velocity model
- Helps seismic imaging:
 - Static correction
 - Pre-stack depth migration

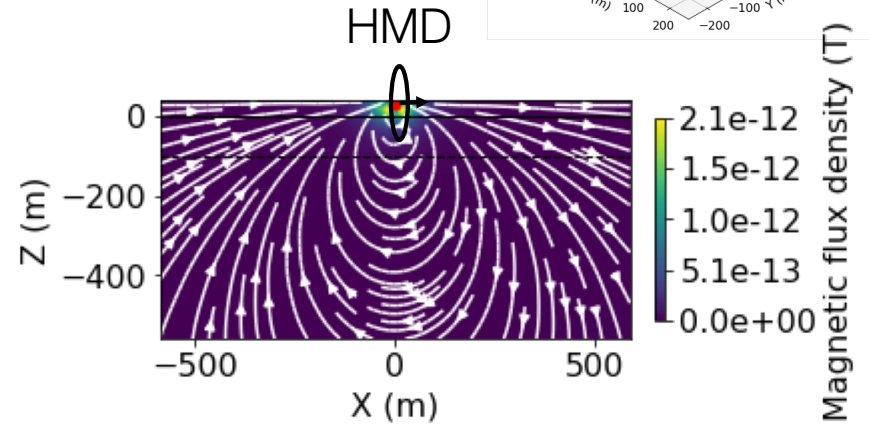
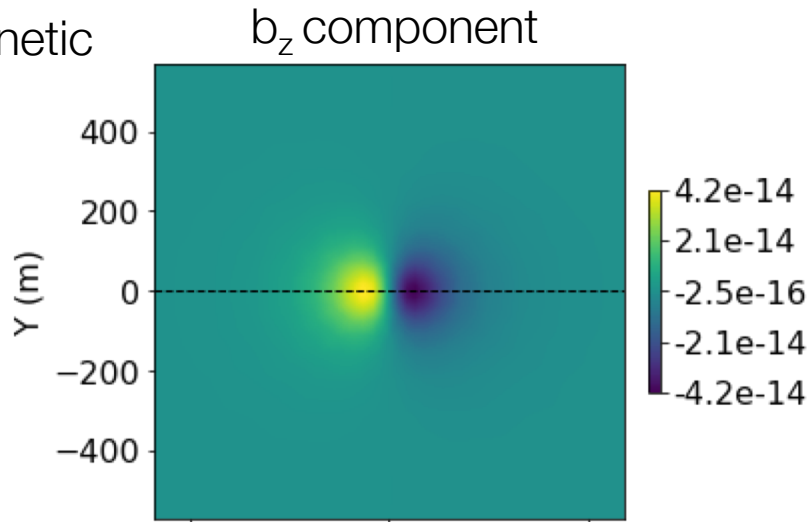
Horizontal Magnetic Dipole (HMD)



- Same physical principles as VMD, but different source geometry
- Focus on magnetic field and currents
- Different coupling for conductive targets

