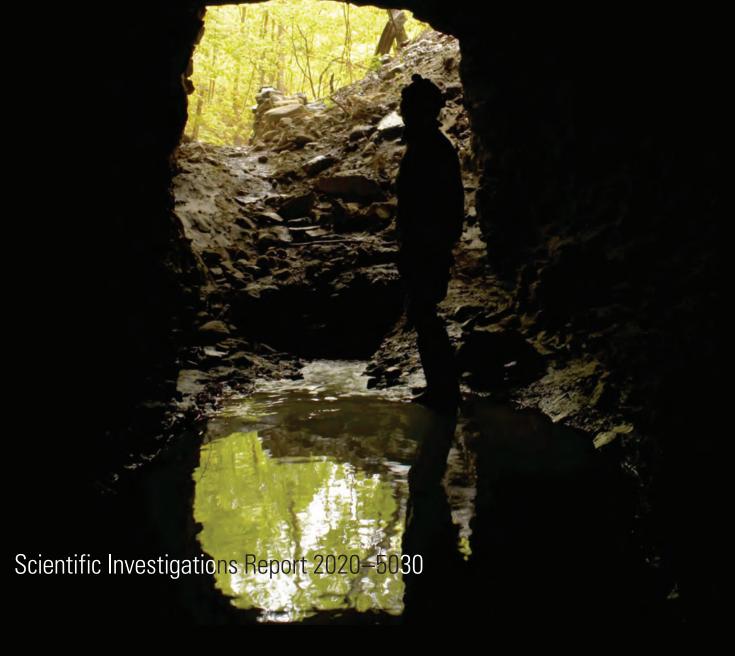


Prepared in cooperation with the New York State Department of Environmental Conservation

Statewide Assessment of Karst Aquifers in New York With an Inventory of Closed-Depression and Focused-Recharge Features



Cover. Front. Entrance to Schoharie Caverns in Schoharie County, New York. Entrance is in the Manlius Limestone near the top of a steep slope. The spring at the cave opening is perched high above nearby creek.
Back. Photograph of a cave passage from McFail's Cave. Photograph by Art Palmer, professor emeritus, State University of New York at Oneonta, New York.

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By William M. Kappel, James E. Reddy, and Jonathan C. Root
Prepared in cooperation with the New York State Department of Environmental Conservation
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Conversion Factors

U.S. customary units to International System of Units

Multiply	Ву	To obtain
inch (in.)	2.54	centimeter (cm)
foot (ft)	0.3048	meter (m)
mile (mi)	1.609	kilometer (km)
yard (yd)	0.9144	meter (m)
acre	0.4047	hectare (ha)
square mile (mi ²)	2.590	square kilometer (km²)
gallon (gal)	3.785	liter (L)

International System of Units to U.S. customary units

Multiply	Ву	To obtain		
centimeter (cm)	0.3937	inch (in.)		
meter (m)	3.281	foot (ft)		

Temperature in degrees Celsius (°C) may be converted to degrees Fahrenheit (°F) as follows: $^{\circ}F = (1.8 \times ^{\circ}C) + 32.$

Datum

Vertical coordinate information is referenced to the North American Vertical Datum of 1988 (NAVD 88).

Horizontal coordinate information is referenced to the North American Datum of 1983 (NAD 83).

Elevation, as used in this report, refers to distance above the vertical datum.

Abbreviations

DEM digital elevation model

DRG digital raster graphic

E. coli Escherichia coli

GIS geographic information system

gSSURGO gridded Soil Survey Geographic database

lidar light detection and ranging

NHD National Hydrography Dataset
NLCD National Land Cover Database

NYSDEC New York State Department of Environmental Conservation

NYSDOH New York State Department of Health

SSURGO Soil Survey Geographic database

TCE trichloroethene

USGS U.S. Geological Survey

Statewide Assessment of Karst Aquifers in New York With an Inventory of Closed-Depression and Focused-Recharge Features

By William M. Kappel, James E. Reddy, and Jonathan C. Root

Abstract

Karst is a landscape formed from the dissolution of soluble rock or rock containing minerals that are easily dissolved from within the rock. The landscape is characterized by sinkholes, caves, losing streams, springs, and underground drainage systems, which rapidly move water through the karst. The two forms of karst in New York State include carbonate karst, which forms in carbonate rock (limestone, marble, and dolostone), and evaporite karst, which forms in rock that contains the evaporite minerals gypsum and halite.

Past and recent studies of karst across the State have shown that areas of focused recharge in karstic carbonate rock allow contaminants to enter aquifer systems with little attenuation. Focused areas of recharge need to be identified to help prevent such contamination from sources on or adjacent to the karst. The New York State Departments of Environmental Conservation and Health are collaborating with the agricultural community to make farmers and farmplanning advisors more aware of karst and how to manage daily farming activities to reduce their impact on surface water and groundwater resources, especially in karst areas. There is also a need to make regulators, planners, and the general public aware of New York's karst resources and to properly protect and manage these resources to protect the quality of groundwater and surface water that can flow into, through, and from karst bedrock.

Using publicly available geospatial data, karst bedrock and closed depressions over or near karst rock were identified across New York. Carbonate, evaporite, and marble geologic units were selected from a statewide 1:250,000-scale bedrock geology dataset. The selected geologic units were intersected with 7.5-minute quadrangle maps to define the study area.

The U.S. Geological Survey has compiled an inventory of closed depressions from statewide digital contour data, scanned 7.5-minute topographic maps known as a digital raster graphics, and light detection and ranging (lidar) digital elevation models. Analysis of the data resulted in the identification of 5,023 closed depressions statewide. The inventory was conducted to eliminate duplication of results from analysis of the three data sources. A series of overlay analyses

was conducted using the closed depressions and thematic data known to be key factors in determining the probability of a closed depression contributing to focused groundwater recharge; the thematic data include bedrock geology, soil type, soil infiltration rate, and land cover.

Though the extent of karst development is important in understanding the interaction between surface water and groundwater in karst terrains, some of the worst cases of groundwater contamination in karst can occur where only minor karst features might be present. The presence of karst—be it a short section of a solutioned fracture or an extensive cave system—requires careful consideration, forward-looking environmental planning, and consistent water-quality protection to preserve New York State's water resources.

Introduction

The New York State Department of Environmental Conservation (NYSDEC) and Department of Health (NYS-DOH) are concerned about groundwater1 contamination in the karst aquifers of New York State (fig. 1), especially relating to the unintended introduction of industrial contaminants (such as volatile organic compounds [VOCs]) or agricultural contaminants (such as liquid manure; Reddy and Kappel, 2010). The NYSDEC and NYSDOH have started collaborating with the agricultural community to make farmers and farm-planning advisors more aware of karst and how to manage daily farming activities to reduce their impact on surface water and groundwater resources, especially in karst areas. To this end, the U.S. Geological Survey (USGS) has compiled an inventory of closed depressions from statewide digital contour data, scanned 7.5-minute topographic maps known as a digital raster graphics, and light detection and ranging (lidar) digital elevation models. This report provides information on the location of karst bedrock and associated closed depressions

¹Words that appear in bold typeface are defined in the glossary of this report.

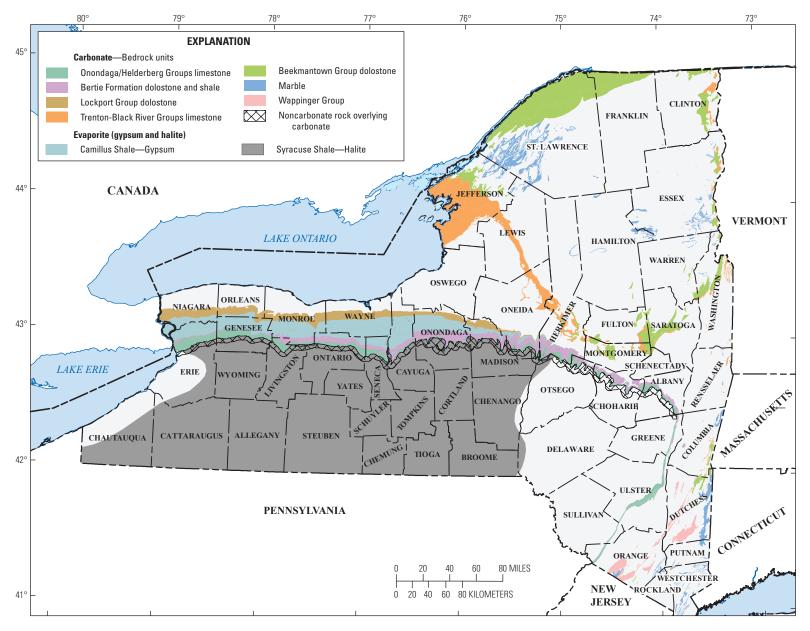


Figure 1. Major geologic units in which karst may be developed in New York State.

that focus surface water recharge to groundwater in karst areas across New York. The report also provides examples of why it is important to carefully manage and protect this resource.

Karst development (karstification) forms highly permeable bedrock aquifers that transmit and yield water from solution-enlarged fractures, bedding planes, and other openings (Olcott, 1995). Karst is a landscape formed from the dissolution of soluble rock or rocks that contain minerals that are easily dissolved from within the rock. The landscape is characterized by sinkholes, caves, losing streams, springs, and underground drainage systems, which rapidly move water through the karst. The two most common types are karst that forms in carbonate rock, such as limestone, marble, and dolostone (carbonate karst), and karst that forms in rock that contains evaporite minerals, such as gypsum or halite (evaporite karst).

The potential for focused recharge in karst (figs. 2 through 5) is greatest where the rocks crop out at or near land surface and where **closed depressions** at the land surface can focus infiltration of surface-water runoff into the aquifer. Karst aquifers are recharged by direct infiltration of precipitation that falls on the karst landscape and by infiltration of surface runoff at closed depressions, sinkholes, and **swallets**. Surface runoff may include runoff from adjacent lowlands underlain by glaciolacustrine deposits or shale-bedrock uplands that are mantled with glacial till.

Evaporite deposits of gypsum and halite associated with some shales are the most soluble of the common rock types with solubilities at standard temperature (20 degrees Celsius) and atmospheric pressure (1 bar) of 2.5 grams per liter (g/L) for gypsum and 360 g/L for halite. Karst features similar to those present in carbonate rock can develop in these evaporite deposits. The solubility of carbonate bedrock is 0.39 g/L for dolostone and 0.0062 g/L for limestone, making the rocks relatively insoluble in relation to evaporite deposits. A major difference between carbonate- and evaporite-rock karst is the time it takes for karst features to develop. Karst in carbonate rock develops over years, decades, or centuries, but karst in evaporite rock can develop over days, weeks, or years (Johnson, 2007). Gypsum deposits are more common than halite deposits and represent most evaporite-type karst in the United States (Dean and Johnson, 1989; Johnson, 2003). The solutioning of gypsum, where gypsum crops out at land surface or is within 100 feet (ft) of the land surface, can form karst features, but halite rarely appears at or near the land surface because it usually is dissolved at depths ranging from tens to hundreds of feet below the land surface (Johnson, 2007). Karst development in the carbonate rock of New York State is dependent on the rock type (limestone versus dolostone), rock structure (bedding dip and fracturing), and potential for focused recharge. Karst can develop in carbonate rock at any depth where favorable conditions are present and as long as there is a continuity of flow from where water enters and eventually exits the soluble rock.



В



Figure 2. A small sinkhole in Genesee County, New York, *A*, with a small stream disappearing into it (a swallet) and *B*, in early spring. As snowmelt runoff exceeded the swallet capacity, the area flooded, and an ice sheet formed. When runoff ceased and all water drained into the swallet, the unsupported ice cover collapsed into it. Photographs courtesy of Paul Richards, State University of New York at Brockport.

4 New York Karst Aquifers With an Inventory of Closed-Depression and Focused-Recharge Features

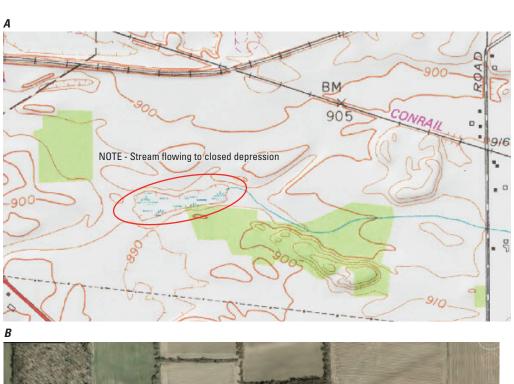
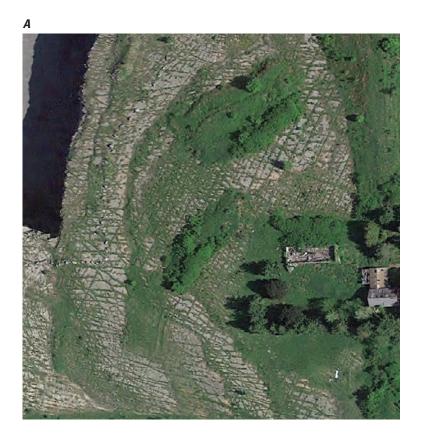






Figure 3. A closed depression near Batavia, Genesee County, New York, from A, a part of a 1:24,000-scale topographic map and B, a Google Earth image showing a closed depression (red ellipse) and C, the interior of the closed depression where surface water infiltrates into the Onondaga Limestone. The high-water line is shown to the left of the wooded area and standing water adjacent to the woods. Tree 'scars' (C) appear where ice formed when the depression was filled with water and the ice damaged the bark layer of several trees.



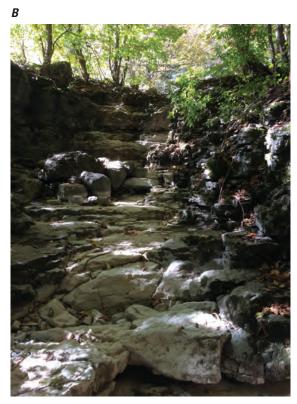




Figure 4. Karst bedrock features in Madison County, New York. *A,* Google Earth image of the fractured bedrock surface at a limestone quarry north of Perryville, New York; quarry floor is in the shadow left of the fractured bedrock surface. *B,* Dry streambed in carbonate bedrock; the streamflow disappears into the bedrock upstream from this location. *C,* Karren pavement with trees growing between fractured and solutioned carbonate bedrock.



Figure 5. Swallet entrance to McFails Cave in Schoharie County, New York, during *A*, the dry season and *B*, a 50-year flood discharge in January 1996; photographs courtesy of Art Palmer, professor emeritus, State University of New York at Oneonta.

Contamination in Karst

The karst landscape creates a situation where surface water and any contamination source can rapidly move from the land surface to an underlying aquifer. These contaminants, after entering the aquifer, can move rapidly with groundwater from areas of focused recharge to points of discharge and to any other intermediary location where groundwater is withdrawn, such as a water-supply well.

Karst aquifers are susceptible to contamination beneath urban areas from varied contaminant sources, including septic-tank effluent, runoff containing metals, oil and grease, solid and liquid wastes, and accidental or intentional dumping of chemical wastes by industrial facilities and homeowners (Veni and others, 2001). Karst aquifers in New York have been contaminated by unintended or improper chemical disposal or spills of contaminants such as petroleum products

or trichloroethene (TCE) and its breakdown residuals. These contaminants, dependent on their nature, can readily move into karst aquifers to cause short- or long-term water-quality issues in regional aquifer systems (fig. 6). The age and degree of dissolution in karst bedrock can also lead to land-surface **subsidence** and infrastructure issues (leaking pipelines and storage tanks or damaged septic systems resulting from subsidence) in rural and urban settings which, in turn, might also affect local and regional water quality.

In rural and agricultural areas, karst aquifers are also subject to environmental degradation from a variety of sources, including natural and chemical fertilizers, pesticides, and herbicides, along with their breakdown products (Veni and others, 2001). In New York, groundwater contamination from liquid manure applications and resulting water-quality issues, such as fecal coliform bacteria (*Escherichia coli* [*E. coli*]) and nitrate, were increasing in the early 2000s, and contamination incidents were reported to the NYSDEC (Reddy and

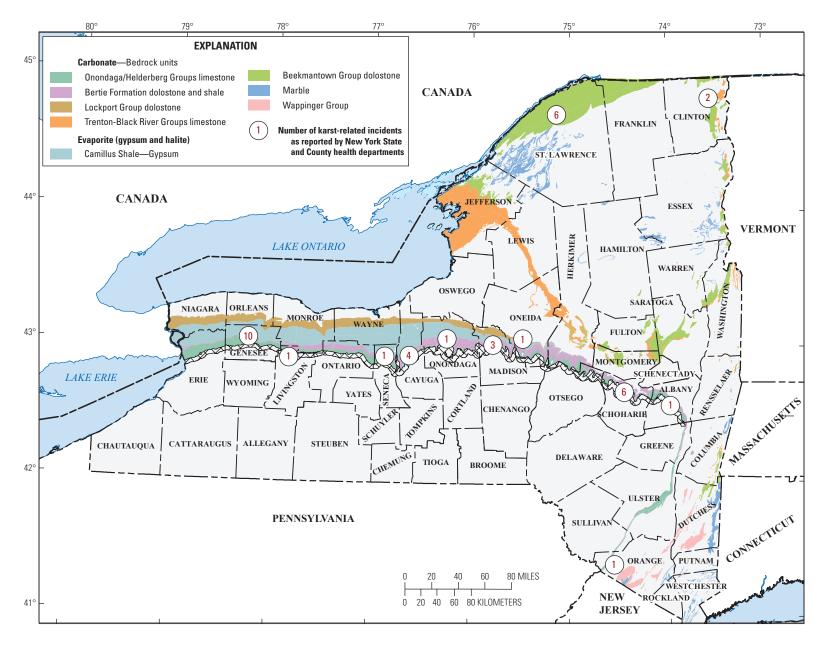


Figure 6. Contamination issues in karst areas in New York, by county, as reported by the New York State Department of Health and county health departments between 2000 and 2016.

Kappel, 2010). Nitrate from fertilizer is highly soluble in water and thus can move readily through the soil and enter the groundwater system. E. coli bacteria are not water soluble, but because of their small size, they can move through larger soil pores and into bedrock fractures, especially solutionwidened fractures in the karstic rock. The high water content and organic nature of liquid manure may enhance movement of some contaminants under certain meteorological and soil conditions. As these rural contamination issues were being reported, the use of bacterial source tracking—using DNA "fingerprints" to determine the source(s) of E. coli pollution identified or ruled out certain bacterial sources, allowing watershed management decisions that are better informed.

New York Bedrock as Affected by Karst and Glacial Processes in New York State

Karstification in New York State is highly variable because of the type of bedrock present; how the bedrock has been affected by tectonic, glacial, and other geological forces; the amount of water that moves through the karst aquifer (focused recharge); the topographic position of the rock; and how long the dissolution process has been active. Most of the bedrock in New York has little in the way of primary permeability and **porosity**—that is, water cannot easily move through or be stored in the bedrock matrix. Secondary permeability in the form of bedrock joints and other fractures was created by continental tectonic forces and stress relief including the loading and unloading of glacial ice. Weathering and dissolution of bedrock fractures in carbonate and evaporite bedrock increase the amount of water moving through the fracture system, which further enhances water movement, widens and lengthens openings, and increases permeability (sometimes termed tertiary permeability) and storage, (Ford and Williams, 2007; Kresic, 2007). In the short term, these processes can create simple, localized recharge to and discharge from a fracture or set of connected fractures over several hundreds or thousands of feet. In the long-term, these processes can create complex regional conduit and cave systems that can extend in three dimensions throughout a carbonate bedrock sequence over many square miles. How these systems develop can mean the difference between diffuse flow systems that may be seasonally unaffected by individual precipitation events and systems that may, under focused-recharge conditions, rapidly transmit surface water to groundwater and then rapidly allow groundwater to reappear as surface water at a downgradient discharge location.