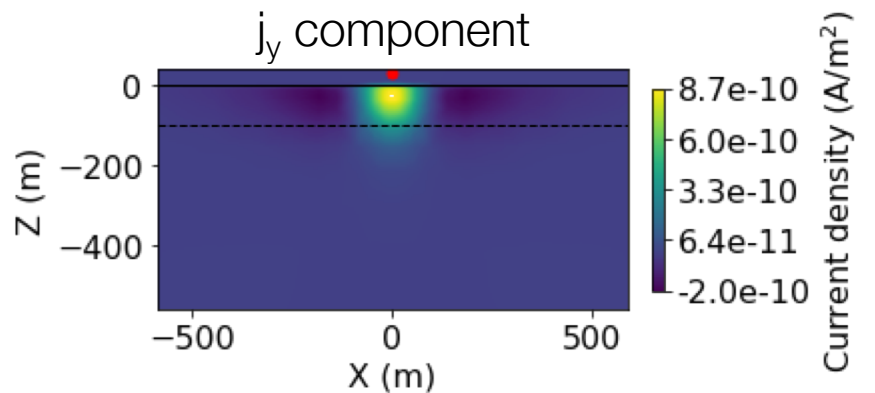
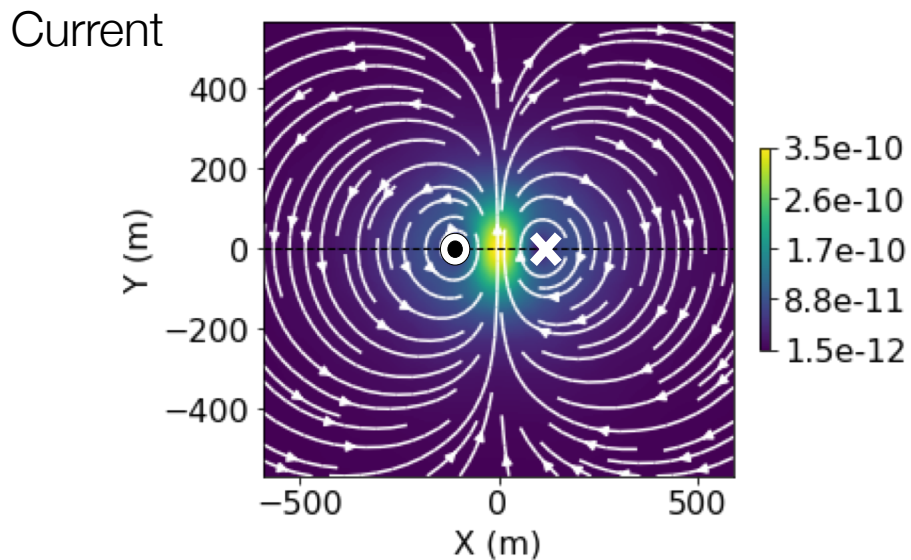
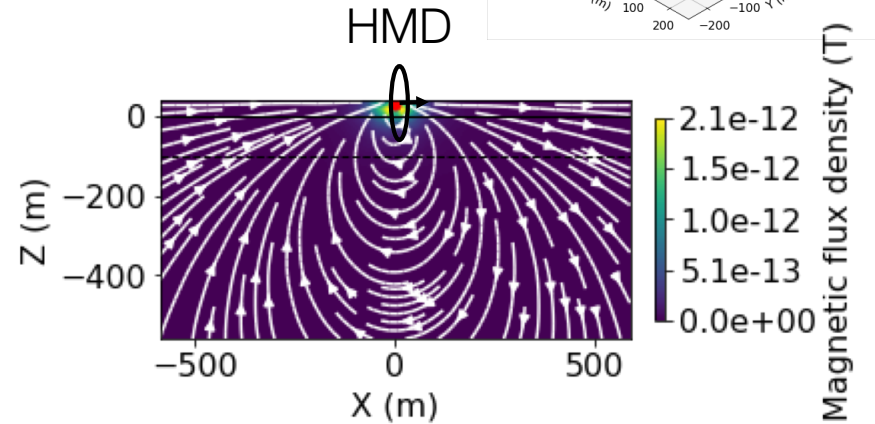
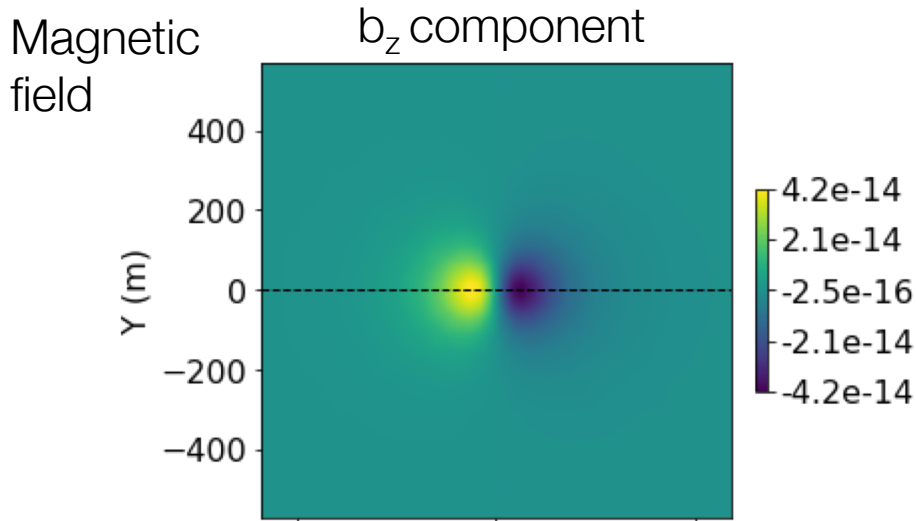
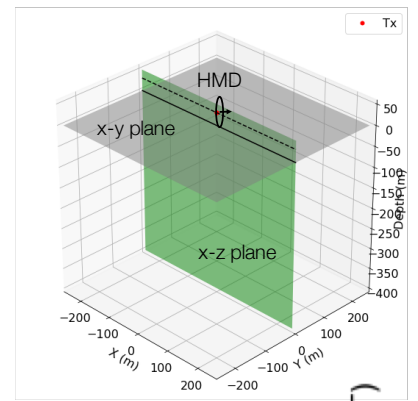
Magnetic field

Magnetic field



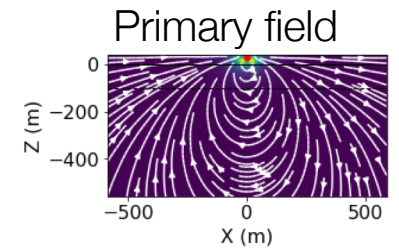
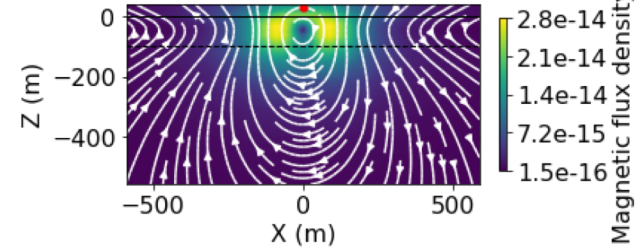
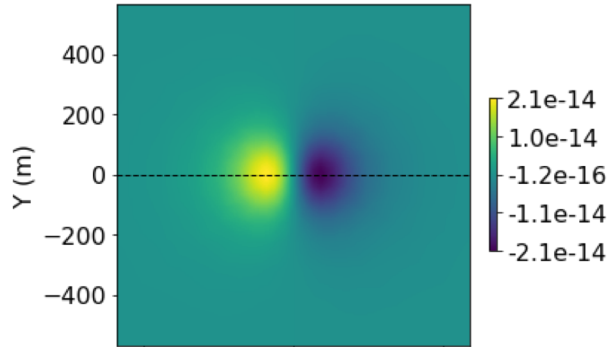
What currents can generate these magnetic fields?