Although the extent of karst development is important for understanding the surface water-groundwater interaction in a karst terrain, "the number of karst features has little to do with the problems of groundwater flow and contamination in

karst. In fact, sparse karst can give a false security. Some of the worst cases of groundwater contamination are in carbonates with only minor karst features" (Art Palmer, professor emeritus, State University of New York at Oneonta, written commun., 2016).

Karst Hydrology: New York Examples

Rapid changes in groundwater levels in karst aquifers highlight the hydraulic effect of focused recharge. For example, in Livingston County in western New York, there are two groundwater wells (LV county designation) about 3 miles (mi) distant from each other, and water-level changes are monitored as part of the statewide groundwater level cooperative monitoring program between the NYSDEC and the USGS. These two wells (U.S. Geological Survey, 2015) are near Caledonia (fig. 7); one well monitors the groundwater level in a surficial sand and gravel aquifer (well LV 330; USGS site number 425833077503901) and the other monitors the groundwater level in the underlying Onondaga Limestone, a karstic bedrock aquifer (well LV 333; USGS site number 425844077532901).

The water levels in the two wells both react to individual precipitation events, but the magnitude of reaction is quite different. The water level in well LV 330 changed only about 0.3 ft between summer and fall 2016, following a summer-long drought and subsequent average amounts of autumn precipitation and recharge. On the other hand, the water-level change in well LV 333 was about 12 ft during the same period (fig. 7). Following one precipitation event in which about 2 inches (in.) of precipitation fell on October 21, 2016, the water level in well LV 330 rose by about 0.20 ft during the following few days, whereas the water level in well LV 333 rose by more than 11 ft and then declined rapidly by more than 3 ft the following day. Knowing how and where water moves into and out of a karst aquifer system is important, as was demonstrated in this example case: the Onondaga Limestone, like many karst aquifers, has limited storage capacity but can rapidly transmit water through bedrock solution-enhanced fractures.

In this same area, the effects of water moving through karst can be seen by comparing two aerial images of a closed depression at a nearby golf course just east of LeRoy and west of Caledonia (fig. 8). In figure 8A, there is a slight trace of a dry streambed and several small sinkholes in the upper photograph, but in figure 8B, the same closed depression is filled with water with a high-water outlet at the upper (northern) end of the pond. Water filled this depression from both surface runoff from the upgradient shale bedrock watershed (which is termed allogenic recharge) as well as groundwater flowing in bedrock fractures that filled then overflowed the fractures to create the standing water body. Reportedly from groundskeepers at the golf course, in most cases, this closed depression fills and drains within a few days.

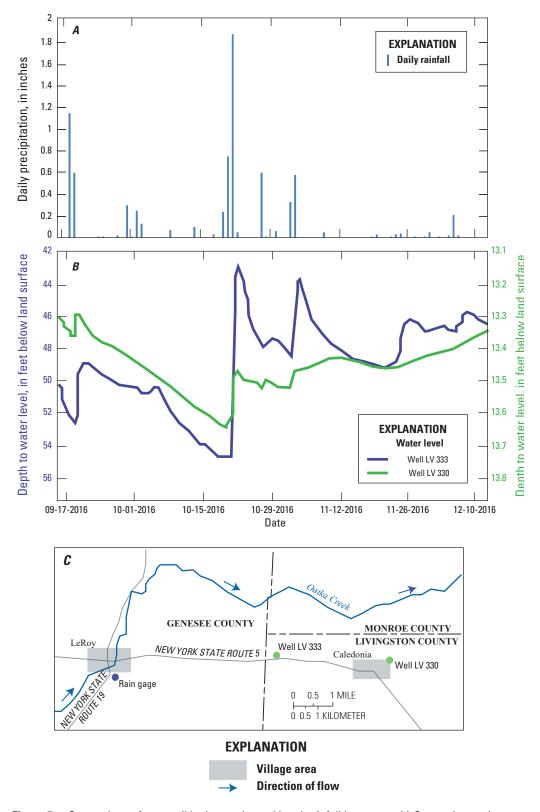


Figure 7. Comparison of two well hydrographs and local rainfall between mid-September and mid-December 2016 in western New York. *A*, Precipitation in LeRoy; *B*, water-level fluctuations at well LV 330 in an unconsolidated sand and gravel aquifer (green line) and well LV 333 in the Onondaga Limestone bedrock aquifer (blue line); *C*, location map for wells LV 330 and LV 333 and LeRoy. Note difference in water-level elevation scales for the two wells.

A similar karst situation occurs further east, near Marcellus in Onondaga County, at the locally so-called "Disappearing Lake." The closed depression is dry in September (fig. 9A) but filled with water in June (fig. 9B) with allogenic recharge from the surrounding steep shale hillsides and the interaction with the underlying karst aquifer (Proett, 1978; van der Hoorne, 2010). There are several small exposures of bedrock on the floor of the depression that can only receive or discharge a certain amount of water to or from the underlying carbonate aquifer. The lake feature may take months to drain but, dependent on local precipitation patterns and water table elevations, this flow can rapidly exceed the ability of the bedrock fractures to accept surface water to the groundwater system or discharge groundwater to the land surface

Near the City of Watertown in Jefferson County in northern New York, a holding pond for surface-water runoff was built over the Black River Limestone (fig. 10). The construction drawings clearly show the pond being constructed between several closed depressions, avoiding what could have been subsidence caused by minor sinkhole development or soil being removed by the fluctuating groundwater levels in the limestone bedrock below the clay. The pond functioned as designed for a number of years but then catastrophically failed, draining part of the volume of the pond. A dike was built to isolate the leak and, upon clearing the clay soil from the floor of the pond, a fracture in the limestone bedrock (fig. 10A) and a small solution-enlarged hole along the bedrock fracture were present where the water had drained into the limestone bedrock (fig. 10B). The hole was sealed, the clay replaced and consolidated, and the pond now functions as designed. Reviewing the location of the closed depressions, the orientation of bedrock fractures in the nearby channel of the Black River (fig. 10C), and the joint in the partly drained pond reveal that the fracture orientation extends through two of the closed depressions on either side of the holding pond.

Although the exact cause of the leak is still undetermined, the continuous saturation of the clay over the bedrock feature may have weakened the clay, while the fluctuating water table may have removed the clay from below, making the surface-water-to-groundwater connection. Additionally, several months before the failure, an earthquake occurred to the east in the Adirondack Mountain region. This earthquake may have affected the bedrock fracture and the overlying clay, causing the leak to slowly increase to the point that the clay seal failed and the pond partly drained.

A. September 2013



B. March 2010



Figure 8. A closed depression (sinkhole-swallet) in Onondaga Limestone at the LeRoy Country Club in LeRoy, Genesee County, New York, in A, March 2010 and B, September 2013. Surface water contributed from a shale bedrock watershed to the south of this view. In March 2010, the depression was filled with surface-water runoff and had a karst water table above land surface, whereas in September 2013, the karst water table was below land surface.

A. September 2009



B. June 2011



Figure 9. Disappearing Lake near Marcellus, Onondaga County, New York. *A*, Aerial image from September 2009, with no standing water in the closed depression; *B*, Aerial image from June 2011 following a wet spring season, where runoff from the surrounding shale hillsides has filled the closed depression and water is slowly draining into the underlying Onondaga Limestone.







Figure 10. Draining of a surface-water holding pond in Watertown, New York. *A,* Bedrock fracture found below approximately 6 feet of glacial clayey soil in which the holding pond was located; *B,* solutional opening along this fracture where surface water drained to groundwater; *C,* orientations of primary bedrock fractures in the Black River, southwest of the holding pond and east of the New York State Route 12 bridge in Watertown.

Karst in Bedrock of New York State

In New York State, karst is present in a number of bedrock lithologies (limestone, marble, dolostone, and shale with evaporite deposits; table 1.1) and is also dependent on the structure of and within the rock (strike and dip of the rock and the presence of joints, fractures, faults, and bedding planes within the rock mass). Karst develops in either gently dipping west-to-east oriented bands in western and central New York or in steeply dipping beds or isolated exposures in south-eastern New York. Weary (2008) characterizes New York's carbonate karst into two major groups, each with a subgroup based on glacial-cover thickness.

The first group of bedrock consists of gently folded and flat-lying carbonate rocks that have not been strongly deformed. One subclass of this group is thinly covered with glacial material and the other subclass has thick (more than 50-ft) unconsolidated glacial materials that tend to obscure karst land-surface features. In this group, where the carbonate-rocks are thick and extensive, cave systems may be long and complex. Where the carbonates are thin and interbedded with noncarbonates, caves are small and short (Weary, 2008). Palmer (1991) indicates that the geometry of caves that follow stratigraphic and bedding planes form meandering branchwork (stream-like) caves and **passages**.

The second group of karst bedrock is composed of folded and faulted carbonate rocks flanking structural deformation zones that may be intensely folded, jointed, and faulted. Caves can be small and simple or long and complex (Weary, 2008). Palmer (1991) indicates that cave patterns in this group show structural control and form networked (angular grid- or road-like) cave patterns as well as vertical connections to the land surface. Here, too, where the glacial deposits are thick (more than 50 ft), the karst features typically are not apparent at land surface.

Cooper and Mylroie (2015) described the geologic age and location of karst bedrock in New York (fig. 1; table 1) as

- Precambrian Grenville Complex marble in the Adirondack Mountain region;
- Precambrian, Cambrian, and Ordovician marble in eastern New York (Wappinger and Chazy Groups);
- Cambrian and Ordovician limestone in northern New York lowlands and the Mohawk Valley (Trenton Group, Black River Group, and Beekmantown Group);
- Silurian and Devonian limestone and dolostone in nearly flat-lying beds from western to east-central New York (Onondaga Limestone, Helderberg Group, and Lockport Group), and within gypsum- and halitebearing shale (Salina Group) in western to central New York. Other carbonate beds are present in the Silurian and Devonian sequence and can be karstic locally (such as the Manlius, Coeymans, Rondout, and Tully

- Limestones and the Akron-Bertie Dolostone), but they are not highlighted here because they are generally associated with the previously mentioned bedrock units.
- Silurian and Devonian limestone and dolostone in structurally deformed and steeply dipping beds along the western side of the Hudson River Valley in eastern New York (Onondaga Limestone and Helderberg Group).

All bedrock in New York has been affected to various degrees through geologic time by continental movement (plate tectonics) and the associated structural deformation including folding, fracturing, and faulting. In an extreme case in the lower Hudson River Valley, bedrock strata have been deformed from a horizontal to a nearly vertical orientation (fig. 11*A*). The Adirondack Mountain dome continues to rise to this day, slowly deforming the sedimentary bedrock that surrounds it.

During the past 2.6 million years, bedrock in New York State has been subjected to periods of continental glaciation (Fullerton, 1980) and intervening periods of erosion and deposition by fluvial processes. During and following each glacial period (1) the bedrock surface was reshaped, (2) surface-water and groundwater drainage patterns were modified, locally and regionally, (3) internal drainage through dissolution of carbonate rock or evaporite deposits was enhanced, and (4) internal bedrock flow and storage were increased, developing secondary and tertiary permeability and porosity. Dependent on the age and degree of karst development, surface water and groundwater interaction in karst has been affected by both glacial and interglacial dissolution processes. For example, some caves predate the most recent glacial episode (Wisconsinan) and were partially filled with glacial sediments that were subsequently re-excavated (fig. 12). Continued dissolution of karstic bedrock today further develops karst systems and highlights the need to understand and protect karst-aquifer water resources in New York State.

The carbonate and evaporite bedrock units in New York have undergone various degrees of karst development, from recent dissolution that may only extend several hundred feet within carbonate bedrock to long-term (several tens to hundreds of thousands of years) development of large karstaquifers that evolved into interconnected cave systems that extend over many square miles. In some places, the karst bedrock is overlain by unconsolidated deposits and bedrock that may be in hydraulic contact with the underlying karstic bedrock, termed subjacent karst. Adjacent and topographically higher nonkarst bedrock can also route surface water and groundwater into a nearby karst aquifer (allogenic recharge). These adjacent areas are part of the focused-recharge watershed of karst aquifers, even though they might not be types of bedrock in which karst would normally form in or flow through. Figures 8 and 9 are examples of shale uplands that drain to a karst feature.

 Table 1.
 Stratigraphic column of New York State bedrock indicating those units in which karst features might be present

[Cells shaded in blue indicate formations prone to carbonate karst development; cells shaded in green indicate formations somewhat prone to carbonate karst development; cells shaded in orange indicate formations somewhat prone to gypsum evaporite karst development; cells shaded in yellow indicate formations somewhat prone to halite or gypsum evaporite karst development; and cells not shaded indicate formations that do not have karst development]

Period	Epoch	Group or formation	West	Cantral		on of group in		St. Lawrence	Formation or member	Generalized lithology
Devonian	Upper	Conewango	X	Central	East	Southeast	Cnampiain	St. Lawrence	Cattaraugus	Shale and sandstone
Devoman	Оррег	Conneaut	X						Ellicott/Chadakoin/Whitesville	Shale
		Canadaway	X	X					Gowanda/Dunkirk/Caneadea/Wellsville	Shale and sandstone
		West Falls-Java	X	X					Hanover/Wiscoy/Angola/Nunda/Rhinestreet	
		west raiis-java	Λ	А	v				, ,	Shale; shale and sandstone
					X				Slide Mountain	Shale, sandstone, conglomera
		Sonyea	X	X					Cashaqua/Middlesex/Rock Stream	Shale; shale and sandstone
				X	X				Glen Aubrey/Lower Walton	Shale and sandstone
		Genesee	X	X					West River/Penn Yan	Shale
			X	X	X				Geneseo/Ithaca/Oneonta	Shale and sandstone
	Middle	Tully Limestone		X					Tully	Limestone
		Hamilton			X				Ashokan/Manorkill	Shale and sandstone
			X	X	X				Moscow/Portland Point	Shale; limestone
					X				Cooperstown	Shale and sandstone
			X	X	X				Centerfield/Ludlowville/Skaneateles/Panther	Shale and limestone
									Mountain	
			X	X	X				Marcellus (Union Springs-Oatka Creek)/	Shale; shale and sandstone
									Mount Marion	
		Onondaga Limestone	X	X	X	X			Onondaga	Limestone
	Lower	Tristates	X		X	X			Bois Blanc/Schoharie	Limestone; shale
	Lowei	Histates	Λ							
				**	X	X			Esopus	Shale
				X	X	X			Oriskany/Glenerie	Sandstone; limestone
		Helderberg			X	X			Alsen/Port Ewen	Limestone; shale
					X	X			Becraft	Limestone
					X	X			Kalkberg/New Scotland	Limestone; shale
				X	X				Manlius/Coeymans	Limestone
				X	X				Chrysler/Rondout	Dolostone
Silurian	Upper	Salina	X	X	X				Akron/Cobleskill	Dolostone; limestone
	**		X	X	X				Bertie	Dolostone with gypsum
			X	X					Camillus	Shale, with gypsum
			X	X	X				Syracuse/Wawarsing/Poxono Island	Shale, with halite, limestone;
			Λ	А	Λ				Syracuse/ wawarshig/1 0x0110 Island	shale
			X	X	X				Vernon/Bloomsburg	Shale, with halite and gypsum shale and sandstone
		Lockport	X	X					Guelph/Eramosa/Sconondoa	Dolostone; limestone and dolostone
					X				Ilion/Shawangunk	Shale and dolostone; conglom erate and sandstone
			X	X					Goat Island/Gasport/Penfield/Sconondoa	Dolostone; limestone and dolostone
						X			Green Pond	Conglomerate
		Clinton	X	X	X	**			DeCew/Glenmark	Dolostone; shale
		Chinon	X	X	X				Burleigh Hill/Rochester/Gates/Herkimer	Shale; dolomite and shale; sandstone
	Lower	_	X	X	X				Irondequoit/Thorold/Rockway/Dawes	Limestone and shale; sand- stone; dolostone
						X			Green Pond	Conglomerate
		Matter	v			Λ				-
		Medina	X						Grimsby/Kodak/Power Glen/Whirlpool	Sandstone; shale and siltstone sandstone
Ordovician	Upper	Queenston Shale	X	X					Queenston	Shale
	**	Lorraine		X	X				Oswego/Pulaski/Frankfort	Sandstone and shale; sand- stone and shale; shale and
	16:27	T		37	37				Hillian/Danian/C Di Fri	siltstone; shale
	Middle	Trenton		X	X				Hillier/Denley/Sugar River/Kings Falls/ Selby	Limestones
				X	X		X		Utica/Snake Hill/Stony Point	Shale; shale and siltstone
							X		Glens Falls/Orwell	Limestones
		Black River		X	X				Watertown/Lowville/Pamelia/Isle la Motte	Limestones
				A	Λ		X		Valcour/Crown Point/Day Point	Limestones
	Lower	Chazy Beekmantown-					X			
	Lower	Wappinger					Х		Tribes Hill/Copake	Dolostone, limestone, and siltstone; limestone
						X			Copake/Rochdale/Halcyon Lake (Hudson Valley area)	Limestone; limestone and dolostone; dolostone and siltstone
Cambrian	Upper-	Wappinger-				X			Briarcliff/Pine Plains (Hudson Valley area)	Dolostone; dolostone, sand-
сашопай	Middle- Lower	Stockbridge New York City				X			Inwood	stone, and shale Marble
		Group								
	Upper- Middle	Beekmantown			X				Tribes Hill/Little Falls/Hoyt/Galway	Dolostone; limestone; sand- stone and dolostone
	Middle						X		Providence Island/Fort Ann/Great Meadows/ Whitehall/Galway	
										dolostone and sandstone
		_						X	Theresa	Dolostone, sandstone, and
	Middle-									

This table can be downloaded as a full-size PDF version at https://doi.org/10.3133/sir20205030.



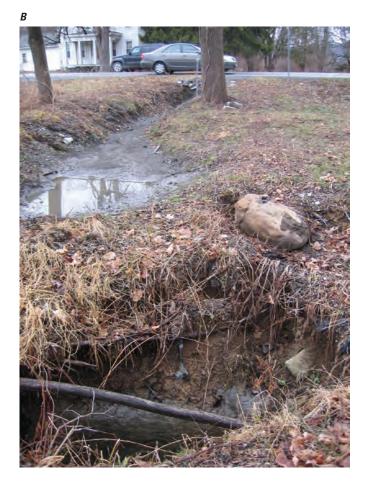


Figure 11. Karst features in the Rondout Valley in Ulster County, New York: *A*, heavily jointed and fractured Onondaga Limestone (note near-vertical bedding orientation of the rock, typical for sedimentary bedrock west of the Hudson River), and *B*, roadside ditch draining into small sinkhole-swallet feature near Stone Ridge, New York. Photographs by Paul Heisig, U.S. Geological Survey.

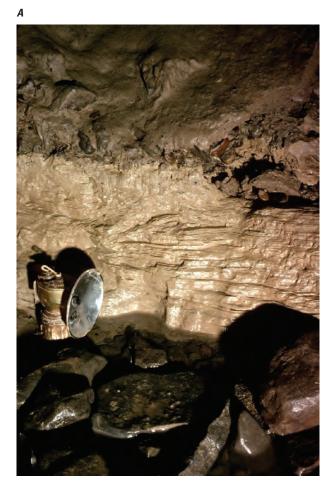




Figure 12. Glacial-related features common in Schoharie Creek Valley caves. *A*, Varved glaciolacustrine clay near the floor of the cave likely deposited during the Wisconsinan glacial period; *B*, glacial boulders trapped in a cave passage. These features indicate the cave was present before and during the Wisconsinan glacial period. Photographs courtesy of Art Palmer, professor emeritus, State University of New York at Oneonta.

Western and Central New York

Most bedrock in western to east-central New York dips to the south at approximately 30 to 50 feet per mile. Zones of allogenic recharge have been identified in central and western New York, topographically upgradient and south of specific karst units, such as the Onondaga Limestone and Lockport Group. Because of glacial erosion and the dip of the regional bedrock, the Onondaga Limestone is typically at or near land surface at its northern extent, creating a bedrock escarpment (cliff). However, south of the escarpment the Onondaga Limestone is buried by progressively thicker glacial deposits—mostly till and glaciolacustrine fine sand, silt, and clay, but locally sand and gravel (Staubitz and Miller, 1987). The Onondaga Limestone is considered a karst aquifer because of the amount of karst development that has occurred along joints and other fractures and nearly horizontal bedding planes. The Onondaga Limestone has been characterized as a principal aquifer in New York, which is defined as "an aquifer known to be highly productive or whose geology suggests abundant potential water supply, but which is not intensively used currently as a source of water supply by major municipal systems" (New York State Department of Environmental Conservation, 2020).

In a similar fashion, the dolostone of the Lockport Group gently dips to the south and has a northern escarpment in western New York, but the escarpment is buried further to the east. At the Niagara escarpment, the bedrock is at land surface but further to the south, it is buried by thick glaciolacustrine silt and clay, especially in western New York (Kappel and Jennings, 2012). Owing to fine-grained sediment cover and the fact that the dolostone is not as easily solutioned as limestone (because dolostone, which is calcium magnesium carbonate, has different mineral characteristics compared with limestone, which is calcium carbonate), the Lockport Group is less likely to develop karst features. Some karst development in the form of sinkholes and springs has been documented in the Lockport Group but typically along bedding planes near the bedrock

surface (Yager and Kappel, 1987). Regionally, the Lockport aquifer system is generally confined to these bedding-plane structures and the weathered bedrock surface.

Numerous carbonate rock quarries are present in both the Onondaga Limestone and Lockport Group dolostone in this part of the State. Products derived from these quarries vary from large to small rip-rap rock, construction-base material, and agricultural lime. These quarries generally are dewatered to some extent to allow for the removal of the rock. The amount of water removed is generally greatest in the spring (especially if the quarry shuts down operation in the winter), but the amount of water removed varies between quarry operations. In general, more dewatering takes place from quarries in limestone as compared to those in dolostone due to the greater potential for solutioned-enhanced permeabilities (Kappel and Jennings, 2012).

Northern New York

In northern New York, the Trenton and Black River Group limestones and Beekmantown Group sandstone and dolostone that surround the Adirondack Mountains (fig. 1) have variable regional bedrock strike and dip. These regional differences in bedrock dip may limit the northern carbonate rocks from having a zone of allogenic recharge similar to the gently dipping, east-west trending carbonate-rock units of central and western New York. The Black River Limestone near Watertown is known for young karst development (shallow solutional fractures, sinkholes, and cave systems), especially where these rocks are adjacent to surface water, such as the Black River (Waller and Ayer, 1975). As an example of this type of karst development, a **fluorescein** dye trace was completed in the Black River Limestone in December 2016 to determine the connection between some small karst features (swallets) and the Black River. The dye took just over 4 hours to move from the swallets to the Black River, a distance of 1,500 ft, and when the dye appeared, its color intensity was not substantially diluted (fig. 13).

The Trenton Limestone, on the other hand, is not known for karst development or for comprising a regionally extensive aquifer system (Waller and Ayer, 1975). The difference between these two units is that the Black River Limestone is generally a massive and pure limestone prone to dissolution at shallow depth and the Trenton Limestone is more argillaceous (shaley), thinner bedded, and less prone to solution.

The Beekmantown Group of northern New York unconformably (former overlying bedrock units not present because of erosion millions of years ago) overlies the Precambrian Shield (Bernstein, 1992); the deep part of the Beekmantown Group is generally a sandstone, but the shallow units are dolostone. The aquifer within the Beekmantown Group is generally present at the weathered bedrock surface and within the relatively shallow fractured bedrock resulting from glacial rebound of the land surface. Fracture density and orientations vary across this system (Cooper and Mylroie, 2015), and

karst development is considered young and slowly developing. Bedrock quarries are present in all carbonate units in this region. Dewatering volumes necessary to maintain quarry operations are dependent on the type of bedrock being mined and the location of the quarry in relation to its regional topographic position within the local water table, with greater dewatering volumes being produced when the quarry is below the local water table.

Eastern and Southeastern New York

In the eastern part of New York, the strata of the Onondaga Limestone and Helderberg Group transition from dipping gently to the south in the Mohawk River Basin to steeply dipping along the western side of the Hudson River Valley. The eastern Mohawk basin has New York's most strongly developed karst, including the oldest known cave systems in the northeastern United States. Lauritzen and Mylroie (2000) have dated stalagmites in some caves to be more than 350,000 years old, indicating cave development through several complete glacial cycles (Cooper and Mylroie, 2015). These caves (figs. 14, 15, and 16) have been explored since the mid-1800s and have been extensively studied since the 1950s (Baker, 1976; Palmer, 1991; Palmer and others, 1991, 2003; Lauritzen and Mylroie 2000; Weremeichik and Mylroie, 2014; Cooper and Mylroie, 2015). It is likely that many more cave systems have yet to be discovered in this area. Numerous large carbonate quarries are present in this part of the State and, depending on their location and elevation within the welldeveloped karst aquifer system, need either substantial or little dewatering to carry on their quarrying operations. At Howe Caverns in Schoharie County, a quarry operation just downgradient from the caverns affected the flow and water level in the caverns until an inadvertent breach of the karst aguifer in the quarry was sealed.

Along the western Hudson River Valley, the tectonic forces that created the valley have caused the sedimentary rocks on the west side to become steeply inclined and fractured. Because of their steep orientation, carbonate rock here is less prone to deep groundwater infiltration as the solution process must have an outlet to the surface to enhance karst development. The age of karst here is varied, dependent on its topographic location. The karst features that have been reported are generally confined to near-surface exposures in these bedrock units (Cooper and Mylroie, 2015). In areas such as Kingston and Rosendale in Ulster County, extensive natural cement mines were developed beginning in the early 1800s. The miners followed veins of the argillaceous (shaley) Rondout Formation limestone that in some places are inclined as much as 70 degrees from horizontal (fig. 17). Other carbonate quarries are present throughout the region, which can have variable amounts of dewatering required during the quarry production season. Evaporite karst is not present in the eastern half of New York because there are no gypsum or halite deposits present in that area (fig. 1).



Figure 13. Dye trace in Black River Limestone, near Watertown, Jefferson County, New York. *A*, Injection of fluorescein dye into small closed depression where water flows laterally into a bedrock fracture (movement indicated by white arrows), and dye flowing *B*, in down-gradient bedrock fracture, *C*, from foundation drainpipe adjacent to the Black River, approximately 4 hours after injection, and *D*, into the Black River about 1,500 feet from dye injection point shown in *A*.

Photograph by W. Kappel, USGS

Photograph by W. Kappel, USGS

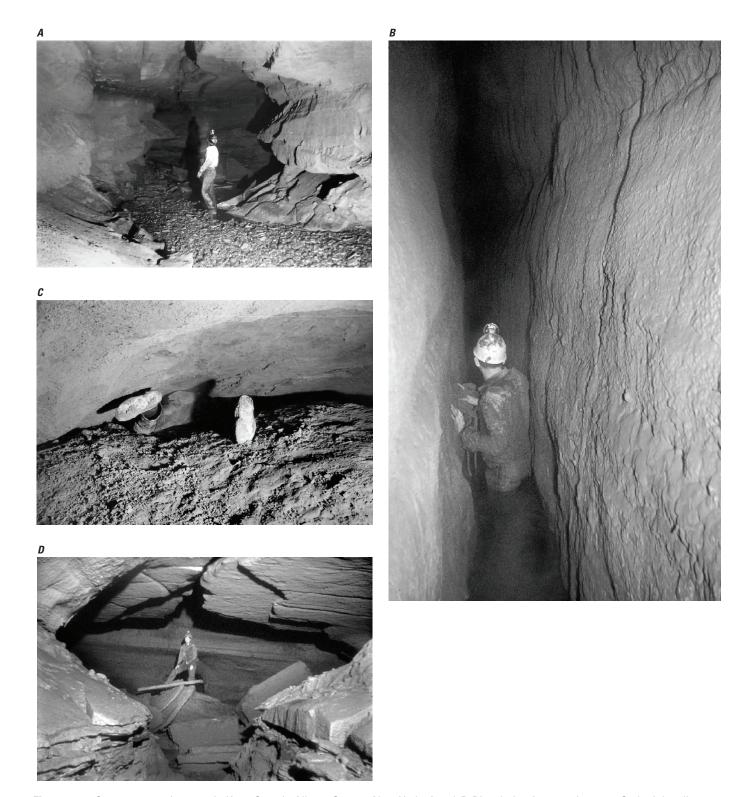


Figure 14. Cave passage features in Knox Cave in Albany County, New York. *A* and *B*, Dissolution features in cave; *C*, glacial sediments found within most cave passages; *D*, Coeymans Limestone in the ceiling and Manlius Limestone below. Photographs courtesy of Art Palmer, professor emeritus, State University of New York at Oneonta.



Figure 15. Features in Knox Cave in Albany County, New York. *A,* Linear passage along major joint in Manlius Limestone and *B,* the Gun Barrel cave feature along the Manlius Limestone (above) and Rondout dolomitic limestone (below) contact. Photographs courtesy of Art Palmer, professor emeritus, State University of New York at Oneonta.

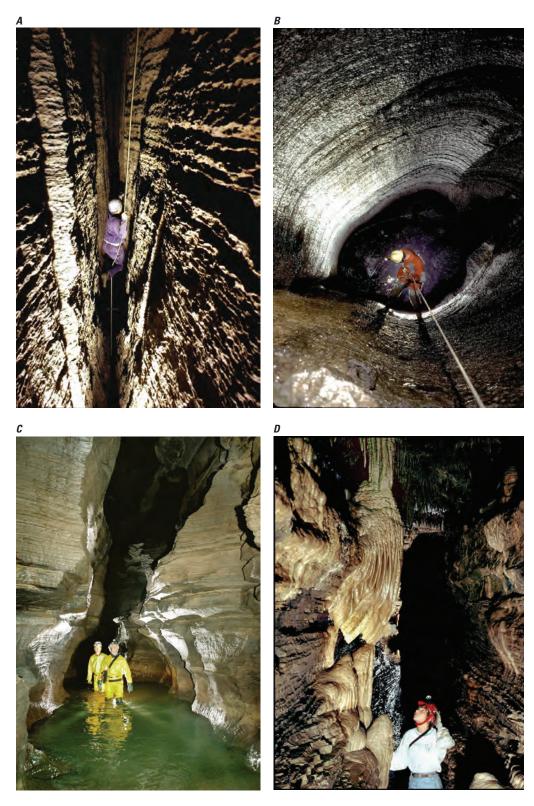


Figure 16. Cave passages from the valley of the Cobleskill Creek, a tributary to Schoharie Creek, in Schoharie County. *A*, Joint-controlled fissure passage in Coeymans Limestone; *B*, 78-foot-deep shaft enlarged by a major sinking stream, also in Coeymans Limestone; *C*, downstream passage in McFails Cave in Manlius Limestone; and *D*, calcite dripstone deposits in Manlius Limestone passageway. Photographs courtesy of Art Palmer, professor emeritus, State University of New York at Oneonta.



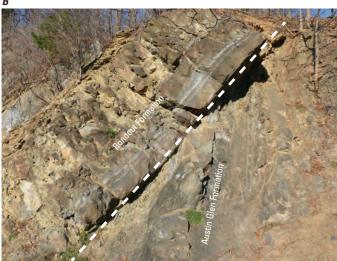


Figure 17. West side of the Hudson River Valley, New York. A, Entrance to a natural cement mine at Rosendale, New York, showing the dip of the bedrock; and B, Taconic unconformity (white dotted line) at Catskill, New York, between mid-Ordovician Austin Glen Formation and late Silurian Rondout Formation. Note the high-angled nature of the dip of both bedrock units. Photographs courtesy of Steven Schimmrich, Ulster Community College, State University of New York.

Marble of the Wappinger Group is present in the downstate region of New York into the New York City area (Inwood Marble) but other marble is present as inclusions in the Precambrian metamorphic rocks of the Adirondack Mountains in northern New York. As a result of the Acadian orogeny in the downstate area, the limestone bedrock in this region is highly deformed but the extent of marble deposits is limited (Cooper and Mylroie, 2015). During the construction of the Catskill and Delaware water-supply conduits from their respective watersheds to New York City, when borings and construction of the conduits crossed under the Rondout valley,

workers encountered glacial and interglacial bedrock valleys, overthrust faults, and numerous caves and clay-filled cavities and joints, about 800 ft below land surface. At its greatest depth, these features were below about 540 ft of unconsolidated material and an additional 260 ft into bedrock. In some cases, the cavities extended for vertical distances of 20 ft or more (Fluhr, 1973). The Inwood Marble in the New York City area was also intersected during conduit construction. In some cases, during the drilling of test borings, the marble bedrock was heavily fractured and eroded to depths greater than 300 ft, and a substantial amount of water had to be grouted or otherwise removed to allow for construction of the conduits (Fluhr and Terenzio, 1984). The deep karst development noted in Fluhr (1973) and Fluhr and Terenzio (1984) likely formed when glaciers were active and the high hydraulic pressure at the glacial ice-bedrock interface created ideal conditions for karst development. During deglaciation, thick sequences of glacial sediments were deposited, especially in the deeper bedrock valleys, possibly slowing karst development unless there was an active flow system to continue the karst development process.

Previous U.S. Geological Survey Karst Studies in New York

In the early 1980s, the USGS, in cooperation with the Erie County Department of Environment and Planning, monitored groundwater-level declines in a 2.2-mi-long by 800-ftwide closed depression area in eastern Erie County (Staubitz and Miller, 1987; fig. 18). More than 60 domestic water wells and several wetlands went dry and several new sinkholes developed or expanded. The water-level decline was caused by the combination of groundwater withdrawal at a limestone quarry and the diversion of that water away from a downgradient sinkhole-swallet in the eastern end of the large closed depression. The routing of the quarry water away from the former receiving swallet to an adjacent surface-water stream removed the allogenic recharge from the closed depression's groundwater watershed west of the quarry and caused the water-level declines noted in many homeowner wells within the depression further to the west.

In 2008, the USGS, in cooperation with the NYSDEC, initiated a pilot study to characterize the recharge settings of the carbonate karst in Genesee County in western New York (fig. 1) after a number of groundwater contamination issues were reported during the previous decade (Reddy and Kappel, 2010). Gently dipping Onondaga Limestone is present as an east-west band across Genesee County. The Onondaga Limestone crops out to the north along an escarpment and a 2- to 5-mi-wide band of the limestone is covered by glacial deposits to the south. A series of 1:24,000-scale maps was constructed and showed areas where focused recharge can most readily infiltrate into the karst carbonate bedrock and potentially contaminate the groundwater-flow system (Reddy

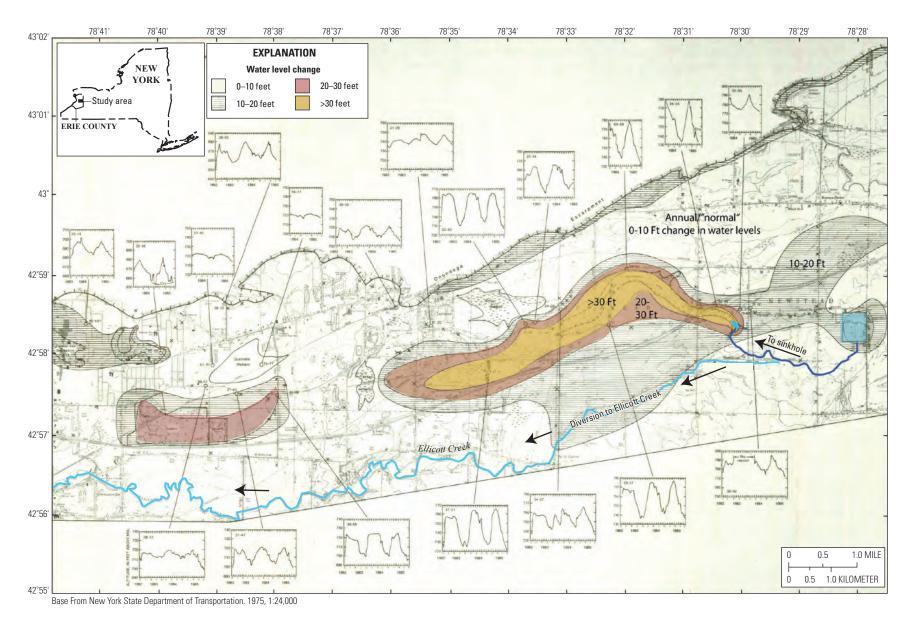


Figure 18. Decline in water levels in a large closed depression, western New York, from April 1981 to October 1985, indicating the effect of diverting quarry dewatering flow from a downstream sinkhole-swallet to Ellicott Creek; modified from Staubitz and Miller (1986, pl. 4). This diversion removed groundwater from the downstream sinkhole, which caused water-level declines greater than 30 feet in some down-gradient homeowner drinking-water wells.

and Kappel, 2010). Evaporite karst (shale with gypsum) north of the Onondaga Limestone escarpment was not characterized in that study. Using information from the USGS study of Genesee County, Cornell University and the NYSDEC developed manure management guidelines for karst areas in the county, including where concentrated animal feedlot operations were applying manure within karst areas of the County (Czymmek and others, 2011).

The same manure-management practices have been encouraged in other parts of State where the risk of applying manure may impact surface-water and groundwater resources. For example, Madison County (in central New York) has utilized these same manure-management criteria and several other counties in New York have considered following these guidelines to better protect their groundwater resources where karst is present (Karl Czymmek, Cornell University, oral commun., February 2017). Any farm that has a concentrated animal feedlot operation permit is required to follow these or previously designated guidelines to protect groundwater (New York State Department of Environmental Conservation, 2017b).

Between 2001 and 2011, the USGS, in cooperation with the U.S. Environmental Protection Agency (EPA), assisted with a study of volatile-organic compound contamination in the carbonate-bedrock aquifer system in Cayuga County between Auburn and Union Springs (a distance of 7 mi; Eckhardt and others, 2011). The USGS collected and compiled geophysical logs from 93 monitor and water-supply wells in the carbonate-bedrock aguifer between Owasco and Cayuga Lakes, two of central New York's Finger Lakes. The USGS analyzed geophysical logs along with core samples from drilled bedrock wells, inspected outcrops of the bedrock to define the stratigraphic units penetrated by wells, evaluated flow zones within the bedrock aquifers, and developed and installed multilevel monitoring systems for the study (U.S. Environmental Protection Agency, 2012). The USGS also assisted the EPA in a similar fashion in western New York near the town of LeRoy in northeastern Genesee County. In 1970, a railroad tanker car derailment spilled thousands of gallons of trichloroethylene into the local carbonate bedrock (Onondaga Limestone), contaminating drinking water wells (U.S. Environmental Protection Agency, 2017).

Karst Development in New York

The development of karst is variable across New York but is generally related to the geologic characteristics of the rock (carbonate versus evaporite), the quantity and quality of water that the rock is exposed to, and the length of time that the karst process has been occurring. The repeated glaciation of New York has assisted with karst development as glaciers have removed overburden and injected water directly into bedrock, enhancing karst development. However, at the same

time, repeated glaciation has destroyed former karst systems by eroding the karstic bedrock and creating a postglacial environment to start the karst-forming process once again.