Magnetic field and Current

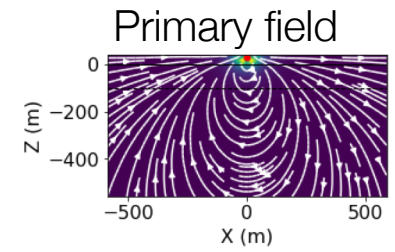


Magnetic field in time

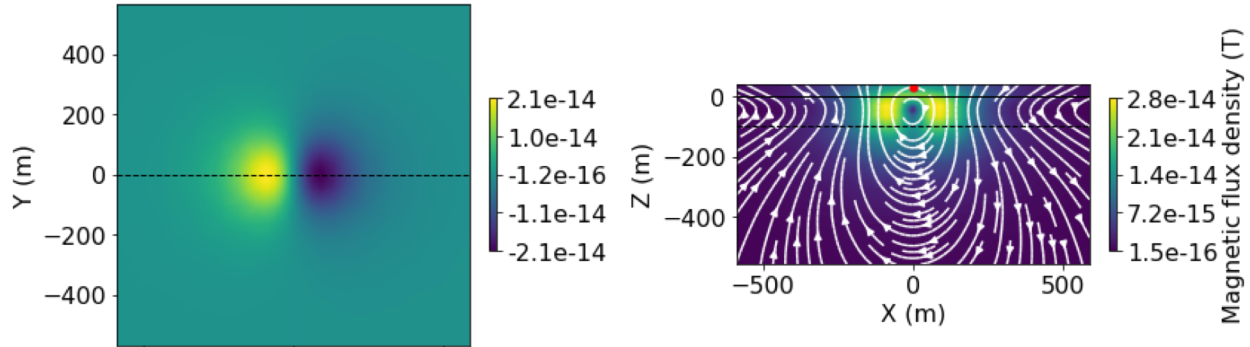
$t=0.03\text{ms}$



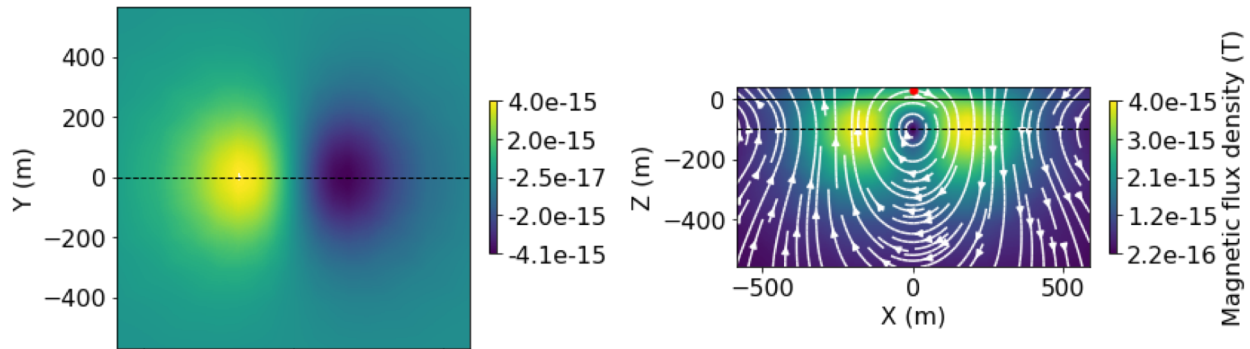
Magnetic field in time



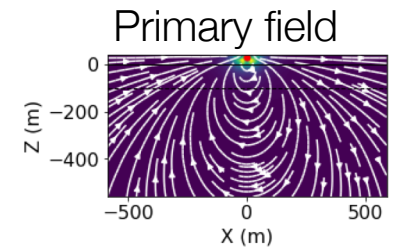
t=0.03ms



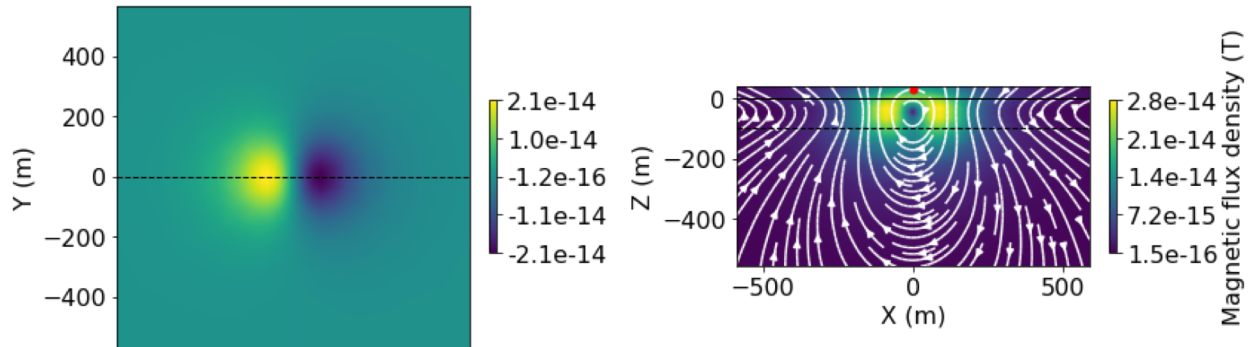
t=0.13ms



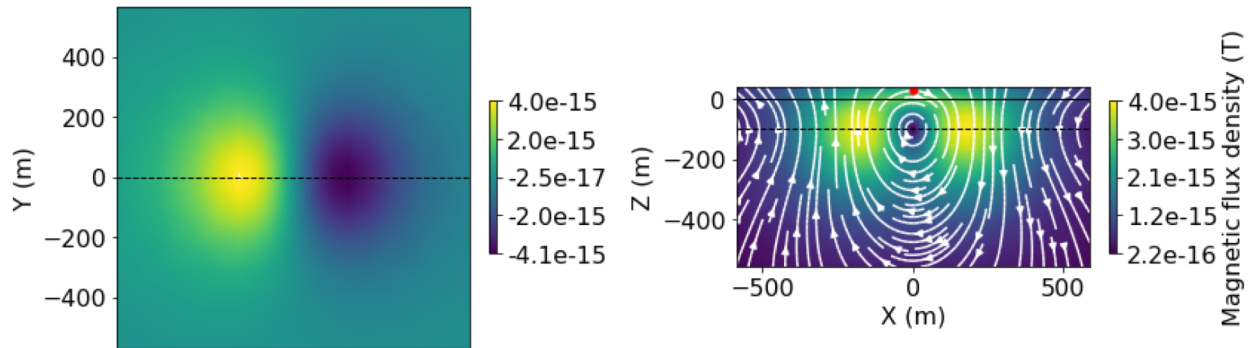
Magnetic field in time



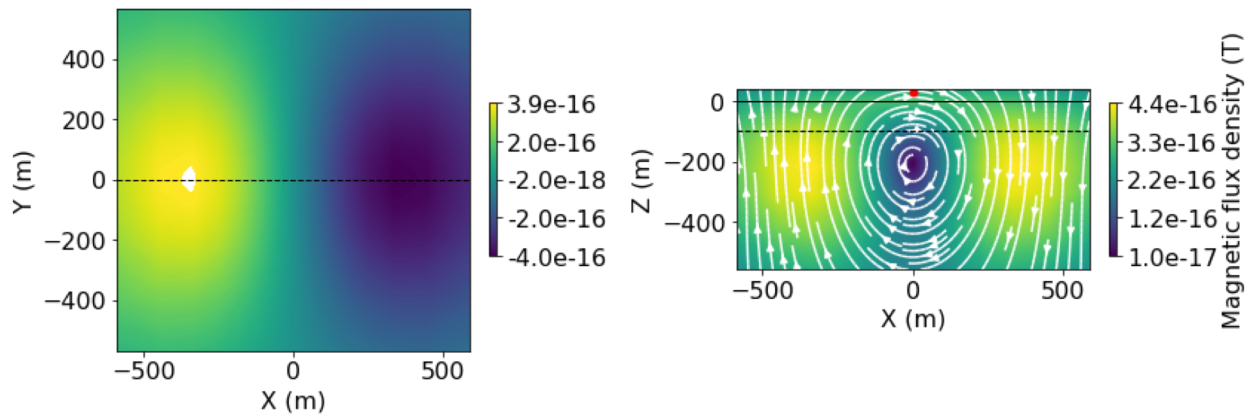
t=0.03ms



t=0.13ms

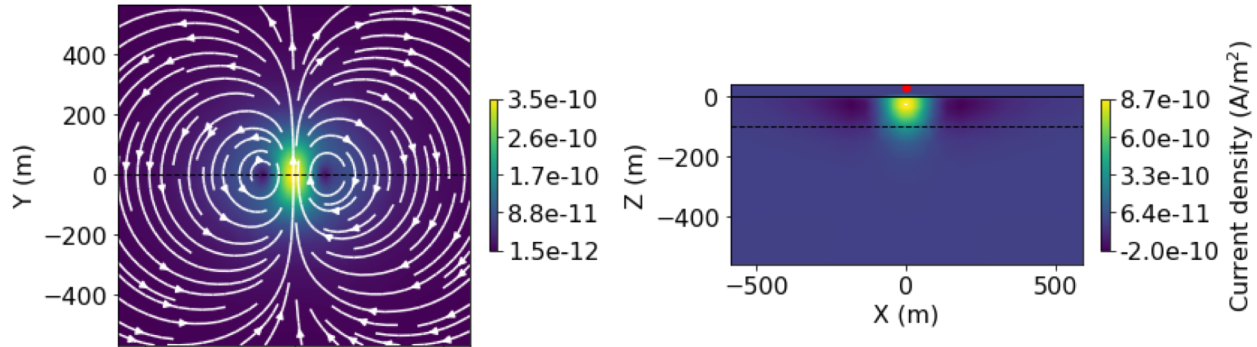