Western and Central New York

In western and central New York, the gently southward dipping Onondaga Limestone and Lockport Group dolostone have young or early karst features, especially near their escarpments. Apparently during glacial advance, former escarpments were heavily eroded back to more competent bedrock, and the new escarpment initially had limited karst development. The land surface downgradient from each escarpment was usually mantled by glacial and postglacial sediments, which limited karst development until glacial rebound occurred, allowing the dissolution of bedrock joints and fractures as these features opened due to the land-surface rebound processes. The Onondaga Limestone is more prone to karst development than the dolostone of the Lockport Group because of its different geochemical composition. Although karst solution features are present in both units, the Onondaga has more well-developed karst flow systems (large closed depressions, early conduit-cave development, and sinkholerelated subsidence) than the Lockport Group (Staubitz and Miller, 1987). For example, in the town of Clarence (Erie County), a bowling alley was built in the early 1960s on a farm property known as "The Sinks" (The Clarence Museum, written commun., February 2017). The building collapsed into a sinkhole in the early 1980s, and reportedly, a section of New York State Route 5 in front of the bowling alley also slightly subsided at the same time, although the subsidence did not close the roadway. The remaining part of the bowling alley building was demolished, the sinkhole filled, and the site today [2020] remains undeveloped. Although both the Onondaga and Lockport bedrock units have some solution channel and modest cave development, these features more commonly occur near their escarpments and along adjacent steep erosional features (such as along the Niagara River gorge), but as indicated in the bowling alley story, solution features (sinkholes, closed depressions) have occurred elsewhere in the karstic rock.

Just north of the Onondaga Limestone escarpment, gypsum deposits are present within the underlying Salina Group shales (fig. 1). Karstification developed where gypsum deposits are close to the land surface and, in some cases, over former gypsum mines. In western New York, thin beds of gypsum (3 to 5 ft thick) have experienced extensive gypsum mining from the 1800s through the late 1990s. The extent of these mines is roughly from the Clarence area in Erie County eastward through parts of Genesee County to the southwestern part of Monroe County, in the Garbutt area. From Garbutt to just east of Syracuse (Onondaga and Madison Counties), gypsum was obtained from shallow open-pit quarries as the gypsum seams are discontinuous in this part of the State and more easily accessed from the land surface.

Where these former subsurface mines were developed, dissolution of the remaining gypsum support pillars or partial collapse of the mine roof has led to some land-surface subsidence and closed depression development. A number of geologic studies were performed in the town of Clarence, but north of the Onondaga Limestone escarpment, between 1972 and 1989 to determine whether residential and commercial development could occur over a part of the former National Gypsum Company mine (HydroTech Ground Water Consultants, 1989). A summary of these various studies indicated that future sinkhole development was possible because several sinkholes were already present over the mine in 1972, but the timing and extent of such development was difficult to quantify. The site over the former gypsum mine and other places nearby have since been fully developed.

In the town of Wheatland in southwestern Monroe County, the 460-acre Oatka Creek Park is maintained for hiking and cross-country running and skiing along 6.5 mi of trail. In the late 1990s, the USGS was asked why the land surface of the park was experiencing subsidence. A surface examination of the park revealed linear subsidence patterns generally aligned in a north-south direction. A search of the town's history (Schmidt, 1953) and a summary of gypsum resources in New York State (Newland, 1929) revealed that there were at least three former gypsum mines along the southern side of Oatka Creek in this part of the county. The mines had exploited gypsum seams that surfaced along the southern side of Oatka Creek near stream level and mine adits (dipping-tothe-south tunnels) were dug to remove the gypsum. In this case, the present-day park was built over the former Garbutt Gypsum Mine, which utilized such adits (Schmidt, 1953). Subsequent to these findings, the Monroe County Parks Department contracted with a private firm to perform a geophysical (electrical resistivity) survey to confirm the location of the adits. Results of the Oatka Creek Park survey indicated "the presence of numerous sinkhole and subsidence features related to either collapsed shafts [that is, subsidence features related to the bedrock collapse over the trend of the adits] or open-pit mining of gypsum near Oatka Creek" (Forrest Environmental Services, Inc., 2001, p. 2).

As a result of the amount of former gypsum mining in Wheatland, a mine subsidence hazard district was created as a zoning overlay. The stated purpose of the district was, in part, "[i]n the interest of the public health, safety, and welfare, these regulations are intended to protect development located in an area in the Town of Wheatland that has been subjected to underground mining, with much of such area [sic] not provided with maps of the undermining. More specifically, the purpose of these supplementary controls is to prevent the erection of habitable structures in an area that could be unsafe due to undermining, to minimize danger to public health..... due to subsidence" (Town of Wheatland, 2015).

In central and eastern New York (east of Rochester), the gypsum beds become thin to nonexistent because gypsum is present only as pockets and lenses (Tennissen, 1952), which

were mined from the land surface. Though these pockets are considered small, apparently some land-surface subsidence has occurred, although such subsidence has been reported in just a few areas. For instance, in Syracuse in Onondaga County, gypsum dissolution may have caused nearly 30 ft of landsurface subsidence since the last period of glaciation, about 12,000 years ago (Phillips, 1955). Such gypsum-dissolution subsidence likely occurred elsewhere within these limited gypsum deposits, but without substantial urban development, land-surface subsidence may not have been noticed other than possible development of isolated sinkholes or closed depressions, some of which may have developed into wetland areas. In the Syracuse area, sinkhole and subsidence have been reported, and in one case, an investigation of the local geology revealed the presence of a heavily weathered shale zone at the bedrock surface and a vertical zone of heavily weathered gypsiferous shale within surrounding unweathered shale that has anhydrite- and gypsum-bearing seams and inclusions (fig. 19) similar to what was indicated by Phillips (1955).

The Salina Group, and in particular the Syracuse Formation, contains beds of halite (rock salt) interbedded with shale (fig. 1). Near the outcrop of the Salina Group at land surface, any halite that may have been present was easily removed by dissolution during and following glacial periods when meltwater was likely forced into the heavily fractured rock by the overburden pressure of the glacier, dissolving halite that interacted with the injected water (Goodman and others, 2011). Any subsidence related to this solution process happened at that time and is of little concern presently. But halite has been and continues to be mined in the subsurface either by solution- or hard-rock-mining techniques in central New York, well to the south of the former outcrop area. There have been cases of land-surface subsidence related to both forms of halite mining. Though the extent of subsidence is locally important—such as subsidence from solution-well mining in the Tully Valley near Syracuse (Hackett and others, 2009) or room-and-pillar mining of halite in the Genesee Valley south of Rochester (Yager and others, 2012)—these types of halite-related subsidence are not extensive across the central and western part of the State.

Northern New York

In northern New York and along the western side of the Adirondack Mountains, strata of the Black River and Trenton Groups are nearly flat lying and have been fractured due to glacial advance and recession and continued tectonic doming of the Adirondack Mountains (Isachsen, 1981). Karst development in this bedrock unit is considered young as it is primarily influenced by recent glaciation and its proximity to flowing surface water and the zone of annual groundwater fluctuation (Waller and Ayer, 1975). In the vicinity of Watertown, the limestone of the Black River Group has karst features and "a vast network of caverns of varying heights and expanses with intricate ramifications and labyrinthine corridors [sic]

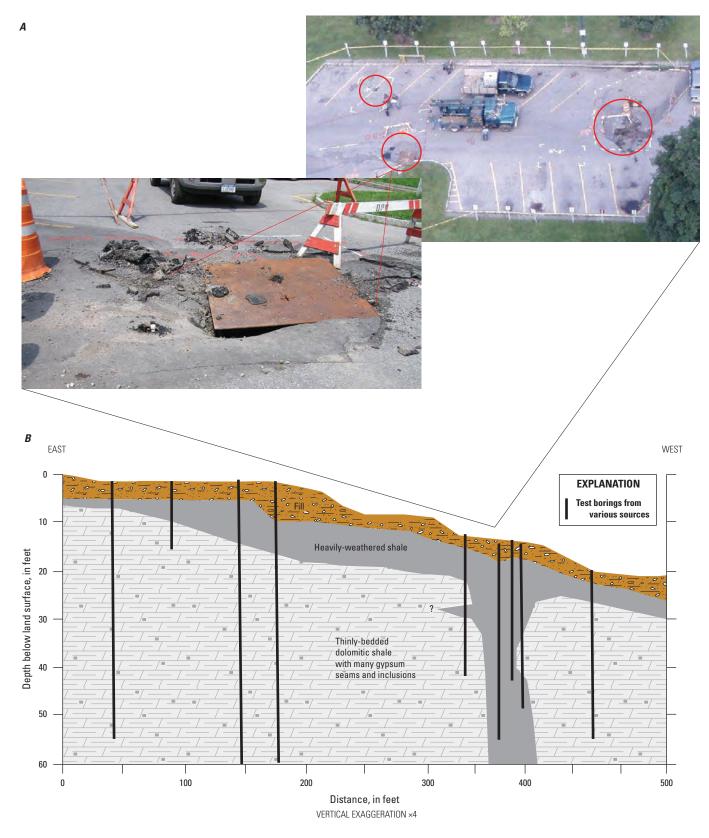


Figure 19. Sinkhole development in evaporite karst bedrock in Syracuse, Onondaga County, New York. *A*, Parking lot subsidence (red circles) in area where deeply solutioned bedrock is present, and closeup of surface area of one of the sinkholes. *B*, Schematic cross section of the depth to heavily weathered shale and more competent shale with gypsum seams and inclusions beneath, and a deep zone of gypsum-solutioned bedrock where continual land-surface subsidence (shown in *A*) has occurred.

occupy the thick and expansive deposit of limestone over which the city of Watertown has been built" (Northern New York Genealogy, 2001). Some of the caves of various sizes and lengths were used to store beer during the summer season, another cave was an attraction for a local amusement park where cave boat tours were given, and several caves in Watertown were considered for use as bomb shelters during the Second World War. Of note is that the entrances to all these caves are presently [2020] sealed.

The Trenton Group, which overlies the Black River Group, does have some vertical fracturing but is less prone to karst development because of its chemical makeup: an argillaceous (shaley) limestone with thin limestone interlayered with shale throughout the Trenton sequence. The Beekmantown Group, primarily in northern New York but also present along the east side of the Adirondack Mountains (fig. 1), is mostly composed of dolostone with interbedded limestone and sandstone. Where the Beekmantown bedrock is near the land surface, it has been weathered and fractured by glacial advance and retreat (glacial rebound), which does enhance some karstic development. This is especially the case where the bedrock is associated with underlying dolostone or permeable sandstone, which allows surface water to move through an upper dolostone layer to a point of discharge or into a deeper bedrock flow system. Some limestone in this region has also been metamorphosed into small bands of marble that also exhibit karstic attributes.

Eastern and Southeastern New York

Karst in eastern New York is well developed because of the great thicknesses of carbonate bedrock (Onondaga and Helderberg Group) and the extended period of time in which the bedrock has been exposed to repeated glaciation and interglacial erosion, especially in the eastern Mohawk River basin (Baker, 1976; Cooper and Mylroie, 2015). Because of the presence of this well-developed karst, several studies have linked karst and the need to protect groundwater from contamination. "These [karst] features complicate the flow of ground water and create special problems regarding water supply. Due to thin soils, there is a potential for widespread ground water contamination and a limited ability to support foundations and building loads. There are few regulatory controls devoted to protecting karst areas" (Schoharie Planning and Development Agency, 2002).

A study of the Cobleskill Plateau by Braun (2000, p. 62–63) indicated that "Of the 40 wells for which chemical analysis was available, 87.5 percent showed some level of contamination from nitrates and 82.5 percent contained phosphates... Of the 43 wells tested for bacteria using a presence or absence test, 81.4 percent showed a presence. These figures indicate a strong degree of interaction of surface water with well water on the limestone plateau. All the wells [tested] on the Cobleskill Plateau are located in an area where dairy farming is the number one industry. One of the wells, which had

high nitrate levels, had raw manure coming through the water system....as reported by the homeowner. The well drew water from 116 ft below the surface [likely the location of the well pump, as the water level in the well was not noted]. Adjacent to the house are crop fields where sinkholes form and are filled with fieldstone... The fieldstones fill the sinkholes, but they do not seal the sinkholes, which have a direct connection with groundwater."

The Helderberg-Cobleskill Plateau region (Baker, 1976; Palmer and others, 2003; Cooper and Mylroie, 2015) is an incised carbonate terrain resulting from repeated periods of glaciation and erosion, which, over time, has created solutionderived fracture conduits and well-developed cave systems (figs. 14, 15, and 16). The glacial and interglacial processes bisected the land surface to create a number of hydraulically isolated carbonate rock "islands" where vertical bedrock joints and sinkholes intersected south-dipping bedding planes to create a classic karst topography: upgradient losing streams and swallets flowing to downgradient springs, some of which might exhibit artesian flow where former, lower spring outlets were blocked by more recent glacial debris (Palmer and others, 2003). The area of Schoharie, Albany and Greene Counties (Helderberg-Cobleskill plateau) has the oldest known and largest number of well-developed, most-explored, and documented cave systems in New York (fig. 20). Many of these cave systems have been mapped, and some cave-catchment areas have been defined using dye tracing techniques.

Further downstate, along the western side of the Hudson River valley, the Onondaga and Helderberg Groups exhibit localized karst development. A study of the New York City water supply conduit flowing from the Rondout Reservoir through the West Branch Tunnel under Wawarsing (fig. 21) indicated a hydraulic connection between the unlined conduit and the overlying bedrock and near-surface unconsolidated aquifers (Stumm and others, 2012). During construction of the West Branch Tunnel, this section of the tunnel was observed to have been highly faulted and weathered, had inflow of more than 9,000 gallons per minute, and required more than 20,000 bags of concrete and steel reinforcements to stabilize the bedrock (Stumm and others, 2012). This condition indicates fracture solutioning and enhanced permeability within the carbonate bedrock and was likely influenced by the active geologic history of the Hudson River valley creating extensive faulting within the bedrock.

Along Interstate 87 (New York State Thruway) near the city of Saugerties in Ulster County, the use of road salt on the interstate highway has apparently affected a number of homeowner wells that were drilled into the steeply dipping Onondaga Limestone. A nearby water line was extended to these homes to resolve the problem (Paul Rubin, HydroQuest, written commun., January 2016).

There is a number of named cave systems in this region (fig. 20) but far fewer than in the same bedrock units just to the north. Further south, the Inwood Marble also exhibits some karst features, although the amount of karst is limited

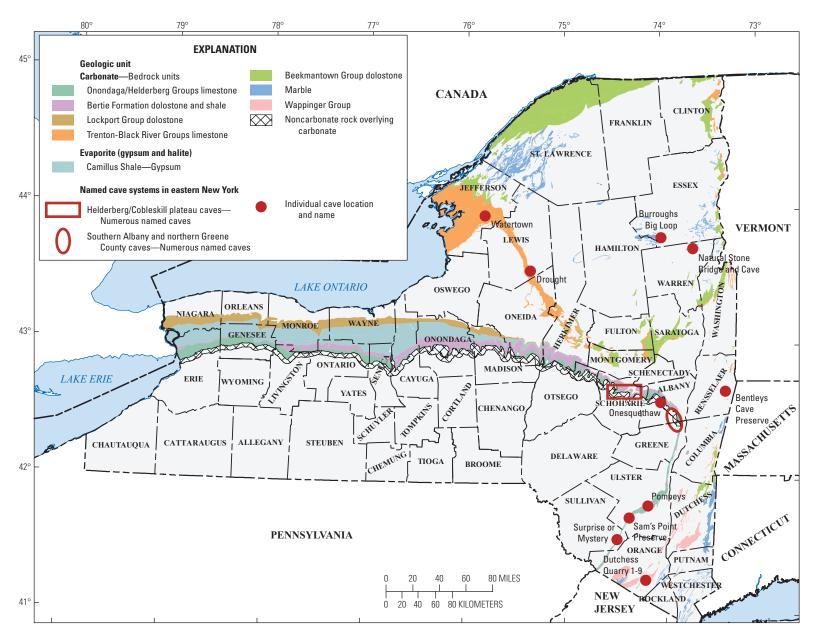


Figure 20. Selected named cave systems in eastern New York.

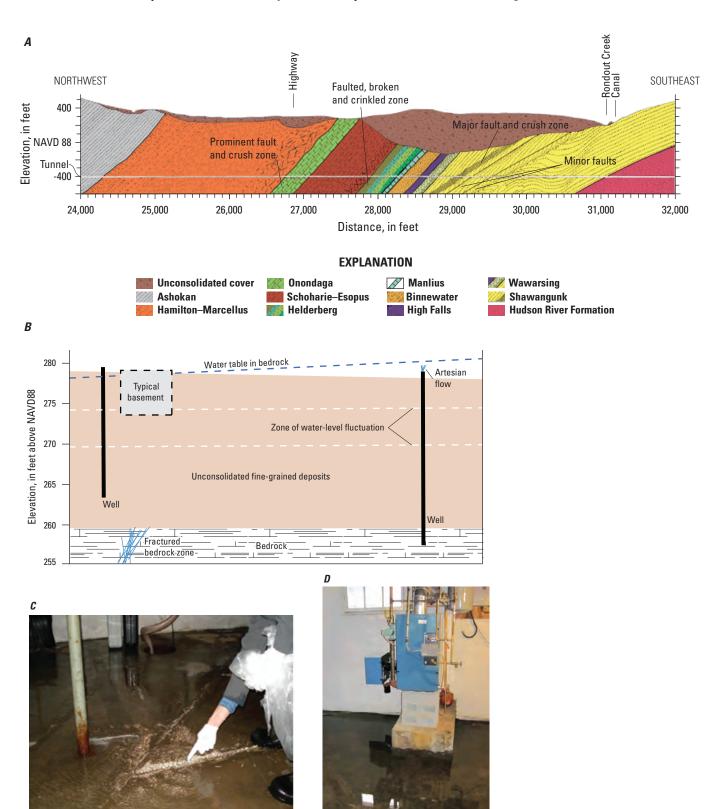


Figure 21. Hydrogeologic conditions at Wawarsing in Ulster County, New York. *A*, Geologic cross section for the Rondout tunnel near Wawarsing, showing bedrock types and fracture-fault locations; modified from Stumm and others (2012); *B*, schematic section of the water-table configuration for unconsolidated and bedrock aquifers. *C* and *D*, Flowing artesian conditions in flooded Wawarsing area basements; from Stumm and others (2012). NAVD 88, North American Vertical Datum of 1988.

by the small extent of the formation and the amount of urban development that has obscured the expression of karst features at land surface.

Closed-Depression and Focused-Recharge Inventory

The objective of the geographic information system (GIS) closed-depression and focused-recharge analyses conducted for this study was to develop a comprehensive inventory of all closed depressions proximal to carbonate, evaporite, or marble units within the State. The distribution of these bedrock units in New York form the basis of the study area. The inventory is

restricted to an area defined by bedrock lithologies where karst is known to be present. However, no inference can be made that a mapped closed depression is karstic in origin. Individual evaluation of each inventoried depression would need to be conducted to determine the processes that created the feature. That level of effort is beyond the scope of this study.