t=0.63ms

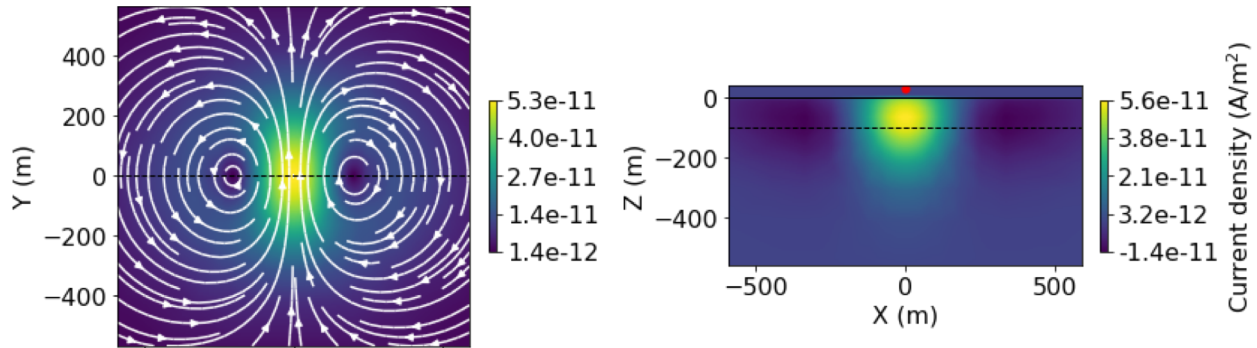


Current in time

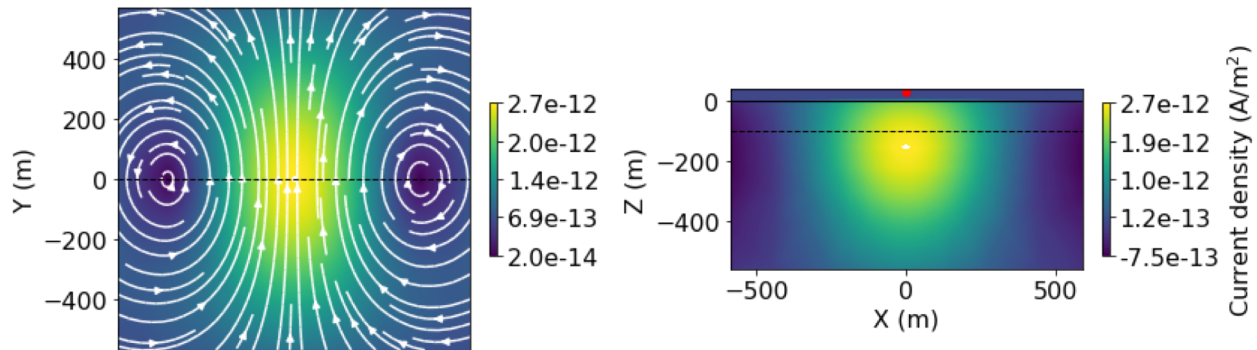
t=0.03ms



t=0.13ms

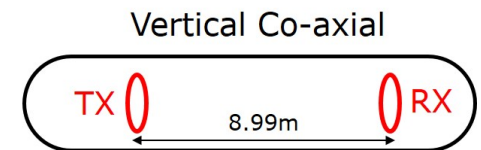
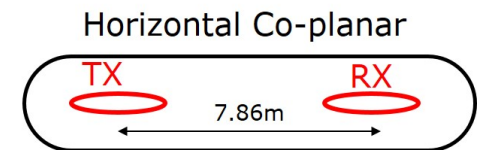
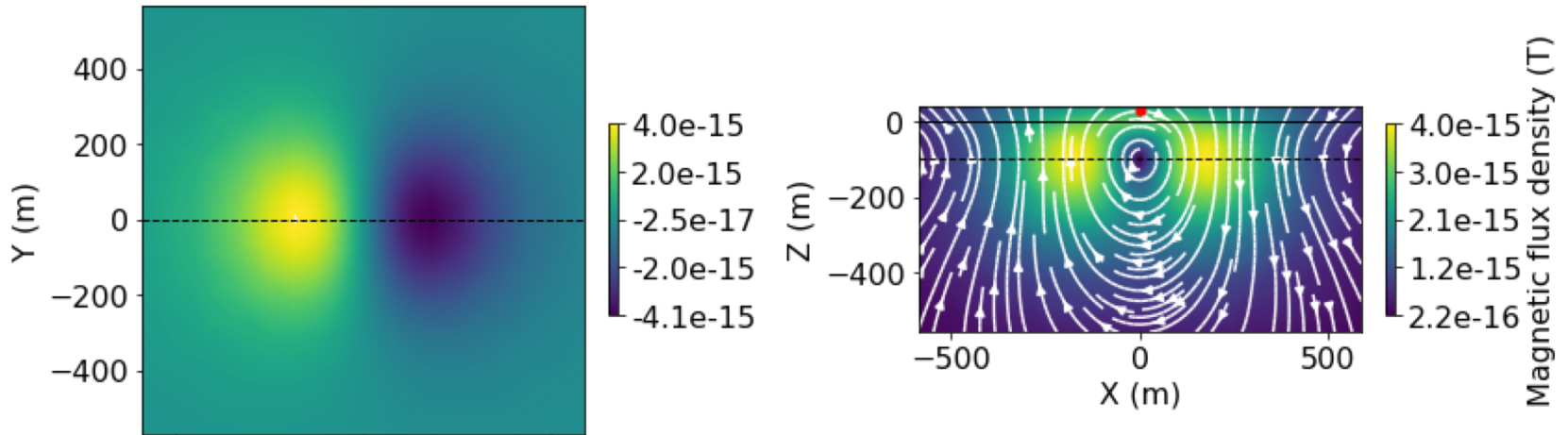


t=0.63ms



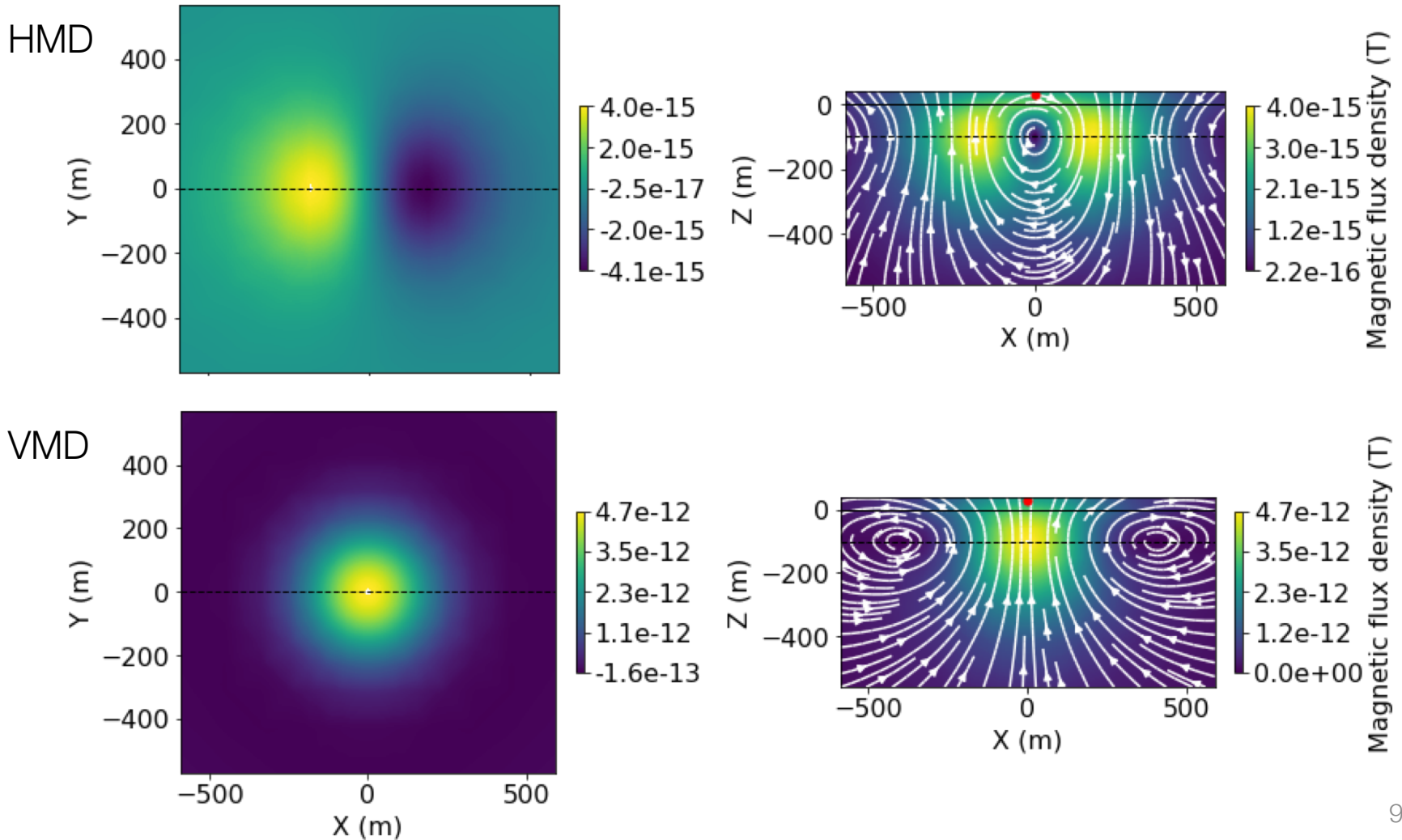
Data: where and what to measure?

Magnetic field at 0.63 ms



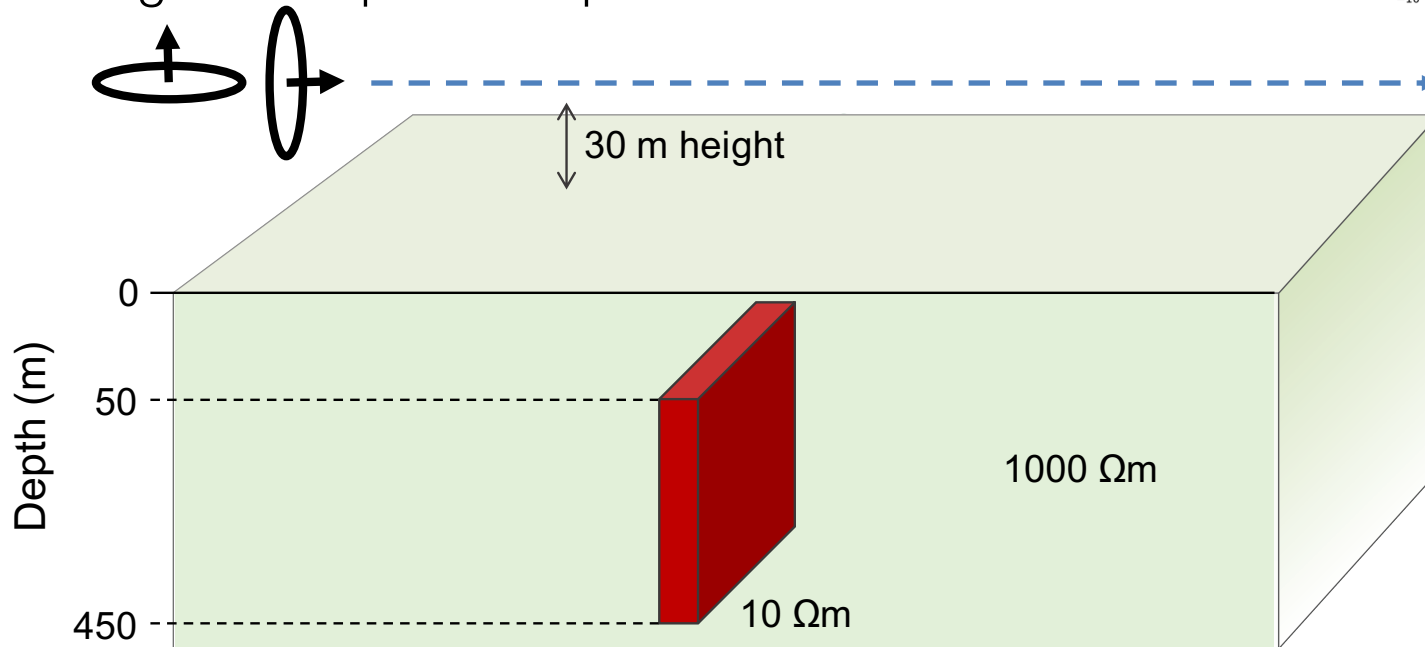
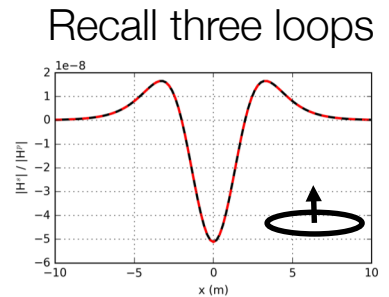
How different from VMD?