A closed depression is defined as an enclosed area lower than the surrounding area that has no surface-drainage outlet and from which water escapes only by evaporation or subsurface drainage (that is, an area of focused recharge). On USGS topographic maps, a closed depression is indicated by a hachured contour line forming a closed loop (Natural Resources Conservation Service, 2017). Examples of closed depressions in a karst area are shown in figure 22. The map representation applies to closed depressions of both natural and anthropogenic origin. Many of the closed depressions

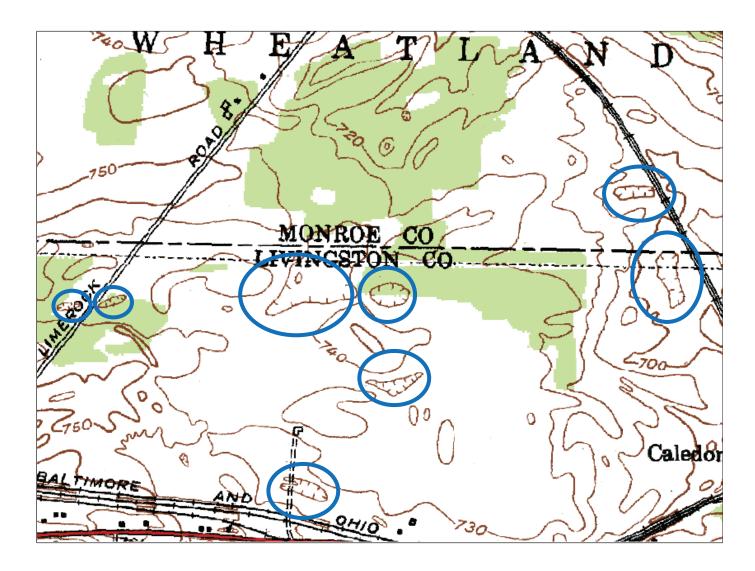


Figure 22. Examples of closed depressions as they appear on a topographic map; modified from U.S. Geological Survey (1950). Closed depressions are indicated within blue circles.

included in this inventory, though not the result of karst processes, are within or adjacent to a karst landscape and could contribute to focused recharge.

Study Area

The study area boundary adopted for the GIS analyses was created by intersecting an index of USGS 1:24,000-scale, 7.5-minute quadrangle maps (U.S. Geological Survey, 2014) with data containing selected bedrock geologic units (table 2; Horton and others, 2017). The final study area boundary covers more than 20,000 square miles, equivalent to 38 percent of New York State and transecting 42 counties (fig. 23).

Data Sources

Data acquired for the GIS analyses can be grouped into two categories: (1) source data for identifying closed depressions and (2) ancillary data used to assign geologic, soil, and land cover characteristics to the inventoried closed depressions (table 3). The spatial relation between ancillary characteristics and closed depressions can be used to assess which closed depressions have a greater potential for contributing focused recharge.

Closed Depressions

The closed depressions were derived from three data sources and were compiled in the following order: (1) digital contour database of closed depressions, (2) digital raster graphic database of closed depressions, and (3) light detection and ranging (lidar) database of closed depressions. All three databases are from Reddy and others (2020). There is no duplication of features among the three databases. Additionally, the closed depressions inventoried were compared with closed depressions mapped in other published geospatial data to eliminate duplication with those datasets. The datasets referenced were New York State Department of Environmental Conservation (2017a) and U.S. Geological Survey (2013).

Table 2. Reclassified bedrock geologic units for karst aguifers in New York.

[Geologic units are from Horton and others (2017). —, not designated]

Major unit 1	Major unit 2	Major unit 3	Formation or unit name			
		Onondaga				
Limestone	Chert	Shale; conglomerate	Glenerie Formation			
Sandstone	Limestone	_	Oriskany Formation			
Limestone	Sandstone	Chert	Onondaga and Bois Blanc Limestones			
Limestone	Chert	_	Onondaga Limestone			
Shale	Limestone	Siltstone	Onondaga Limestone			
Limestone	Dolostone (dolomite)	Sandstone; shale; conglomerate	Undifferentiated Lower Devonian and Silurian rocks			
Siltstone	Limestone	Sandstone; dolostone (dolomite); shale	Rondout Formation			
		Akron				
Limestone	Dolostone (dolomite)	_	Helderberg Group			
Dolostone (dolomite)	Shale	_	Akron Dolostone			
Dolostone (dolomite)	Shale	Limestone	Cobleskill Limestone			
Limestone	Shale	Dolostone (dolomite)	Cobleskill Limestone			
Dolostone (dolomite)	Shale	Evaporite	Syracuse Formation			
Camillus						
Shale	Dolostone (dolomite)	Evaporite; black shale	Camillus, Syracuse, and Vernon Formations			
Shale	Dolostone (dolomite)	Evaporite	Camillus and Syracuse Formations			
Shale	Black shale	_	Vernon Shale			

Table 2. Reclassified bedrock geologic units for karst aquifers in New York.—Continued

[Geologic units are from Horton and others (2017). —, not designated]

Major unit 1	Major unit 2	Major unit 3	Formation or unit name
		Lockport	
Black shale	Dolostone (dolomite)		Ilion Shale
Dolostone (dolomite)	Limestone	_	Lockport Group
		Trenton-Black R	iver
Shale	Limestone	Dolostone (dolomite); conglomerate	Poultney Formation ("A" Member)
Limestone	_	_	Balmville Limestone
Limestone	Dolostone (dolomite)	Shale; arkose; chert	Black River Group
Carbonate	Melange	_	Cambrian through Middle Ordovician carbonate rock
Limestone	Siltstone	_	Valcour, Crown Point, and Day Point Limestones
Limestone	Shale	_	Trenton Group
Limestone		_	Trenton and Black River Groups, undivided
Limestone	Dolostone (dolomite)	_	Trenton and Black River Groups, undivided
		Beekmantown	1
Dolostone (dolomite)	Limestone	Sandstone; quartzite	Beekmantown Group
Sandstone	_	_	Potsdam Sandstone
Dolostone (dolomite)	Sandstone	Shale	Theresa (Galway) Formation
Dolostone (dolomite)	Shale	_	Briarcliff Dolostone and Pine Plains Formation
Dolostone (dolomite)	Limestone	Chert; siltstone	Beekmantown Group
Dolostone (dolomite)	Sandstone	_	Theresa Formation
Limestone	Dolostone (dolomite)	Chert	Copake and Halcyon Lake Formations, Rochdale Limestone
		Marble	
Marble	Schist	Granulite; quartzite	Inwood Marble
Schist	Marble	Calc-silicate rock	Manhattan Formation ("A" Member)
Marble	_	_	Stockbridge Formation
Marble	Calc-silicate rock	Amphibolite; granulite; gneiss; quartzite; schist	Undifferentiated dolomitic and calcitic marbles interlayered with significant amounts of calc-silicate rock
Marble	Calc-silicate rock	Amphibolite	Calcitic and dolomitic marble
Gneiss	Marble	Calc-silicate rock	Rusty and gray biotite-quartz-feldspar gneiss
		Wappinger	
Shale	Conglomerate	Limestone	Germantown Formation
Dolostone (dolomite)	Shale	_	Stissing Formation
Limestone	Dolostone (dolomite)	Sandstone; shale; conglomerate	Undifferentiated Lower Devonian and Silurian rocks
Limestone	Dolostone (dolomite)	Shale; chert	Wappinger Group

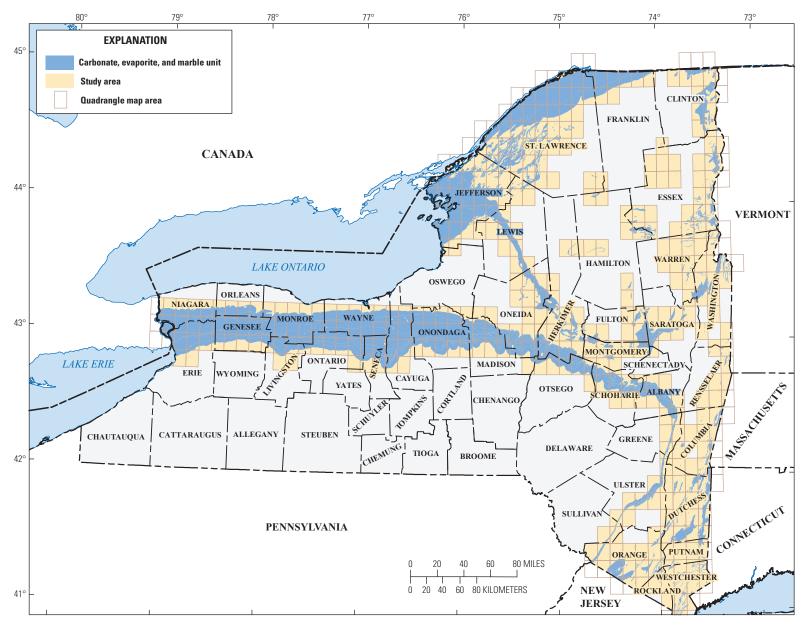


Figure 23. Study area boundary and selected U.S. Geological Survey 7.5-minute quadrangle maps adopted for geographic information system (GIS) analyses of karst aquifers in New York. Selected bedrock geologic units are shown.

Table 3. Sources for data acquired to conduct geographic information system analyses for karst aguifers in New York.

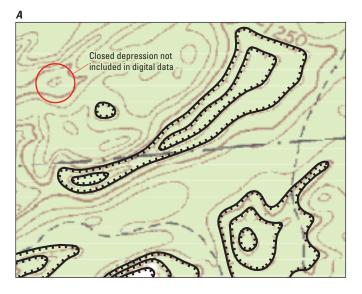
[lidar, light detection and ranging; DEM, digital elevation model; gSSURGO, gridded Soil Survey Geographic database; NLCD 2011, 2011 National Land Cover Database]

Data Data type		Source		
		Closed depression identification data		
Digital contours	Vector	Tyler and Greenlee (2012)		
Digital raster graphic images	Raster 1:24,000-scale, 7.5-minute topographic maps (U.S. Geological Survey, 2020)			
Lidar-derived bare-earth DEMs	Raster	New York State Office of Information Technology Services (2017)		
		Ancillary data		
Bedrock geology	Vector	Horton and others (2017)		
Soils	Raster	gSSURGO for New York (Natural Resources Conservation Service, 2014)		
Land cover Raster NLCD 2011 (U.S. Geological Survey, 2015)		NLCD 2011 (U.S. Geological Survey, 2015)		

The digital contour database of closed depressions contains features derived from data associated with Tyler and Greenlee (2012). The source data are a statewide contour dataset that was generated from the National Elevation Dataset (U.S. Geological Survey, 2017a) and the National Hydrography Dataset (U.S. Geological Survey, 2017b) in a fully automated process. The process created digitally produced contours that match, as closely as possible, the cartographic contours that appear on the 1:24,000-scale, 7.5-minute USGS topographic maps. An exact match was not possible because the digitally produced contours are based on newer, more accurate horizontal and vertical datums than those used to create the contours appearing on the topographic maps

(Tyler and Greenlee, 2012). Furthermore, deletion of very small polygonal contours was a processing step used in the development of the digital contours (Tyler and Greenlee, 2012). Within the study area, the discrepancy between digitally produced closed depression contours and closed depression contours appearing on 7.5-minute topographic maps is substantial. The digital contour lines representing closed depressions were converted to polygons and added to the inventory (fig. 24).

The features included in the digital raster graphic database of closed depressions were compiled to address the discrepancy between digitally produced closed depression contours and closed depression contours appearing on



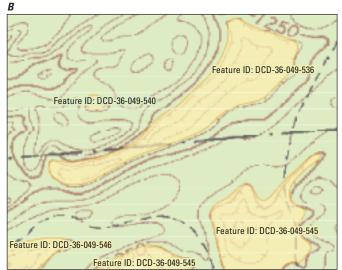


Figure 24. Closed depressions converted from line to polygon features. *A,* Closed depression contours appearing on the topographic map and in the digital contour dataset (black hachured lines; Tyler and Greenlee, 2012). An example of a closed depression that appears on the topographic map but is not included in the digital contour dataset (closed depression circled in red); *B,* closed depression contour lines from the digital contour dataset converted to polygons and assigned feature identification numbers for the digital contour database (Reddy and others, 2020).

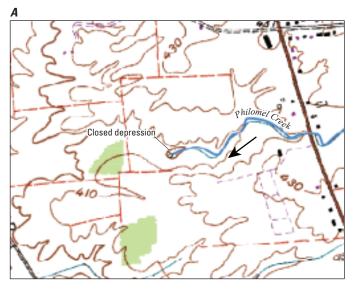
7.5-minute topographic maps. The features in the digital raster graphic database of closed depressions represent the locations of closed depressions that appear on the 7.5-minute topographic maps but were not represented in the digital contour dataset because of the previously described processing limitations. These closed depressions were digitized as feature centroids from an assemblage of approximately 650 digital raster graphic (DRG) images of scanned USGS 1:24,000-scale topographic maps. A georeferenced DRG was added as a background layer in a GIS. A dataset of disconnected stream reaches imported from the National Hydrography Dataset (U.S. Geological Survey, 2017b) for New York was used to assist with identification of closed depression features appearing on the DRG images (fig. 25).

The lidar database of closed depressions contains polygon features generated from analysis of 1- and 2-meter (m; 3.28- and 6.56-ft) horizontal resolution lidar-derived bareearth digital elevation models (DEMs; New York State Office of Information Technology Services, 2017). The DEMs were made from multiple collection projects conducted between 2005 and 2017. The vertical accuracy of the lidar projects ranged from a 4.0- to 18.5-centimeter (cm; 1.6- to 7.3-in.) root mean square error, and the horizontal accuracy of each project had less than 2-m root mean square error. At the time of analysis [2017], lidar data existed for approximately 85 percent of the study area (fig. 26). The DEMs were processed to identify closed depressions not represented in the digital contour database or the digital raster graphic database. The processes used followed the methods described by Doctor and Young (2013) using the ArcHydro Tools data modeling toolbox and by Wall and others (2015) using the Esri toolbox Hydrocutter.

A lidar-derived bare-earth DEM was created for each 7.5-minute quadrangle map within the extent of the lidar coverage. Each DEM quadrangle was reconditioned using ArcGIS version 10.5.0 (Esri, 2016) to produce a polygon dataset of possible closed depressions. Reconditioning involved the removal of artificial dams along manmade features, such as streets and railroads, where surface water flows through drainage structures that are hidden to the DEM. Mapped inventories of culverts and bridges are not complete as these features are numerous, often undocumented, or on private land; therefore, it is necessary to develop a method to restore connectivity of flow paths. The hidden flow paths need to be added to the DEM, artificially lowering the DEM at locations of drainage structures, thereby extending flow paths through artificial dams and restoring connectivity of surface-water flow. To do so, methods developed by Doctor and Young (2013) and Wall and others (2015) were incorporated and combined.

Doctor and Young (2013) recondition DEMs by digitally creating flow paths through road and railroad embankments. The reconditioned DEM is run through the ArcGIS Fill tool (part of the Spatial Analyst tools) to identify cells that are lower than all its adjacent neighbors and raise the value of that cell to the lowest value of an adjacent cell. A depth raster is then generated by subtracting the reconditioned DEM from the filled DEM. Cells within the depth raster that satisfy a minimum depth threshold are extracted and converted to polygon features having area characteristics. A final closed depression dataset is then generated that satisfies both minimum depth and area thresholds.

Wall and others (2015) developed the Hydrocutter toolbox to aide in connecting streams through roads and railroad embankments. Two separate tools, Hydro and Cutter, are



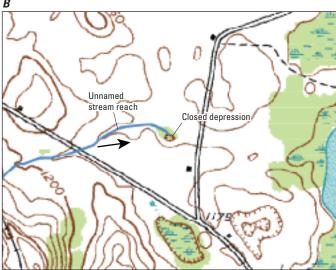


Figure 25. Closed depressions not represented in the digital contour data because of processing limitations that were added to the digital raster graphic database (Reddy and others, 2020) from analysis of digital raster graphic (DRG) images and disconnected stream reaches from the National Hydrography Dataset (NHD; U.S. Geological Survey, 2017b). *A*, Closed depression near Watertown, Jefferson County; *B*, Closed depression near Pine Woods, Madison County.

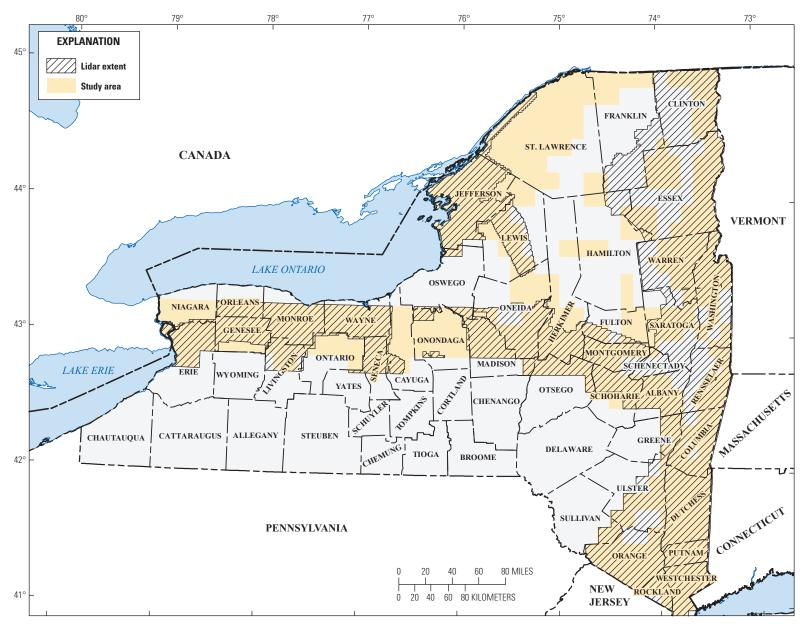


Figure 26. Extent of light detection and ranging (lidar)-derived digital elevation models within the karst aquifer study area in New York.

included with the toolbox, which can be used iteratively to increase the precision of closed depression detections. Hydro allows flow measures to be manipulated and expose flow paths that do not exhibit a surficial, geomorphic expression such as a stream channel. Because of the scope of this study, Hydro was not used because flow paths from the National Hydrography Dataset (U.S. Geological Survey, 2017b) were created with adequate flow conditions to produce closed depressions of the desired scale. Cutter intersects and buffers roads and railroads with flow paths to generate a reconditioned raster DEM, which is then filled. A depth raster is generated by subtracting the reconditioned DEM from the filled DEM, and a final closed depression dataset is produced that satisfies both minimum depth and area thresholds. Results from the two methods were combined into one dataset.

The number of possible depressions identified per DEM quadrangle is determined by the minimum depth and area thresholds applied during data processing. Thresholds used for this statewide study were set at 1-m (3-ft) depth and 4,047-square meter (m²; 1-acre) area; these threshold values are greater than what is typically used for lidar-based sinkhole identification studies. For the purposes of this study, the use of lidar was primarily intended to identify closed depressions that were not represented in the digital raster graphic database, in the same manner that the DRG images were used to identify closed depressions not represented in the digital contour database. For that reason, the threshold values were based on random sampling of DRG-derived closed depressions within the study area and represent the approximate mean geometric characteristics of the closed depressions sampled. Figures 27 and 28 provide examples of lidar-derived anthropogenic and

naturally formed closed depressions inventoried using these thresholds. For ongoing and planned large-scale, county-based assessments in New York, the thresholds will be reduced to 10- and 30-cm (3.9- and 11.8-in.) depth and 100-m² (0.025-acre) area.

Aerial imagery from Pictometry International (Pictometry International Corp., 2015) and Google Earth (Google, 2015) was used to photographically verify each lidar-derived closed depression. This process provided the opportunity to delete any remaining artificial depressions and to include a brief description of each feature. The lidar-derived features were grouped into a six-category classification scheme (table 4).

Ancillary Data

Multiple thematic datasets can be used to assess the probability that a closed depression contributes to focused recharge. For this study, the thematic data selected are bedrock type, soil type, soil infiltration rate, and land cover. For each closed depression that was inventoried, a series of overlay analyses were used to determine the predominant bedrock type, soil type, and infiltration rate within a 250-m (820-ft) buffer around the feature. Land cover percentages within a 250-m (820-ft) buffer around each feature were also computed. The results from the overlay analyses for each closed depression are summarized in the "Results" section of this report; results are recorded as feature attributes to the data included in Reddy and others (2020).