Magnetic field at 0.63 ms

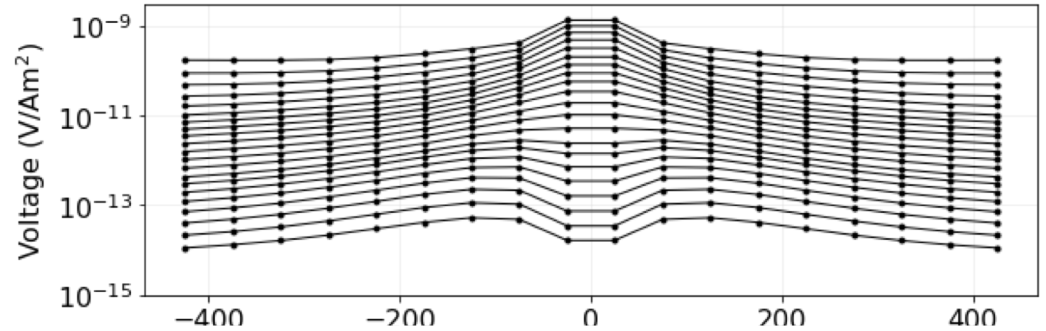
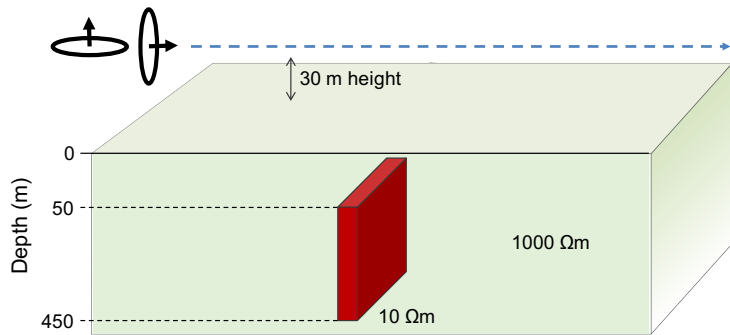


Coupling to a vertical plate

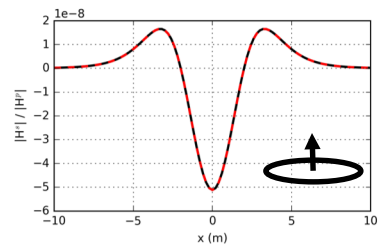
- Assume coincident loop case (Tx and Rx are coincident)
 - Both VMD and HMD
- Consider a profile line data (multiple time channels)
- Imagine how profile response will look like?



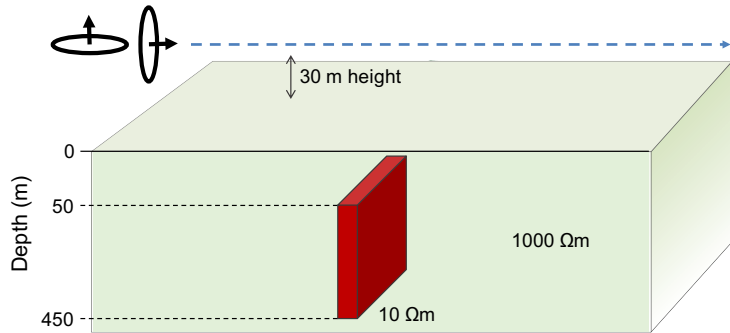
VMD vs. HMD: profile line data



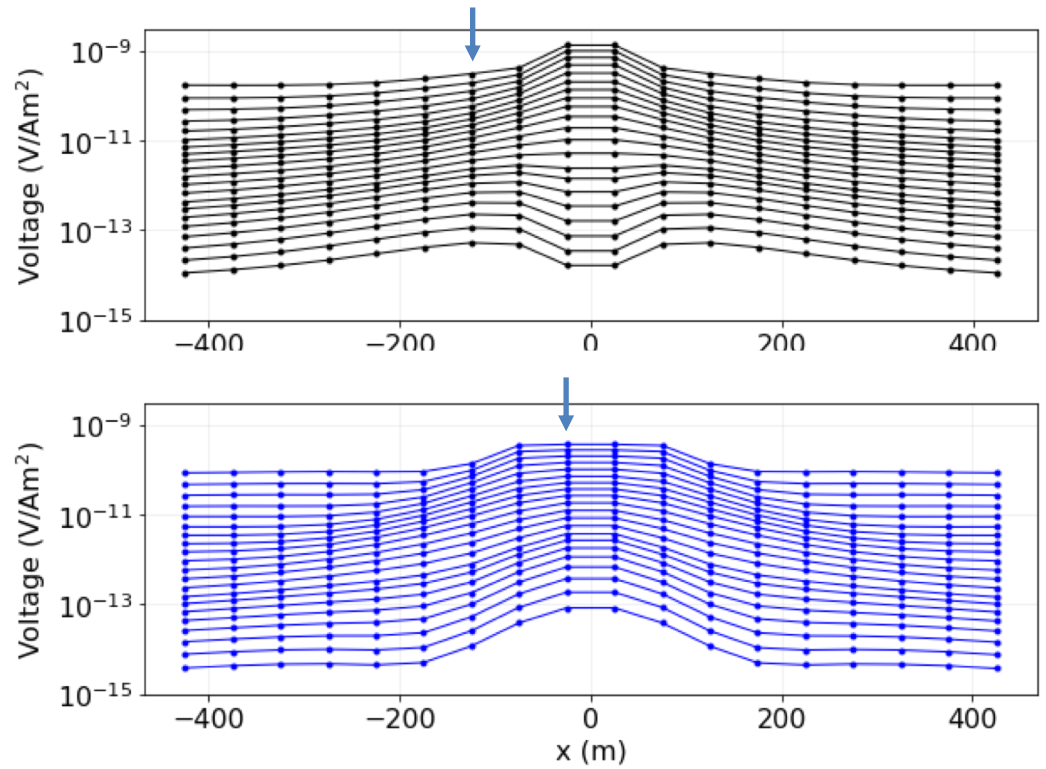
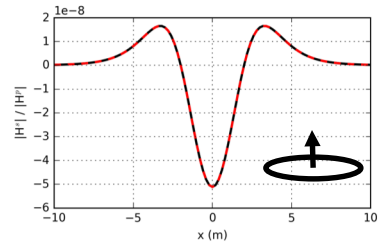
Recall three loops



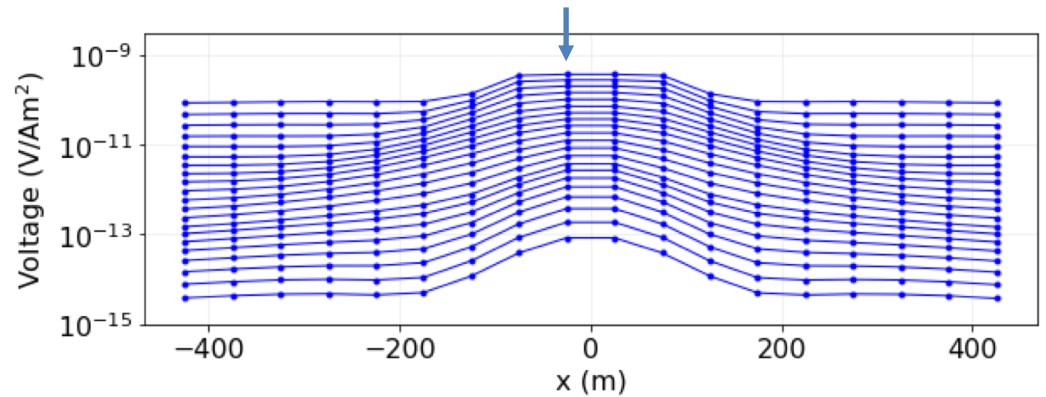
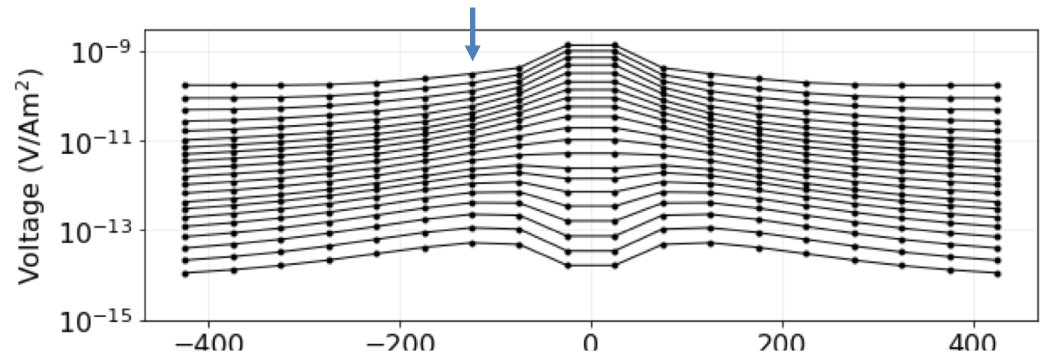
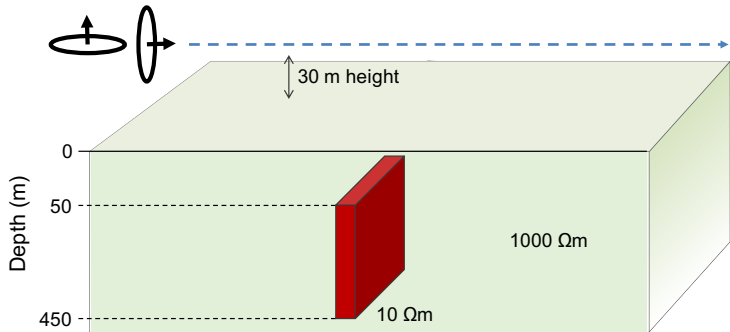
VMD vs. HMD: profile line data



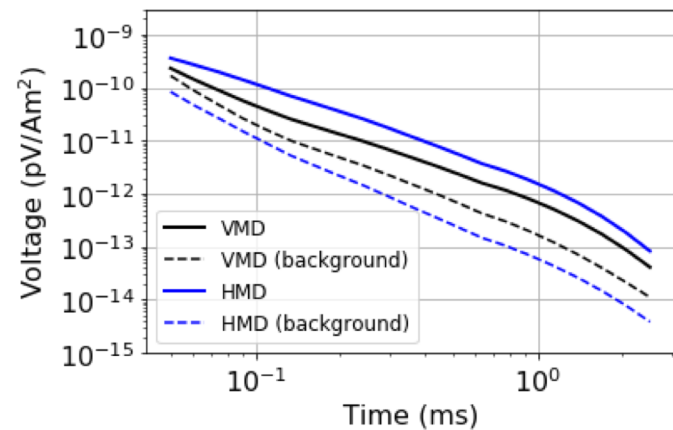
Recall three loops



VMD vs. HMD: time decays



 HMD is better coupled to the vertical conductor



End of inductive source (TDEM)

Setup

Time Domain EM

- Vertical Magnetic Dipole
- Propagation with Time
- Effects of Background Conductivity
- Transmitters and receivers
- Decay Curves
- Case History: Groundwater, Minerals, Hydrocarbons
- Horizontal magnetic Dipole

Next up: Frequency domain EM