The bedrock-type attribute was derived from the State Geologic Map Compilation Geodatabase of the Conterminous United States (Horton and others, 2017). Bedrock geologic

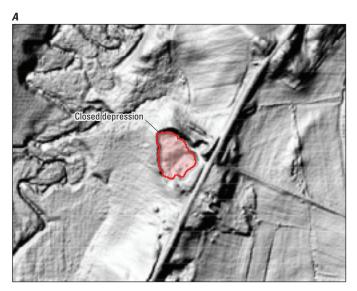




Figure 27. A closed depression near Argyle, Washington County, New York, identified using light detection and ranging (lidar)-derived digital elevation models (DEMs). This feature (inactive gravel pit) was not represented in other data sources; the feature as it appears on *A*, lidar-based hill-shade relief imagery and *B*, Google Earth imagery. The feature is divided between agricultural and developed land cover classes, underlain by noncarbonate bedrock (Canajoharie Shale) and glacial outwash soil with a high infiltration rate.

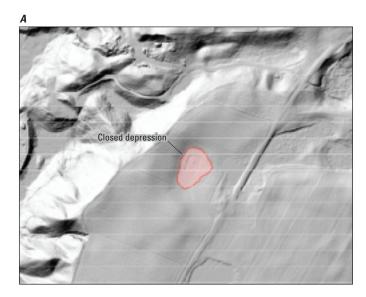




Figure 28. A closed depression near Ludlow Corners, Oneida County, New York, identified using light detection and ranging (lidar)-derived digital elevation models (DEMs); this feature was not represented in other data sources. The probable sinkhole as it appears on *A*, lidar-based hill-shade relief imagery and *B*, Google Earth imagery. The feature is located in an agricultural field, underlain by carbonate bedrock and glacial outwash soil with a high infiltration rate.

units were selected based on their established association with karst development; these units were then grouped into eight categories (table 2). A 2-mi-wide area of noncarbonate bedrock extending from Erie County to Greene County was included in the bedrock-type overlay analyses. In this area, which is south of the Onondaga and Helderberg Groups limestone, carbonate bedrock is overlain by shale. This represents an area of potential allogenic recharge. Surface water in this area flows from shale onto the adjacent carbonate bedrock and contributes to the potential for focused recharge (fig. 29).

Soil type was derived from the gridded Soil Survey Geographic database data for New York (Natural Resources Conservation Service, 2014). Based on research at Cornell University (Czymmek and others, 2004) and the U.S. Department of Agriculture (Stephen Page, Natural Resources Conservation Service, written commun., May 2017), 182 soil types in New York were identified as having characteristics that potentially present an increased risk for groundwater contamination. For this study, the 182 soil types were grouped into five classes (table 5). Three of the five soil classes

Table 4. Classification scheme for light detection and ranging (lidar)-derived closed depressions in New York.

[FeatCode, feature code]

Feat- Code	Origin	Description
1	Natural	Definite closed depression, natural in origin, for example, a closed depression in agricultural land exhibiting sinkhole signature.
2	Anthropogenic	Definite closed depression, anthropogenic in origin, for example, a small gravel pit or a surface water detention basin.
3	Natural and anthropogenic	Probable closed depression, includes natural and anthropogenic in origin, for example, shallow depression in agricultural land or probable detention basin at industrial park.
4	Natural and anthropogenic	Possible closed depression, includes natural and anthropogenic in origin, same as FeatCode 3 but with less certainty.
5	Anthropogenic	Impeded areas where surface water drainage is impeded, resulting in standing liquid, not a true closed depression, anthropogenic in origin, for example, surface water that collects along a roadbed near a culvert.
6	Natural and anthropogenic	Indeterminate, lidar-derived features that could not be verified through photographic analysis, for example, closed depressions in heavily wooded areas.

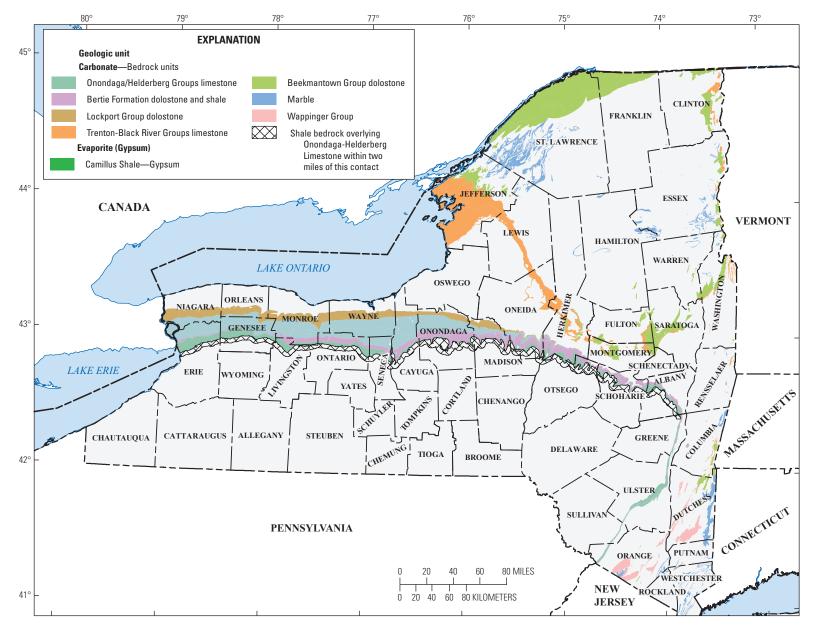


Figure 29. Shale bedrock in New York that may contribute allogenic recharge to the adjacent carbonate-rock aquifer.

Table 5. Classification scheme for soils in New York.

[Soil classifications are from the Soil Survey Geographic database (SSURGO; Natural Resources Conservation Service, 2019). in., inch]

	SSURGO	map	unit	name
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Soil class 1, depth less than 20 in. over carbonate bedrock

Benson

Farmington

Galoo

Gouverneur

Hannawa

Joliet

Kings Falls

Rockland, limestone

Rubbleland (Genesee County)

Ruse

Summerville

Soil class 2, depth less than 40 in. over carbonate bedrock

Aurora

Brockport

Camillus

Chaumont

Chippeny

Galway

Groton variant (Jefferson Co.)

Guff

Guffin

Lairdsville

Lockport

Madalin variant (Montgomery and Schenectady Co.)

Matoon

Neckrock

Nehasne

Newstead

Ogdensburg Palatine

Riga

Sun variant (Monroe Co.)

Varick

Wassaic

Wilpoint

Soil class 3, depth range 20 to 60 in. over carbonate bedrock

Allis

Angola

Appleton, limestone substratum (Monroe, Orleans)

Cayuga, limestone substratum (Niagara)

Table 5. Classification scheme for soils in New York.—Continued

[Soil classifications are from the Soil Survey Geographic database (SSURGO; Natural Resources Conservation Service, 2019). in., inch]

SSURGO map unit name

Cazenovia, bedrock substratum (Monroe, Niagara, Orleans, Wayne)

Collamer, bedrock substratum (Jefferson, Seneca)

Gardenisle

Hilton, bedrock substratum (Livingston, Niagara, Orleans, Wayne)

Honeoye, limestone substratum (Monroe)

Kendaia, bedrock substratum (Orleans)

Lakemont, shale substratum (Orleans, Wayne)

Lima, limestone substratum (Monroe)

Lyons, rock substratum (Orleans)

Mohawk, shale substratum (Herkimer)

Niagara, bedrock substratum (Jefferson)

Nuhi

Ontario, bedrock substratum/moderately shallow (Cayuga, Niagara,

Orleans, Seneca)

Ontario, stony (Genesee: OsB, OsC)

Ovid, bedrock substratum (Monroe, Niagara, Orleans)

Redwater

Stockbridge (Ulster)

Uwihreh

Yunenyeti

Soil class 4, well-drained glacial outwash bedrock

Adams

Agawam

Allagash

Allard

Allard variant

Altmar

Alton

Atsion

Au Gres

Barbour

Berryland

Bonaparte

Carver

Castile

Champlain

Chenango

Colonie

Colosse

Colton

Constable

Copake

[Soil classifications are from the Soil Survey Geographic database (SSURGO; Natural Resources Conservation Service, 2019). in., inch]

Table 5. Classification scheme for soils in New York.—Continued

[Soil classifications are from the Soil Survey Geographic database (SSURGO; Natural Resources Conservation Service, 2019). in., inch]

SSURGO map unit name	SSURGO map unit name		
Covert	Phelps		
Croghan	Philo		
Deerfield	Pipestone		
Deford	Pits, gravel and sand		
Deinache	Plainfield		
Deposit	Plymouth		
Duane	Podunk		
Duneland	Pompton		
Duxbury	Pootatuck		
Elnora	Raypol		
Enfield	Rippowam		
Fahey	Riverhead		
Fluvaquents	Rumney		
Fredon	Scarboro		
Gougeville	Sciota		
Granby	Searsport		
Grattan	Stafford		
Groton	Sudbury		
Halsey	Suncook		
Haven	Trestle		
Hempstead	Trout River		
Hinckley	Tunkhannock		
Homer	Udifluvents		
Hoosic	Udipsamments		
Howard	Waddington		
Jebavy	Wainola		
Junius	Wallace		
Kars	Walpole		
Knickerbocker	Wampsville		
Linlithgo	Wappinger		
Middlebury	Wareham		
Mooers	Windsor		
Naumburg	Wyalusing		
Ninigret	Soil class 5, depth less than 20 in. over noncarbonate bedrock		
Oakville	Abram		
Occum	Arnot		
Occum variant	Canaan		
Olean	Couchsachraga		
Ondawa	Glover		
Otisville	Halcott		
Palmyra	Hawksnest		
Pawling	Hogback		

Table 5. Classification scheme for soils in New York.—Continued

[Soil classifications are from the Soil Survey Geographic database (SSURGO; Natural Resources Conservation Service, 2019). in., inch]

SSURGO map unit name				
Hollis				
Holyoke				
Insula				
Irona				
Kearsarge				
Knob Lock				
Lordstown variant (Cortland Co.)				
Lyman				
Nassau				
Quetico				
Ricker				
Rock Land				
Rock Outcrop				
Skylight				
Taconic				
Topknot				
Tor				
Torull				
Tuller				
Woodstock				
Wotalf				

represent thin soils over carbonate bedrock, one represents thin soils over noncarbonate bedrock, and the fifth represents soils over well-drained glacial materials.

Infiltration rate is an important soil characteristic to consider when assessing the probability that a closed depression contributes to focused recharge. A soil infiltration rate dataset was derived from the gridded Soil Survey Geographic database for New York (Natural Resources Conservation Service, 2014). The dataset identifies areas in New York with soils that have high and moderate infiltration rates (table 6).

The 2011 National Land Cover Database (NLCD; Homer and others, 2015) was used to compute land cover percentages within a 250-m (820-ft) buffer around each closed depression. The NLCD 2011 data for New York is a 15-class land cover classification system. For this study, the NLCD 2011 data were reclassified to a five-class system (table 7). The land cover percentages were computed in ArcGIS using the Feature Statistics-To-Table tool available in the National Water-Quality Assessment Area-Characterization toolbox (Price and others, 2010)

Table 6. Classification scheme for soil infiltration rate in New York.

[gSSURGO, gridded Soil Survey Geographic database (Natural Resources Conservation Service, 2014); HydroGrp, hydrological group; >, greater than]

HydroGrp gSSURGO soil group		Description		
1	Group A	> 1.42 inches per hour		
2	Group B	0.57 to 1.42 inches per hour		

Table 7. Land cover classifications for New York.

[Classifications are from the National Land Cover Database for 2011 (Homer and others, 2015)]

Class code	Classification			
	Open water			
11	Open water			
	Developed			
21	Developed, open space			
22	Developed, low intensity			
23	Developed, medium intensity			
24	Developed, high intensity			
31	Barren land			
Forested				
41	Deciduous forest			
42	Evergreen forest			
43	Mixed forest			
52	Shrub/scrub			
	Agriculture			
71	Grassland/herbaceous			
81	Pasture/hay			
82	Cultivated crops			
	Wetland			
90	Woody wetlands			
95	Emergent herbaceous wetlands			

Results

A total of 5,023 closed depressions were inventoried for this study (table 8). The digital contour database contains 2,032 closed depressions (fig. 30), the digital raster graphic database contains 2,477 closed depressions (fig. 31), and the lidar database contains 514 large closed depressions (fig. 32). Several areas of clustered closed depressions are evident when the findings from the three databases are combined (fig. 33). Some of the clusters occur in areas underlain by noncarbonate bedrock.

Table 8. Closed depressions in New York, by data source.

[lidar, light detection and ranging]

Data source	Data type	Number of closed depres- sions
Digital contour database	Polygon	2,032
Digital raster graphic database	Point	2,477
Lidar database	Polygon	514

In terms of total number of closed depressions per county, Dutchess, Jefferson, and St. Lawrence rated highest with more than 300 in each county (fig. 34). In terms of concentration of closed depressions per unit study area for

each county, Monroe, Ontario, and Oswego are rated highest with greater than 0.2 depression per square kilometer (fig. 35). Comparing the locations of the closed depressions with the selected bedrock geologic units indicated that 51 percent of the features inventoried lie outside the extent of the selected units (fig. 36); 10 percent of the closed depressions lie within the Beekmantown Group, and only 1 percent, within the Wappinger Group. The number of depressions within the remaining limestone, dolostone, and marble units ranges from 4 to 9 percent.

From the overlay analysis between closed depressions and soil class (fig. 37) it was determined that most of the closed depressions lie outside the extent of the five soil classes; 35 percent of the depressions lie within well-drained soils on sand-and-gravel deposits, which might include alluvium, alluvial fan, glaciofluvial, and glaciodeltaic deposits. The number of closed depressions within the remaining four soil classes ranges from 1 to 8 percent.

Twenty-three percent of the closed depressions are on soils with high infiltration rates (fig. 38), and 6 percent are on soils with moderate infiltration rates. Most of the closed depressions lie outside the extent of the high- to moderate-infiltration-rate soils.

Overlaying closed depressions and land cover indicated that half of the closed depressions inventoried lie within the forested land cover class (fig. 39), and 28 percent are in agricultural areas. The wetland land cover class accounts for 12 percent of the depressions; 9 percent are in areas classified as developed, and 1 percent of the closed depressions lie within the open water land cover class (these are typically small ponds or marshes).

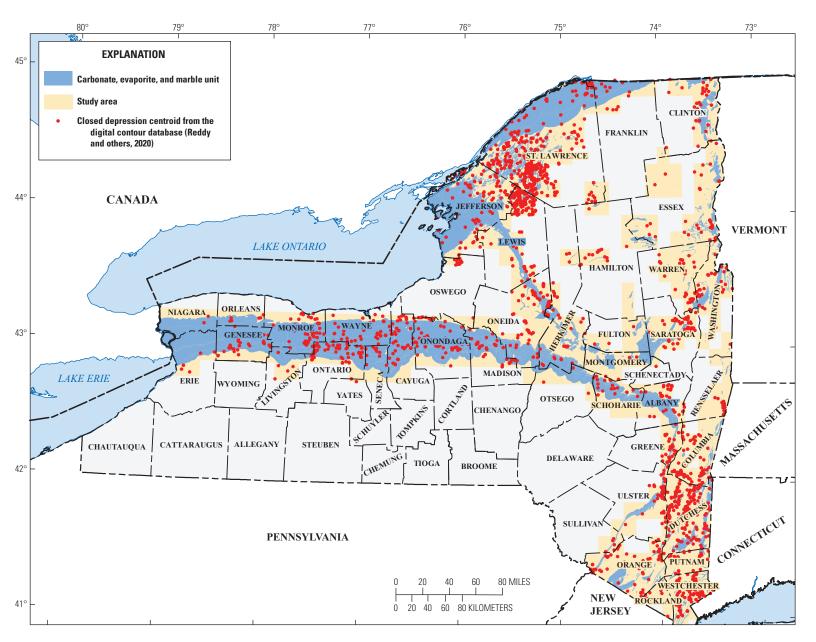


Figure 30. Distribution of closed depressions in New York in the digital contour database (Reddy and others, 2020).

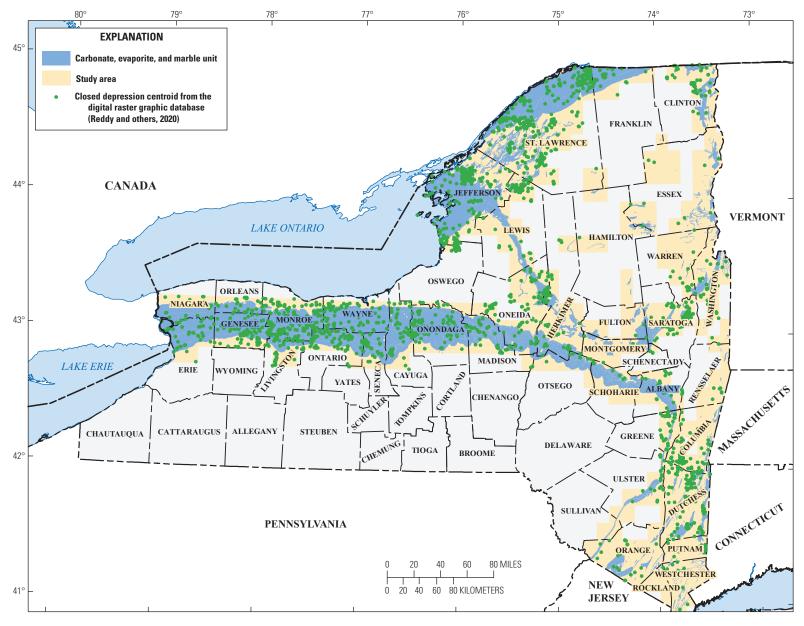


Figure 31. Distribution of closed depressions in New York in the digital raster graphic database (Reddy and others, 2020).

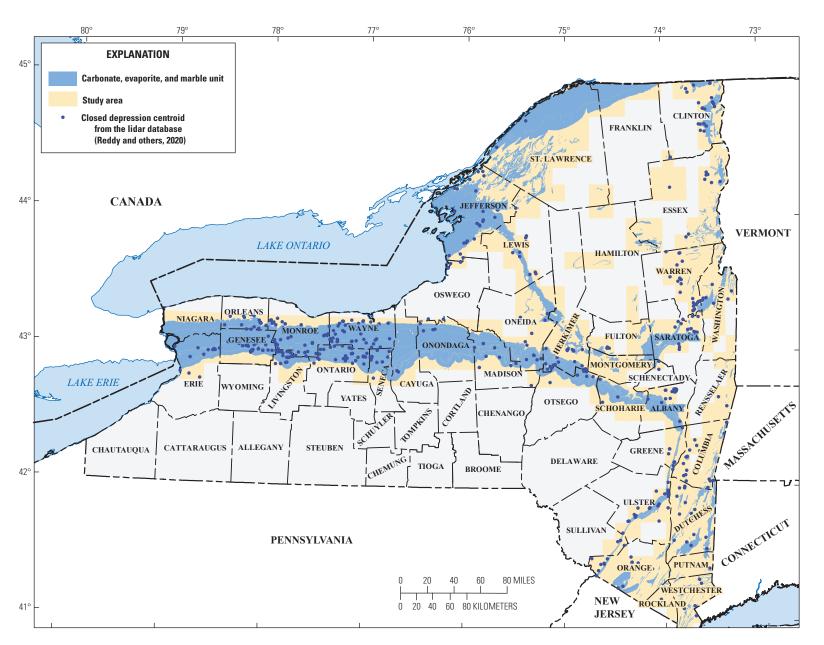


Figure 32. Distribution of closed depressions in New York in the light detection and ranging (lidar) database (Reddy and others, 2020).

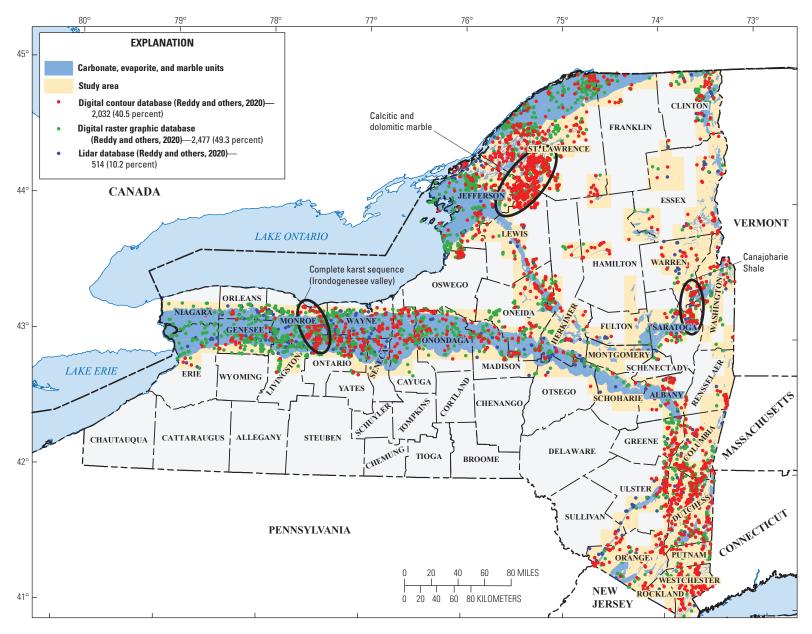


Figure 33. Distribution of all closed depressions in New York, from the digital contour, digital raster graphic, and light detection and ranging (lidar) databases (Reddy and others, 2020).

Figure 34. Concentration of closed depressions in New York, by source and quantity for each county. Dutchess, Jefferson, and St. Lawrence Counties (hachured) rate highest with more than 300 depressions in each county. >, greater than.

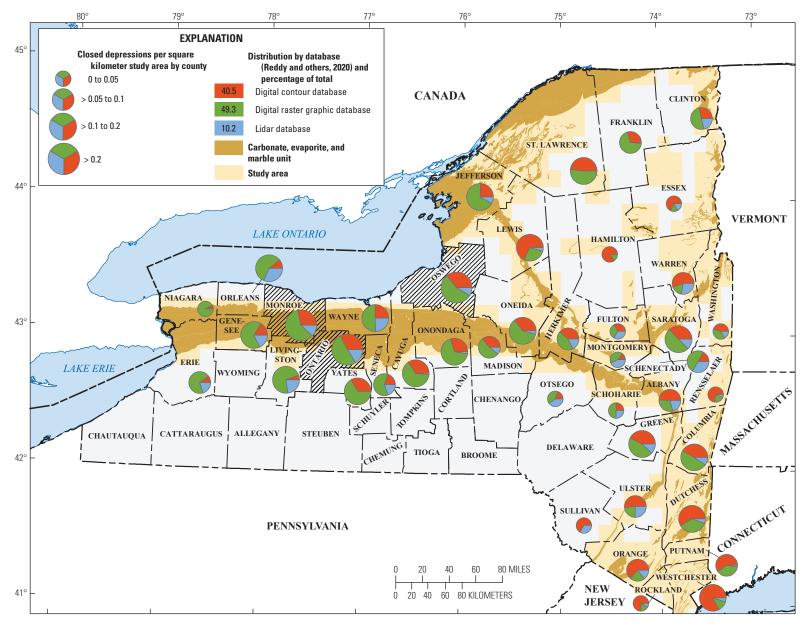


Figure 35. Concentration of closed depressions in New York, per square kilometer of study area for each county. The counties with the highest concentration of closed depressions (hachured) are Monroe, Ontario, and Oswego. >, greater than.

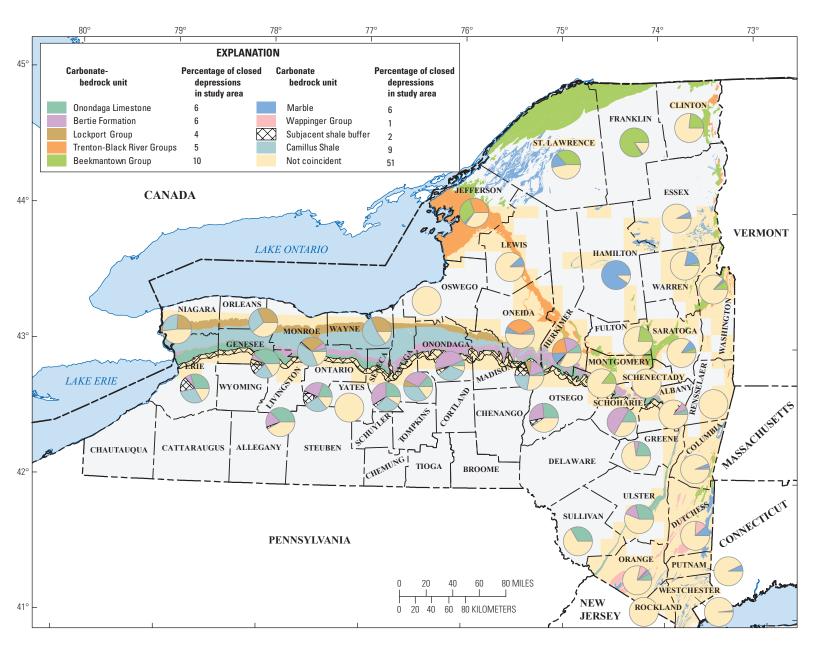


Figure 36. Bedrock geology and distribution of closed depressions in New York, by county.

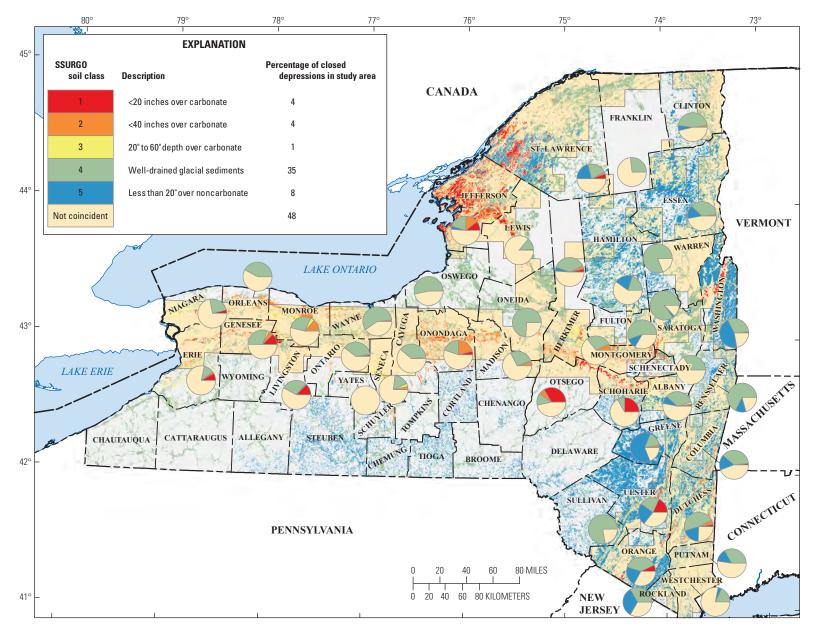


Figure 37. Distribution of soils and closed depressions in New York, by county. SSURGO, Soil Survey Geographic database (Natural Resources Conservation Service, 2019); <, less than.

Figure 38. Distribution of soil infiltration rate and closed depressions in New York, by county. SSURGO, Soil Survey Geographic database (Natural Resources Conservation Service, 2019); >, greater than.

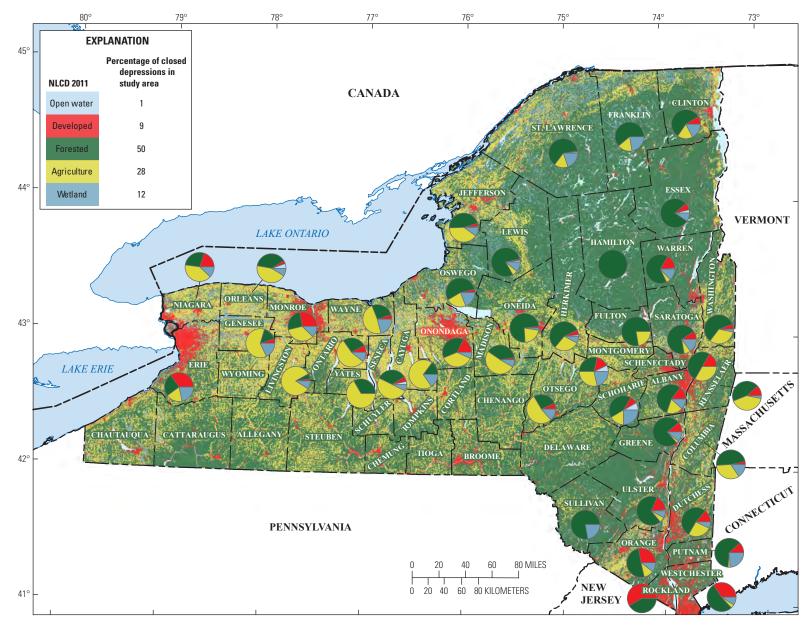


Figure 39. Distribution of land cover and closed depressions in New York, by county. NLCD 2011, National Land Cover Database for 2011 (Homer and others, 2015; U.S. Geological Survey, 2015).

Summary

Karst is a landscape formed from the dissolution of soluble rock or rock containing minerals that are easily dissolved from within the rock. The landscape is characterized by sinkholes, caves, losing streams, springs, and underground drainage systems which rapidly move water through the karst. The two forms of karst in New York State include carbonate karst, which forms in carbonate rock (limestone, marble, and dolostone), and evaporite karst, which forms in rock that contains evaporite minerals (gypsum and halite). Closed depressions that are created at the land surface by karstification of the underlying bedrock or through other natural or anthropogenic processes focus surface-water recharge to underlying bedrock.

Past and recent studies of karst bedrock across the State have shown that areas of focused recharge in karstic carbonate rock allows contaminants to enter aquifer systems with little attenuation. Focused areas of recharge need to be identified to help prevent such contamination from sources on or adjacent to karst. The New York State Department of Environmental Conservation has begun to work with the agricultural community to make farmers and farm-planning advisors more aware of karst and how to manage their daily activities to reduce their impact on surface-water and groundwater resources, especially in karst areas. Where evaporite karst is present, in addition to limited attenuation of any contaminant within a karst system, there is also a greater potential for land-surface subsidence due to the more-aggressive dissolution of evaporite minerals (gypsum or halite) if close to the land surface or as a result of mining of these evaporite minerals.

The U.S. Geological Survey has compiled an inventory of closed depressions from statewide digital contour data, scanned 7.5-minute topographic maps known as a digital raster graphics, and light detection and ranging (lidar) digital elevation models. The identification of where karst rock is present in New York and the location of closed depressions over or near karst rock was completed using publicly available geospatial data and Esri ArcGIS software. Carbonate, evaporite, and marble geologic units were selected from a statewide 1:250,000-scale bedrock geology dataset. The selected geologic units were intersected with an index of 7.5-minute quadrangle maps. Quadrangle maps that overlay the selected geologic units form the basis of the study area.

Analysis of the data resulted in the identification of 5,023 closed depressions. The results are publicly available in three databases, one for each source of elevation data. The inventory was conducted such that there is no duplication of results from analysis of the three data sources. The closed depression features in the digital contour database and the lidar database are represented as polygon features. Points are used to represent the closed depressions in the digital raster graphic database.

A series of overlay analyses were conducted using the closed depressions and thematic data known to be key factors in determining the probability of a closed depression contributing to focused groundwater recharge. The thematic data

layers used for the analyses were bedrock geology, soil type, soil infiltration rate, and land cover. Results from the overlay analyses were added as attributes to the data tables associated with the three publicly available databases.

Though the extent of karst development is important in understanding the surface water—groundwater interaction in karst terrains, some of the worst cases of groundwater contamination are in karst where only minor karst features might be present. The presence of karst, be it a short section of a solutioned fracture or an extensive cave system requires careful consideration, forward-looking environmental planning, and consistent water-quality protection to preserve New York State's water resources.

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Glossary of Selected Karst Terminology

This glossary was adapted from Monroe (1970) and Field (2002). The referenced glossaries are more extensive, as the terms shown here are related to just those used in this report. The etymologic origin for some of the terminology is shown in brackets.

Α

allogenic recharge Recharge from adjacent insoluble bedrock into a soluble karst aquifer.

aquifer A groundwater reservoir; pervious rock that is completely saturated and will yield water to a well or spring.

В

bare karst See naked karst

bedding plane A plane or surface that separates two strata of differing characteristics.

bicarbonate A salt containing the radical bicarbonate (HCO_3^{-1}) , such as calcium bicarbonate $[Ca(HCO_3)_2]$.

blowing cave A cave out of which or into which a current of air flows intermittently.

C

calcite A mineral composed of calcium carbonate (CaCO₃), such as aragonite, but differing in crystal form; the principal constituent of limestone.

carbonate A salt or ester of carbonic acid; a compound containing the radical carbonate (CO_3^{-2}), such as $CaCO_3$. Also, a rock consisting mainly of carbonate minerals, such as limestone or dolostone.

cave A natural underground room or series of rooms and passages large enough to be entered by a person; generally formed by solution of limestone. Also, a similar artificial opening.

cave system An underground network of passages, chambers, or other cavities. Also, the caves in a given area related to each other hydrologically, whether continuous or discontinuous from a single opening.

chert Light-cream or gray to black rock composed of silica (silicon dioxide $[SiO_2]$), found occurring as nodules or layers in limestone, or as a replacement of limestone.

clint [country of origin, United Kingdom] Slabs of limestone, parallel to the bedding, forming a fractured limestone pavement. Widened joints, or grikes, isolate individual limestone slabs.

closed depression A general term for any enclosed topographic basin having no external drainage, regardless of origin or size.

conduit A subterranean stream course filled completely with water and always under hydrostatic pressure.

D

doline A basin- or funnel-shaped hollow in limestone, ranging in diameter from a few meters to a kilometer, and in depth from a few to several hundred meters. Some dolines are gentle grassy hollows; others are rocky cliff-bounded basins. Closed depressions receiving a stream are known as swallow holes or stream sinks. In the United States most dolines are referred to as sinks or sinkholes.

dolomite A mineral composed of calcium magnesium carbonate $[CaMg(CO_3)_2]$.

dolostone Rock chiefly composed of the mineral dolomite.

dolomitization The process whereby limestone becomes dolomite by the substitution of magnesium carbonate for part of the original calcium carbonate.

Ε

epiphreatic zone The zone in a cave system, immediately above the phreatic (saturated) zone, affected morphologically and hydrologically by floods too large for the cave to absorb at once. Also known as epiphreas.

exsurgence The point at which an underground stream reaches the surface if the stream has no known surface headwaters.

F

fluorescein A reddish-yellow crystalline compound that imparts a brilliant green fluorescent color to water in very dilute solutions; used to label groundwater for identification of an emergence.

formation The fundamental unit in rock, a stratigraphic classification, consisting of a distinctive mappable body of rock.

G

grike [country of origin, United Kingdom] A vertical or subvertical cleft in a limestone pavement developed by solution along a joint or system of crisscrossing joints. Grikes separate clints from one another.

groundwater The part of the subsurface water that is in the zone of saturation. Used loosely by some to refer to any water beneath the surface. *See also* phreatic water.

gypsum A mineral composed of hydrous calcium sulfate $(CaSO_4 \cdot 2H_2O)$.

Н

hydraulic conductivity The ability of a rock unit to conduct water under specified conditions.

I

inflow cave Cave into which a stream flows or formerly entered.

influent cave See inflow cave

intake or recharge area The surface area in which water is absorbed into an aquifer eventually to reach the zone of saturation.

K

karren [country of origin, Germany] Channels or furrows caused by solution on massive bare limestone surfaces; they vary in depth from a few millimeters to more than 1 meter and are separated by ridges. In modern usage, the terms are general, describing the total complex of superficial solution forms (grikes and clints) found on compact pure limestone.

karst [originally, the German form of the Slavic word kras or krs, meaning a bleak waterless place; the German name for a district east of Trieste having such terrain] A terrain, generally underlain by limestone, in which the topography is chiefly formed by the dissolving of rock, and which is commonly characterized by karren, closed depressions, subterranean drainage, and caves.

karst plain A plain on which closed depressions, subterranean drainage, and other karst features may be developed. Also called a karst plateau.

karst topography Topography dominated by features of solutional origin.

karstic Occasionally used as the adjective form of karst.
karstification Action by water, mainly chemical but also mechanical, that produces features of a karst topography, including such surface features as dolines, karren, and mogotes and such subsurface features as caves and shafts.

L

limestone A sedimentary rock consisting chiefly of calcium carbonate.

limestone pavement A bare plane surface of limestone, parallel to the bedding, commonly divided into blocks (clints, flachkarren) by solutionally widened joints (grikes, kluftkarren), and pitted by solution pans.

M

marble Limestone recrystallized and hardened by heat and pressure. Commercially, any limestone that will take a high polish

marl Unconsolidated sedimentary rock consisting largely of calcium carbonate and clay; usage varies from calcareous clay to earthy limestone; in some parts of the United States, the term has been used for any unconsolidated sedimentary rock containing fossil shells.

N

naked karst Karst topography having much exposed bedrock.

P

paleokarst A karstified rock or area that has been buried by later sediments; in some places, ancient caves have been completely filled by the later sediments.

passage In a cave, the opening between rooms or chambers.

permeability The ability of a rock or soil to permit water or other fluids to pass through it.

phreatic water [from the Greek word $\varphi \rho \epsilon \alpha \rho$, meaning well] Water in the zone of saturation, below the water table.

phreatic zone The region below the water table, in which rock is saturated with water.

piezometric surface The imaginary surface to which water from a given aquifer will rise under its full static head. Also called potentiometric surface.

pipe Small cylindrical hole in unconsolidated sediments caused by removal of fine material by water.

piping Formation of a passage by water under pressure in the form of conduits through permeable materials, when the hydraulic head exceeds a certain critical value.

polje [Slavic word meaning field] In areas of karst topography, a very large closed depression, in some places several kilometers long and wide, having a flat floor either of bare limestone or covered by alluvium, and surrounded by generally steep walls of limestone.

ponor [Slavic] Hole in the bottom or side of a closed depression through which water passes to or from an underground channel. *Compare to* swallet, swallow hole.

porosity The ratio of the aggregate volume of interstices in a rock or soil to its total volume; generally stated as a percentage.

pseudokarst Karstlike terrain produced by a process other than the dissolving of rock, such as the rough surface above a lava field, where the ceilings of lava tubes have collapsed. Features of pseudokarst include lava tunnels, lava tubes, lava stalactites, and lava stalagmites.

R

recharge area See intake area

resurgence A point at which an underground stream reaches the surface and becomes a surface stream. In European literature, the term is reserved for the reemergence of a stream that has earlier sunk upstream; the term resurgence is applied to a stream without known surface headwaters.

rhodamine A family of related chemical compounds (fluorone dyes) used to label groundwater to determine the rate and direction of flow and transport. Rhodamine dyes fluoresce and can be detected with a fluorometer.

S

sink See sinkhole

sinkhole [country of origin, United States] General terms for closed depressions. May be basin, funnel, or cylindrical shaped. *See also* closed depression; doline; swallet; swallow hole.

solution pipe A vertical cylindrical hole attributable to solution, often without surface expression, filled with debris, such as sand, clay, rock chips, and bones.

solution scarp Escarpment formed by more active solution of lower area or by corrosional undercutting of the base of the escarpment.

solution subsidence Any subsidence caused by solution of underlying rock but particularly the subsidence of parts of a formation into hollows or pockets of an immediately underlying soluble formation.

speleogenesis The varied processes by which caves are formed.

spring Any natural discharge of water from rock or soil onto the surface of the land or into a body of surface water.

subjacent karst Karst landscape in noncarbonate rocks due to presence of karstified rocks beneath the surface formation.

subsidence The lowering of the surface of the ground because of removal of support. Caused in karst areas by subterranean solution or collapse of caves.

swallet [country of origin, United Kingdom] A place where water disappears underground in a limestone region. A swallow hole generally implies water loss in a closed depression or blind valley, whereas a swallet may refer to water loss into alluvium at a streambed, even though there is no depression.

swallow hole See swallet

T

travertine Calcium carbonate, light in color and generally concretionary and compact, deposited from solution in ground and surface waters. Extremely porous or cellular varieties are known as calcareous tufa, sinter, or spring deposit. Compact banded varieties, capable of taking a polish, are called onyx marble or cave onyx.

١

vadose water Water in the zone of aeration, above the zone of saturation (water table).

vug A small cavity in rock usually lined with crystals.

W

water table The upper boundary of an unconfined zone of saturation, along which the hydrostatic pressure is equal to the atmospheric pressure.

Ζ

zone of aeration The zone in permeable soil or rock that is above the zone saturated with water, the zone of vadose water.

zone of saturation The zone in permeable soil or rock that is saturated with water; the phreatic zone.

Appendix 1. Characteristics of Caves in New York

Table 1.1. Characteristics of caves in New York, by county.

County	Geology	Karst features present	Cave size	Cave fre- quency	Cave hydrology	Other solution- or karst-related features	Comments ¹
Albany	Helderberg, Onondaga limestones, and talus	Sinkholes, solu- tion channels, springs, caves, and caverns	Mostly small, some large (long) cav- erns present	Extensive	Wet and season- ally dry		Clarksville 7.5-minute USGS quad has nearly 170 identifiable karst features and this count does not include solutionally-enlarged joints. The overwhelming majority of the county's caves are small. This high number represents many small caves found in isolated knolls and valleys formed as the result of folding and faulting commonly found throughout much of the southern part of the county. In fact, in sheer numbers, most of the county's caves are found from Thacher State Park and south. This is that part of the county that shows increasing structural complexity the farther south one goes, entering what Marshak (1990) calls the Hudson Valley fold-thrust belt.
Allegany Bronx	No karst Inwood marble	Caves (very small) not verified, possible solution channels	Small	Uncommon	Unknown	_	
Broome	Talus	Caves possible but not verified	Small	Uncommon	Unknown	_	
Cattaraugus	Fractured-rock caves in conglomerates	Fracture caves, un- known solution features	Very small	Uncommon	Dry	_	Olean Rock City, Salamanca Rock City, Thunder Rocks, Bear Cave Ledge
Cayuga	Onondaga Limestone	Sinkholes, solu- tion channels, springs, a few caves (historic reports)	Small	Present	Unknown for caves, solu- tion channels with variable water levels.		Robert W. Carroll Jr. reports sinkholes in the Auburn area that during the mid-1960s defied efforts to fill them. [Study of contaminant plume flowing downdip in the Onondaga Limestone, from Auburn area to Union Springs (Cayuga Lake), New York (Eckhardt and others, 2011; U.S. Environmental Protection Agency, 2019)]*
Chautauqua	Fractured-rock caves in conglomerates	Fracture caves	Small	Present	Dry?	_	Panama Rocks There are stories of sub- merged limestone caves in Lake Erie, but nothing has ever been verified.

 Table 1.1.
 Characteristics of caves in New York, by county.—Continued

County	Geology	Karst features present	Cave size	Cave fre- quency	Cave hydrology	Other solution- or karst-related features	Comments ¹
Chemung	No karst	_	_	_	_	_	
Chenango	No karst	_		_	_	_	
Clinton	Talus, fracture caves, lime- stone of the Chazy Group, Trenton/Black River Group limestone, and Beekmantown Group	Talus caves generally found in gneiss	Small	Present	Mostly dry, Some wet caves adjacent to and in Lake Champlain		On Valcour Island there are at least two maze caves each exceeding 400 ft in length. The passages are typically tear-drop in shape. These caves are at lake level with 85 percent or more of the passage below water. Indeed, there is evidence for substantial karstification below water. Whether this represents ongoing processes or are remnants of lower lake levels is not clear at this time. There are also some fracture caves formed in the Potsdam Sandstone in the area of Ausable Chasm.
Columbia	Marble, possibly limestone	Marble caves, un- known solution features	Small	Present	Dry	_	No bottom pond drains to marble karst adjacent to pond.
Cortland	No karst		_	_	_	_	
Delaware	No karst	_	_	_	_	_	
Dutchess	Marble	Few caves, some rock shelters, un- known solution features	Small	Present	Unknown	_	
Erie	Onondaga Limestone, Camillus shale (gypsum)	Sinkholes, solu- tion channels, springs, few small caves	Small	Present	Dry caves, solu- tion channels saturated, flow to west	Former gypsum mines north of the Onondaga Limestone escarp- ment may have caused minimal subsidence issues as mine openings collapse.	Diversion of quarry discharge away from nearby sinkhole caused extensive dewatering of homeowner wells in 1980s along closed depression-solution channel over an approximate 4-mi length (Staubitz and Miller, 1987). Onondaga Limestone is sometimes used to 'drain wells' for surface runoff and shallow soil dewatering.*

Table 1.1. Characteristics of caves in New York, by county.—Continued

County	Geology	Karst features present	Cave size	Cave fre- quency	Cave hydrology	Other solution- or karst-related features	Comments ¹
Essex	Talus, marble, fracture caves, limestone?	Talus caves, some solution features	General small talus caves, although a few exceed 1,000 ft in length. Few marble caves, and few fracture caves	Extensive	Dry	_	The longest known solutional caves are Burroughs Cave (2,000+ ft) and Big Loop Cave (1,150 ft). Burroughs Cave is the longest known marble cave in New York. There are some short fracture caves (most are <20 ft long) in the county. Most notable is Pitchoff Mountain Cave formed in anorthosite. There are also a few fracture caves in Potsdam Sandstone in the east part of the county, near Lake Champlain.
Franklin	Talus, fracture caves	Unknown solution features	Small	Present	Dry	_	
Fulton	Talus, Limestone	Some sinkholes, so- lution channels, springs	Small	Uncommon	Dry	_	The Little Falls Dolomite and Ordovician-aged limestones crop out in the county, but no caves are known.
Genesee	Onondaga Limestone, Camillus shale (gypsum)	Sinkholes, solution channels, springs, a few small caves	Small	Sinkholes, solution channels, and springs present, caves un- common	Solution chan- nels intermit- tently wet	The Camillus Shale has a thin bed of gypsum (3 to 5 ft thick) which has been mined north of the Onondaga Escarpment. It has been reported that small solutional conduits drained into these mines and some land- surface subsidence (bulk or sinkhole subsidence fea- tures) has also oc- curred as the mine	One of the largest sinking streams in New York, Oatka Creek, is found there, and has a drainage area of 139 mi. On the east side of the County the Onondaga Limestone and other bedrock units have been displaced to the north by the Clarendon-Linden Fault and there are several karst springs along and near the fault. [Also see Reddy and Kappel (2010) for a study of karst in the Onondaga Limestone in Genesee County.]*

openings collapse.

 Table 1.1.
 Characteristics of caves in New York, by county.—Continued

County	Geology	Karst features present	Cave size	Cave fre- quency	Cave hydrology	Other solution- or karst-related features	Comments ¹
Greene	Helderberg and Onondaga Limestones, a few fissure caves in Catskill region in con- glomerates	Few sinkholes and solution chan- nels, some caves	Small	Present	Unknown	_	These limestones are all found in the Hudson Valley fold-thrust belt. The rocks are largely folded and faulted and are compressed in an east-west direction. As a result, many of the limestone caves are generally oriented along the strike of the major faults or down locally, in steeply dipping rock.
Hamilton	Marble, talus, and fracture caves. Limestone?	Caves, mostly short, in talus, although some may be solu- tional.	Small	Present	Unknown	_	
Herkimer	Trenton- Black River Limestone, marble, talus, fracture, and two tufa caves.	Caves, generally small, sinkholes, short solution channels	Small	Present	Unknown	_	Of interest is the chyle hole. This feature, a large occluded sink that takes a stream, has given its name to similar features in parts of Herkimer, Oneida, and Otsego Counties. Thus, locals often refer to sinks as chyle holes. The Van Hornesville Tufa Caves are one of two known groups of caves formed in calcareous tufa in the State.
Jefferson	Trenton-Black River Group Limestone, marble, talus fracture caves and two tufa caves.	Caves, solu- tion channels, sinkholes, losing/ gaining streams	Mostly small, a few larger (long) cav- erns present	Common	Intermittently to nearly always wet	_	All but a few of these caves are formed in the Ordovician-aged Trenton-Black River Group limestones. The exception is Natural Bridge Caverns formed in pre-Cambrian marbles (often called the Grenville Marble). At 2,000 ft, this is the longest known marble cave in New York.
Kings	No karst	— —	— C 11			_	
Lewis	Trenton-Black River Group limestone, marble	Solution channels, some caves	Small cases mostly	Present	Unknown	_	There is one known large cave—Drought Cave (2,000 ft)—close to the Oneida County line just north of Boonville. This low number may be because few people are looking. The Sugar River sinks into the its limestone channel.

Table 1.1. Characteristics of caves in New York, by county.—Continued

County	Geology	Karst features present	Cave size	Cave fre- quency	Cave hydrology	Other solution- or karst-related features	Comments ¹
Livingston	Onondaga Limestone	Solution chan- nels, sinkholes, springs, losing/ gaining streams	Unknown	Uncommon	Unknown	A few phantom lakes and ponds occur in the Onondaga Limestone where it crops out near land surface - variable solution channel capacities cause these intermittent water bodies to quickly appear and disappear.	Karst is found near Caledonia where one of the State's largest karst springs, Big Spring, is located. It has a low flow of 5 ft ³ /s. [A railroad tank car derailment in 1970 caused a large amount of TCE to flow into the Onondaga Limestone and flowed eastward, but it was never detected (at any level of concern) in Big Spring which feeds a fish hatchery (U.S. Environmental Protection Agency, 2017).]*
Madison	Helderberg and Onondaga Limestone	Solution chan- nels, springs, cascading water in wells, losing/ gaining streams	None known	Uncommon	Unknown	_	Some caves are reported near Chittenango/Chittenango Falls State Park. [Cascading water in drinking water wells has caused groundwater/ surface water interaction, causing drinking-water well, water-quality issues.]*
Monroe	Onondaga Limestone, Camillus shale (gypsum)	Solution chan- nels, sinkholes, springs	None known	Uncommon	Unknown	Gypsum mines north of the Onondaga Limestone escarp- ment have caused subsidence issues, such as in Monroe County's Oatka Creek Park.	
Montgomery	Dolostone, lime- stone	Fractured (faulted) rock	Small	Uncommon	Unknown	_	
Nassau	No karst	_	_			_	
New York (Manhattan)	Inwood marble	Very small caves, unknown karst features	Small	Uncommon	Unknown	The type locality of the Inwood Marble is Inwood Park at the north end of Manhattan Island. There are three small caves here and two in Central Park.	

 Table 1.1.
 Characteristics of caves in New York, by county.—Continued

County	Geology	Karst features present	Cave size	Cave fre- quency	Cave hydrology	Other solution- or karst-related features	Comments ¹
Niagara	Lockport Dolostone	Some sinkholes, solution chan- nels, and springs known, a few caves	Small	Uncommon	Dry?	The few karst features that are found in the upper weathered Lockport Dolostone are usually associated with solution channels: some vertical, some flowing along discrete bedding planes in the Lockport Dolostone. The large caves appear to be found near the Lockport Escarpment.	Lockport Dolostone unit has some large caves on either side of the Niagara River gorge and along its escarpment. Lockport Cave may have had as much as 2 mi of passage, but the cave has not been found in recent times. Of the approximately 40 caves, 7 are historically well documented, but have since been destroyed. Five other caves are rumored or lost.
Oneida	Trenton Limestone (north) Helderberg- Onondaga Limestone (south) and tufa	Sinkholes, solution channels and springs, but few caves	Small	Uncommon	Unknown	There is a calcareous tufa deposit northwest of Waterville at Forge Hollow that contains one of two known groups of New York caves formed in such material. The Unadilla River rises from a limestone spring. There are reports of caves upstream from Pixley Falls State Park, where the Lansing Kill apparently sinks into a fissure resurging 0.5 mi downstream.	

Table 1.1. Characteristics of caves in New York, by county.—Continued

County	Geology	Karst features present	Cave size	Cave fre- quency	Cave hydrology	Other solution- or karst-related features	Comments ¹
Onondaga	Onondaga and Manlius Limestone, Camillus Shale (gypsum)	Few sinkholes, minor caves, fracture- openings, evapo- rite subsidence	Very small	Uncommon	Unknown	One small cave under I-81 in south Syracuse documented by New York State Department of Transportation, and several fracture- opening caves adjacent to Rock Cut Road quarry escarpment.	See Robertson (1970) for information on caves in Onondaga County. [Subsidence of about 30 ft attributed to gypsum dissolution (Phillips, 1955). City infrastructure study also looked at water main breaks in relation to several anthropogenic factors and gypsum. Engineering studies have documented infrastructure subsidence in the Camillus region of the City.]* French (1860) suggests gypsum caves near Jamesville.
Ontario	Onondaga Limestone	Few sinkholes, solution channels noted	None known	Uncommon	Unknown	·	
Orange	Marble, Helderberg and Onondaga Limestone	Some solution features may be present	Small	Uncommon	Unknown		The best-known caves are probably Dutchess Quarry Caves 1 through 9. These are actually only four small caves, but all have great archeological significance.
Orleans	No karst	_	_	_	_	_	_
Oswego	No karst	_			_	_	
Otsego	Helderberg and Onondaga Limestone, and facture openings in shale	Sinkholes, solu- tion channels, springs, possible losing/gaining streams	Generally small	Present	Mixed		The Helderberg and Onondaga limestones are exposed in irregular bands along the northern edge of the county. There is much karst to be explored and one suspects that the lack of caves is from a lack of looking and nothing more. Summit Lake is fed by some large karst springs to the east which resurge at the base of the Onondaga limestone. Two fracture caves formed in shale near Cooperstown.
Putnam	Marble	Unknown solution features	Very small	Uncommon	Unknown	_	A scattering of marble slivers in the Taconic mélange are found throughout Putnam County though they are more common in the eastern half of the county. None of the 10 or so small caves are notable.
Queens	No karst	_	_	_	_	_	

 Table 1.1.
 Characteristics of caves in New York, by county.—Continued

County	Geology	Karst features present	Cave size	Cave fre- quency	Cave hydrology	Other solution- or karst-related features	Comments ¹
Rensselaer	Marble, tectonic fracture	Unknown solution features	Very small	Uncommon	Unknown	_	Snow Hole is a small tectonic cave formed in slate and phyllite. It generally contains snow and ice well into summer. Country Club Cave is a gravity-slip block cave formed in shale.
Richmond	No karst	_	_	_	_	_	
Rockland	Marble, fracture caves	Unknown solution features	Very small	Uncommon	Unknown	_	Rockland includes some marble outcrops and a few small caves. The best- known caves in the county, however, are small fissure caves formed in igne- ous and other metamorphic rocks.
Saint Lawrence	Talus, marble, Potsdam sand- stone (fracture) caves	Sinkholes, some so- lution channels, springs	Small	Common	Unknown	_	The predominant cave type in the county is talus, with about 70 known solutional caves, and there are fracture caves in the Potsdam Sandstone in the northern part of the county.
Saratoga	Beekmantown Group Limestone, dolostone	Few sinkholes, some solution features	Very small	Uncommon	Unknown		All known karst lie in the Ordovician- aged limestones that ring the Adirondacks. There is karst in Saratoga Springs and stories about caves in Congress Park and under the new Skidmore Campus. Two new cave discoveries have come to light between Saratoga and Glens Falls, and others west of Saratoga.
Schenectady	Dolostone, lime- stone	Fractured (faulted) rock	Small	Uncommon	Unknown	_	

Table 1.1. Characteristics of caves in New York, by county.—Continued

County	Geology	Karst features present	Cave size	Cave fre- quency	Cave hydrology	Other solution- or karst-related features	Comments ¹
Schoharie	Helderberg and Onondaga Limestone	Many sinkholes, solution channels, springs, caves and caverns, losing/gaining streams	Many lengths	Extensive	Mixed		There are three major cave areas in the county: the Cobleskill Plateau, Terrace Mountain, and Barton Hill. The Cobleskill Plateau is bounded on the south by the Cobleskill Creek, on the east by the Schoharie Creek, on the north by the Helderberg escarpment, and on the west by the younger Devonian formations, most notably the Esopus Shale. There is a fourth area that will likely yield large caves in the future, but now has only a few small caves. The area is bounded on the north and west by Kings Creek, on the south by the Fox Creek, and on the east by the Albany County line. In the northwest corner of the county are sinkholes thousands of feet across that have yet to be checked out.
Schuyler	No karst			_		_	•
Seneca	No karst	_	_	_	_	—	
Steuben	No karst	_	_	_	_	_	
Suffolk	No karst	_	_	_	_	_	
Sullivan	Limestone	Unknown solution features	One long cave	Uncommon	Unknown	_	Limestone is found only along the southeast edge of Sullivan County. In that relatively small, steeply dipping exposure there is one large cave: Surprise or Mystery Cave with 9,975 ft of passage.
Tioga	No karst	_	_	_	_	_	1 0
Tompkins	Tully Limestone/ Moscow shale	Solution channels at base of Tully Limestone	Small	Uncommon	Seasonally wet	_	The Tully Limestone crops out north of Ithaca. The caves, which are generally joint-controlled, are formed at the contact between the Tully limestone and the underlying Moscow shale with the majority of the passage in the shale. [Likely other Tully/Moscow caves are present where this contact is exposed along the edges of the Finger Lakes.]*

Table 1.1. Characteristics of caves in New York, by county.—Continued

County	Geology	Karst features present	Cave size	Cave fre- quency	Cave hydrology	Other solution- or karst-related features	Comments ¹
Ulster	Helderberg and Onondaga Limestone, talus, fracture caves	Unknown solution features	Mostly small	Present	Seasonally wet		The Helderberg and Onondaga limestones crop out as bands that run south to Kingston and then southwest until they leave the county near Ellenville. In the area from Saugerties through Kingston and down beyond Hurley, the beds are highly folded and faulted. To the southwest the severe deformation ceases, but the beds are still steeply dip to the northwest. The talus and fissure caves are formed in the Shawangunk Grit.
Warren	Marble, talus	Unknown solution features	Small	Present	Unknown	_	Warren County's caves are half solutional and half talus. As with the counties east of the Hudson, Warren County has isolated lenses of marble in which caves have formed. In the southeast corner of the county there is limestone, but no caves are reported.
Washington	Marble, talus, limestone(?)	Unknown solution features	Small	Uncommon	Unknown	_	
Wayne	No karst	_	_	_		_	
Westchester	Marble, talus	Unknown solution features	Small	Uncommon	Unknown	_	A scattering of marble slivers in the Taconic mélange are found throughout Westchester County though they are more common in the eastern half of the county.
Wyoming	No karst	_		_	_	_	
Yates	No karst			_	_		

¹Comments are from Thom Engle of the Northeastern Cave Conservancy and the National Speleological Society, except where indicated by an asterisk; comments with an asterisk are by William M. Kappel.

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